

# The Kazhdumi Formation (Lower Cretaceous, upper Aptian–upper Albian) in the Zagros Basin, Iran

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## ABSTRACT

In the Zagros Basin (southwestern Iran), a 270 m thick section of shales, marls, marly limestones and limestones of the Kazhdumi Formation has yielded a diverse ammonite fauna which allowed for a detailed biostratigraphic zonation. Aptian ammonites of the families Deshayesitidae and Parahoplitidae occur indicating a latest early to late Aptian age. Albian Douvilleiceratidae, Lyelliceratidae and Brancoceratidae indicate early, middle and late Albian age. Their range and generic and specific determinations are tabulated and illustrated.

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## 1. Introduction

The Zagros fold-thrust belt in southwestern Iran exhibits a Triassic to Cenozoic succession accumulated on a Paleozoic platform (James and Wynd, 1965; Kheradpir, 1975; Khalili, 1976; Setudehnia, 1978; Szabo and Kheradpir, 1978; Rahaghi, 1984; see Fig. 1). The succession contains a number of mega-sequences which define an elongate, asymmetrical, dynamically fluctuating foreland basin during the Mesozoic (Alavi, 2004). The Lower Cretaceous deposits in the area studied are included in the Khami and Bangestan Groups (Fig. 1). The Bangestan group (Motiei, 1993) includes the Lower Cretaceous Kazhdumi Formation and the overlying Sarvak Formation (Fig. 2), the faunas of which indicate a late Aptian to late middle Albian (Early Cretaceous) for the

Kazhdumi Formation and a late Albian to early Turonian age for the Sarvak Formation (Fig. 2).

Hydrocarbon sources are important in the Zagros belt. Both the 'Neocomian' Garau and the Aptian–Albian Kazhdumi formations are well known reservoir and source rocks (Bordenave and Hegre, 2005; Bordenave and Hegre, 2010) in black shale intervals deposited in deep marine environments (James and Wynd, 1965; Ghazban, 2007; Heydari, 2008). Black shale facies are widespread in the Tethys and Atlantic-Pacific oceanic areas at that time (Arthur et al., 1990; Leckie et al., 2002; Baudin, 2005). The Kazhdumi Formation is the major source for hydrocarbons of the Asmari and Sarvak reservoirs (Bordenave and Burwood, 1990, 1995).

The present work contributes greater precision to the biostratigraphy and ammonite zonation of the Aptian and lower Albian of the Zagros belt. Collignon (1981) studied ammonite fauna of Kazhdumi Formation in Fars and Khuzestan. Based on calcareous nannofossils, Nowrouzi (2005) attributed an Aptian–early Albian age to the Kazhdumi Formation. Kennedy et al. (2009) discussed species discrimination in the Early Cretaceous (Albian) ammonite genus *Knemiceras* (Buch, 1848/1850). Bulot (2010) carried out a systematic study of Aptian and Albian ammonites from

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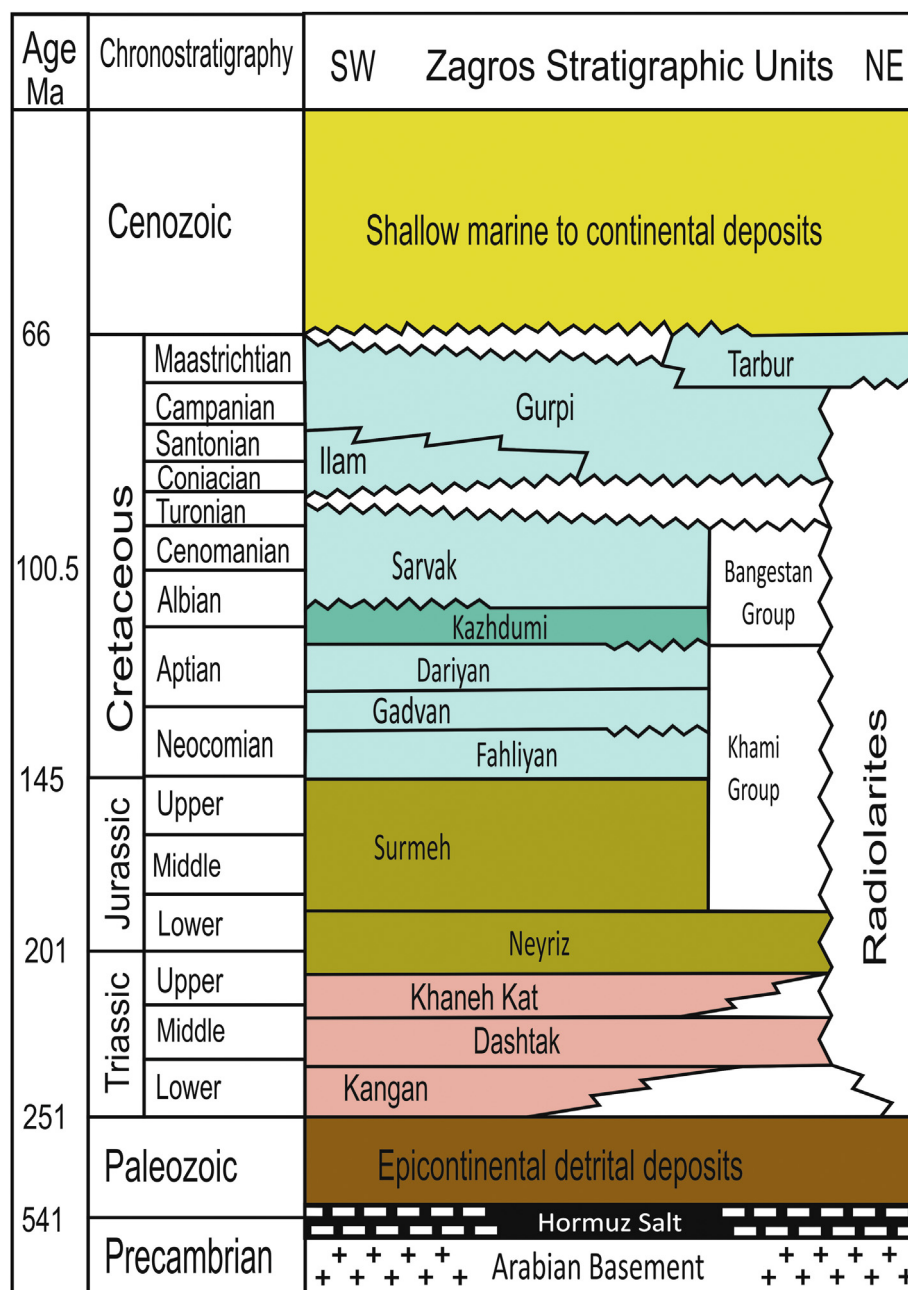


Fig. 1. Simplified table of rock units in southwestern Iran, based on numerous references.

southwestern Iran, including the Kazhdumi Formation. [Nazeri Tahroodi et al. \(2013\)](#) and [Asadi et al. \(2016\)](#) also reported Albian Kazhdumi ammonites in the Bushehr area (near to Persian Gulf). [Abdollahi et al. \(2014\)](#) worked on a foraminifera assemblage in the Behbahan area and suggested a late Aptian–middle Albian age for the Kazhdumi Formation there. [Sharifi et al. \(2019\)](#) had a preliminary discussion on the Aptian–Albian boundary in the Kazhdumi Formation.

In the present study, ammonites were collected bed-by-bed during field work in 2013 in the section of the Kazhdumi Formation at Tange Maghar, situated on the north east flank of the Bangestan anticline ([Fig. 2](#)). The section is illustrated photographically in [Figs. 4–6](#). The ammonite occurrences are listed and illustrated ([Figs. 8–12](#)). The aim of this study is to discuss the lower Aptian–upper Albian biostratigraphy of the studied section. The

ammonite biostratigraphy of the Zagros basin is reviewed, updated and compared with the standard Mediterranean ammonite zonation ([Reboulet et al., 2014, 2018](#)) and other biostratigraphic schemes in southern Tethys. The studied and illustrated specimens are kept at the Department of Geology, University of Birjand, Iran, under label KTM (Kazhdumi Tange Maghar).

## 2. The Kazhdumi Formation in the Tange Maghar section

The name Kazhdumi is taken from the Kazhdumi Castel in the Dezful embayment. The type section is located in the southwestern flank of Kuh Mish, 7 km away from northeast Gachsaran (Dugon-badan) ([Fig. 3](#)).

The Kazhdumi Formation in the type section is a 210 m thick series of calcareous shale, shaly limestone, clay and sandy siltstone

## Geological map of Study area

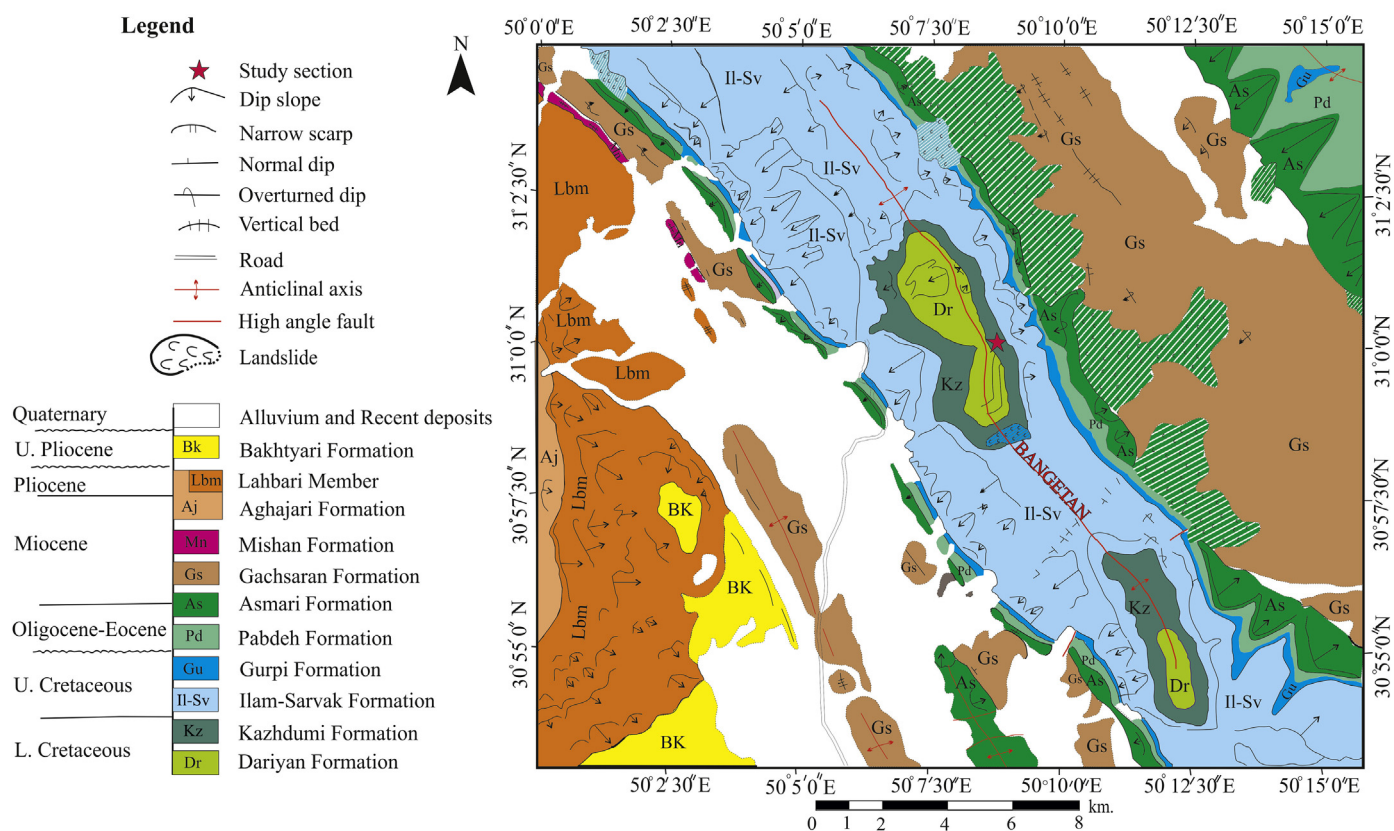


Fig. 2. Geological maps of part of the Bangestan Anticline (after Macleod, 1970 and Macleod and Akbari, 1970).

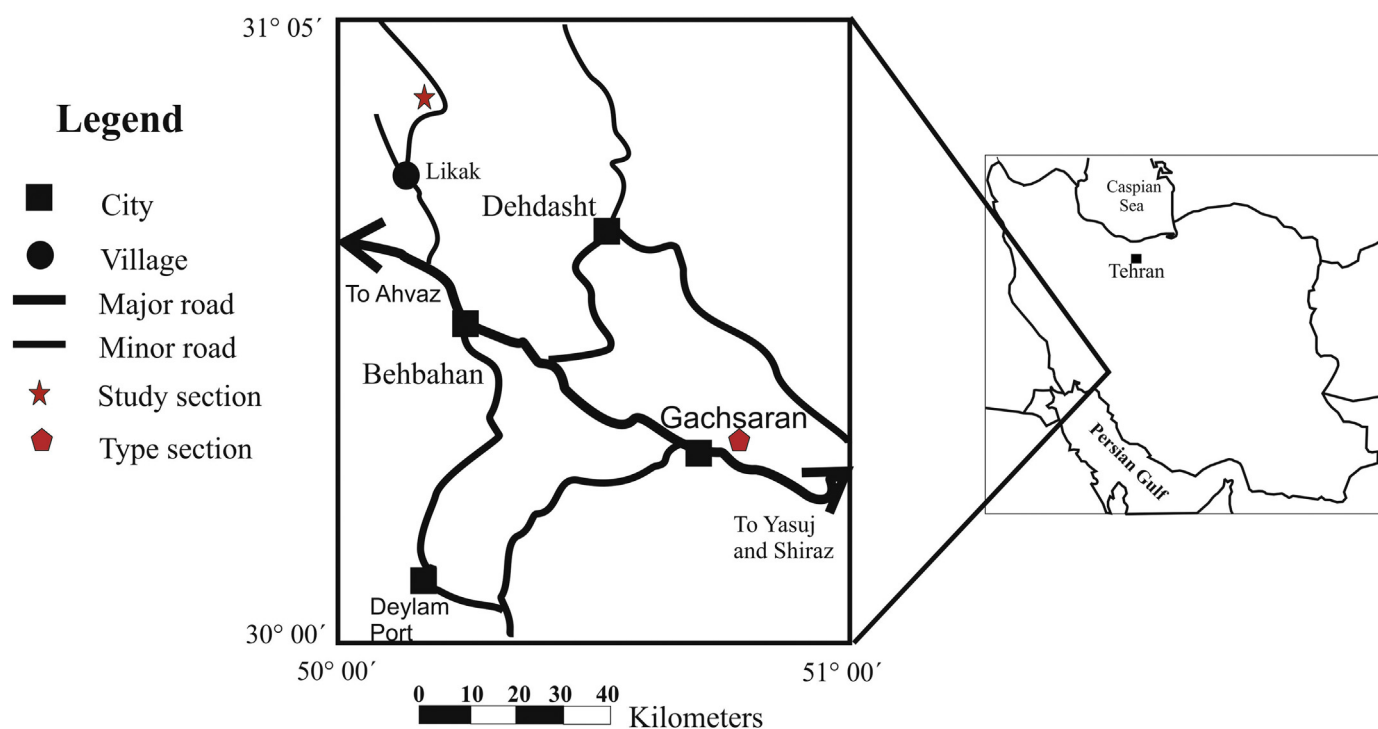


Fig. 3. Geographical position of the study and type sections of Kazhdumi Formation.



which can be divided into three units. Some limestone beds are dark brown to green in colour and contain bitumen and pyrite. A sandstone horizon occurs in the middle part of the formation. The lower contact with the Dariyan Formation, is a red bed and the upper boundary with the Sarvak Formation is a conformable and gradual change to marly limestone and limestone beds (James and Wynd, 1965). The formation includes bitumen-bearing black shales and dark clay limestone in the Dezful Embayment and northeastern Fars. It inter-fingers with the Burgan Formation and Omar Nahar in Kuwait and southwestern Iraq. In other areas of the Persian Gulf, the formation exhibits shallow-water sedimentary conditions with laterite and iron oxides zones and sandstone and silt layers. The Kazhdumi Formation laterally grades into the Garu Formation in central and southwestern Lorestan (Motiei, 1993).

The studied section is located in Tange Maghar, at 50°8'5"N and 31°01'28"E (Fig. 3). Exposed strata and their position are indicated on the geological map of the Zagros Basin (Fig. 2).

### 2.1. Unit 1

Toward the end of the Aptian, a major regression saw the termination of thick limestone bed deposition of the Dariyan Formation, followed by a low amplitude transgression marked by a sudden influx of clastics of the Kazhdumi Formation. The basal unit of the Kazhdumi Formation is composed of grey to light grey thin bedded shales with interbeds of limestone and marl (Figs. 4 and 5). Beds are brown grey to cream with iron oxide nodules containing brachiopods and nautiloids. The ammonite fauna indicates a range from the latest early Aptian with *Chelonicer* sp. and *Dufrenoyia* aff. *furcata* (Sowerby, 1836); a late Aptian age with *Epicheloniceras* sp., *Parahoplites* sp. and *Hypacanthoplites* sp.; an early Albian age with *Douvilleicer* aff. *mammillatum* (Quenstedt, 1849), *D. mammillatum* aff. *aequinodum* (Quenstedt, 1849); and latest early Albian age with

*Lyelliceras* aff. *pseudolyelli* (d'Orbigny, 1841). The biostratigraphic distribution of the ammonites is shown in Fig. 7, and the specimens are illustrated in Figs. 8 and 9.

### 2.2. Unit 2

The unit is composed of dark grey thin bedded to laminated shales with few intercalations of nodular marly limestone. It has yielded the following ammonites; *Oxytropidoceras* (*Mirapelia*) aff. *buarquianum* (White, 1887), *O. (Mirapelia)* aff. *mirapelianum* (d'Orbigny, 1850) and *O. (Mirapelia)* cf. *mirapelianum* (d'Orbigny, 1850) (Figs. 7, 9 and 10), indicating an early-middle Albian age.

### 2.3. Unit 3

The unit consists of dark to grey thin bedded shales, with thin to thick bedded limestones (Fig. 5). It is overlain conformably by the Sarvak Formation and shows a lithology which changes gradually to marly limestone and limestone beds (Fig. 6). The following ammonites were found in this unit; *Hystero* sp., *Oxytropidoceras* sp., *Oxytropidoceras* (*Venezolicer*) aff. *bituberculatum* (Collignon, 1966), *O. (Venezolicer)* cf. *karsteni* (Stieler, 1920) and *O. (Venezolicer)* sp. (Figs. 7, 11 and 12) indicating a late middle to early late Albian age

## 3. Biozonation and biostratigraphy

There are two principle ammonite faunal provinces in the Early Cretaceous in respect of Europe and the Middle East. Most of Europe comes within a well-defined province characterized by hoplitid ammonites. Southern Europe and the Middle East come within a southern Tethyan faunal province characterized by a more diverse ammonite fauna. Separate ammonite zonal schemes have

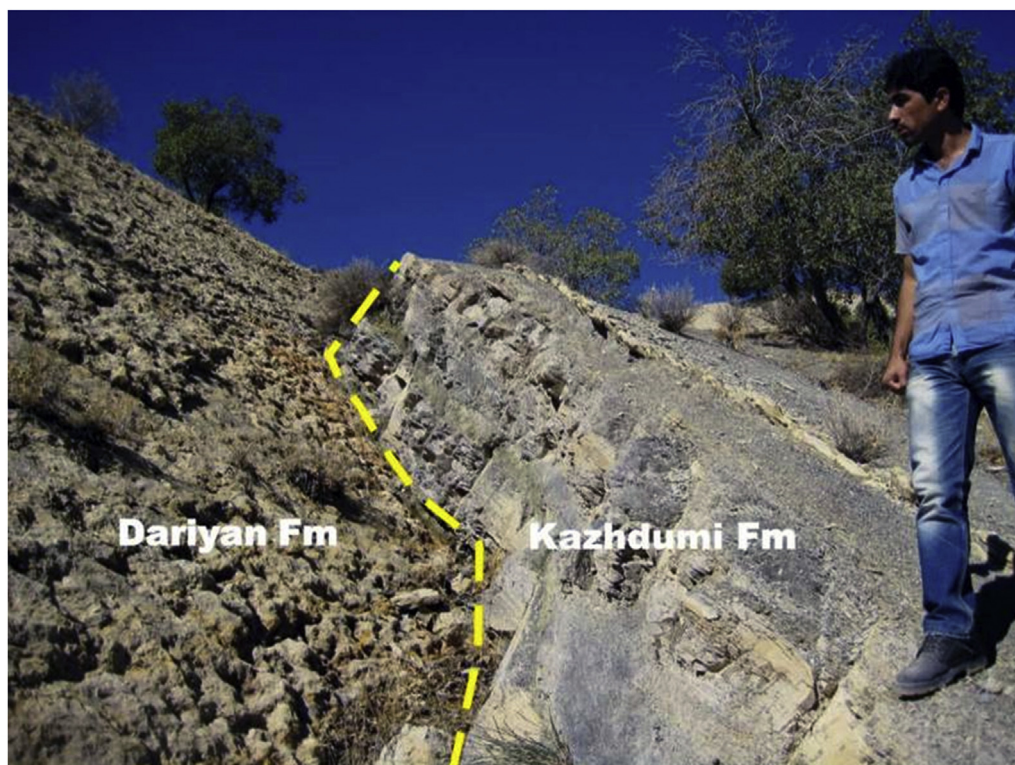


Fig. 4. Lower boundary of the Kazhdumi Formation overlying the Dariyan Formation.





Fig. 5. Upper part of the Kazhdumi Formation (Unit 3).

been developed in these two provinces and migration of Tethyan elements northward at intervals permit some correlation between the two. The Lower Cretaceous of Iran falls within the Tethyan faunal province. We use here the ammonite zonal scheme for the Aptian and Albian of the Tethyan Ammonite Faunal Province given in [Reboulet et al. \(2018\)](#) (Table 1). Ammonite zones positively identified in the Kazhdumi Formation and the ranges of the collected ammonites are tabulated in [Fig. 7](#).

The Cretaceous stages, their boundaries and index faunas in Europe have been discussed by [Birkelund et al. \(1984\)](#) and [Hancock \(1991\)](#). More recently the Barremian, Aptian and Albian stages were discussed by [Rawson \(1996\)](#), [Erba \(1996\)](#), [Hart et al. \(1996\)](#) and [Moullade et al. \(2011\)](#), respectively.

The Aptian and Albian ammonites have been studied in the past and recent years. Aptian ammonite faunas have been extensively studied in England ([Casey, 1960, 1961, 1980](#)), Germany ([Brinkmann, 1937; Kemper, 1967, 1971, 1973, 1982, 1995](#)), France ([Breistroffer,](#)

[1947](#)), Romania ([Avram, 1999](#)), Russia ([Bogdanova, 1983, 1991](#)), Central and North East of Iran ([Seyed-Emami, 1977; Seyed-Emami and Immel, 1995 and 1996; Raisossadat, 2004 and 2006; Mosavinia et al., 2007 and 2014](#)), Caucasus ([Enson, 2009](#)), Spain ([Moreno-Bedmar et al., 2010 and 2017; Martín-Martín et al., 2013](#)), Mexico ([Moreno-Bedmar et al., 2013, 2015; Barragán et al., 2016](#)), Argentina ([Medina and Riccardi, 2005](#)), Chile ([Pérez et al., 1999](#)), Peru and Colombia ([Robert et al., 2002, 2009](#)), Egypt ([Aly and Abdel Gawad, 2001; Aly, 2006](#)), Morocco ([Luber et al., 2019; Giraud et al., 2021](#)), and Tunisia ([Latil, 2011; Ben Chaabane et al., 2019; Latil et al., 2021](#)).

[Casey \(1961\)](#) introduced a biozonation for the Aptian stage that has been quoted in numerous papers in the last few decades. The study was based on the sequences in southeast England, but is applicable much more widely with some local modifications (e.g. [Kemper, 1976](#)). Faunal differentiation between the English and German faunas on the one hand, and those of the southeast of





Fig. 6. Boundary between the Kazhdumi and overlying Sarvak Formation.

France on the other, is much less marked than earlier in the Cretaceous (Rawson, 1981; Hoedemaeker, 1990). Hence the current zonation for the Mediterranean region (Hoedemaeker and Rawson, 2000) uses many of the same genera, and even some of the same species.

Ammonite faunas of the Albian were studied in France (Breistroffer, 1947; Latil, 1994; Kennedy and Latil, 2007; Rey et al., 2013; and Matrimon, 2010), Spain (Lopez-Horgue et al., 2009), Germany (Brinkmann, 1937; Kemper, 1982), and in England (Spath, 1923–1943; Casey, 1965). Owen (1971, 1973, 1988, 1996, 1999) focused on the Albian biostratigraphy and palaeobiogeography of the European area. Albian ammonite faunas and their biostratigraphy were described from the Carpathians (Vasicek, 1995), Crimea (Marcinowski and Naidin, 1976; Marcinowski and Wiedmann, 1985), Hungary (Szives, 2007) and former Soviet Union (Druschitz and Kudryutzeva, 1960; Saveliev, 1973, 1992; Mikhailova and Saveliev, 1989; Baraboshkin, 1996), Iran (Seyed-Emami, 1980 and 1995), Venezuela (Renz, 1982) and Madagascar (Collignon, 1962 and 1966).

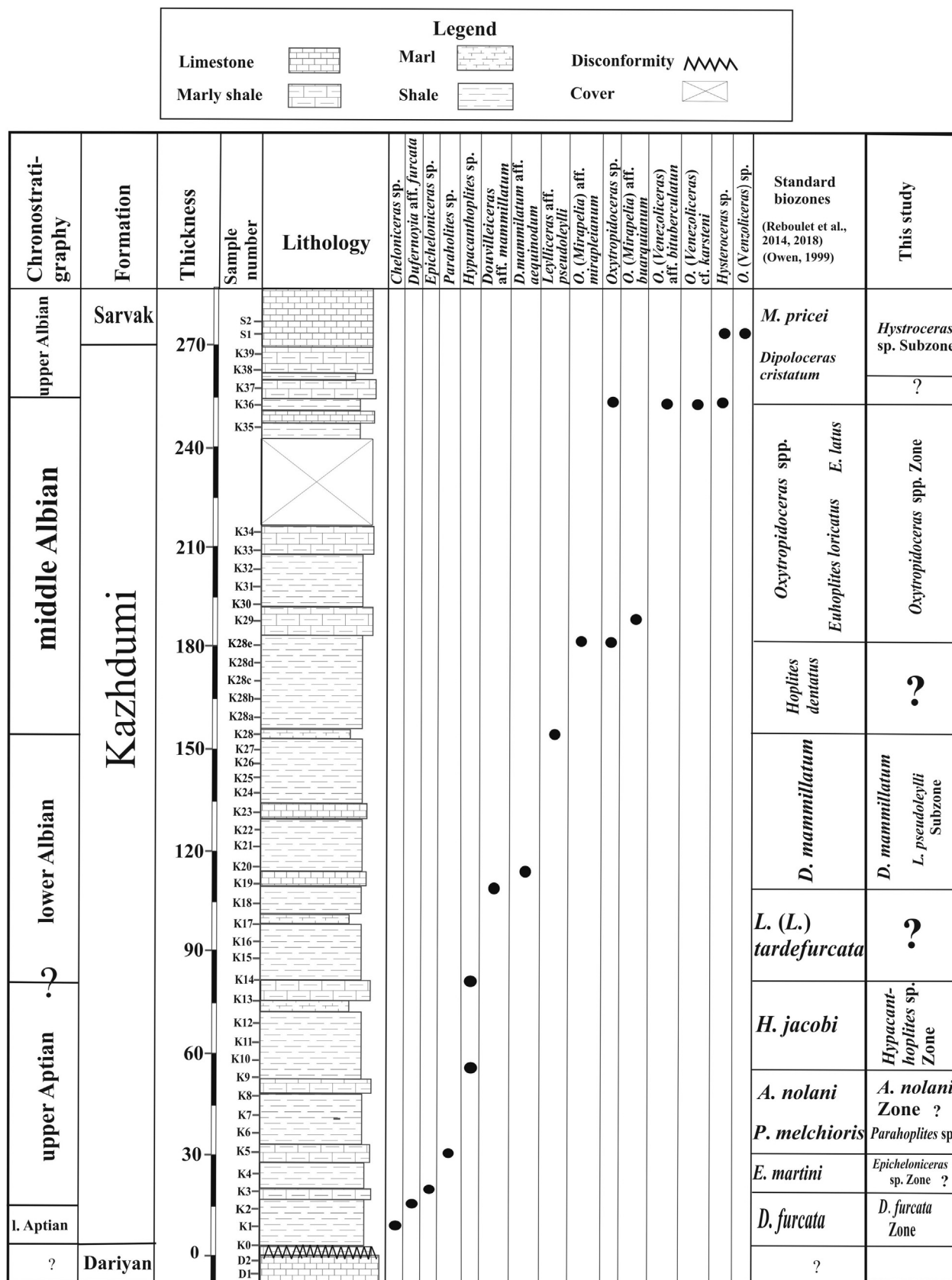
More recent work could be cited, for example in Cuba (Barragán et al., 2011), South Africa (Kennedy and Klinger, 2008), Pakistan (Kennedy and Fatmi, 2014), Ecuador (Bulot et al., 2005), Tunisia (Latil, 2005; Chihaoui et al., 2010), Algeria (Latil, 2011), Japan (Obata and Futakami, 1992 and, 2009; Kawabe et al., 2003; Takashima et al., 2007).

In a discussion about the Aptian–Albian boundary, Casey (1999) suggested that the current use of the first occurrence of *Leymeriella* to mark the base of the Albian is unsatisfactory because of its limited geographical distribution. He proposed that the first occurrence of *Hypacanthoplites* should be used instead, because of the world-wide distribution of the *Acanthoplites*-

*Hypacanthoplites* lineage, which would put the base of the Albian at the base of the *Hypacanthoplites jacobi* Zone. However, he mentioned that early *Hypacanthoplites* can only be distinguished easily from *Acanthoplites* in microconchs and juvenile macroconchs. This is complicated by the fact that *Acanthoplites* and *Hypacanthoplites* species are very often recorded together within the same levels in the upper Aptian (e.g. Druschitz and Kudryutzeva, 1960; Collignon, 1962; Casey, 1965; Mandov and Nikolov, 1992). Hence Casey's proposal raises new difficulties in recognizing the boundary.

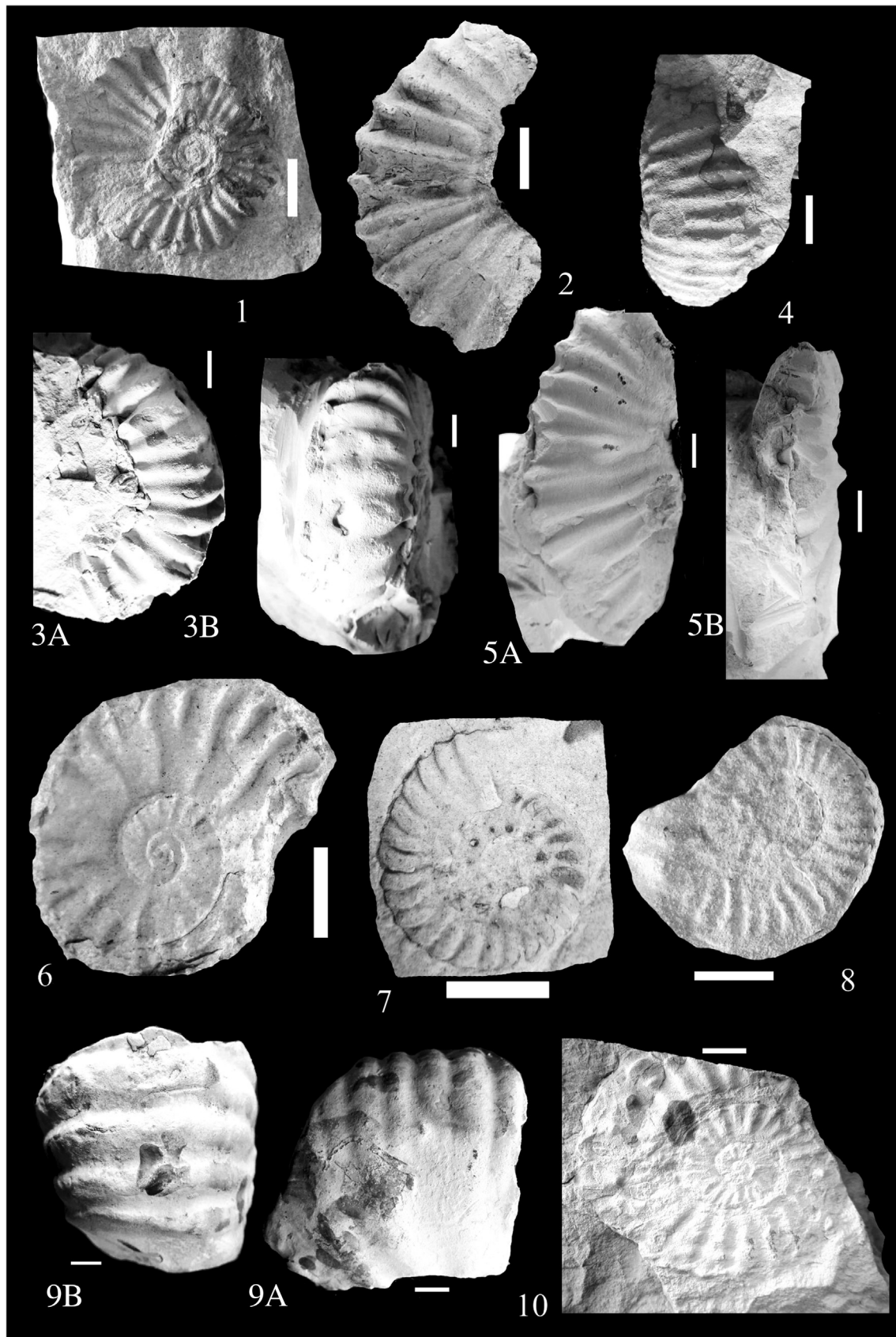
Kennedy et al. (2000) also discussed the Aptian–Albian boundary and suggested two sections in France as potential global boundary stratotype sections. In a first attempt, they put the *Hypacanthoplites jacobi*, *Proleymeriella schrammeni* and *Leymeriella germanica* zones in the upper Aptian and defined the base of the Albian stage at the first appearance of *Leymeriella tardefurcata*. After Kennedy et al. (2014, 2017), the GSSP for the base of the Albian is defined in the col de Pré-Guittard section (SE France), with the first occurrence of the planktonic foraminifera *Microhedbergella renilaevs*. This makes it very difficult to recognize the Aptian–Albian boundary on the basis of the ammonites.

Workshops of the Kilian Group (formerly the Lower Cretaceous Cephalopod Team, a working group of the Subcommittee on Cretaceous Stratigraphy (SCS) of IUGS) focused on ammonite biozonation of the Lower Cretaceous in the Mediterranean Province (Hoedemaeker and Bulot, 1990; Hoedemaeker et al., 1993, 1995; Rawson et al., 1999; Hoedemaeker and Rawson, 2000; Hoedemaeker et al., 2003; Reboulet et al., 2006, 2009, 2011, 2014, 2018) (Table 1).

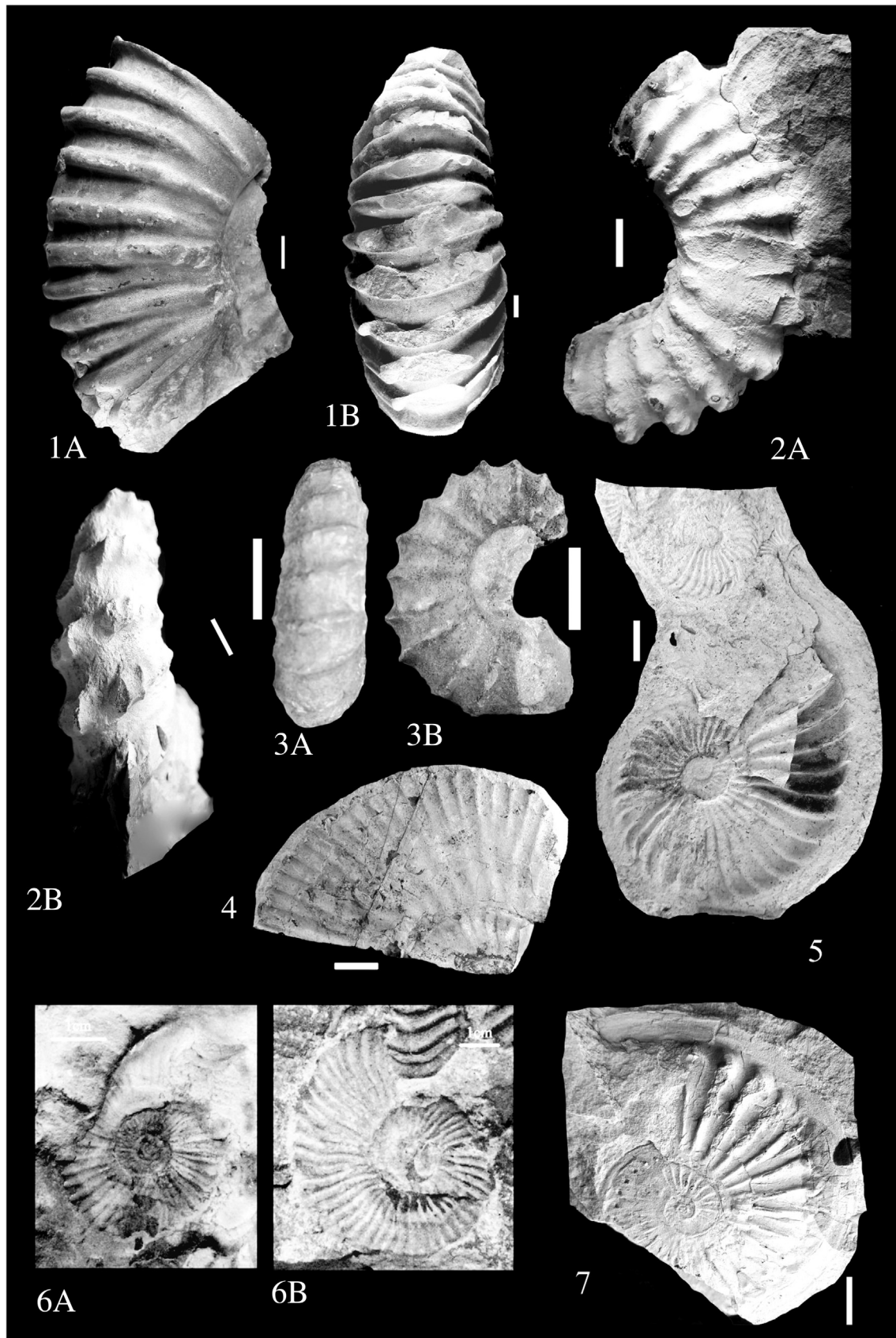


**Fig. 7.** Stratigraphical column, ammonite distribution and proposed biozones of the Kazhdumi Formation in Tange Maghar, Zagros basin.



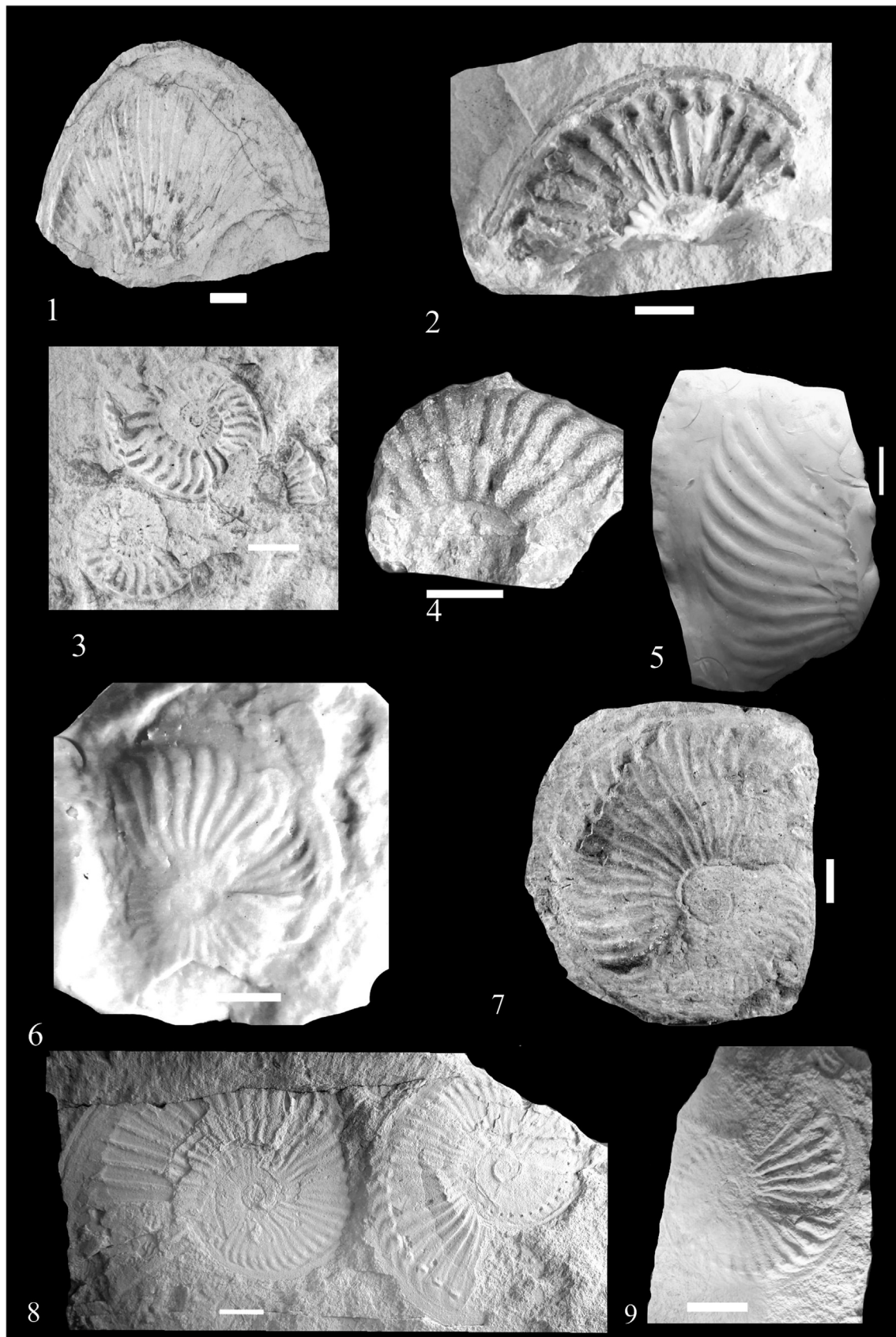


**Fig. 8.** 1, *Chelonicer* sp.?, k1a; 2, *Chelonicer* sp., k1b; 3A, B, *Chelonicer* sp., k1c; 4, *Chelonicer* sp., k1d; 5A, B, *Dufrenoyia* aff. *furcata* (Sowerby, 1836), k2a; 6, *D.* aff. *furcata* (Sowerby, 1836), k2b; 7, *Hystero*ceras sp., k37a; 8, *Hypacanthoplites* sp., k14a; 9A, B, *Hypacanthoplites* sp., k14b; 10, *Lyelliceras* aff. *pseudolyelli* (d'Orbigny, 1841), k28-2; all scale bars: 1 cm.

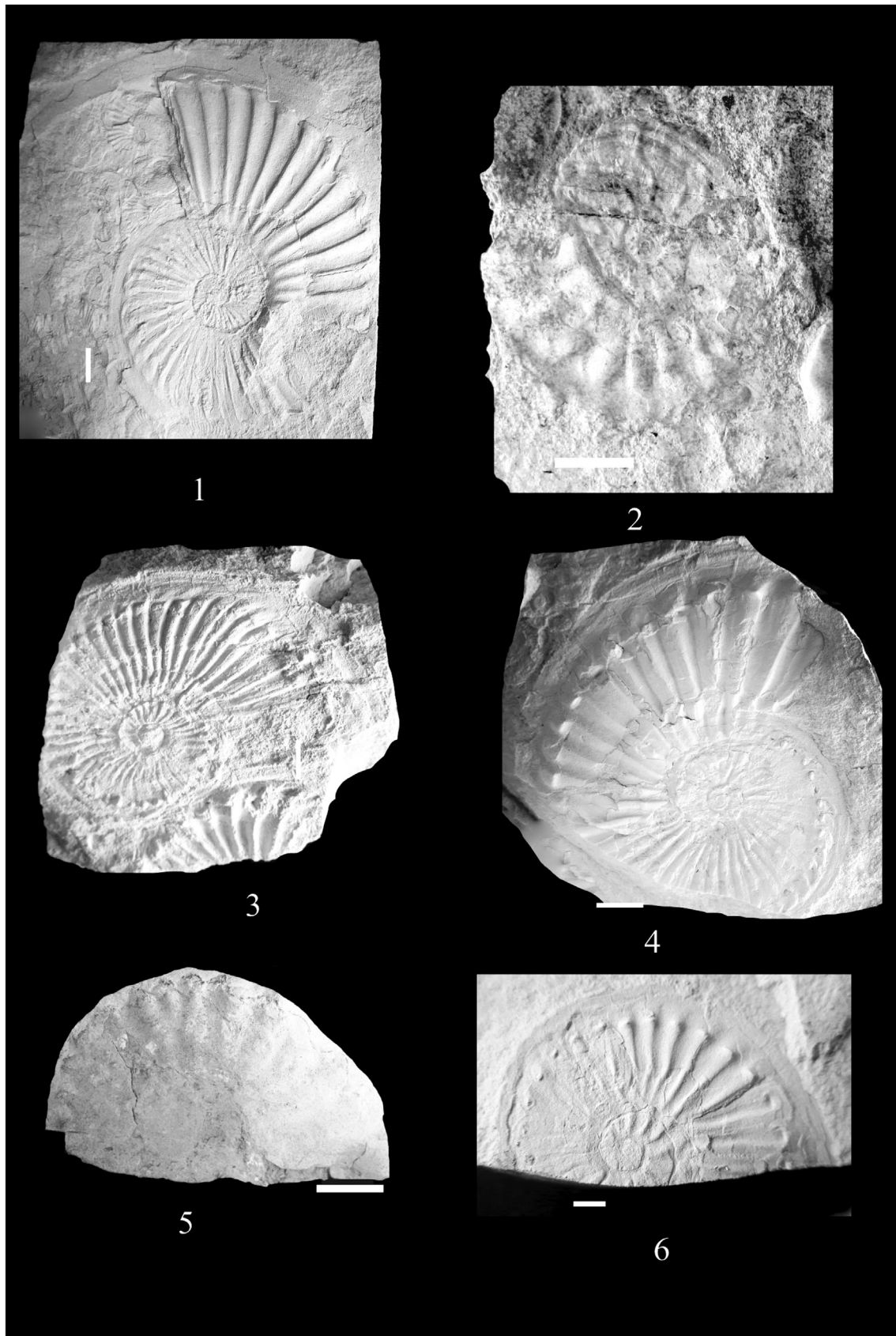


**Fig. 9.** 1A, B, *Hypacanthoplites* sp., k10a; 2A, B, *Douvilleiceras* aff. *mammillatum* (Schlotheim, 1813), k19a; 3A, B, *D. mammillatum* aff. *aequinodum* (Quenstedt, 1849), k20a; 4, *Parahoplites* sp., k5a; 5, *Oxytropidoceras* (*Mirapelina*) aff. *mirapelianum* (d'Orbigny, 1850), k28e-4; 6A, B, *Nolaniceras* aff. *nolani* (Seunes, 1887) (Kamyabi Shadan, 2014); 7, *Oxytropidoceras* (*Venezoliceras*) cf. *bituberculatum* (Collignon, 1966), k36c, all scale bars: 1 cm.



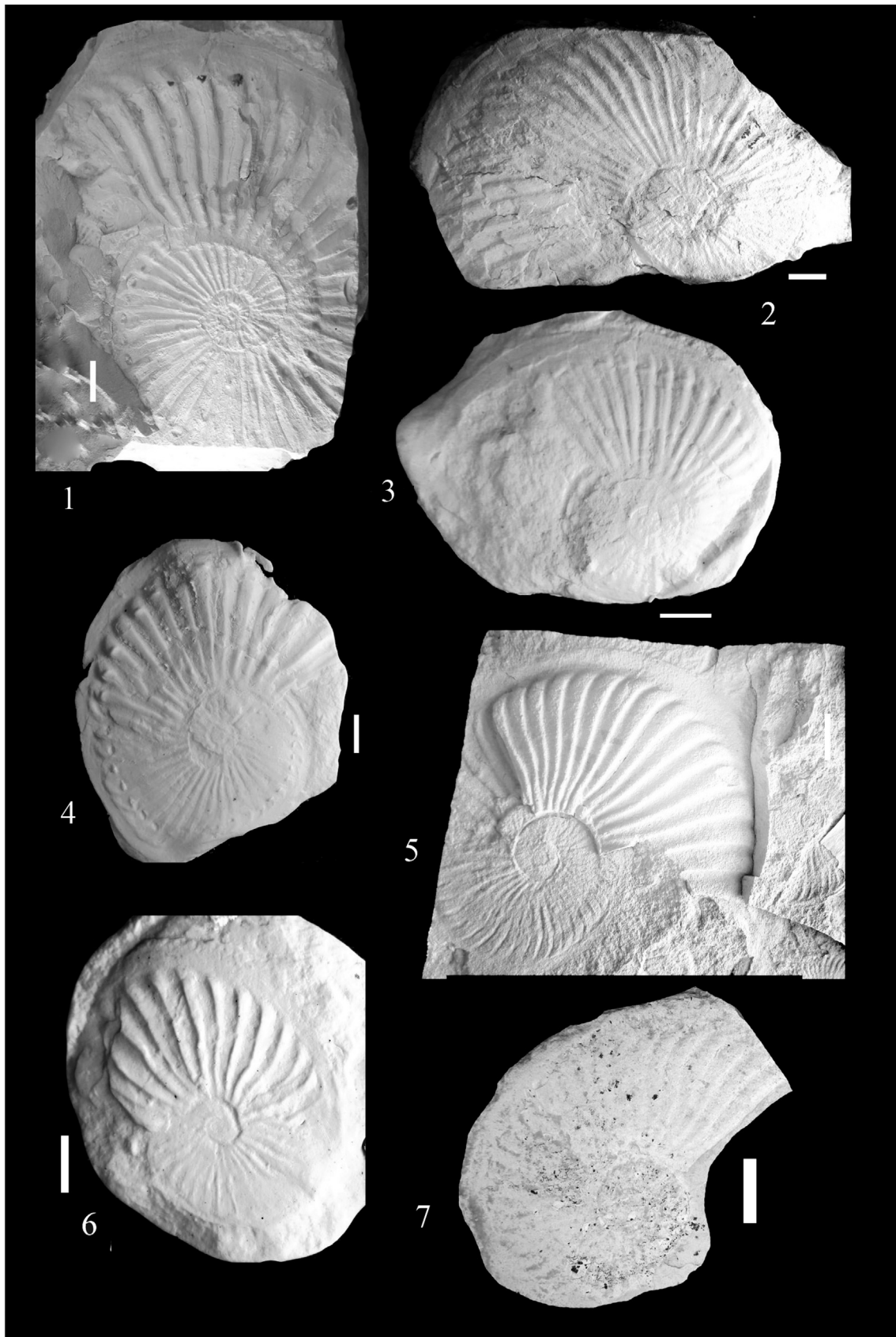


**Fig. 10.** 1, *Oxytropidoceras* (*Oxytropidoceras*) sp., k36a; 2, *O.* (*Venezoliceras*) sp., S1; 3, *Hysterocheras* sp. k36g; 4, *Hysterocheras* sp. S1; 5, *Oxytropidoceras* (*Mirapelia*) aff. *buarquianum* (White, 1887), k30a; 6, *O.* (*Mirapelia*) aff. *mirapelianum* (d'Orbigny, 1850), k28e-1; 7, *O.* (*Mirapelia*) aff. *mirapelianum* (d'Orbigny, 1850), k28e-2; 8, *O.* (*Venezoliceras*) cf. *bituberculatum* Collignon, 1966, k36h; 9, *O.* (*Venezoliceras*) cf. *bituberculatum* Collignon, 1966, k36d, all scale bars: 1 cm.



**Fig. 11.** 1, *Oxytropidoceras* (*Mirapelia*) aff. *buarquianum* (White, 1887), k30b; 2, *Lyelliceras* aff. *pseudolyelli* (d'Orbigny, 1841), k28-1; 3, *Oxytropidoceras* (*Venezoliceras*) cf. *karsteni* (Stieler, 1920), k36i; 4, *O.* (*Venezoliceras*) aff. *bituberculatum* Collignon, 1966, k36j; 5, *Epicheloniceras* sp., k3a; 6, *Oxytropidoceras* (*Venezoliceras*) aff. *bituberculatum* Collignon, 1966, k36e, all scale bars: 1 cm.





**Fig. 12.** 1, *Oxytropidoceras* (*Venezoliceras*) aff. *bituberculatum* Collignon, 1966, k36k; 2, *O.* (*Mirapelia*) aff. *buarquianum* (White, 1887), k30c; 3, *O.* (*Venezoliceras*) cf. *bituberculatum* Collignon, 1966, k36f; 4, *O.* (*Venezoliceras*) cf. *bituberculatum* Collignon, 1966, k36b; 5, *O.* (*Mirapelia*) aff. *buarquianum* (White, 1887), k30d; 6, *O.* (*Mirapelia*) sp., k29a; 7, *O.* (*Mirapelia*) cf. *mirapelianum* (d'Orbigny, 1850), k28e-3, all scale bars: 1 cm.

**Table 1**  
Aptian and Albian biozonation (Reboulet et al., 2018).

| Stage  |        | Zones                                    | Subzones                               |
|--------|--------|------------------------------------------|----------------------------------------|
| Albian | upper  | <i>Mortonicer</i> <i>briacense</i>       |                                        |
|        |        | <i>Mortonicer</i> <i>perinflatum</i>     |                                        |
|        |        | <i>Mortonicer</i> <i>rostratum</i>       |                                        |
|        |        | <i>Mortonicer</i> <i>fallax</i>          |                                        |
|        |        | <i>Mortonicer</i> <i>inflatum</i>        |                                        |
|        |        | <i>Mortonicer</i> <i>pricei</i>          |                                        |
| Aptian | middle | <i>Dipoloceras</i> <i>cristatum</i>      |                                        |
|        |        | <i>Euhoplites</i> <i>lautus</i>          |                                        |
|        |        | <i>Euhoplites</i> <i>loricatus</i>       |                                        |
|        |        | <i>Hoplites</i> <i>dentatus</i>          | <i>Hoplites</i> <i>spathi</i>          |
|        |        |                                          | <i>Leylliceras</i> <i>lyelli</i>       |
|        |        |                                          | <i>Leylliceras</i> <i>pseudolyelli</i> |
|        | lower  | <i>Douvilleiceras</i> <i>mammillatum</i> |                                        |
|        |        | <i>Leymeriella</i> <i>tardefurcata</i>   |                                        |
|        | upper  | <i>Hypacanthoplites</i> <i>jacobi</i>    |                                        |
|        |        | <i>Acanthoplites</i> <i>nolani</i>       | <i>Diadochoceras</i>                   |
|        |        |                                          | <i>nodosocostatum</i>                  |
|        |        | <i>Parahoplites</i> <i>melchioris</i>    |                                        |
|        | lower  | <i>Epicheloniceras</i> <i>martini</i>    | <i>Epicheloniceras</i> <i>buxtorfi</i> |
|        |        |                                          | <i>Epicheloniceras</i> <i>gracile</i>  |
|        |        |                                          | <i>Epicheloniceras</i> <i>debile</i>   |
|        |        | <i>Dufrenoyia</i> <i>furcata</i>         | <i>Dufrenoyia</i> <i>dufrenoyi</i>     |
|        |        |                                          | <i>Dufrenoyia</i> <i>furcata</i>       |
|        |        | <i>Deshayesites</i> <i>deshayesi</i>     | <i>Deshayesites</i> <i>grandis</i>     |
|        |        | <i>Deshayesites</i> <i>forbesi</i>       | <i>Roloboceras</i> <i>hambrovi</i>     |
|        |        | <i>Deshayesites</i> <i>oglanlensis</i>   | <i>Deshayesites</i> <i>luppovi</i>     |

### 3.1. Lower upper Aptian

#### 3.1.1. *Dufrenoyia* aff. *furcata* Zone

This zone is in the lowermost part of the Kazhdumi Formation indicated by the presence of *Chelonicer* sp. and *Dufrenoyia* aff. *furcata*. It is probably the equivalent of the *Dufrenoyia furcata* Zone, the uppermost zone of the lower Aptian (Reboulet et al., 2018). The thickness of deposits of this zone is about 18 m. The occurrence of the *furcata* Zone is geographically widespread in the Tethyan Province, stretching eastward from the Americas to include western Europe.

#### 3.1.2. *Epicheloniceras* sp. Zone

The lowermost upper Aptian ammonite zone of *Epicheloniceras martini* (the *Epicheloniceras martinioides* Zone of Casey 1961) is indicated by the occurrence of *Epicheloniceras* sp. This zone and some of the subzones established by Casey (1961) are of widespread occurrence ranging from North Africa (Memmi 1999) to Mangyschlak (Kazakhstan) (Kopaevich et al., 1999) and the Caucasus (Kotetishvili et al., 2000) in the Tethyan area. It occupies the 10 m interval in the Kazhdumi section below the occurrence of *Parahoplites* sp.

#### 3.1.3. *Parahoplites* sp. and “*Acanthoplites nolani*” Zones

The two zones, if fully present, cannot be separated in the absence of *Acanthoplites*. Together they occupy a thickness of about 25 m in the Kazhdumi section. The occurrence of *Parahoplites* sp. indicates the *Parahoplites melchioris* Zone of the Mediterranean Tethyan Province (Reboulet et al., 2018). Casey (1961) recognised two subzones within his *Parahoplites nutfieldiensis* Zone, a lower one in which *P. nutfieldiensis* is common and an upper one characterised by *Parahoplites cunningtoni*. The English succession indicates that *P. melchioris* is restricted to this upper subzone. The upper boundary of the *Parahoplites* sp. Zone in the Kazhdumi Formation cannot be determined. It should be noted that the systematic and stratigraphic position of “*Acanthoplites*” *nolani*

(*Nolanicer* of Casey) has been questioned by Bulot et al. (2014). However, it is clear that there is an interval between that characterised by *Parahoplites* and the first general occurrence of *Hypacanthoplites*, which is occupied by the genus *Nolanicer* (Kamyabi Shadan, 2014, Fig. 9, 6A,B) characterised by *Nolanicer* aff. *nolani* as indicated by Ruffell and Owen (1995), marking the evolutionary trend from *Parahoplites* to *Hypacanthoplites*.

#### 3.1.4. *Hypacanthoplites* sp. Zone

The occurrence of *Hypacanthoplites* in the Kazhdumi Formation indicates the presence of the uppermost ammonite zone of the Aptian. The thickness of this zone in the section is 20 m. It indicates the *Hypacanthoplites jacobi* Zone of authors (Reboulet et al., 2018). However, as Mutterlose et al. (2003) indicated, the occurrences of *Hypacanthoplites jacobi* in Germany is up to lower Albian *Leymeriella acuticostata* Subzone (the *Leymeriella tardefurcata* Zone of Russian etc. authors). A lower interval characterised by *Hypacanthoplites* is present but might correspond to the lowermost zone of the Albian characterised by *Proleymeriella*. The problem of determining the succession in areas in which the early forms of *leymeriellids* are absent but are characterised by early *Hypacanthoplites*, causes problems in determining the global Aptian–Albian boundary. This situation has been discussed by Casey (1999), Kennedy et al. (2000) and Owen (2002).

### 3.2. Lower Albian

#### 3.2.1. *Douvilleiceras mammillatum* and *Leylliceras pseudolyelli* Zones

These two uppermost lower Albian ammonite zones are represented in the Kazhdumi section by ammonites but cannot be separated vertically. The base of combined zone is defined by occurrence of *Douvilleiceras* aff. *mammillatum* in sample number 19, followed shortly above by *Douvilleiceras mammillatum* aff. *aequidum*; the latter species is also reported by Bulot (2010) from this section. The combined thickness is about 19 m. We use here the



conventional restricted concept of the *mammillatum* Zone used by Reboulet et al. (2018) and not the *mammillatum* Superzone recognised by Owen (1999). *Douvilleiceras* occurs already in the *acuticostata* Subzone of the *Leymeriella tardefurcata* Zone (Kennedy and Kollmann, 1979; Owen, 1999; Kennedy et al., 2000) and ranges into the lower middle Albian *Lyelliceras lyelli* Subzone.

Following Amedro (1992), Latil (1994) recognised the *Lyelliceras pseudolyelli* Subzone at the summit of the lower Albian, characterised by transitions from *Tegoceras* to the lowermost middle Albian *Lyelliceras lyelli*. Both intervals are of geographically widespread occurrence from northern South America eastward to Africa and Madagascar to India and northward into the European faunal Province.

### 3.3. Middle Albian

#### 3.3.1. *Lyelliceras lyelli* Zone

Not proven by ammonites in the Kazhdumi section, but might be present in the interval between 145 and 180 m. The Zone is represented to the south in Zululand (Kennedy and Klinger, 2008) and Madagascar (Collignon, 1963) as well as to the north.

#### 3.3.2. *Oxytropidoceras* spp. Zone

The genus *Oxytropidoceras* (if *Mirapelia* is assumed as a subgenus of *Oxytropidoceras*) appears above the base of lower Albian, but *Oxytropidoceras* s. str. is appearing at the base of middle Albian, and is extinct in the upper Albian, above the *Dipoloceras cristatum* Zone. Many middle Albian species belonging to the genus *Oxytropidoceras* have been used as index species for regional biozonations, in California (Amédro and Robaszynski, 2005), Texas (Young, 1966), Brazil (Bengtson, 1983), Madagascar (Collignon, 1963 and 1978), Angola (Cooper, 1982; Tavares et al., 2007).

Owen (1996) introduced the *Oxytropidoceras* spp. Zone. Bulot and Latil (In Reboulet et al., 2011) have proposed to subdivide the middle Albian 'into a lower *L. lyelli* Zone and an upper *Oxytropidoceras* spp. Zone. This division was first agreed for the "Tethyan Province" during the London meeting (Rawson et al., 1999) to be used in parallel with the hoplitid zones of the "European Province". But subsequent meetings made no reference to the proposal and continued using only the "European Province" zonation'. Gale et al. (2011) recognise a *Lyelliceras lyelli* Zone and a *Oxytropidoceras* (*O.*) *roissyanum* Zone as possible divisions of the middle Albian in southeast France.

On this basis, an *Oxytropidoceras* spp. Zone is suggested for the interval between the first and last occurrences of the genus *Oxytropidoceras* in the studied section, assigning it to the upper middle Albian.

### 3.4. Upper Albian

There appears to be a major hiatus at the top of the middle Albian strata, before the lower upper Albian deposits. This hiatus suggests the *Dipoloceras cristatum* tectonic event of possible global extent indicated by Owen (1971).

#### 3.4.1. *Hysterocheras* sp. Subzone

*Hysterocheras* species has been used to characterize biozone and/or subzone in the lower part of upper Albian. According to Lopez-Horgue et al. (2009), the *Hysterocheras varicosum* Zone includes the *Hysterocheras choffati*, *Hysterocheras orbigny*, and *Hysterocheras varicosum* Subzones, overlying the *Dipoloceras cristatum* Zone, and underlying the *Mortoniceras inflatum* Zone (Owen, 1999; Owen and Mutterlose, 2006; Wiedmann and Owen, 2007) or is considered as an equivalent of the *Mortoniceras pricei* zone (Reboulet et al., 2014).

This biozone is known in the uppermost part of the Kazhdumi Formation and the lowermost part of the Sarvak Formation.

## 4. Conclusions

Representatives of different genera and species of families Brancoceratidae Spath, 1923, Lyelliceratidae Spath, 1921, Douvilleiceratidae Parona and Bonarelli, 1897, Deshayesitidae Stoyanow, 1949, Parahoplitidae Spath, 1922 are reported from the Kazhdumi Formation in the Tange Maghar. These are as follows:

*Chelonicer* sp., *Douvilleiceras* aff. *mammillatum*, *D. mammillatum* aff. *aequinodum*, *Dufrenoyia* aff. *furcata*, *Epicheloniceras* sp., *Hypacanthoplites* sp., *Hysterocheras* sp., *Lyelliceras* aff. *pseudolyelli*, *Oxytropidoceras* (*Mirapelia*) aff. *buarquianum*, *O.* (*Mirapelia*) aff. *mirapelianum*, *O.* (*Mirapelia*) cf. *mirapelianum*, *Oxytropidoceras* sp., *O.* (*Venezolicer*) aff. *bituberculatum*, *O.* (*Venezolicer*) cf. *karsteni*, and *O.* (*Venezolicer*) sp.

The following biozones are suggested on the basis of the ammonite assemblages:

*Dufrenoyia* aff. *furcata* Zone (uppermost lower Aptian).  
*Epicheloniceras* sp. Zone (lowermost upper Aptian).  
*Parahoplites* sp. and "*Acanthoplites nolani*" zones (upper Aptian).  
*Hypacanthoplites* sp. Zone (uppermost Aptian–lowermost lower Albian).

*Douvilleiceras mammillatum* and *Lyelliceras pseudolyelli* zones (uppermost lower Albian).

*Oxytropidoceras* spp. Zone (middle Albian).  
*Hysterocheras* sp. Subzone (lower upper Albian).

An early Aptian to late Albian age is proposed for the Kazhdumi Formation. Correlation of the Hoplitid biozones between the Boreal and north Tethyan realm, and the south Tethyan domain is discussed.

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