Advances and Challenges in understanding the amplitude in noise correlations

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Scale of our problem





Local - Regional Sedimentary Basins 10-100's km

Regional - Global Western US 100-1000's km



Local Scale Building 10's meters





Scale of our problem



250

3

-118.5° -118° -117.5° -117° **Regional - Global** 34.5° Freq: 0.125–0.025 Hz Quake T: 8 – 40 sec **Local Scale** RUS RIO BBR Prop Vel: 3 – 4 km/s Freq: 0.5–5 Hz 34° 8 s Love wave phase velocity **Prop Vel 100 – 200 m/s** LAF RPV 33.5° Local – Regional Freq: 0.1 – 0.5 Hz T: 2 – 10 sec Prop Vel: ~ 3 km/s 235 2.65 3.05 3.15 3.25 3.35 3.45 3.55 3.65 3.75 Lin et al. GJI (2008)

Velocity (km/s)



Lower mantle

Of course, noise has gone global

655

665

9.5

10 10.5 11

9.5

Vp (km/s)

10.5





600

650

8.5

9



Prieto, Science (2012)







Amplification and Attenuation

Tomographic Imaging of Elastic Structure

Higher resolution with increase data coverage

Seismic attenuation tomography has lagged behind

More difficult. Amplitude of seismic waves affected by 3D velocity, multipathing, scattering, and source.



Why is this important?



Amplification and Attenuation

- Recent ground motion simulations suggest wave guide by sedimentary basins and large amplitudes in the LA Basin.
- Basin for major cities (Tokyo, LA, Mexico City, Bogotá, Colombia)



How can we validate these simulations?





Amplification and Attenuation

Attenuation provides relevant physical properties of Earth.

Very sensitive to temperature Presence of fluids (melt or water) Tectonic activity (E,W US)



[*Lawrence & Wysession.*, 2005: AGU Monograph; *Karato*, 2003: AGU Monograph]

How can we validate these simulations? Monography Can we use amplitudes to monitor changes in the media?





Is there reliable amplitude information in noise correlations?

Geometrical Spreading Basin Amplification Attenuation

YES

. . .

(under some conditions) STILL DEBATED HOW

... substantial and competing amplitude anomalies due to elastic and anelastic variations ... (Savage et al., 2010)





Is there reliable amplitude information in noise correlations?

- 1. First suggestions on amplitude information
- 2. Some theoretical questions
- 3. A numerical example
- 4. Two recent success stories
- 5. Always improving, how can we do better?

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Weaver and Lobkis, 2001 Larose et al., 2007 Prieto and Beroza, 2008 *Matzel, 2008 Prieto et al., 2009 Taylor et al., 2009* Cupillard and Capdeville, 2010 Lawrence and Prieto, 2011 *Lin et al., 2011; 2012* Weaver, 2011a; 2011b; 2013

Prieto et al., 2011









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Prieto et al. CRG (2011)



Based on seismic observations, we suggested

Aki (1957) SPAC Method



Freq: Re[
$$\gamma_{AB}(\omega, r)$$
] = $\left\langle \frac{u_A u_B^*}{\langle |u_A| \rangle \langle |u_B| \rangle} \right\rangle$ = $J_0\left(\frac{\omega r}{C}\right)$

C = Phase Velocity

Colin de Verdière (2006)

Sime:
$$\frac{\partial}{\partial \tau} \langle C_{AB}(\tau) \rangle = \frac{-\sigma^2}{4a} (G_{AB}(\tau) - G_{AB}(-\tau))$$

$$\operatorname{Re}[\gamma_{AB}(\omega,r)] = J_0\left(\frac{\omega r}{C}\right) \exp(-\alpha r)$$



Amplitude information present, but not simple $e^{\alpha r}$

Inference of attenuation from comparison of empirical coherencies to $Jo(\omega r/c) \exp(-\alpha r)$ is problematic (Weaver, CRG).

Site factors, noise intensity (directivity) and attenuation all have an effect on amplitude measurements







Behavior of $e^{-\alpha r}$ only in special cases

$$\hat{C}_{xy}^{E} = \frac{1}{I_0(\alpha r_{xy})} J_0\left(\frac{\omega r_{xy}}{c}\right).$$

Uniform distribution of farfield sources

$$\hat{C}_{xy}^{E} = \frac{A_{\theta}^{2}}{\overline{A}^{2}}e^{-\alpha r_{xy}}J_{0}\left(\frac{\omega r_{xy}}{c}\right).$$

Including far and near-field source





SPAC method in disipative media (Nakahara, GJI, 2013)

Analysis of the SPAC expressions for 2-D cases shows that the conjecture of *Prieto et al.* (2009) is not strict but approximately good for small attenuation.

$$C_{1,2}(\mathbf{r},\omega) = J_0(k_0 r) \quad \text{for} \quad 2\text{-}\mathbf{D},$$

Non-dissipative media

$$C_{1,2}(\mathbf{r},\omega) = \frac{\sin(k_0r)}{k_0r}$$
 for 3-D.

$$C_{1,2}(\mathbf{r},\omega) \approx \exp(-\kappa r) \left[J_0(k_0 r) \left(1 + \frac{2\kappa}{\pi k_0} \right) + \frac{\kappa}{2k_0} Y_0(k_0 r) \right].$$

$$C_{1,2}(\mathbf{r},\omega) = \frac{\sin k_0 r}{k_0 r} \exp(-\kappa r).$$

Dissipative, assuming large separations and small attenuation





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Lawrence et al. In prep.



Do simulations support our claim?

We generate

- random array of sensors (red)
- random location of "noise" sources
- 3 months of synthetic data



Lawrence et al. In prep.



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$$u(\mathbf{x},\mathbf{s}) = \int_{A} F(\mathbf{s},\omega) e^{i\omega t} e^{-ikr_{\mathbf{s}\mathbf{x}}} e^{-\alpha(\omega)r_{\mathbf{s}\mathbf{x}}} r^{-0.5} dA$$

 $F(s, \omega)$ – source at location s r_{sx} – source receiver distance

- α attenuation coefficient
- A integration over sources



Lawrence et al. In prep.





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$$\overline{\gamma}_{\mathbf{xy}}(\omega) = \left\langle \frac{\sum_{j=1}^{N_s} F(\mathbf{s}_j, \omega) e^{i\omega t} e^{-ikr_{\mathbf{s}_j\mathbf{x}}} e^{-\alpha(\omega)r_{\mathbf{s}_j\mathbf{x}}} r^{-0.5} \sum_{j=1}^{N_s} F(\mathbf{s}_j, \omega) e^{i\omega t} e^{-ikr_{\mathbf{s}_j\mathbf{y}}} e^{-\alpha(\omega)r_{\mathbf{s}_j\mathbf{x}}} r^{-0.5}} \left| \frac{\left| \sum_{j=1}^{N_s} F(\mathbf{s}_j, \omega) e^{i\omega t} e^{-ikr_{\mathbf{s}_j\mathbf{x}}} e^{-\alpha(\omega)r_{\mathbf{s}_j\mathbf{x}}} r^{-0.5}} \right| \right| \left| \sum_{j=1}^{N_s} F(\mathbf{s}_j, \omega) e^{i\omega t} e^{-ikr_{\mathbf{s}_j\mathbf{y}}} e^{-\alpha(\omega)r_{\mathbf{s}_j\mathbf{x}}} r^{-0.5}} \right| \right\rangle}{\left\langle \left| \sum_{j=1}^{N_s} F(\mathbf{s}_j, \omega) e^{i\omega t} e^{-ikr_{\mathbf{s}_j\mathbf{x}}} e^{-\alpha(\omega)r_{\mathbf{s}_j\mathbf{x}}} r^{-0.5}} \right| \right\rangle} \right\rangle} \right\rangle$$



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Accurate phase velocity and amplitude decay observed in simulated cross-correlations.







Simulated coherencies match Bessel function for elastic case





Simulated coherencies match attenuated Bessel function better





Does not always work, but does on realistic source distributions





Attenuation measured, is the interstation attenuation





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 \bigcirc



(Zhang and Yang., 2013)

Amplitude decay – line-array examples

18 sec

geometric spreading corrected clear EGF, yet **biased** attenuation







Amplitude decay – line-array examples

18 sec geometric spreading corrected clear EGF, yet **biased** attenuation

(Zhang and Yang., 2013)

May this bias be due to noise intensity? noise directivity?







Amplitude decay – line-array examples



(Zhang and Yang., 2013)

 $\left[\right]$

Attenuation from noise – C³











(Zhang and Yang., 2013)



Attenuation from noise – C³



12 sec

 $\alpha (10^{-3})$

Amplitude decay – line-array examples









ambient seismic field recorded at stations





Weak coherent ambient seismic field recorded at stations

Extract impulse response → Path Effects

Convert surface impulse responseto buried double-couple response

➔ Point Source Virtual Earthquake









Depth and mechanism correction

Isotropic radiation from surface source

Azimuth-dependent radiation from a deep source

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Validation against real data









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Standard practice suggests that we can

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$$H1 = \frac{Gx}{Gy} \qquad \qquad H1 = \frac{G_x \bullet conj(G_y)}{G_{xx}}$$

But, every H1 function is based on two-station correlations.

In a multiple Input/Output system (Bendat and Piersol, 2010)

 $G_{2y}(f) = H_1(f)G_{21} + H_2G_{22}$

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Transfer Function is based on coherent signal

$$H_1(f) = \frac{G_{1y}(f) \left[1 - \frac{G_{12}G_{2y}}{G_{22}G_{1y}}\right]}{G_{11} \left[1 - \gamma_{12}^2(f)\right]}$$

 γ_{12} – Coherency between signal at stations 1 and 2

 G_{xy} – X-Y Cross-Spec G_{yy} – Power Spectrum H_1 – Transfer function

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Ambient-noise Green's functions can be used for Attenuation Tomography Long-Period Ground Motion Predictions.

But, we can do better using improved signal processing

- Correct directivity and source intensity
- Apply C3 methods
- Impulse response functions, with coherency constraints.

THANK YOU !