Multi-scale/signal Seismic imaging of Fault Zone Environments

Yehuda Ben-Zion, University of Southern California

with D. Zigone, Z. Ross, A. Allam, H. Qui, P. Share, F. Vernon, G. Hillers, M. Campillo, P. Roux & others



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Why image fault zones?

- Derivation of earthquake source properties.
- Evolutionary processes on long (tectonic) and short (e.g., precursory) timescales.
- Static/dynamic stress fields (e.g. from internal structure).
- Brittle rock rheology (e.g. from observing & monitoring rock damage).
- Elements of FZ structure (bimaterial interfaces and damage zones) can control future (and reflect past) earthquake rupture properties.



Rockwell and Ben-Zion (2007)

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- Elements of FZ structure (bimaterial interfaces and damage zones) can control future (and reflect past) earthquake rupture properties
- FZs control crustal fluid flow: hydrology, oil, sub-surface storage, minerals, etc.



Hauksson et al. (2015)

Focus in this talk: The San Jacinto Fault Zone in southern CA









Data: Earthquake waveforms, ambient seismic noise (+ small "Besty" gunshots)



100m

JF

33°15'

100m

Seismic Velocity Structures in the Southern California Plate Boundary Environment from Doubledifference Earthquake Tomography (Allam and Ben-Zion, GJI, 2012)





Horizontal resolution 2–3 km over depth section 2–15 km

Broad low velocity fault zone layers (3-6 km wide) are:

-Seen clearly along the SJFZ -Offset to the NW in the central section

-Prominent generally in top 5 km -Follow overall "flower" shape with depth

-Larger reductions of Vs (up to 40% in top 3-5 km) than Vp

Clear velocity contrasts (bimaterial interfaces?) in different sections; polarity flips NW of SJB.

 SAF: San Andreas Fault
 SJFZ: San Jacinto Fault Zone
 EF: Elsinore Fault

 SJB: San Jacinto Basin
 AZ: Anza
 TF: Trifurcation An

 CC: Coyote Creek Fault
 SH: Superstion Hills Fault
 IF: Imperial Fault

EF: Elsinore Fault ST: Salton Trough CP: Cajon Pass SBB: San Bernardino Basin TF: Trifurcation Area BR: Buck Ridge Fault CL: Clark Fault IF: Imperial Fault

Additional detailed results on complex SJFZ regions given by Allam, Ben-Zion, Kurzon, Vernon (GJI, 2014)

Noise-based Imaging of Southern California Plate Boundary Area Dimitri Zigone, Yehuda Ben-Zion, Michel Campillo, Phillipe Roux (2015)



Cross-correlations : processing



Surface Waves Green's functions and Velocity Measurements



Use of the correlation tensor

Rayleigh wave:

- Frequency-time analysis of ZZ, RR,
 ZR, RZ on both the causal and acausal parts; 8 possible measurements
- Multiply the results
- Make dispersion measurements
- Love wave:
 - Uses of TT causal and acausal.
 - Only two measurements



Frequency-time dispersion analysis (Levshin et al, 1989)

Propagation of Rayleigh and Love waves



Group Velocity Maps

- Inversion of velocity from dispersion curves (Barmin et al. 2001).
- 1.5 km grid ; use only cells with >3 measurements

Number of paths per cell (used as weight in the inversion)













VS Maps (linearized Inversion of Hermann and Ammon, 2002)



Vs from Rayleigh waves

Horizontal resolution 2–3 km over depth section 1–7 km

Broad low velocity fault zone layers (3-6 km wide)

Clear velocity contrasts in different sections; the polarity flips NW of SJB.

 $2-\theta$ anisotropy with fast directions parallel to simple fault sections and random in complex areas







Summary of "large scale" earthquake- and noise-based imaging results

The TomoDD and Noise-based images resolve multiple important structural features of the plate boundary region around the SJFZ with horizontal scale of 2-3 km over the depth section 0.5-15 km.

Internal components of the SJFZ are imaged with fault zone head and trapped waves (next slides).

The structure in the top 0.5 km is imaged (up to the surface) with high-frequency noise (1-200 Hz). (following slides).

Fault zone head and trapped waves



Imaging bimaterial fault interface with Head Waves



The head waves are <u>first arrivals</u> at stations on the slower side of the fault with normal distance $x < x_c = r \tan \left[\cos^{-1}(\alpha_2/\alpha_1)\right]$ and have <u>opposite polarity</u> than the direct *P* wave (Ben-Zion, 1989, 1990)

Fault zone head and trapped waves



Fohrmann, Igel, Jahnke, Ben-Zion (2004)

Trapped waves depends strongly on $N = r/[W \tan(\theta_c)] = r/[W \tan(\sin^{-1}(\beta_2/\beta_1))]$

Internal structure of the San Jacinto fault zone at Jackass Flat (JF) from data recorded by a dense

linear array (Qiu, H., Y. Ben-Zion, Z. Ross, P.-E. Share and F. Vernon, 2015)



S velocity reduction ~50%,

Q-value ~20.

Top: synthetic waveform fits.

Bottom: parameter-space results of genetic inversion algorithm.

Rupture zone of the 1992 Landers, CA, earthquake (Peng et al., GJI, 2003)



width ~200 m, depth ~3.5 km, S velocity reduction ~50%, Q-value ~15.

Trapped noise: paper & poster by G. Hillers et al.



Similar results are obtained from analyses of trapped waves also at the:

Karadere-Duzce branch of the NAF (Ben-Zion et al. 2003)

•Parkfield section of the San Andreas Fault (Li et al., 1990; Lewis and Ben-Zion, 2010)

•Calico fault, ECSZ, CA (Yang et al., 2011)

•Rupture zones of the 1994 Kobe and 2009 L'Aquila events (e.g., Mizuno & Nishigami 1998; Caldernoni *et al.* 2013)

Detailed analyses of subsets of data imply <u>strong</u> <u>along-strike variations and discontinuities of the</u> <u>trapping structures</u>! (Lewis & Ben-Zion, GJI, 2010)

Characterization of the San Jacinto Fault Zone northwest of the trifurcation area from dense linear Black Burn (BB) array data (Share, P.-E., Y. Ben-Zion, Z. Ross, H. Qiu and F. Vernon, 2015)



contrast 4-5% over top 15 km

(can be twice as large in the top 7.5-10 km)

Fault zone head (green) and direct P (red) waves at 3 stations of the BB array from 35 events.



Bimaterial interface along the Hayward fault

There is a continuous bimaterial interface for ~80km along the Hayward fault with 3-8% velocity contrast



Allam, Ben-Zion & Peng (2014)

Similar results are obtained from analysis of fault zone head waves also at:

Different sections of the San Andreas fault (Ben-Zion and Malin 1991; McGuire and Ben-Zion 2005)

Calaveras fault (Zhao and Peng 2008)

Different sections of the North Anatolian fault (Bulut et al. 2013; Bita et al. 2015)

Garze-Yushu fault in Tibet (Yang et al. 2015)

Eastern CA shear zone (Hough et a. 1994)

Imaging the top few 100 m with correlations of high-frequency noise across SJFZ array DW

(Zigone, Ben-Zion, Campillo, Hillers, Roux, Vernon, 2015)



- Low velocity zone around the surface trace of the fault

- Flower structure with depth

- Vs30 ~ 250 m/s similar to borehole observations

poster by D. Zigone et al.





Sage Brush Flats - Clark Fault







Basic data features and results from a spatially-dense seismic array on the San Jacinto fault zone

Ben-Zion, Vernon, Ozakin, Zigone, Ross, Meng, White, Reyes, Hollis and Barklage (GJI 2015)



946 events with $-0.7 \le M_L \le 3.4$ were detected in the >4 weeks deployment period by the standalone stations (black circles)

120 events (90 of which undetected by the ANZA network) with median and minimum size -0.29 and -1.58 were detected in 24 hrs by the dense network (red circles)

Data at 2 stations in Julian day 146



Seismograms from one event (yellow star in map)



Seismograms from another event (pink star in map)





⁹⁴⁶ events with $-0.7 \le M_L \le 3.4$ were detected in the >4 weeks deployment period by the standalone stations (black circles)

120 events (90 of which undetected by the ANZA network) with median and minimum size -0.29 and -1.58 were detected in 24 hrs by the dense network (red circles)

Times of 120 events in JD 146: vertical lines



across the entire array

Beamforming analysis of repeating sources in JD 146





Event location with matched field processing and back projection

Original location (yellow star in previous plot): black asterisk

MFP location: white asterisk **Median noise levels during the experiment** (note zones with amplified motion)





Snapshot of ground motion produced by local earthquake



Trapped fault zone waves from Betsy gunshots

animations

Imaging the shallow material with high-frequency noise in row 13 of the array



poster by D. Zigone et al.

Conclusions

- High resolution imaging of fault zone environments requires using different signals (body waves, head and trapped waves, scattered waves, anisotropy, surface waves, ...) and techniques (travel time and waveform tomography, noise correlations, ...).
- Earthquake data provide detailed information on the seismogenic sections (depth 3-15 km).
- Ambient noise data provide detailed information on the shallower structure (up to the surface).
- Large fault zones have hierarchical flower-type damaged zones and bimaterial interfaces.
- The damage zones have ~100m wide cores with intense damage (e.g. $\Delta\beta$ ~ 30-50%, Q ~ 10-30) that act as seismic trapping structures. They have considerable along-strike variability & discontinuities over length scales of ~5 km and are surrounded by ~3-5 km zones of lower damage that produce fault-related anisotropy, elevated scattering and motion amplification.
- Prominent bimaterial interfaces extend to the bottom of the seismogenic zone and are continuous over 10s to 100 km. There are also shallow bimaterial interfaces (e.g. at the edge of damage zones)
- The damage zones get re-activated during earthquakes (strong co-seismic velocity reduction of 30-40 % in the top few 100's of m of the crust) followed by log(t) healing.
- The damage zones are often asymmetric w.r.t. the principle slip zone, which may reflect preferred propagation direction of earthquake ruptures.

Analysis of spatially-dense array:

- \cdot Can be used to detect many more smaller quakes
- \cdot Locate sources of noise and small quakes
- \cdot Study coherency of waves and fine structural details
- \cdot The shallow crust sustains numerous ongoing small failures; it is far more dynamic than usually assumed

