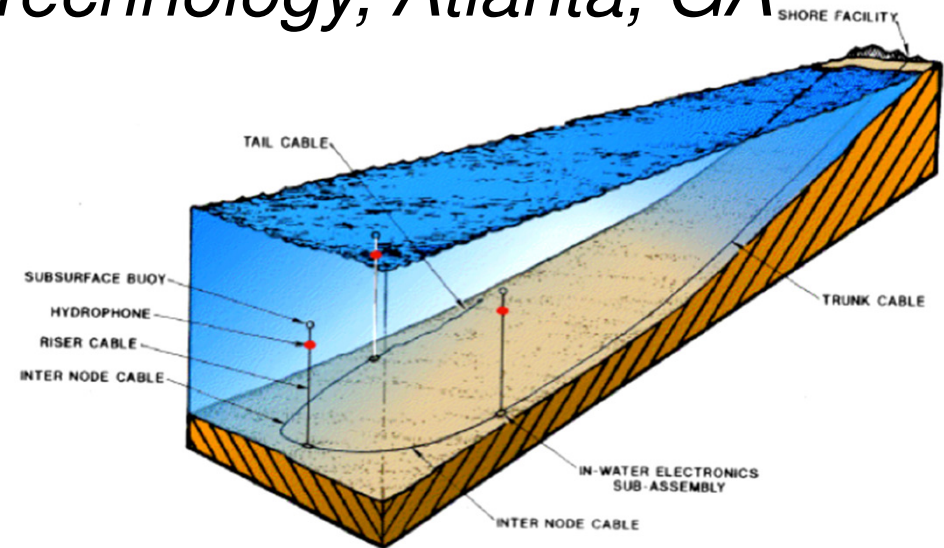
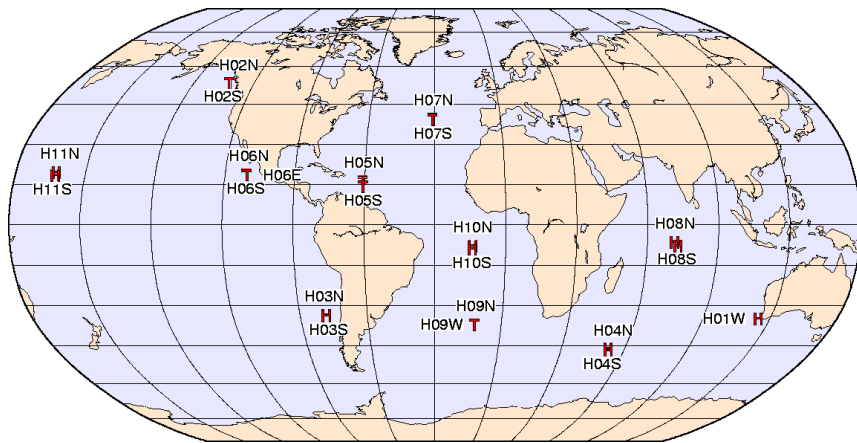


# Passive acoustic thermometry of the deep water sound channel using ambient noise .

Karim G. Sabra

*School of Mechanical Engineering,  
Georgia Institute of Technology, Atlanta, GA*



**Sponsored by the Office of Naval Research  
code 32**

# Presentation Outline

1. **Deep water acoustics & low-frequency noise sources**
2. **Monitoring with ambient noise: Background.**
3. **Passive acoustic thermometry of the deep water sound channel using low-frequency ambient noise**
4. **Monitoring with ambient noise: Optimization.**



**Shane Lani**

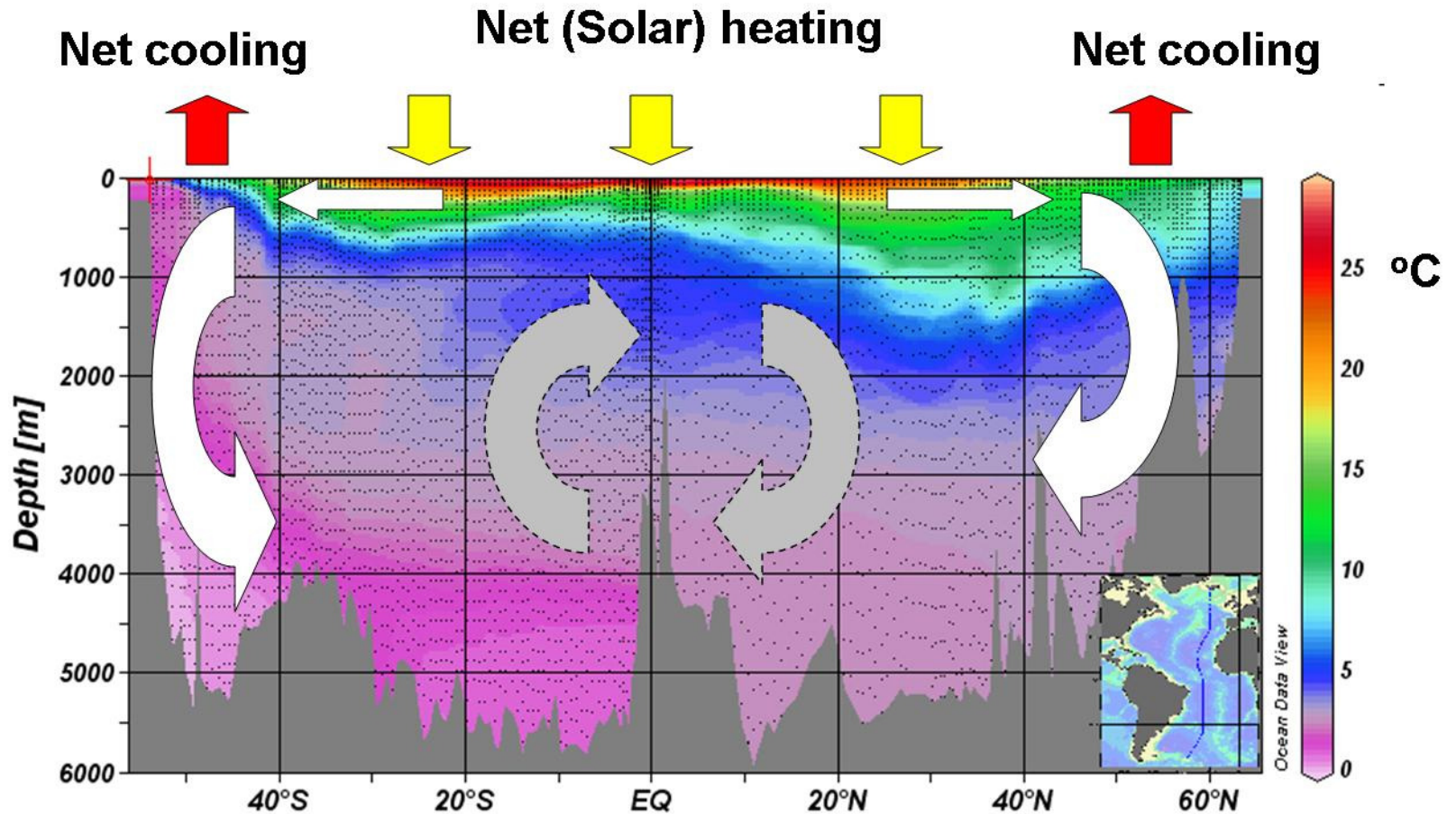
In collaboration with  
Bill Kuperman &  
Philippe Roux



**Katherine Woolfe**

# Oceans are heat sinks

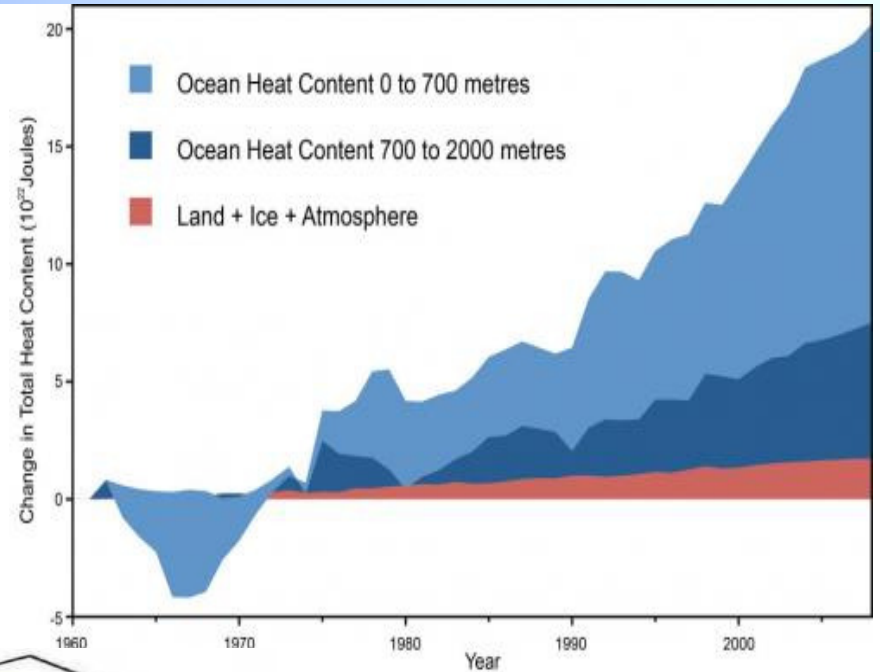
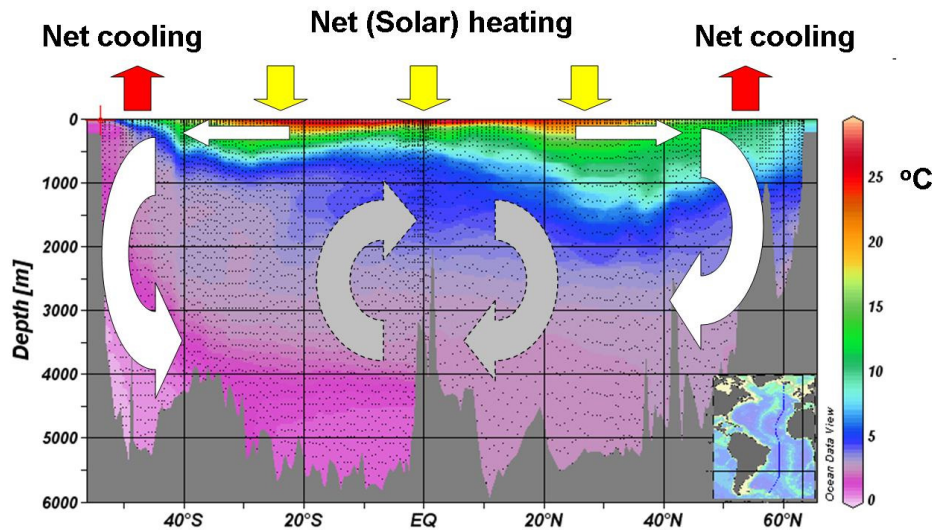
## Average Ocean Temperature Distribution with Depth





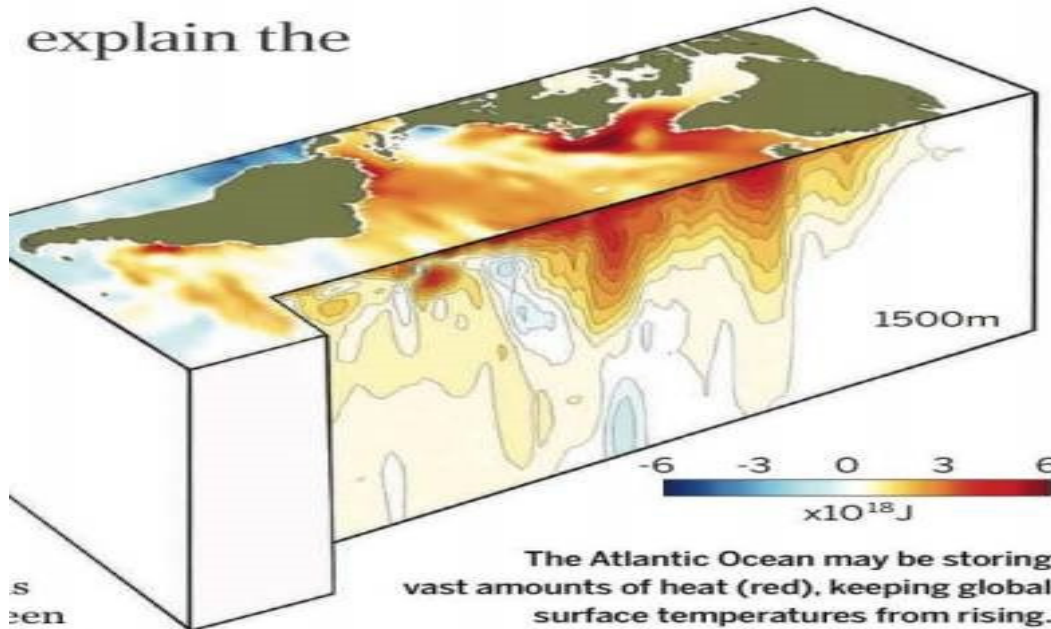
# Oceans are Heat sinks

## Average Ocean Temperature Distribution with Depth



NOAA

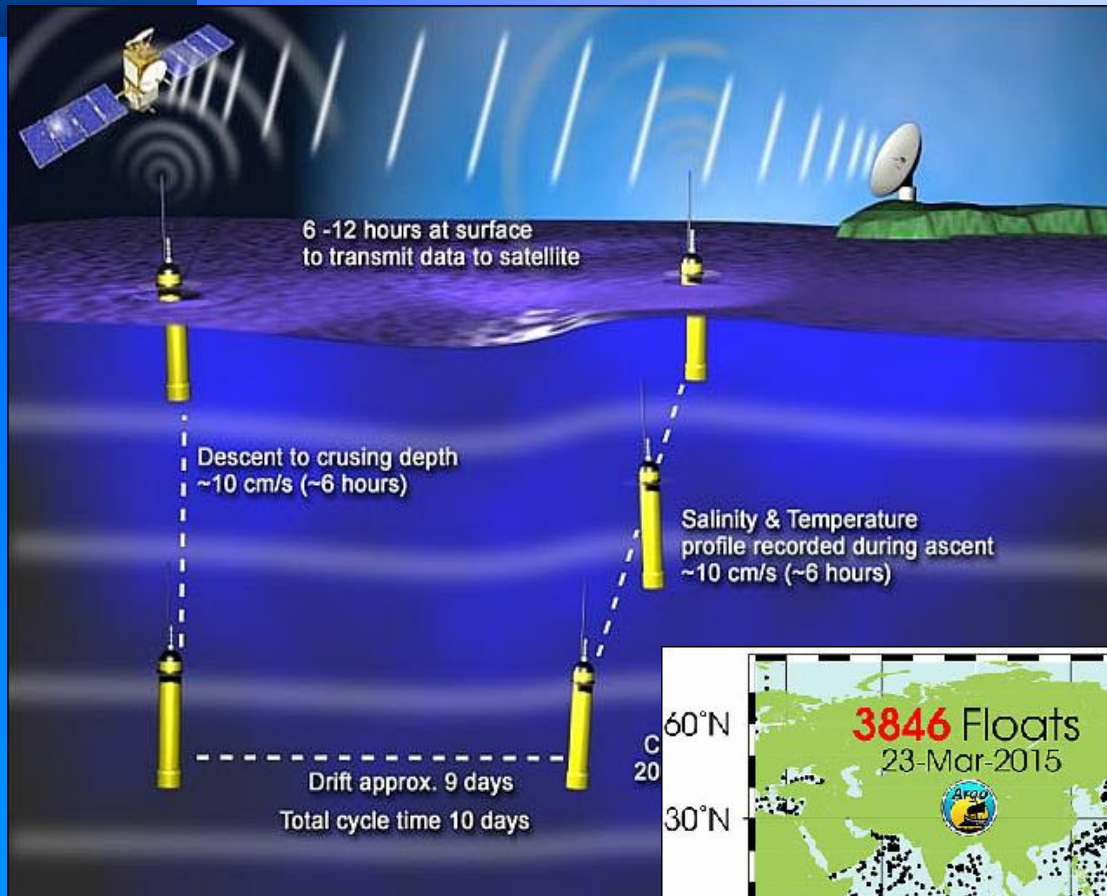
explain the



Science 22 August 2014: Vol. 345 no. 6199 pp. 860-861 Is Atlantic holding Earth's missing heat?

& Variations across ocean bassins...

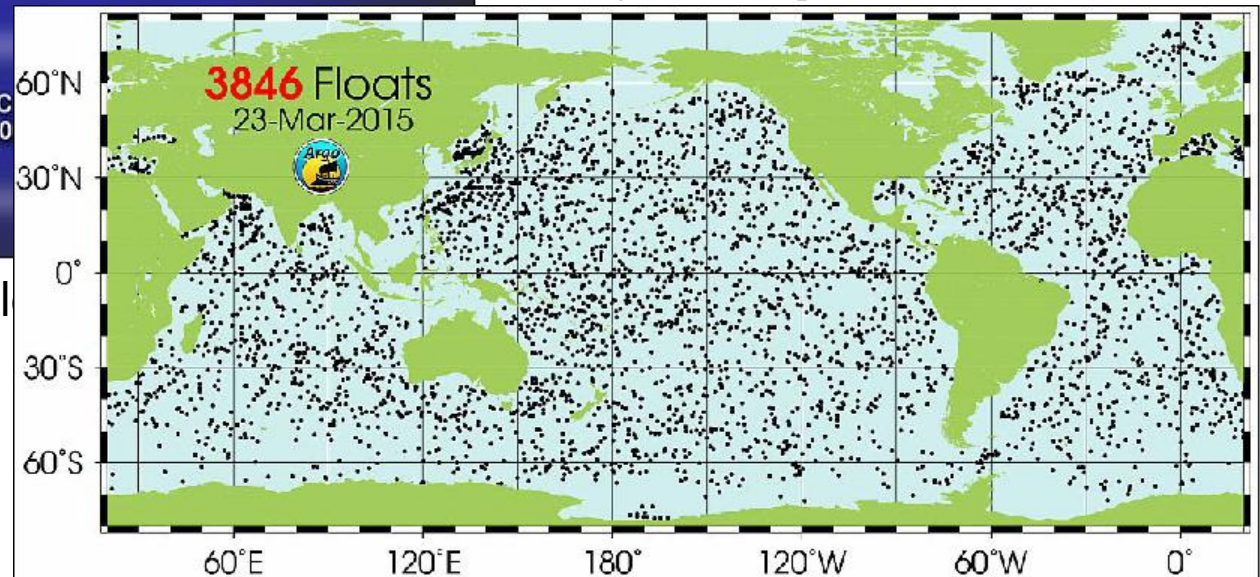
# Measuring deep ocean temperatures



- Satellite observations are most suited for sensing ocean temperature “close” to surface
- Deep oceans most commonly sense with profiling floats.

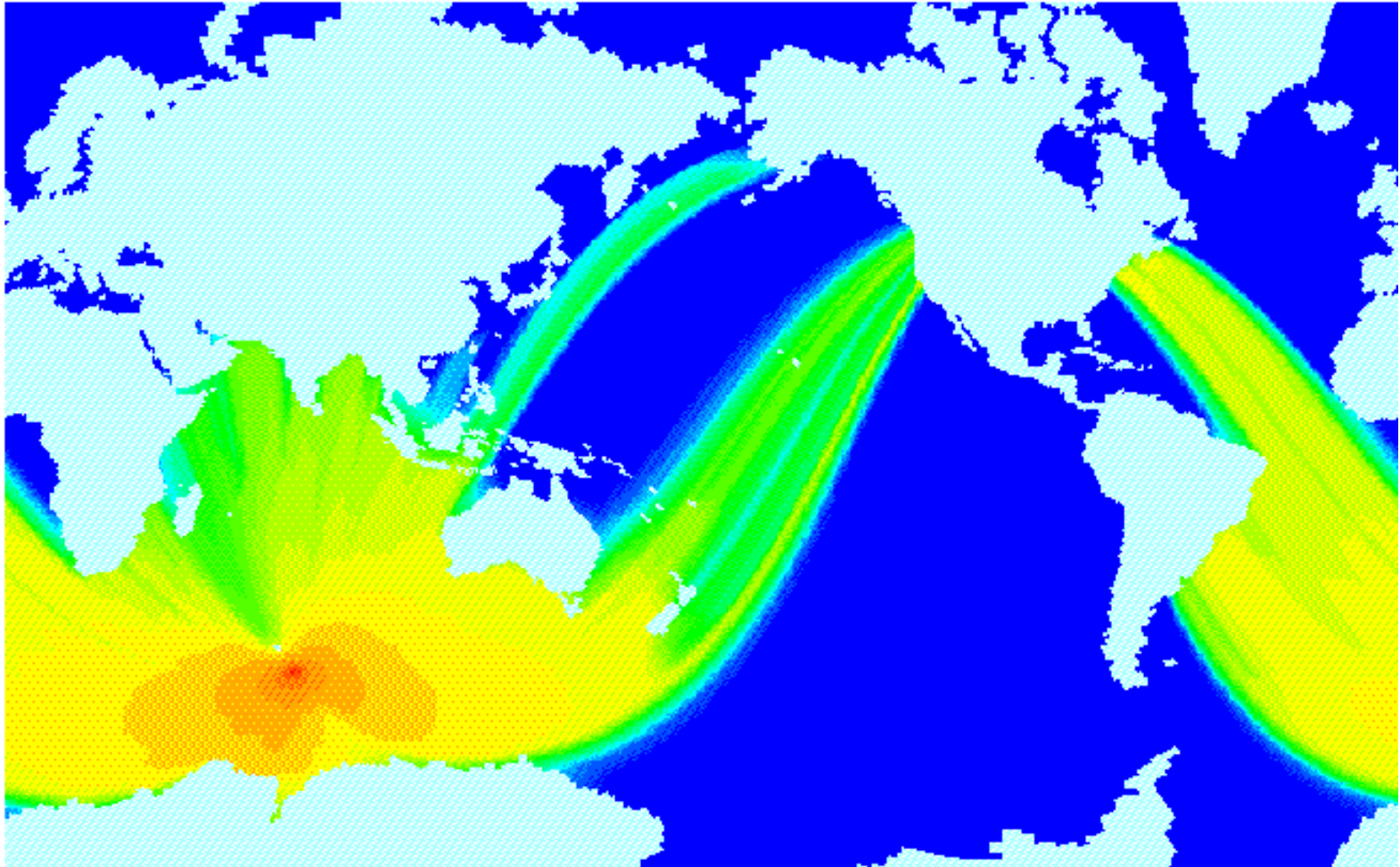
[image credit: AIC (Argo Information Center), UCSD]

Schematic view of the SOLO float (image credit: UCSD)





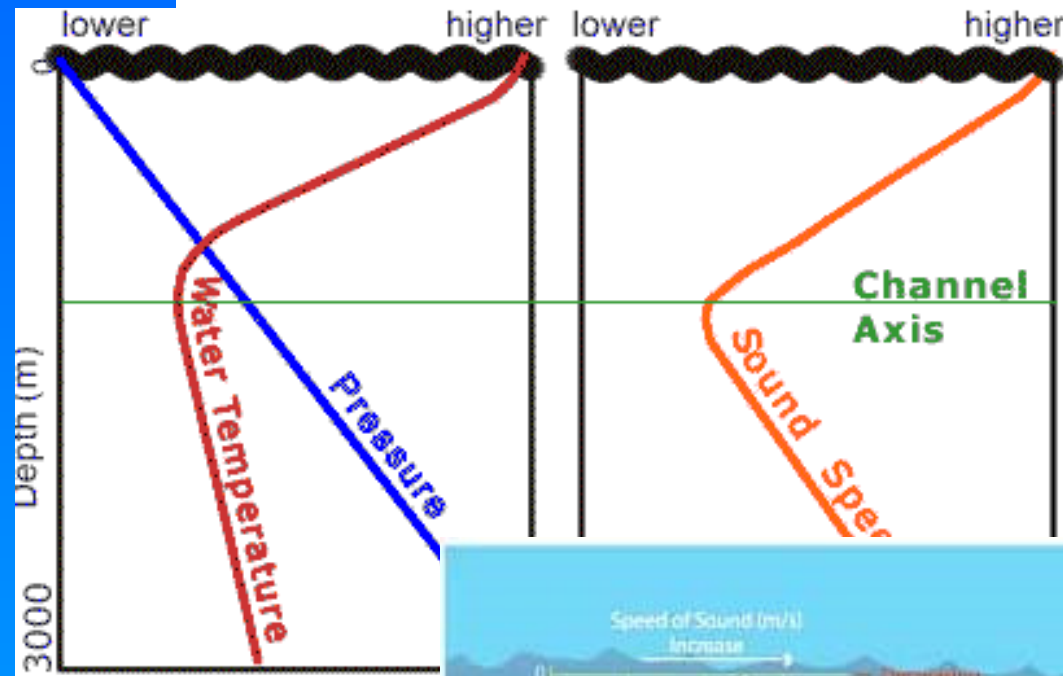
# Sound Propagation Around the Globe



(Collins et al., JASA 97(3), p. 1567)

Basis for Acoustic measurements of ocean temperatures:  
“Acoustic thermometry & Acoustic tomography..”

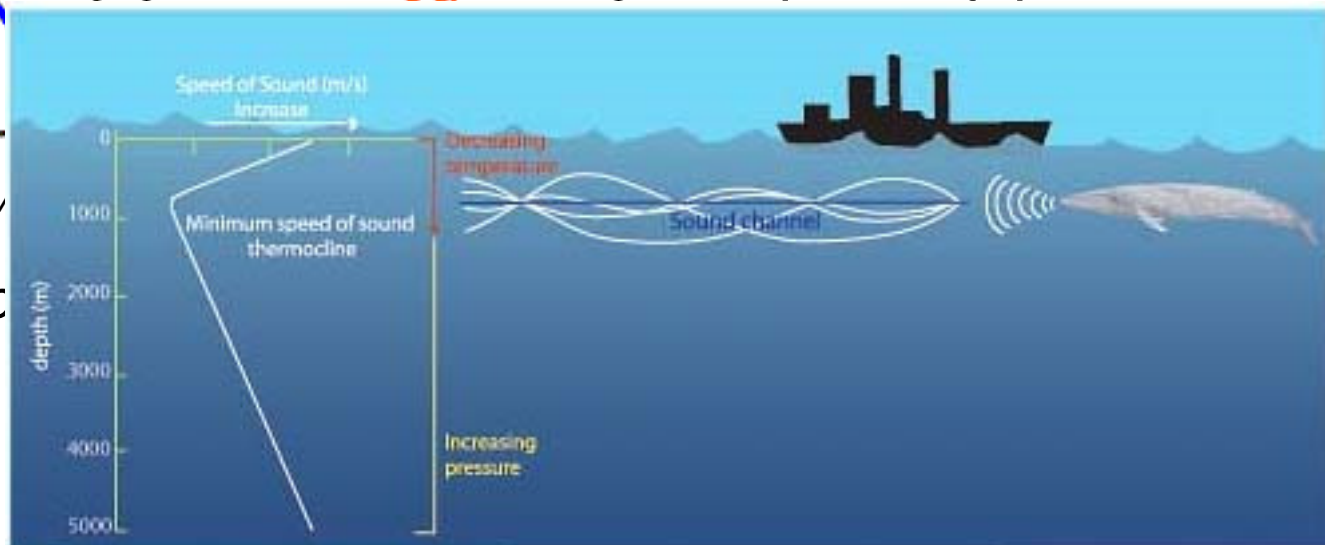
# Generic Sound Speed Structure



- Ocean sound-speed depends on:
  - Temperature, (T)
  - Salinity, (S)
  - Depth, (z)
- Speed of sound increases with all

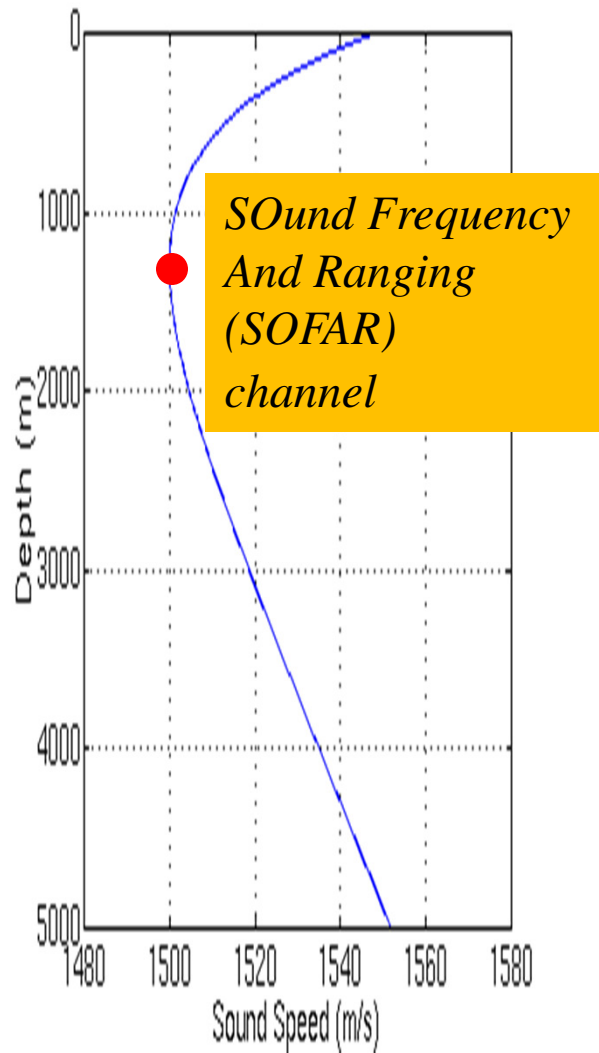
<http://www.pmel.noaa.gov/v>

*Note: In polar regions at the surface*

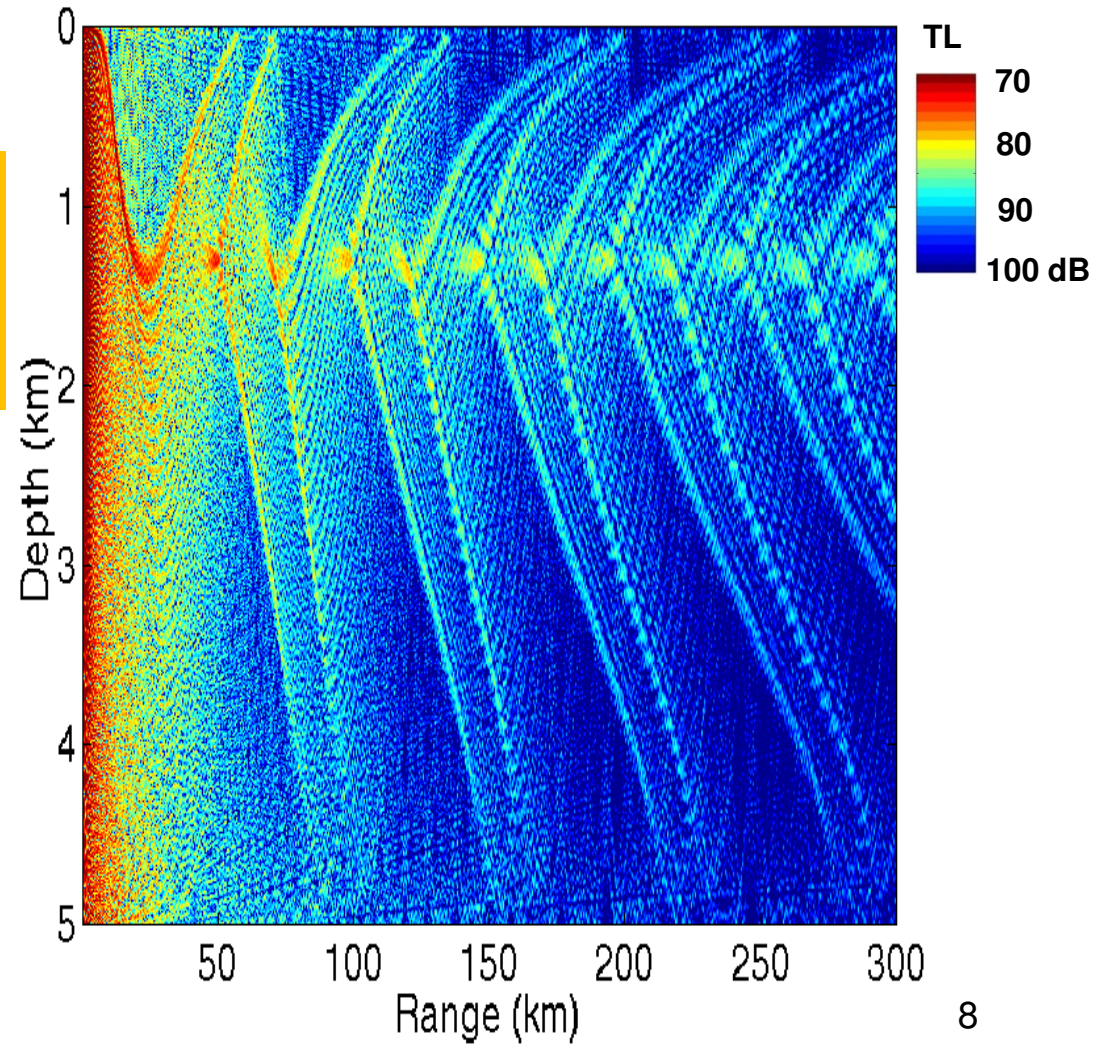


# Deep sound channel propagation

Sound Speed profile



Pressure (R = 300km, SD = 1300m)



70-80Hz- Normal Modes simulation

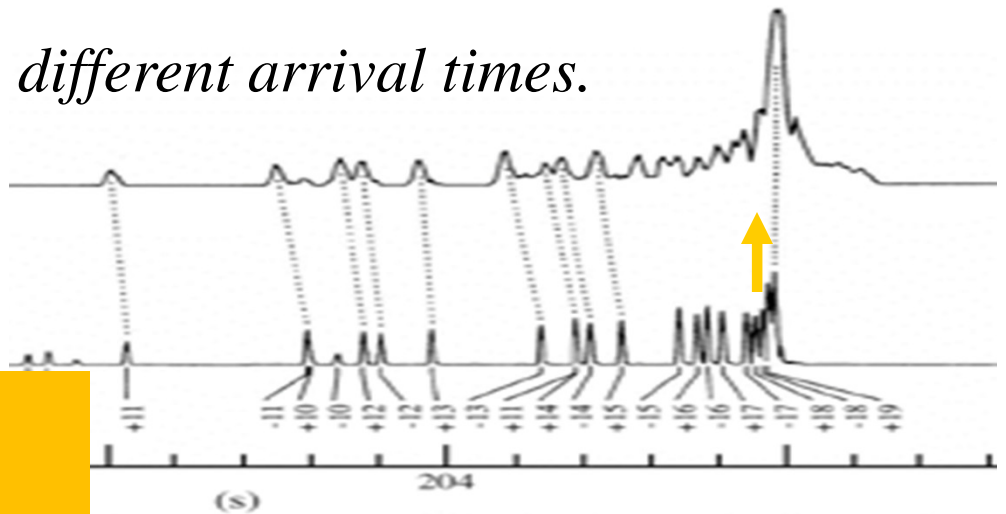
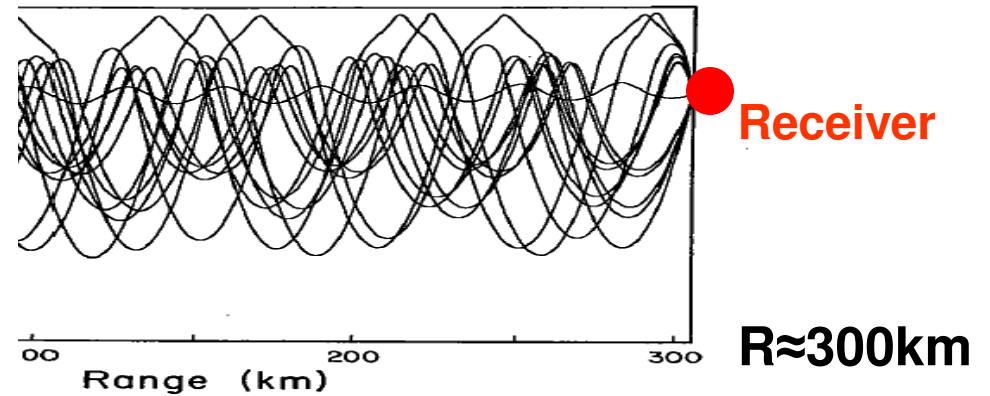


# Acoustic Thermometry

*Different paths have different (group) speeds....*



HLF-5 acoustic source (250 Hz)  
courtesy of Scripps Institution of  
Oceanography.

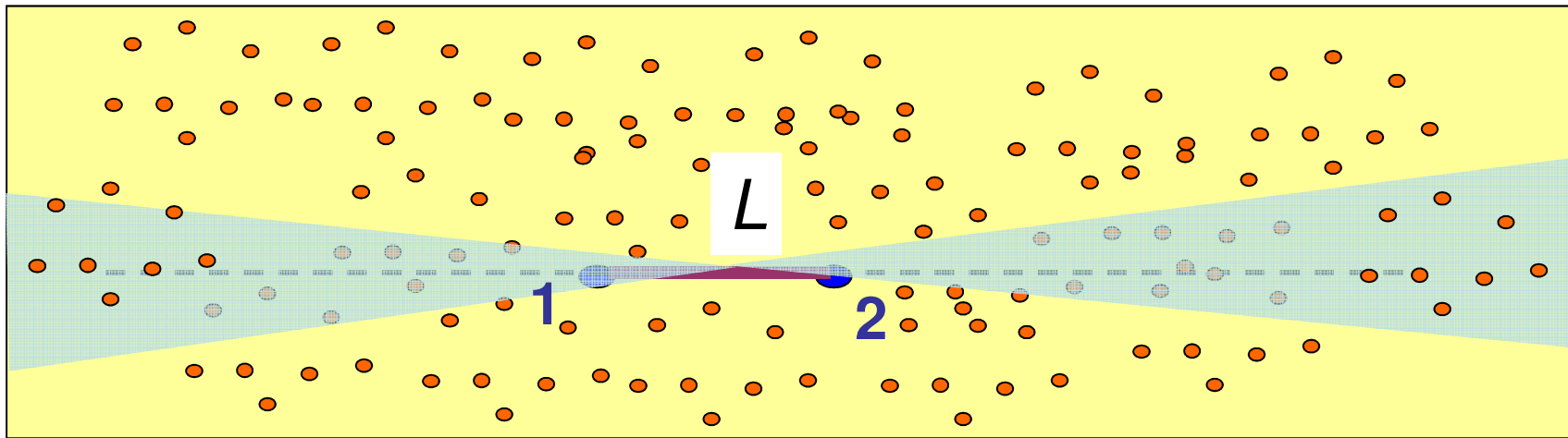


# Presentation Outline

1. **Deep water acoustics & low-frequency noise sources**
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4. **Monitoring with ambient noise: Optimization.**

# Theory: Free space

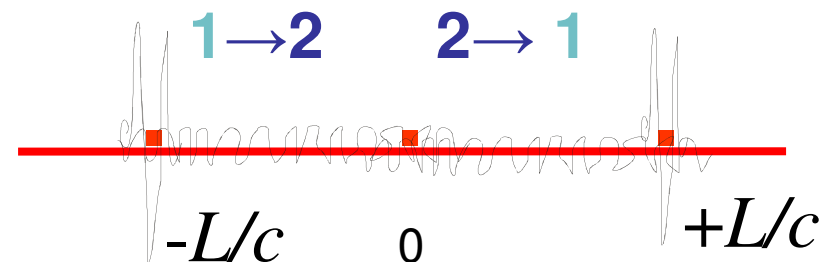
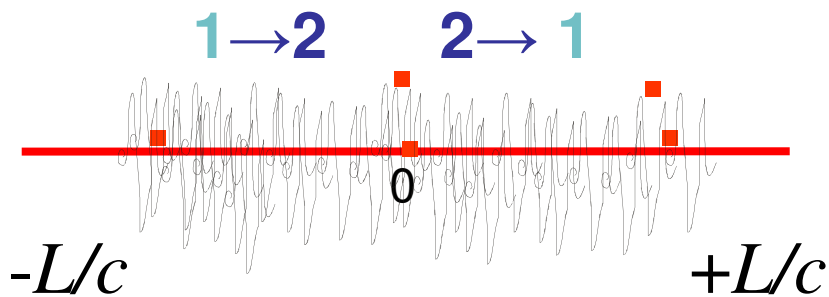
$$C_{12}(\tau) = \int_{-\infty}^{\infty} P(\mathbf{r}_1, t) P(\mathbf{r}_2, t + \tau) dt.$$



Isotropic distribution of uncorrelated random noise sources

*Short-time  $C_{12}(\tau)$*

*Average  $C_{12}(\tau)$  over duration  $T$*



$$SNR = \frac{\text{Coherent Noise amplitude}}{\text{Incoherent Noise amplitude}}$$

$$\approx \sqrt{(B \cdot T) / (k \cdot L)}$$

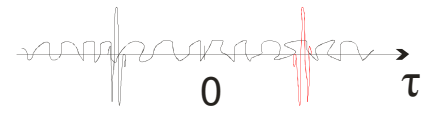
Arrival time rms error:

$$\sigma_t = [\omega_0 \sqrt{SNR}]^{-1}$$



# Practical Limitations...

$$C_{12}(t) \approx F(t) \otimes (G_{12}(t) - G_{21}(-t))$$



2→1      1→2

Limiting Factors:

1. Spatial Distribution
2. Temporal Distribution
3. Power spectrum
4. Source Directivity

due to the noise sources

5. Attenuation

due to the medium

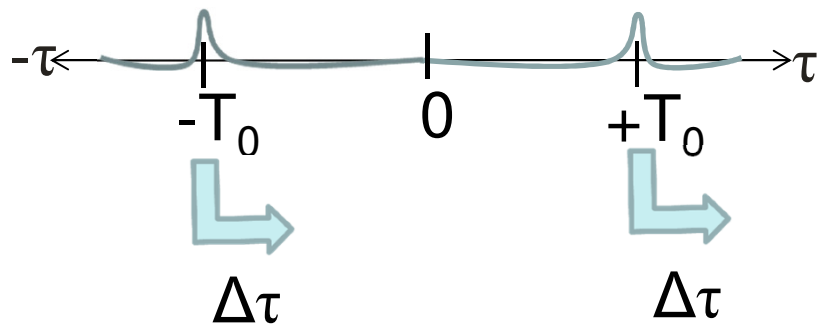
6. Fluctuations

$$T_{\text{statistical}} < T_{\text{recording}} < T_{\text{environmental}}$$

On short time-scales ( $< T_{\text{environmental}}$ ), the cross-correlation process is stationary

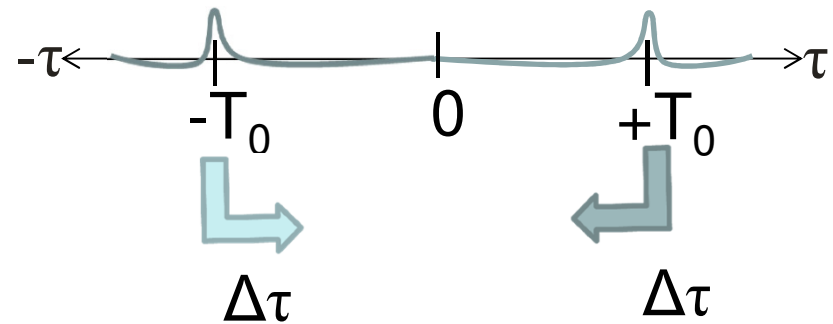
# Passive Ocean monitoring ?

Anti-symmetric time shift



Due to “non-reciprocal”  
environmental changes  
OR clock-drift between  
receivers 1 and 2

Symmetric time shift



Due to “reciprocal”  
environmental changes

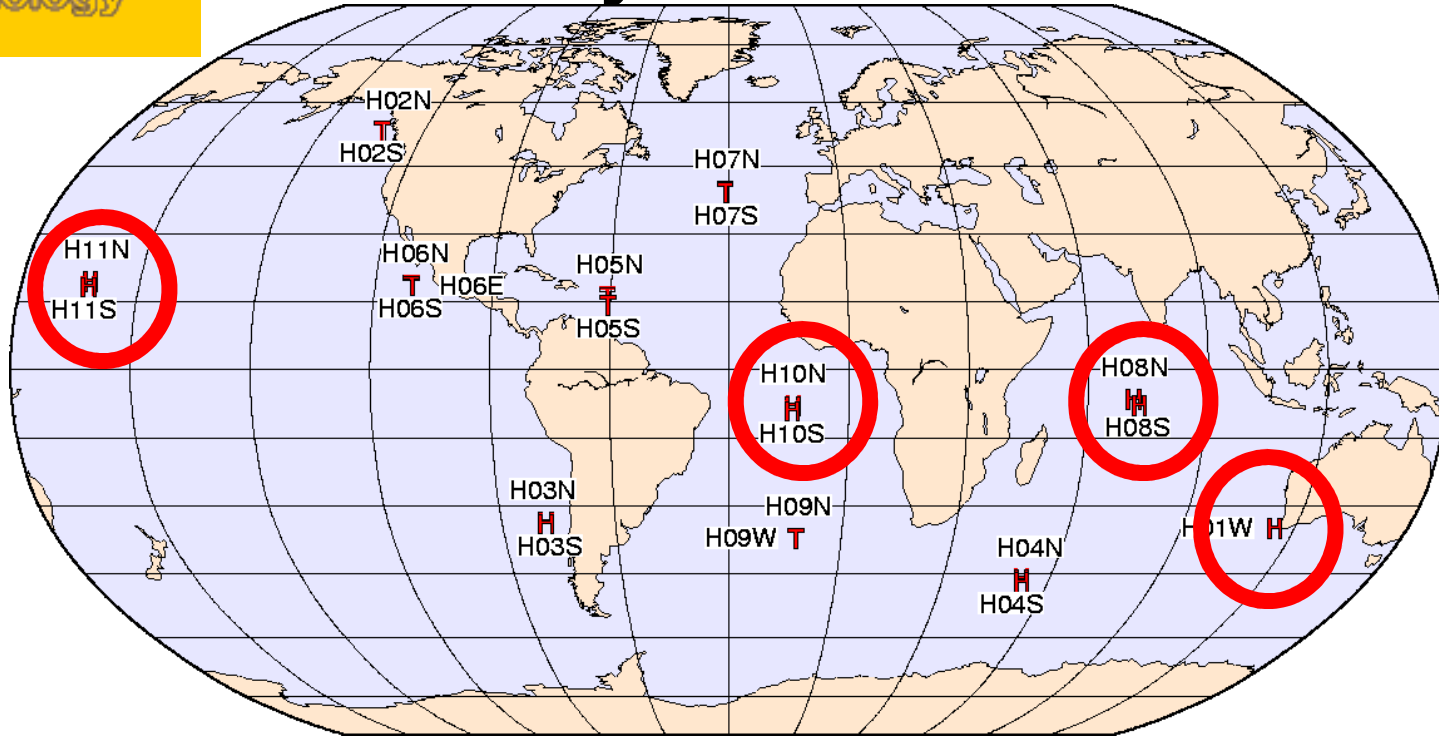
$$\Delta c / C_0 = (\Delta \tau / T_0)$$

# Presentation Outline

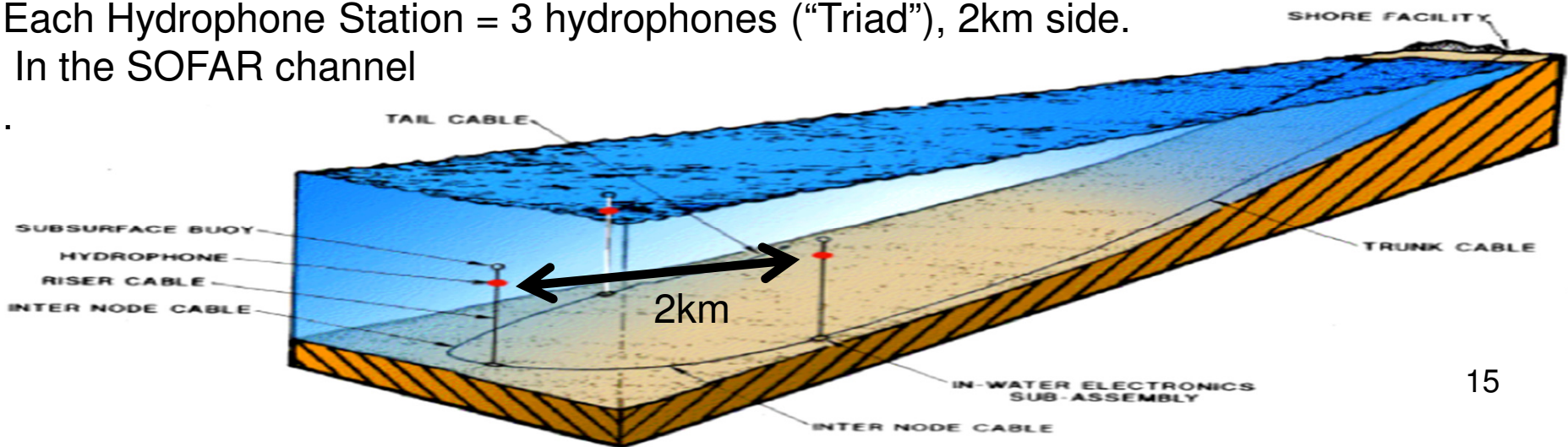
- 1. Deep water acoustics & low-frequency noise sources**
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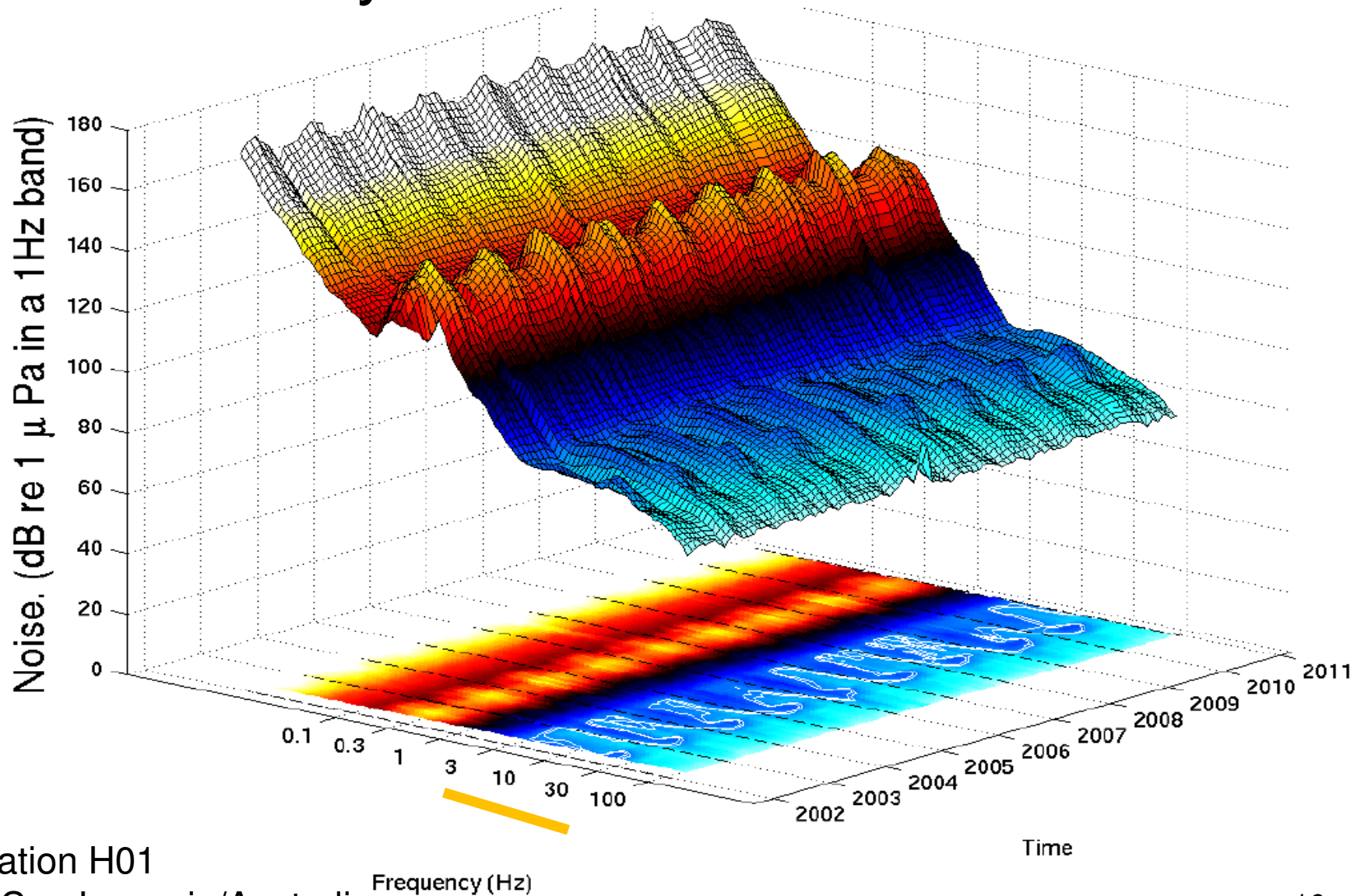
# IMS Hydroacoustic Station



Each Hydrophone Station = 3 hydrophones ("Triad"), 2km side.  
In the SOFAR channel



# Ambient noise measured by hydroacoustic station



Station H01  
@Cap Leeuwin/Australia

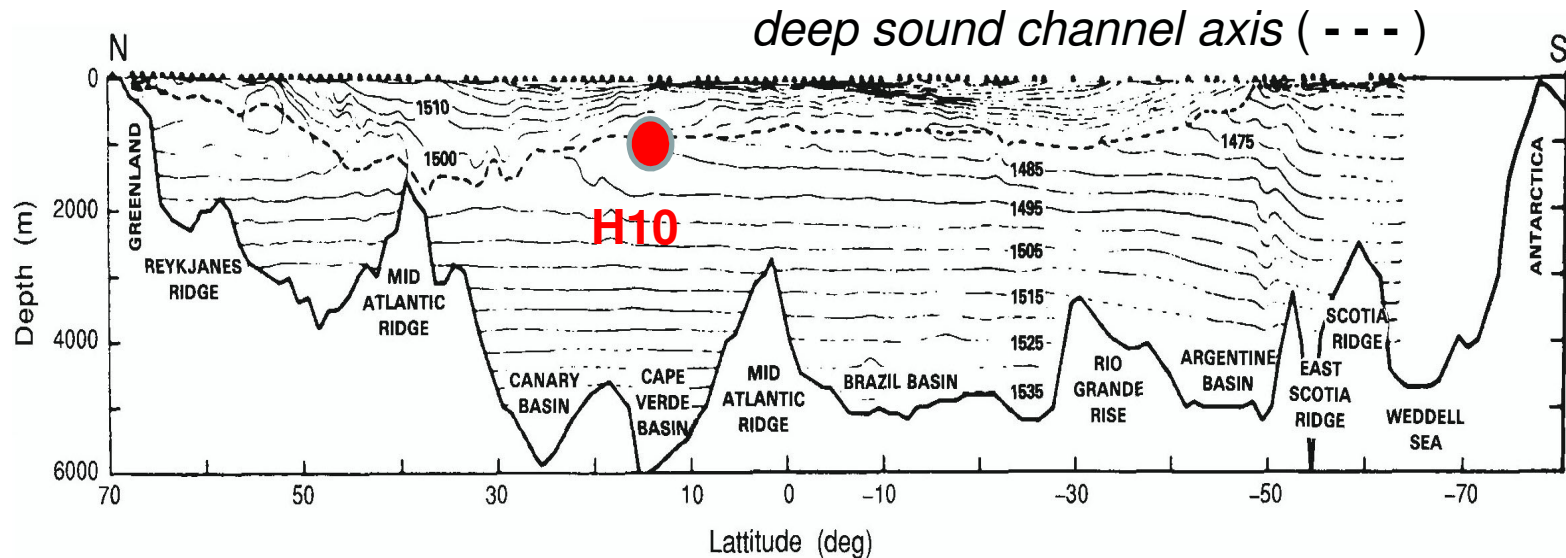
Frequency (Hz)

Time

*PRIOR et al., 2010*

# Global Sound Speed Structure

North-South Atlantic along 30.5°N

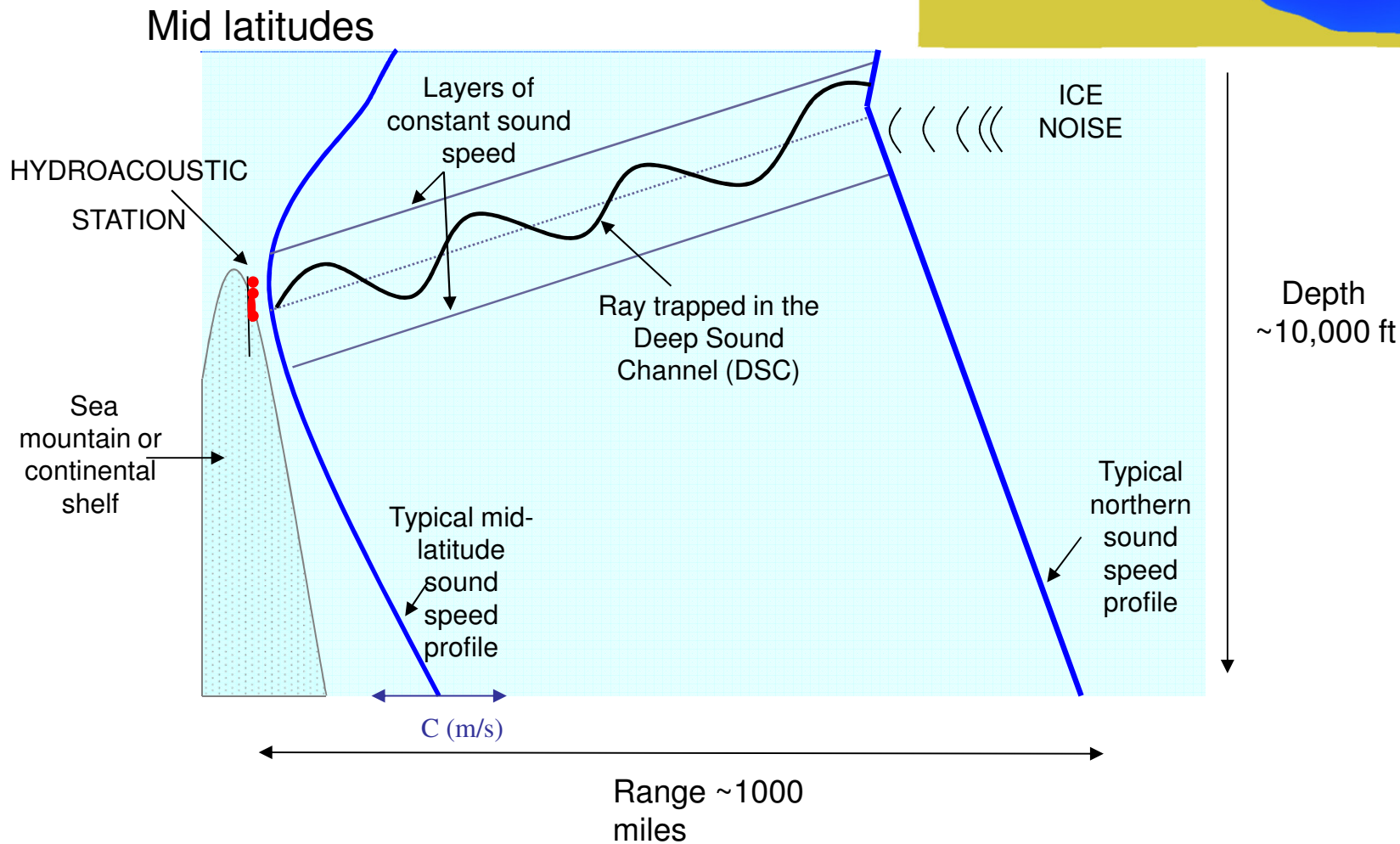
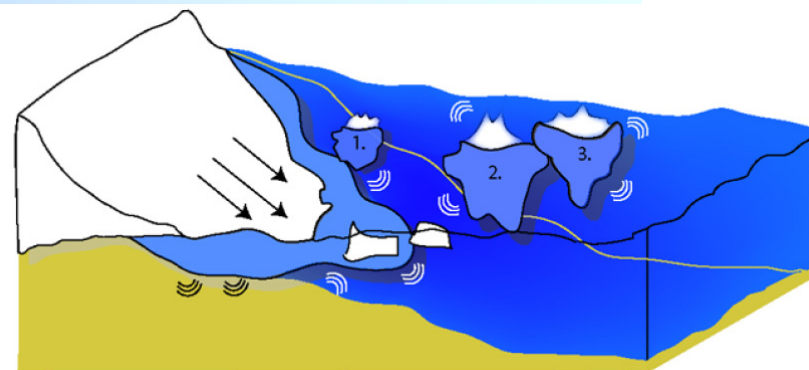


*From Northrup and Colborn, JGR, 1974*

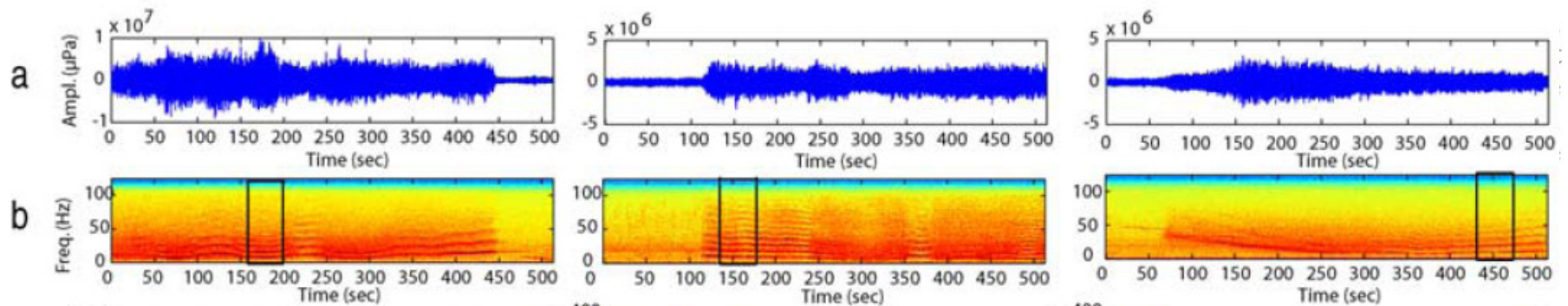
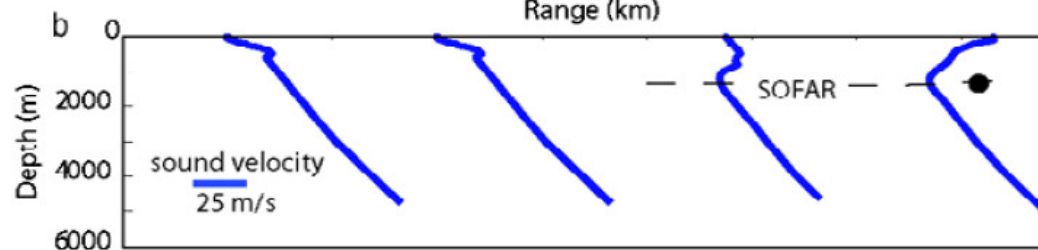
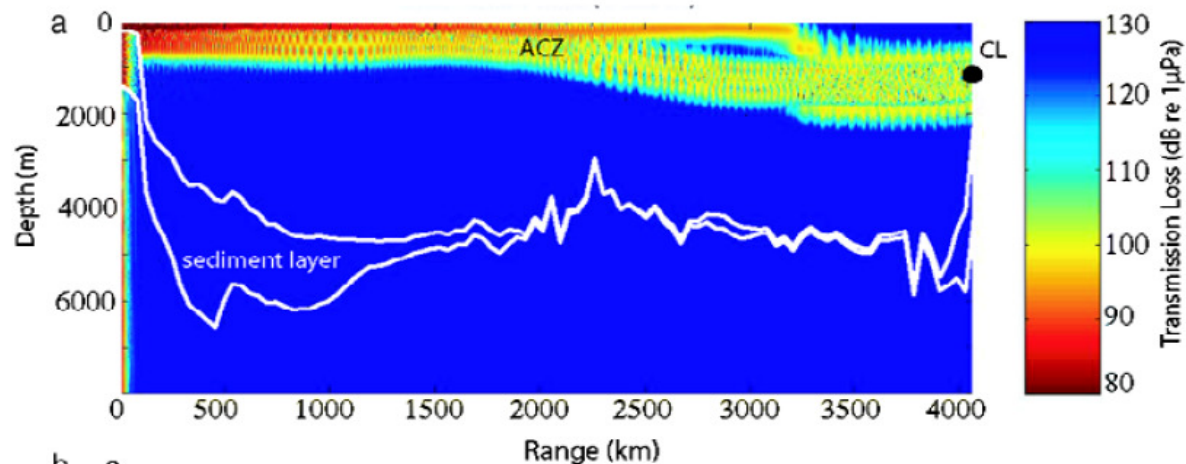
- Variability of the upper-ocean (<1km) sound speed structure.
- Stability of the deep isothermal layer.
- Axis of the deep sound channel becomes shallower towards both poles and eventually reaches the surface.



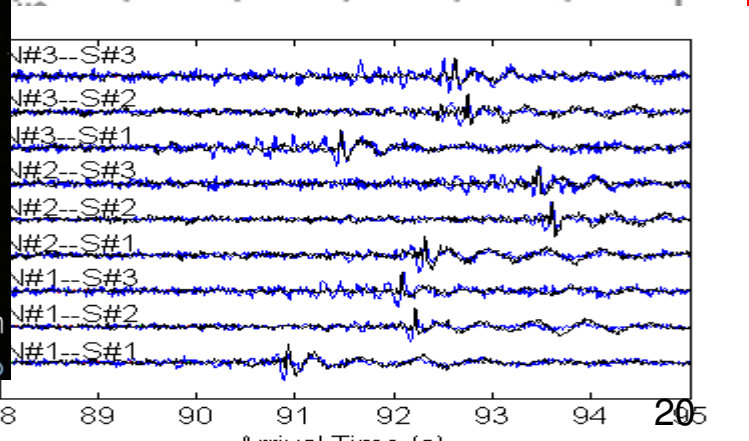
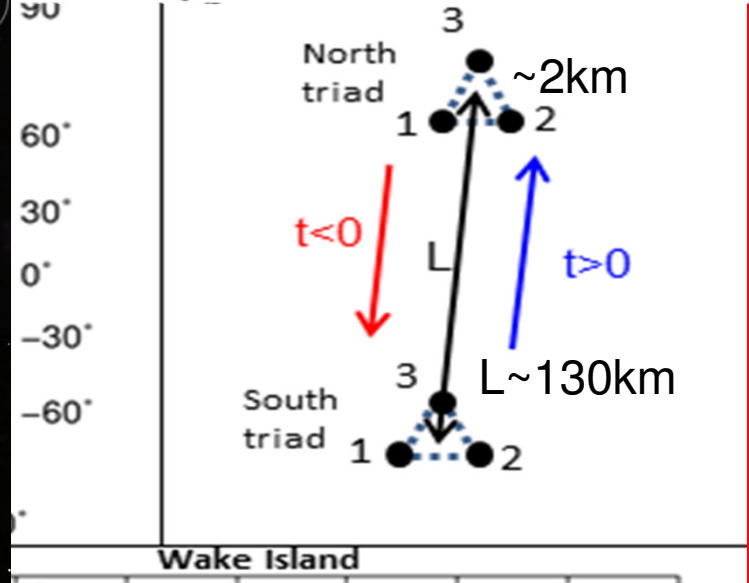
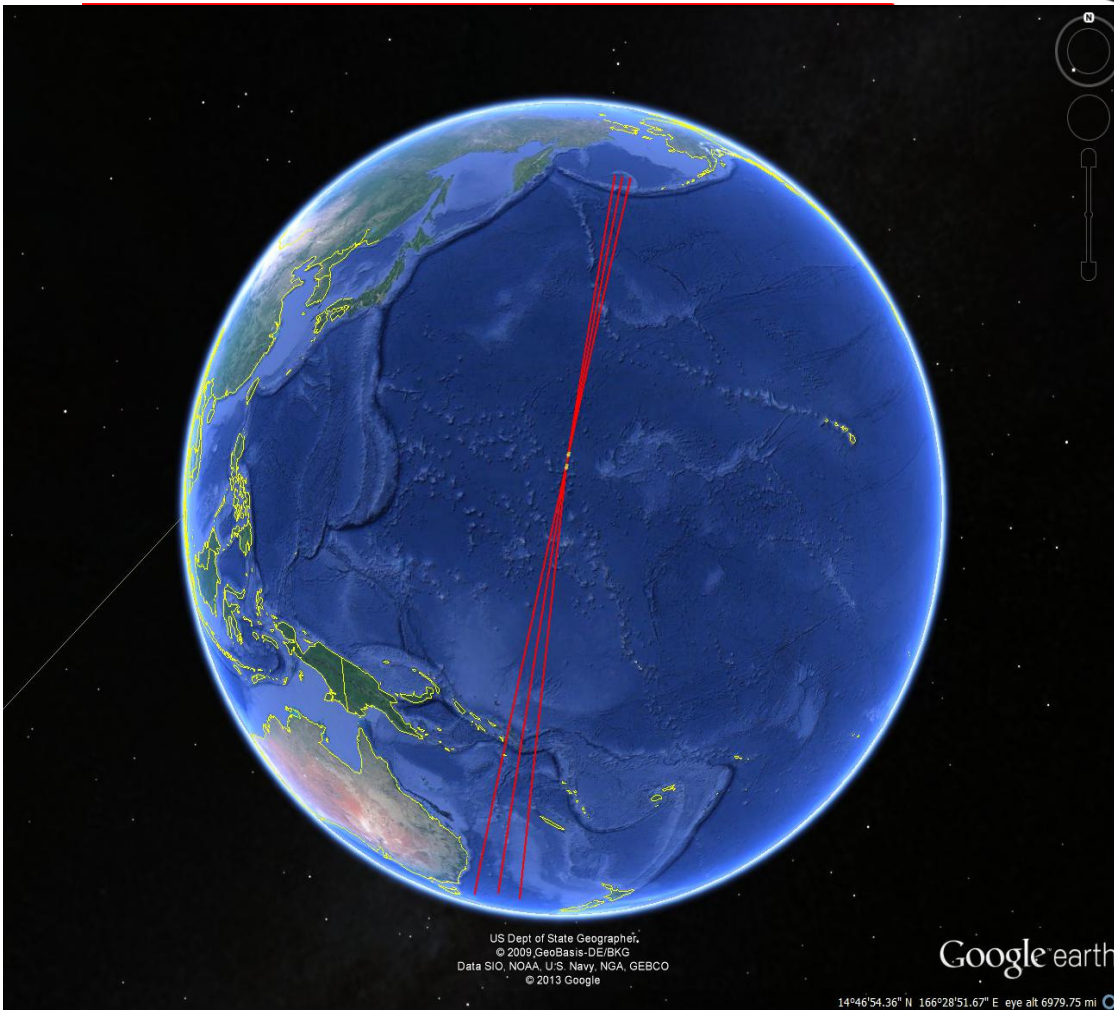
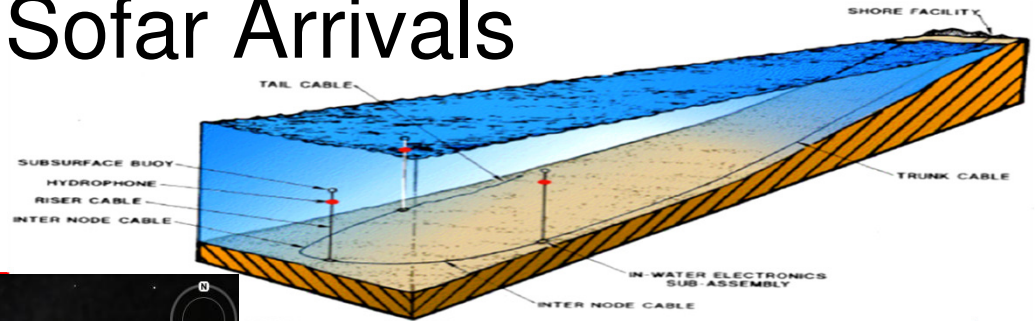
# Ice noise propagation



# Ice-noise long-range propagation



# Coherent Sofar Arrivals

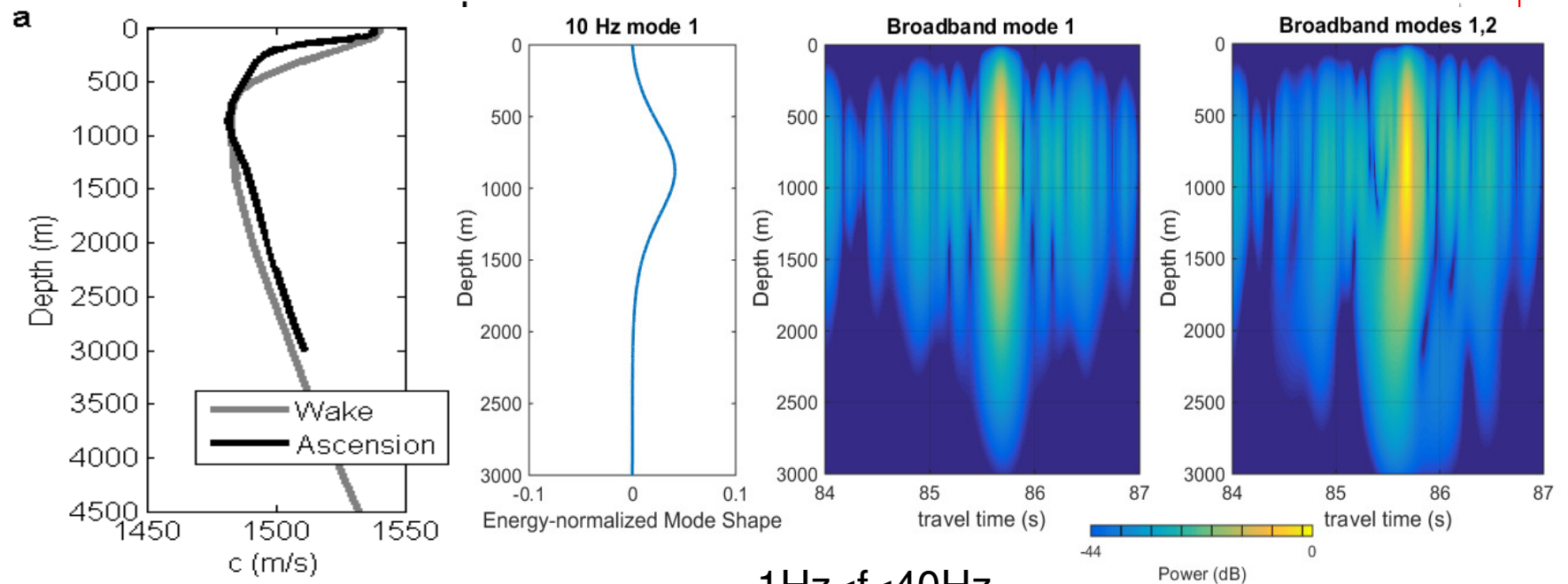
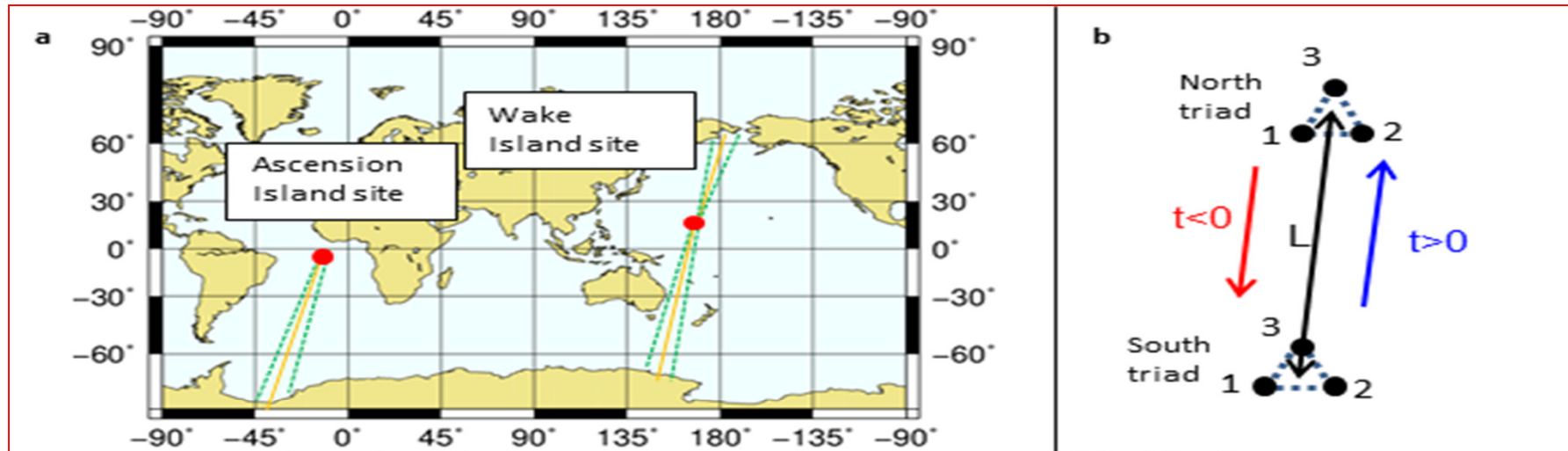


1 year averaging,  $1\text{Hz} < f < 40\text{Hz}$

Woolfe et al., GRL 2015



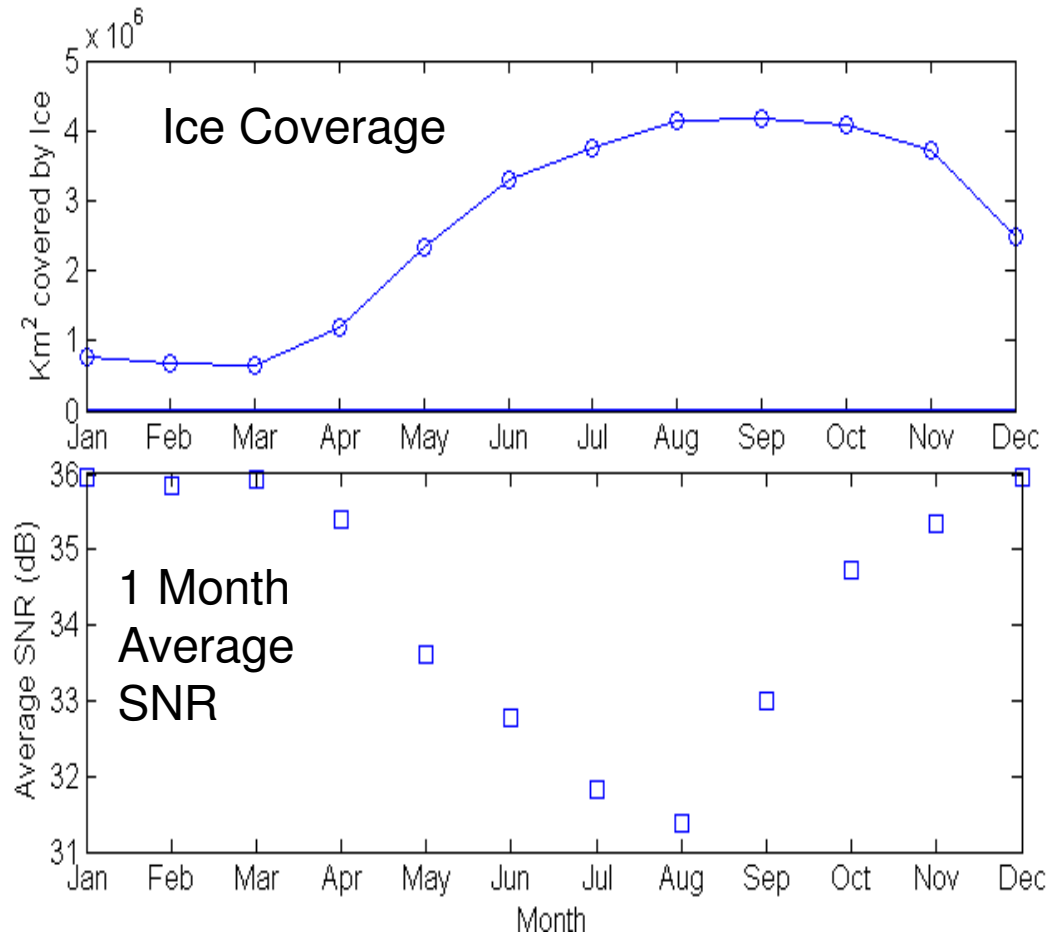
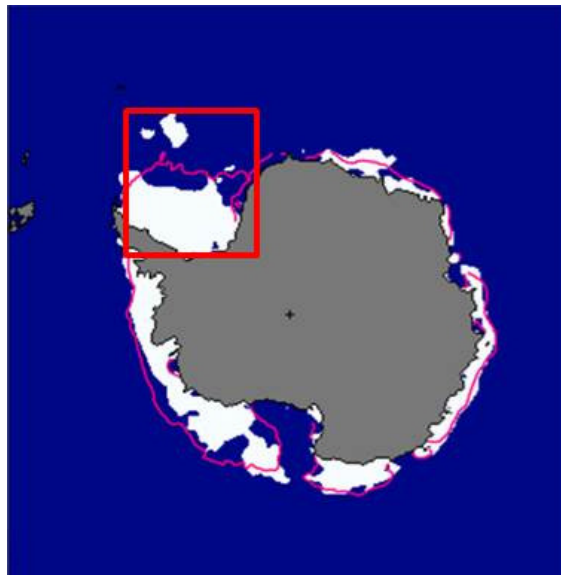
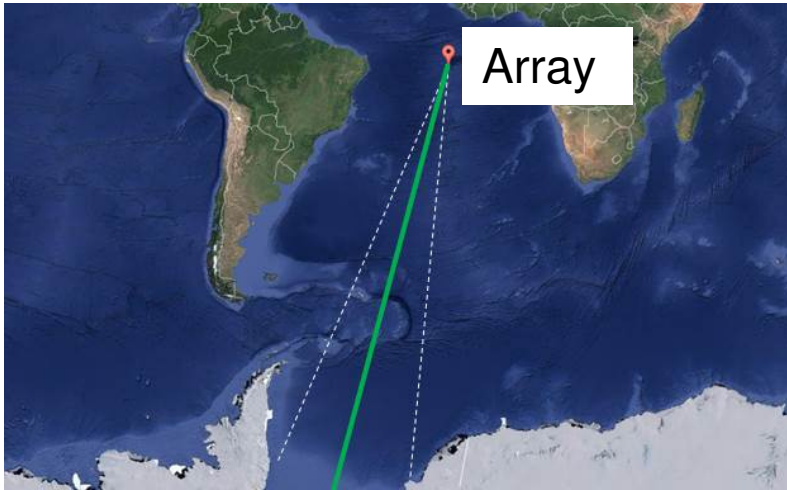
# SOFAR arrival between Triads



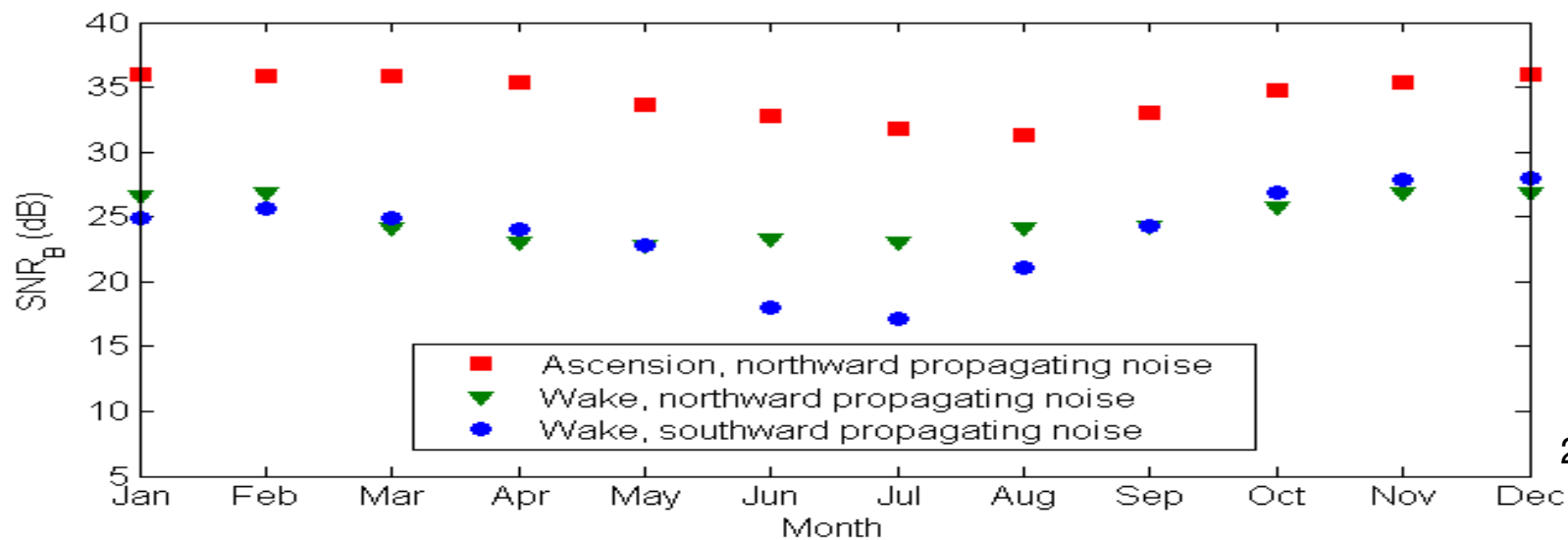
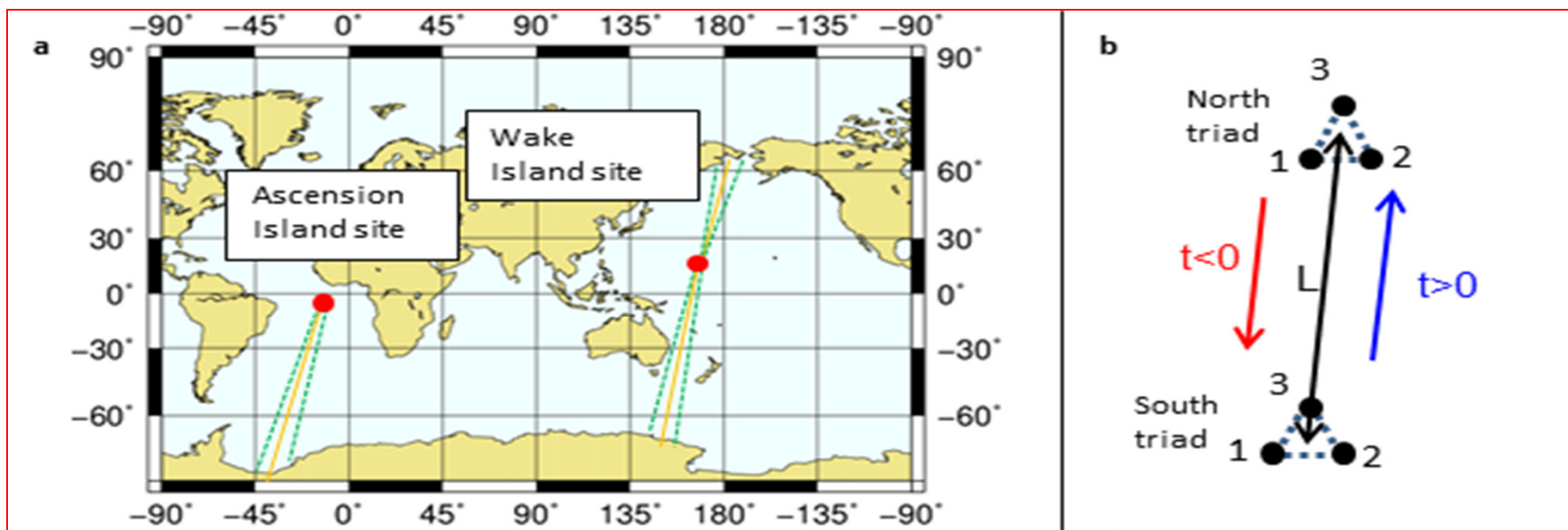
1 Hz < f < 40 Hz



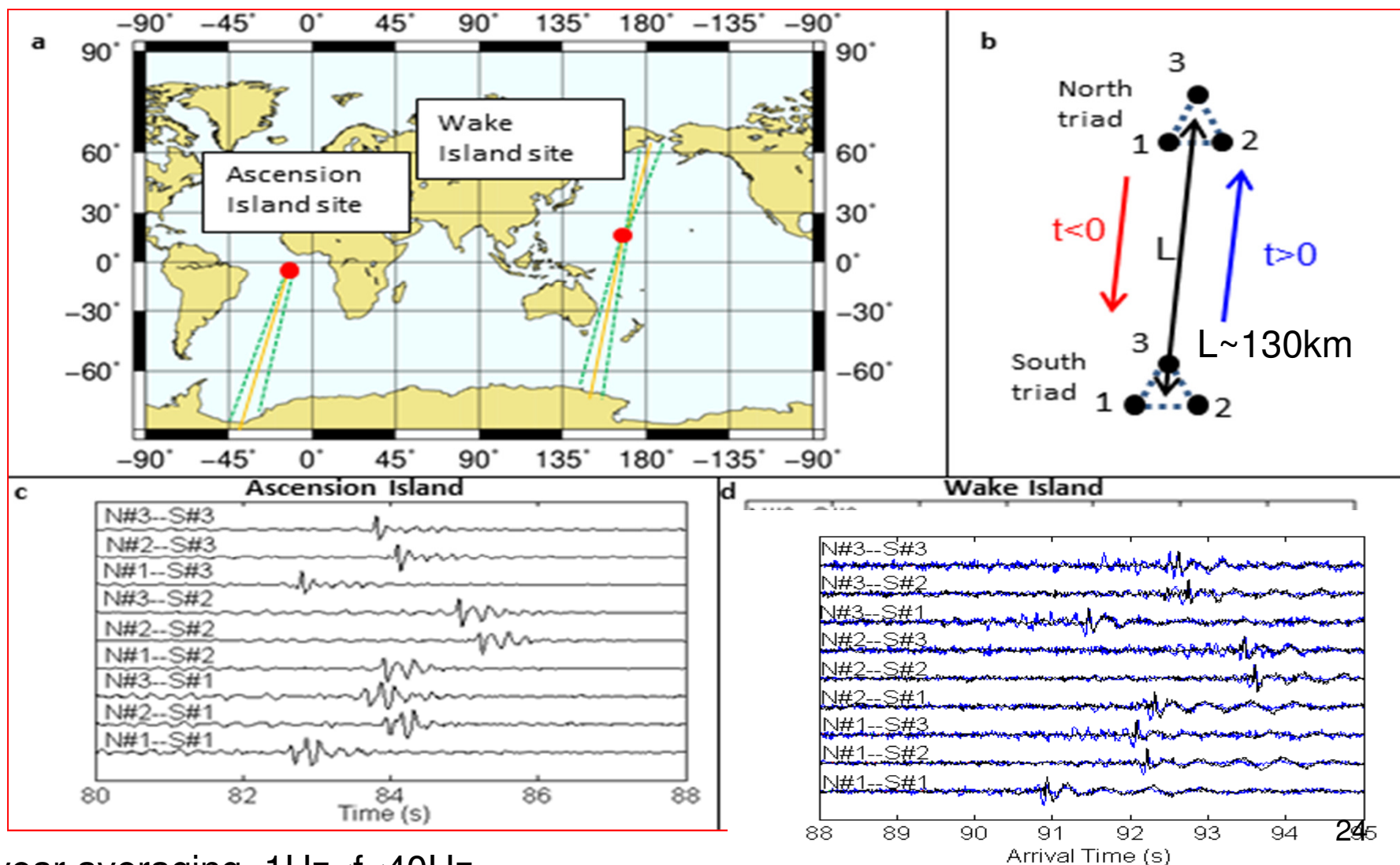
# Polar origin of the coherent noise



# Polar origin of the coherent noise

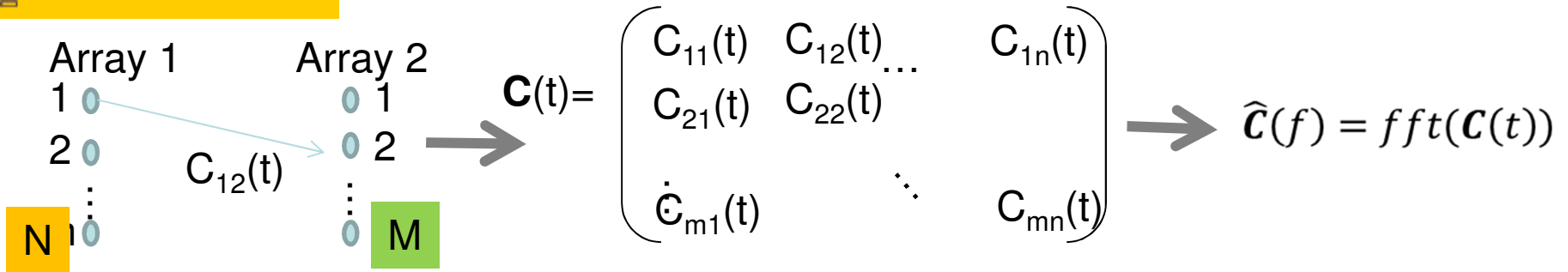


# Coherent Sofar Arrivals



1 year averaging,  $1 \text{ Hz} < f < 40 \text{ Hz}$

# Beamforming of Noise Correlations

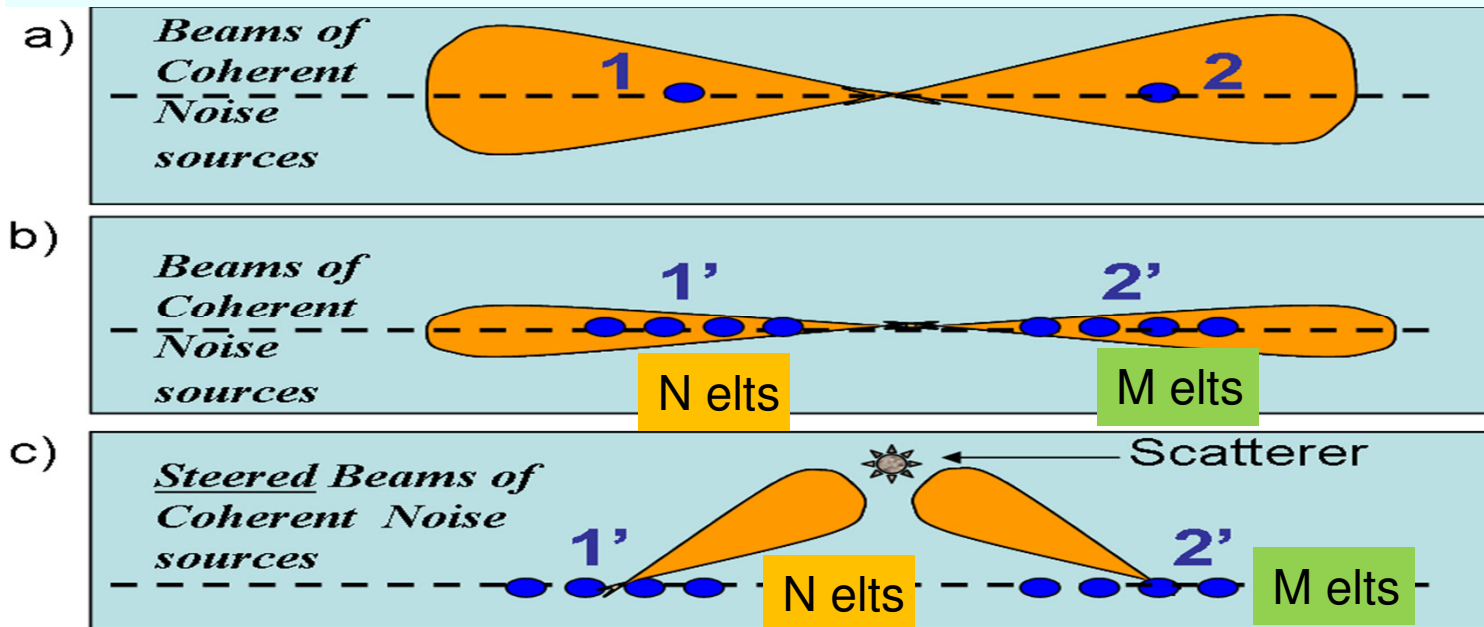


$$\hat{\mathbf{B}}(f) = \mathbf{W}_1^H(f) \hat{\mathbf{C}}(f) \mathbf{W}_2(f)$$

Steering vectors (i.e. weights)

Provides a theoretical SNR gain of a factor of  $\sqrt{NM}$

Goal: Increase SNR of cross-correlations using sensor arrays

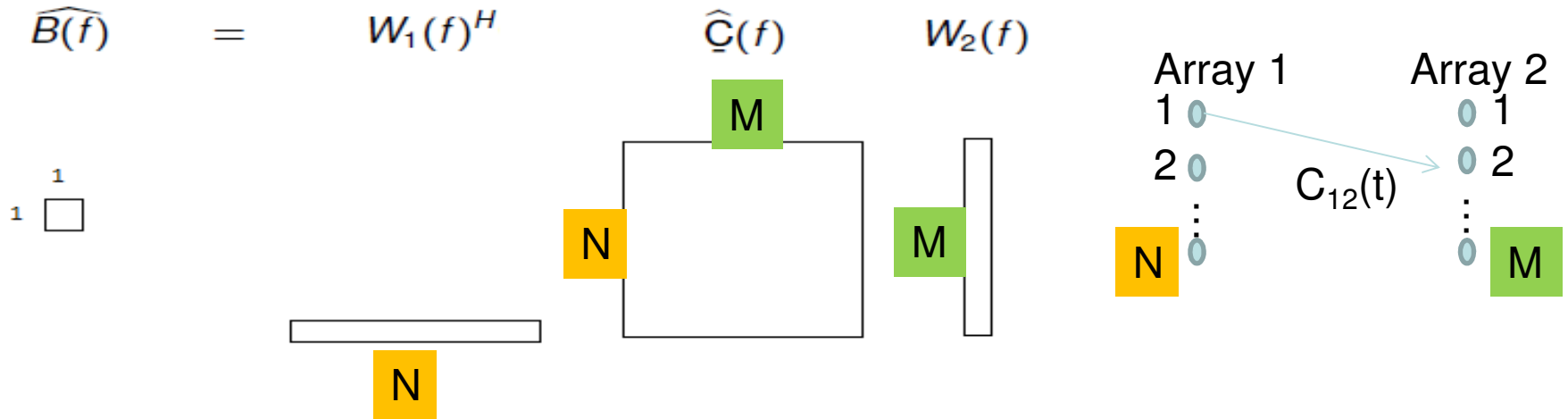


(Leroy et. al. JASA, 2012)



# Beamforming formulation

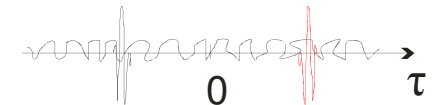
$$\widehat{B}(f) = W_1^H(f) \widehat{C}(f) W_2(f)$$



- $\widehat{B}(f)$  beamformer output (scalar)
- $\widehat{C}(f)$  noise cross-covariance matrix (size  $N \times M$ )
- $W_1(f)$  array weight vector (size  $N \times 1$ )
- $W_2(f)$  array weight vector (size  $M \times 1$ )

Advantage: enhances the SNR (“Signal” over “Noise” Ratio)

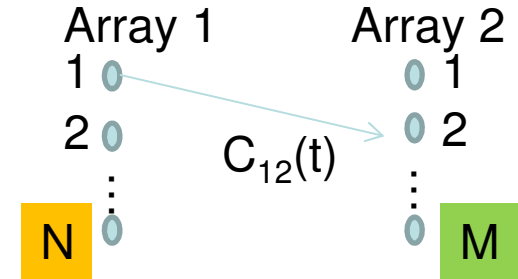
$$\text{SNR} = \frac{\text{Coherent Noise amplitude}}{\text{Incoherent Noise amplitude}} \approx \sqrt{(B \cdot T) \cdot (N \cdot M)}$$



# Beamforming Noise Correlations for monitoring

$$\hat{B}(f) = W_1^H(f) \hat{C}(f) W_2(f)$$

Cross-correlations obtained using Short averaging time



How to select the Array steering vectors  $W_1(f)$  and  $W_2(f)$  ?

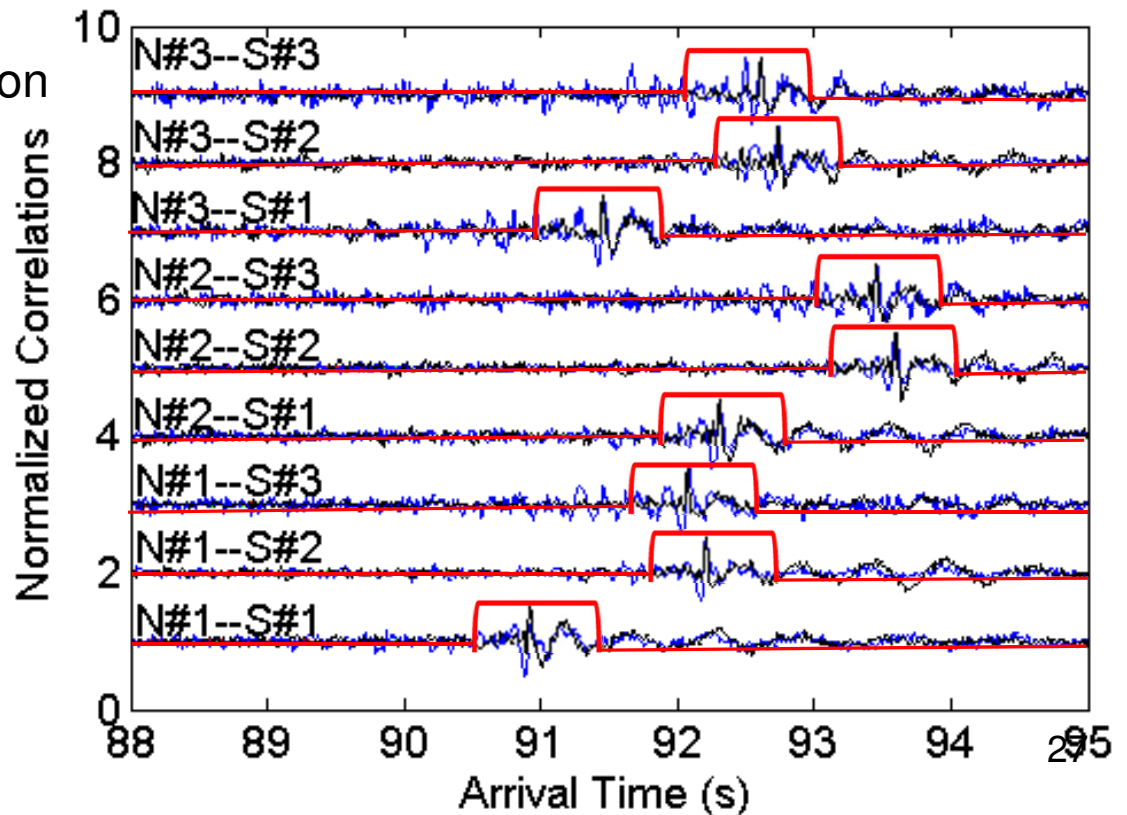
- Model-based (Rays- DBF, Mode)
- Data-derived

Using Singular Value Decomposition

$$\hat{C}_{\text{ref}}(f) = U(f) \Sigma(f) V(f)^H$$

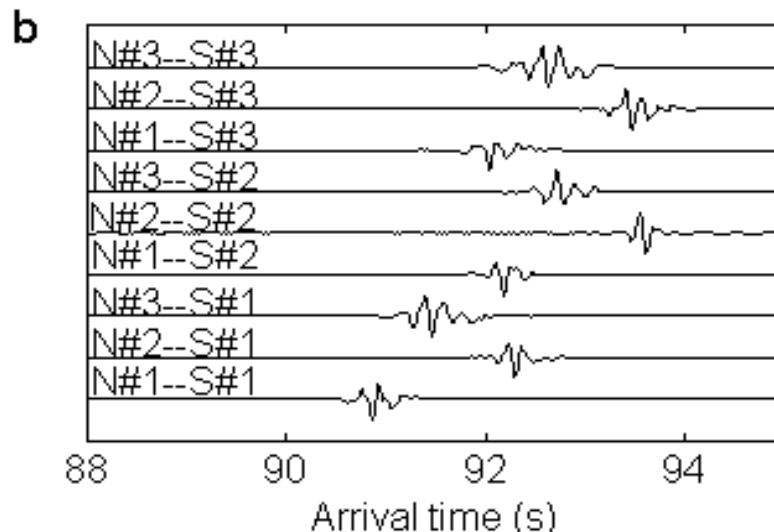
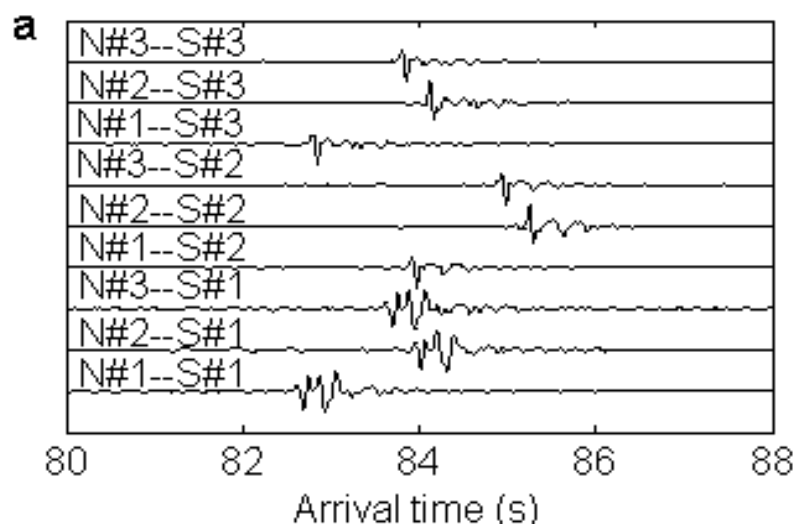
$$\sim \sigma_1 U_1(f) V_1(f)^H$$

$\uparrow$   
 $W_1(f)$ 
 $\uparrow$   
 $W_2(f)$

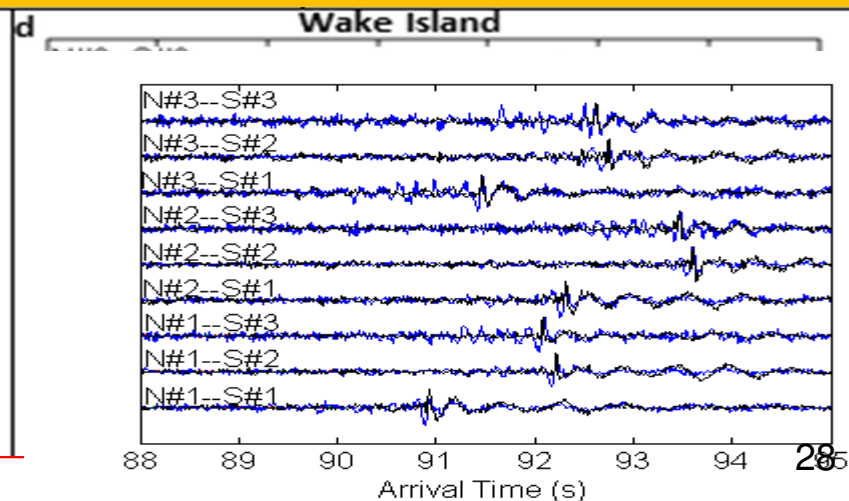
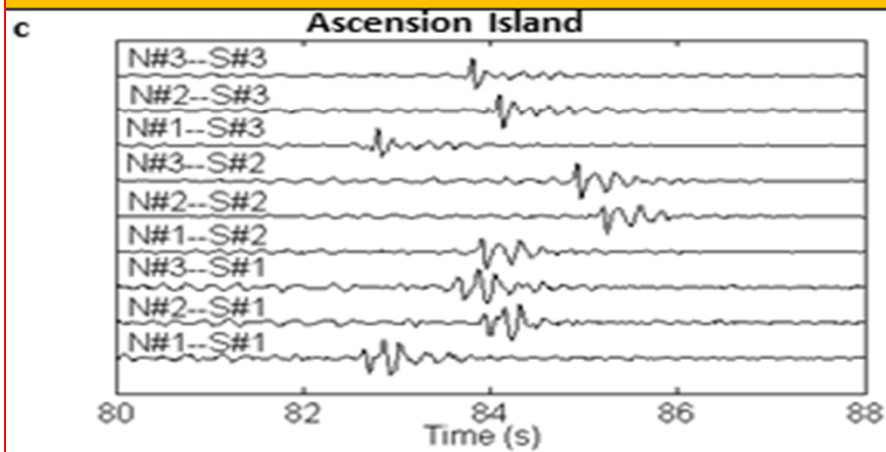


# Spatio-temporal filter for Coherent Arrivals

inverse Fourier transform of the first principal component  $\sigma_1(f)W_1(f)W_2(f)^H$



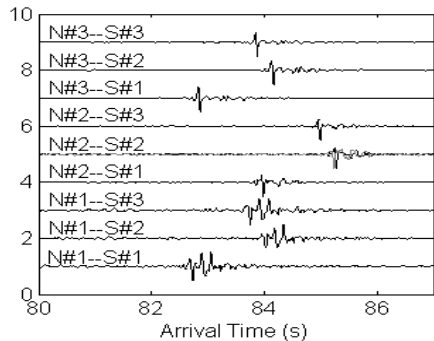
## Original 1 year average Correlations



1 year averaging,  $1\text{Hz} < f < 40\text{Hz}$

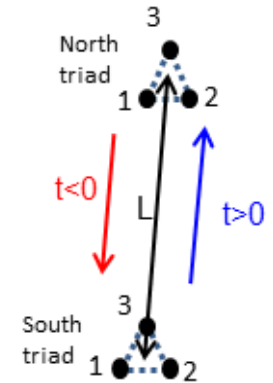
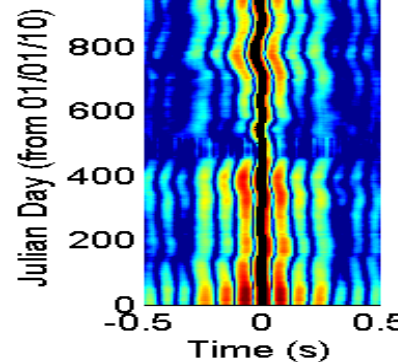
# Passive thermometry of the deep ocean

Result: Trackshifts in arrival time of beamformed cross-correlations (1 week average) over multiple years



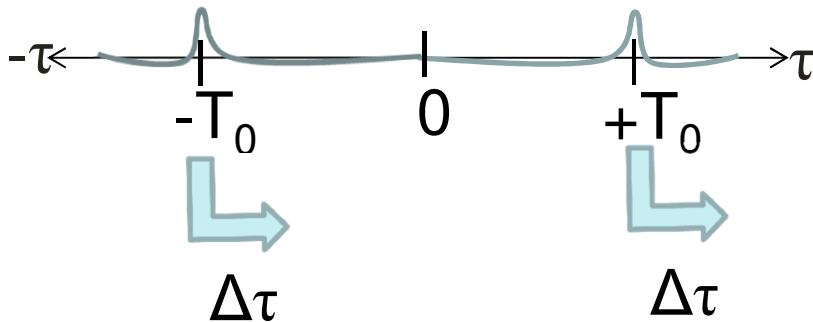
Adaptive Beamformer

Beamformer Output B(t)



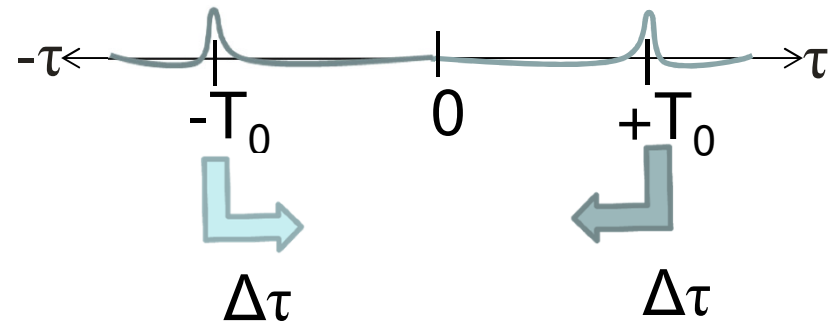
$$\hat{B}(f) = W_1^H(f) \hat{C}(f) W_2(f)$$

Anti-symmetric time shift



“non-reciprocal” changes OR clock-drift between receivers 1 and 2

Symmetric time shift

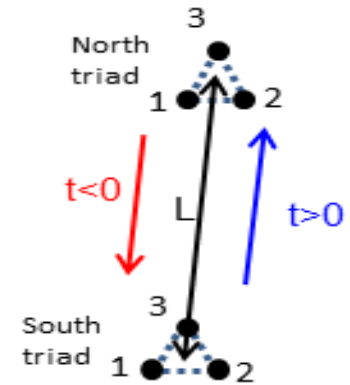
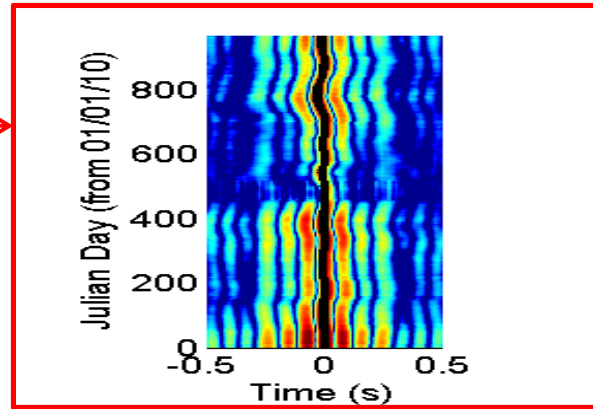
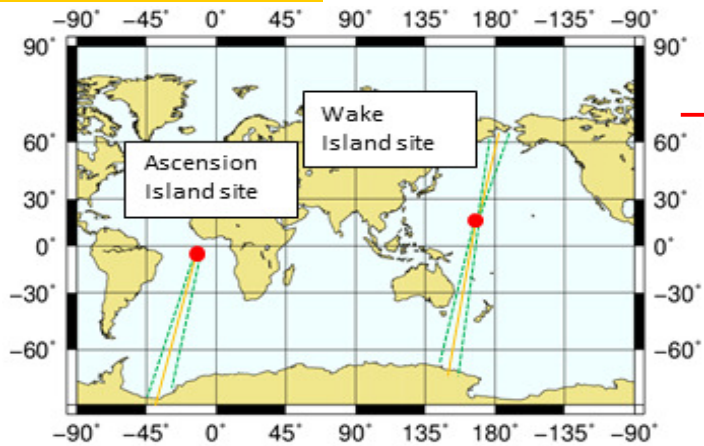


reciprocal” changes

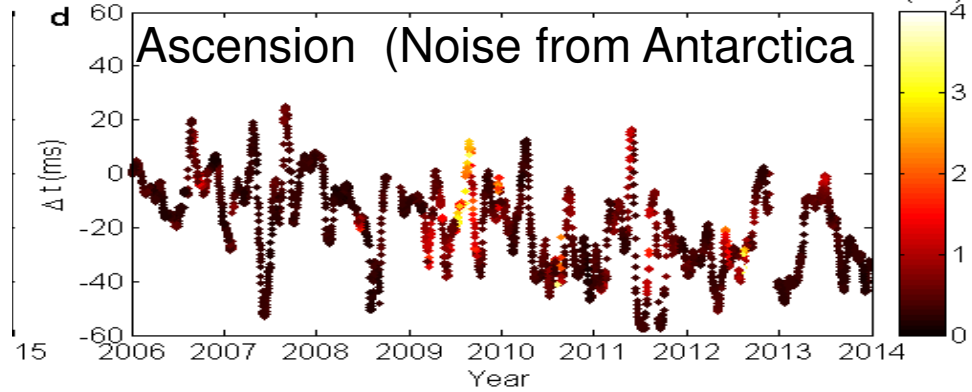
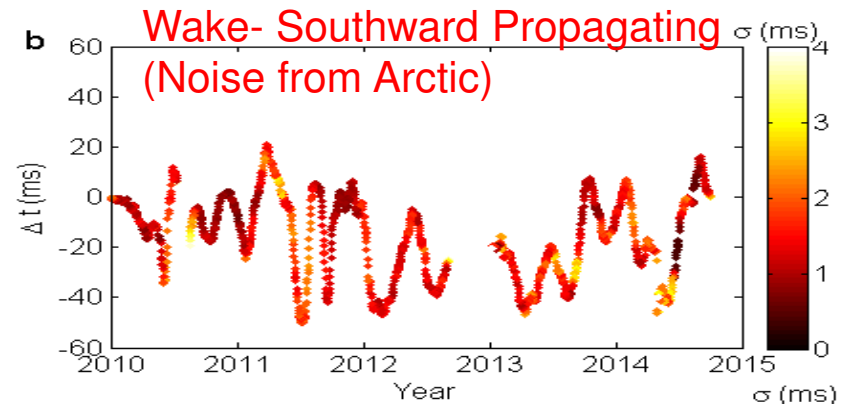
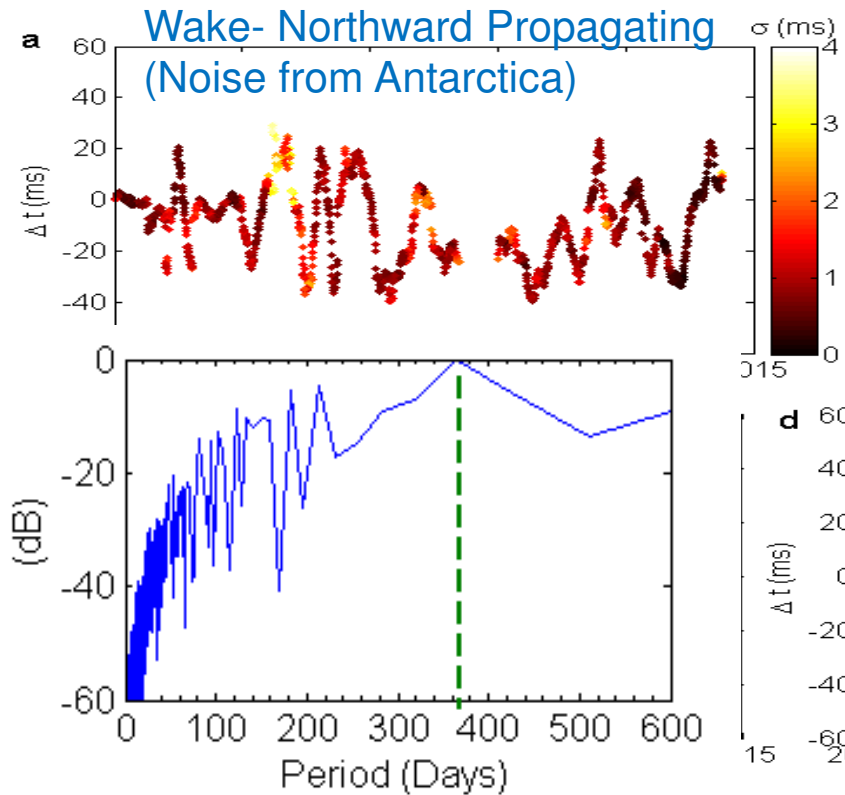
$$\Delta c / C_0 = (\Delta \tau / T_0)$$



# Passive thermometry of the deep ocean



$$\sigma_t = [\omega_0 \sqrt{SNR}]^{-1}$$

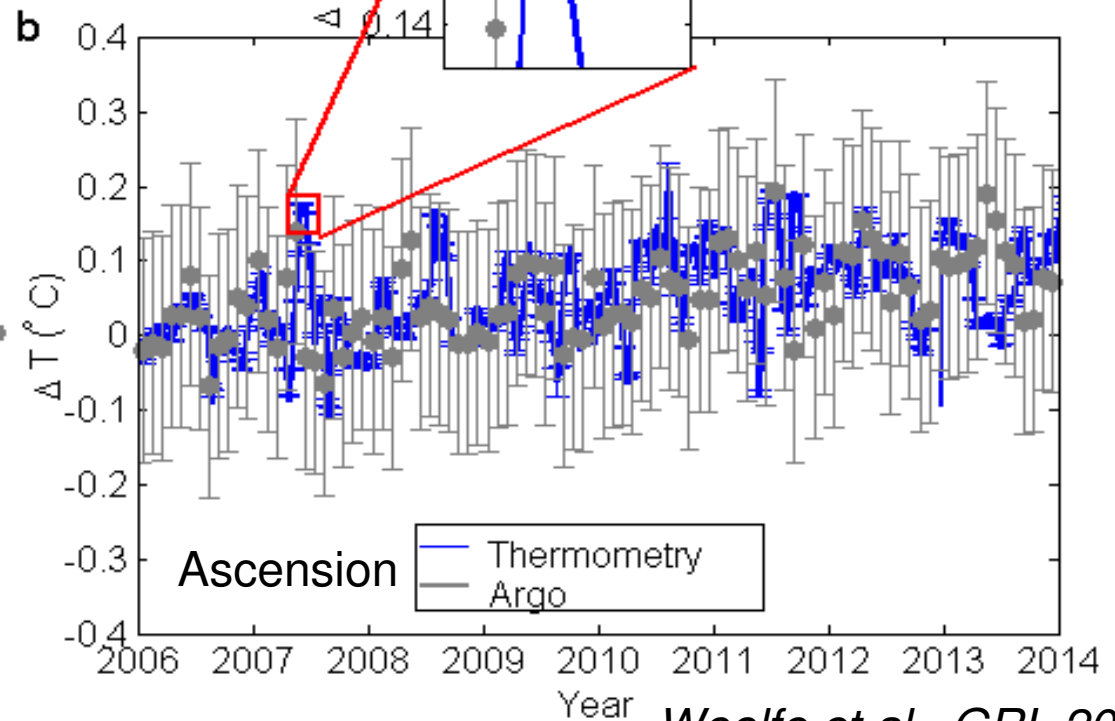
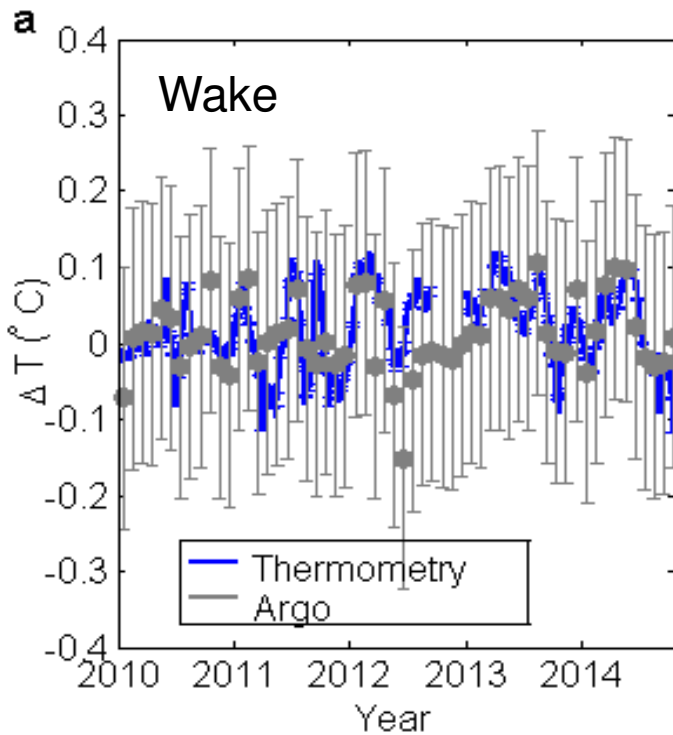
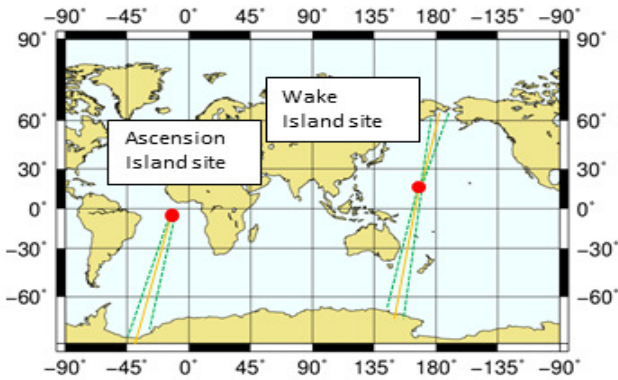


# Passive thermometry of the deep ocean

Comparison with ARGO temperature data (Roemmich, D. and J. Gilson, 2009)

Estimating Temperature changes:  $\frac{\Delta t}{t_0} = -\frac{\Delta c}{c_0} = \alpha \Delta T \left(1 + \frac{\mu\beta}{\alpha}\right)$  (Munk et al. 2009)

From baseline sound speed profile

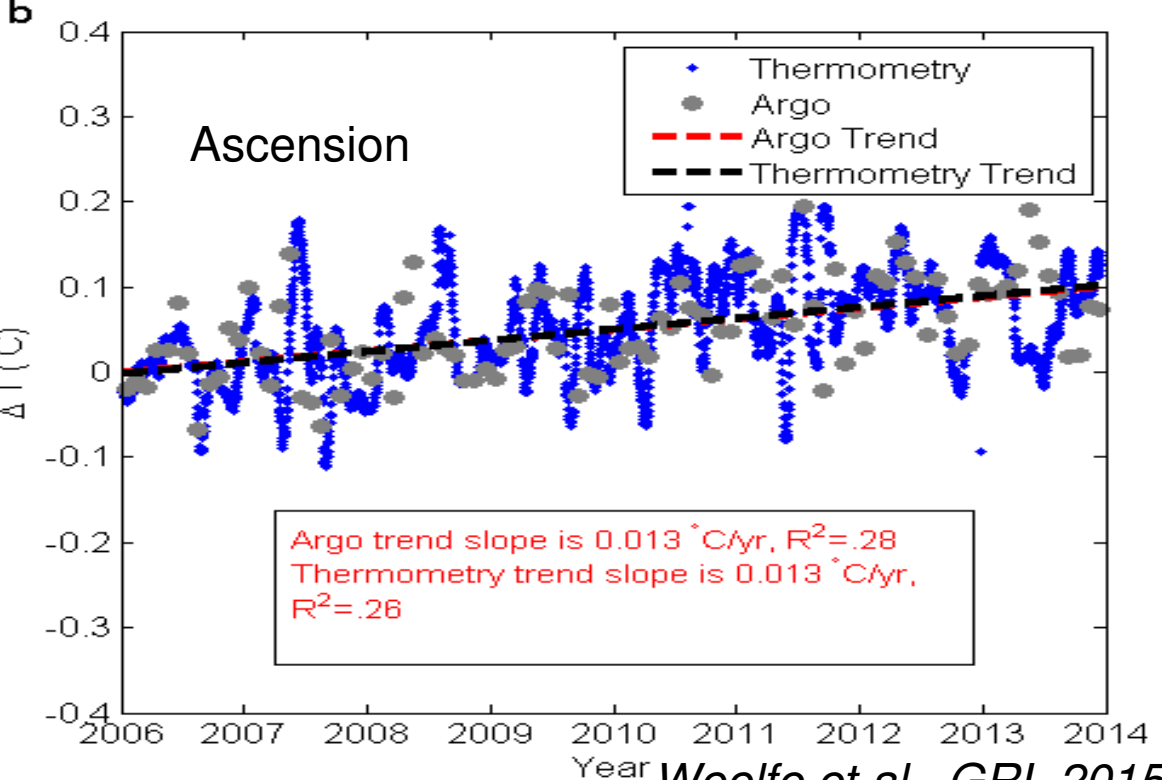
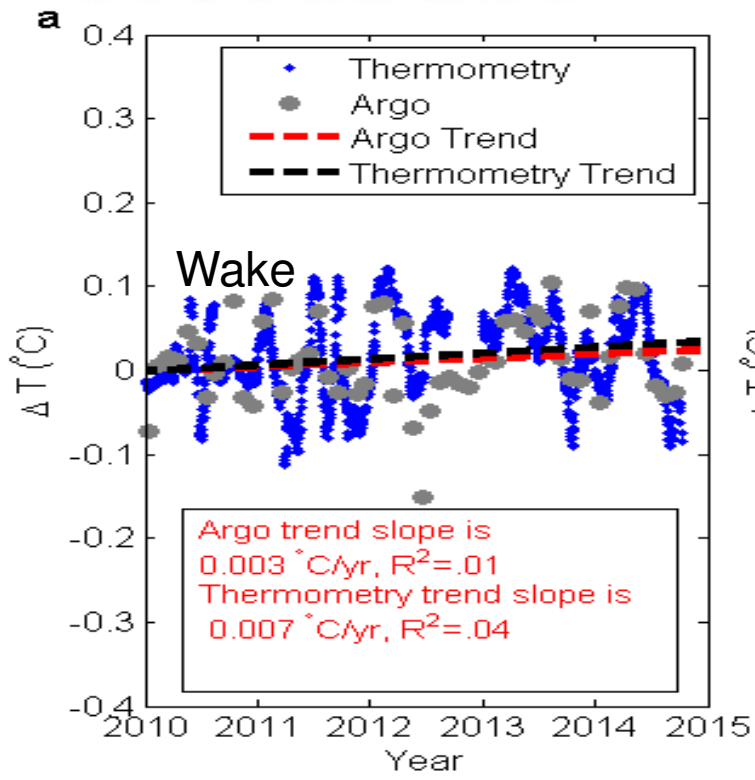
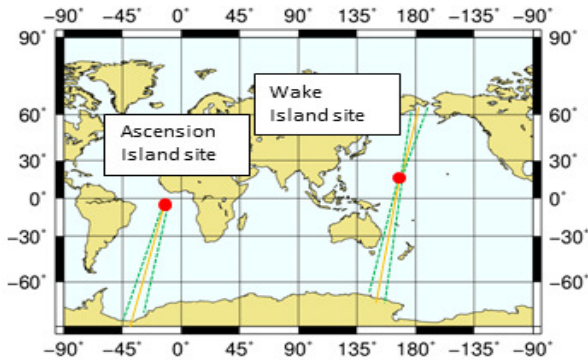


# Passive thermometry of the deep ocean

Comparison with ARGO temperature data (Roemmich, D. and J. Gilson, 2009)

Estimating Temperature changes:  $\frac{\Delta t}{t_0} = -\frac{\Delta c}{c_0} = \alpha \Delta T \left(1 + \frac{\mu \beta}{\alpha}\right)$  (Munk et al. 2009)

From baseline sound speed profile



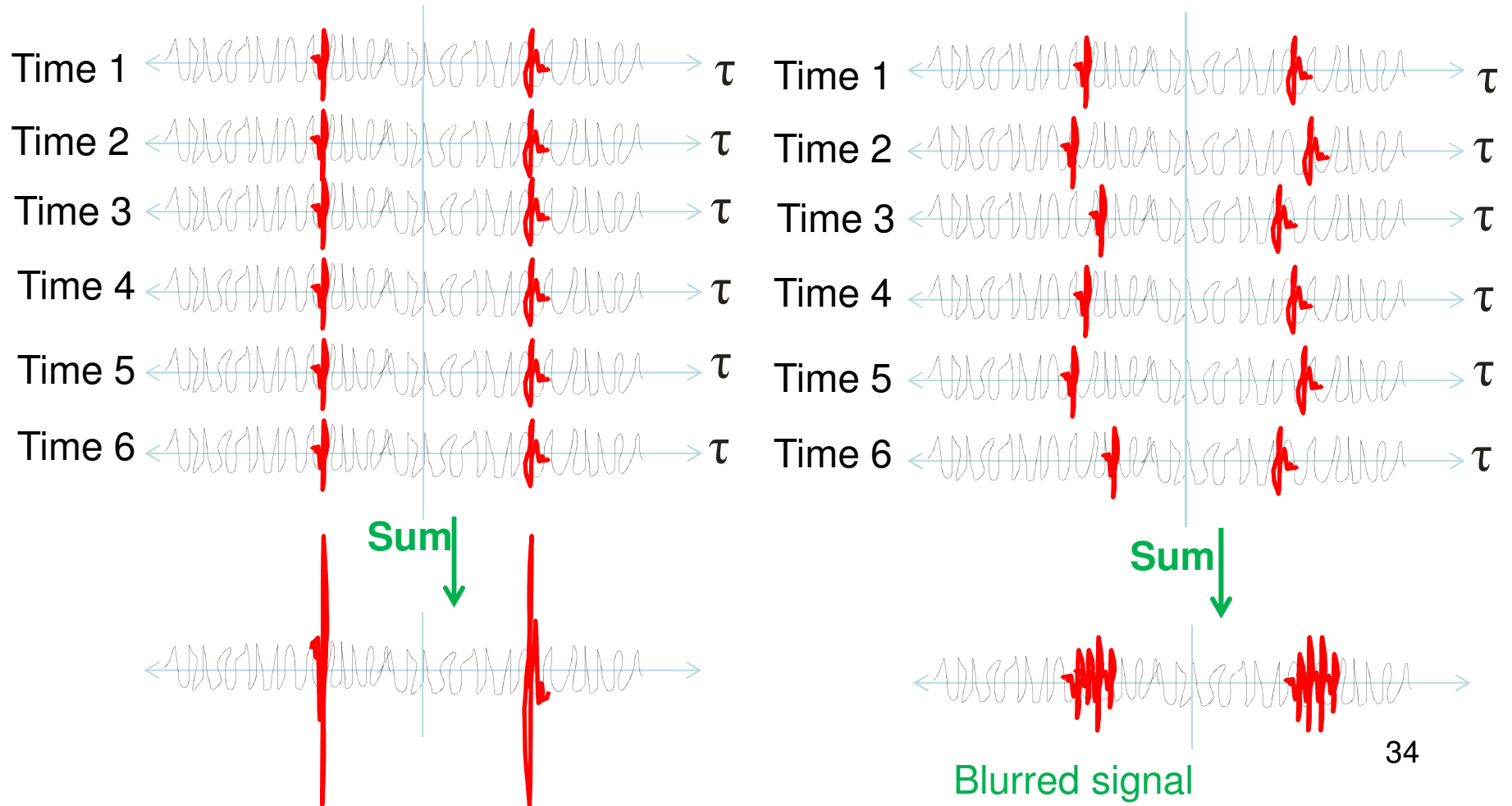
# Presentation Outline

- 1. Deep water acoustics & low-frequency noise sources**
- 2. Monitoring with ambient noise: Background.**
- 3. Passive acoustic thermometry of the deep water sound channel using low-frequency ambient noise**
- 4. Monitoring with ambient noise: Optimization.**

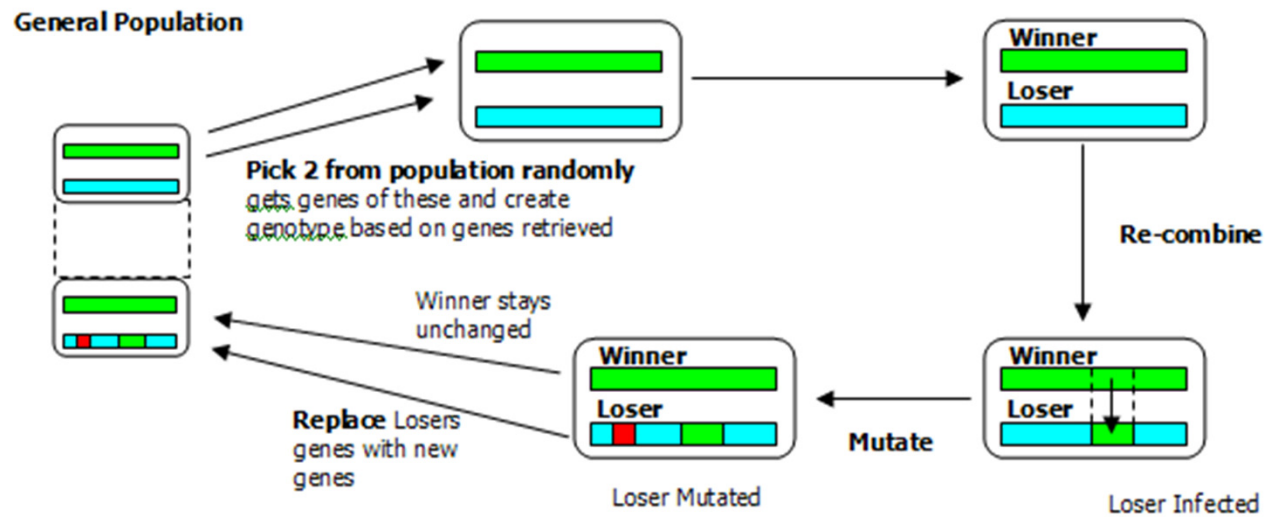
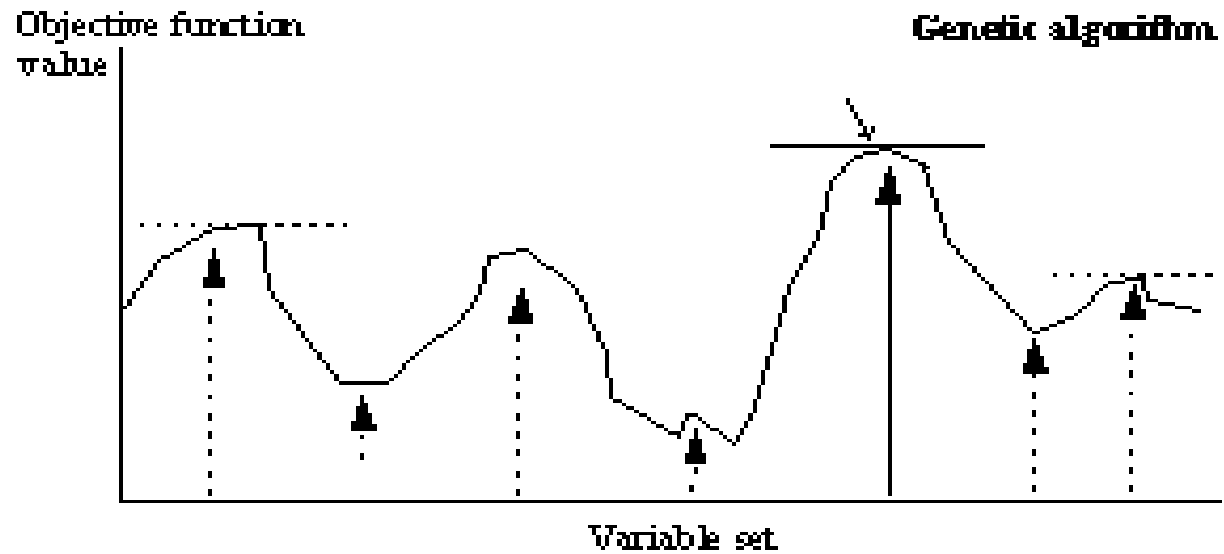


# Enhancing the emergence rate of coherent arrivals using optimization

How can the amount of time needed to extract a coherent arrival be minimized?

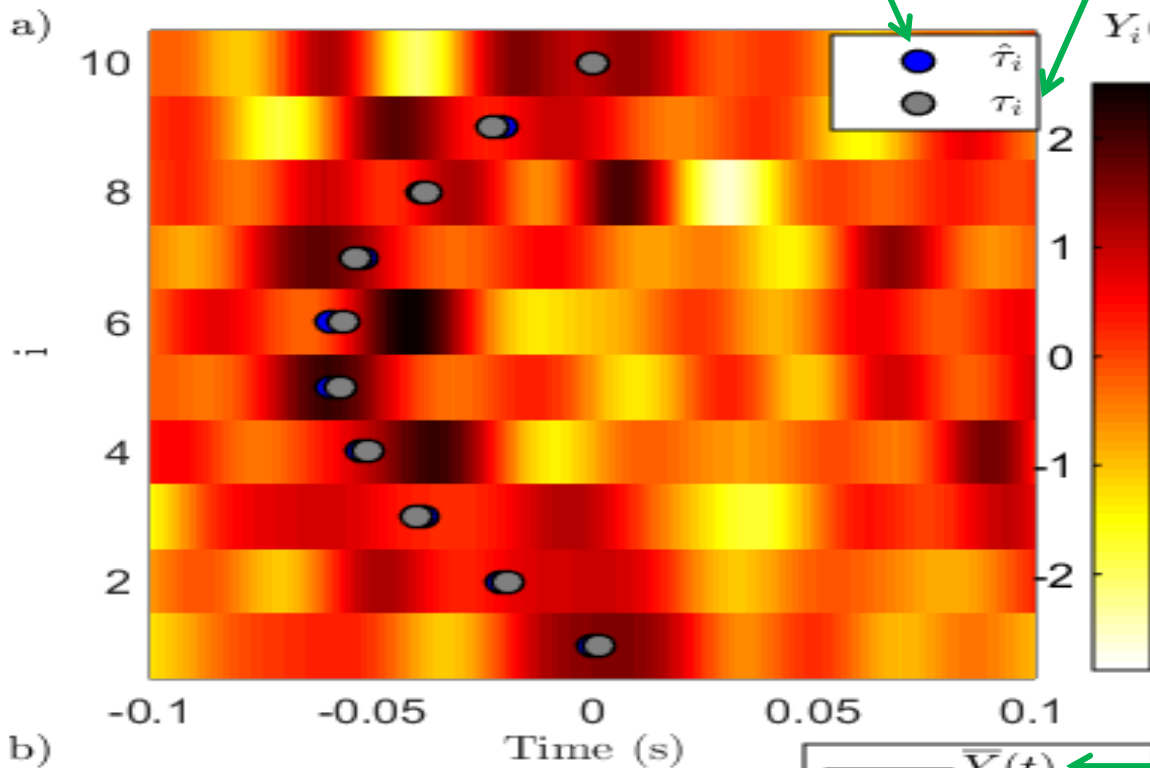


# Principles of Genetic Algorithm.



# Optimization – "delay search": Numerical results

True Delays      Estimated Delays

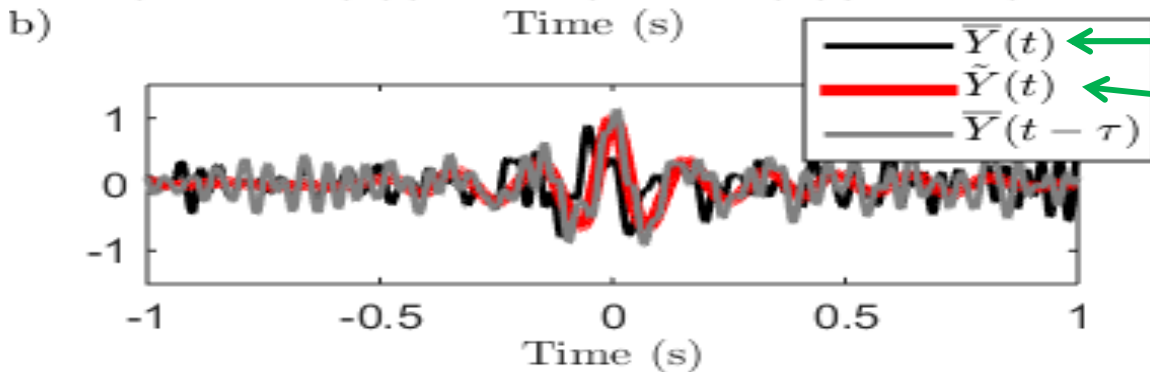


$$Y_i(t) = \tilde{Y}(t - \hat{\tau}_i) + N_i(t)$$

Shifted Reference waveform

Cost function for GA optimization

$$P(\tau_i, i = 1..N) = \left( \left[ \sum_{i=1}^N Y_i(t - \tau_i) \right] * \tilde{Y}(t) \right) \Big|_{t=0}$$



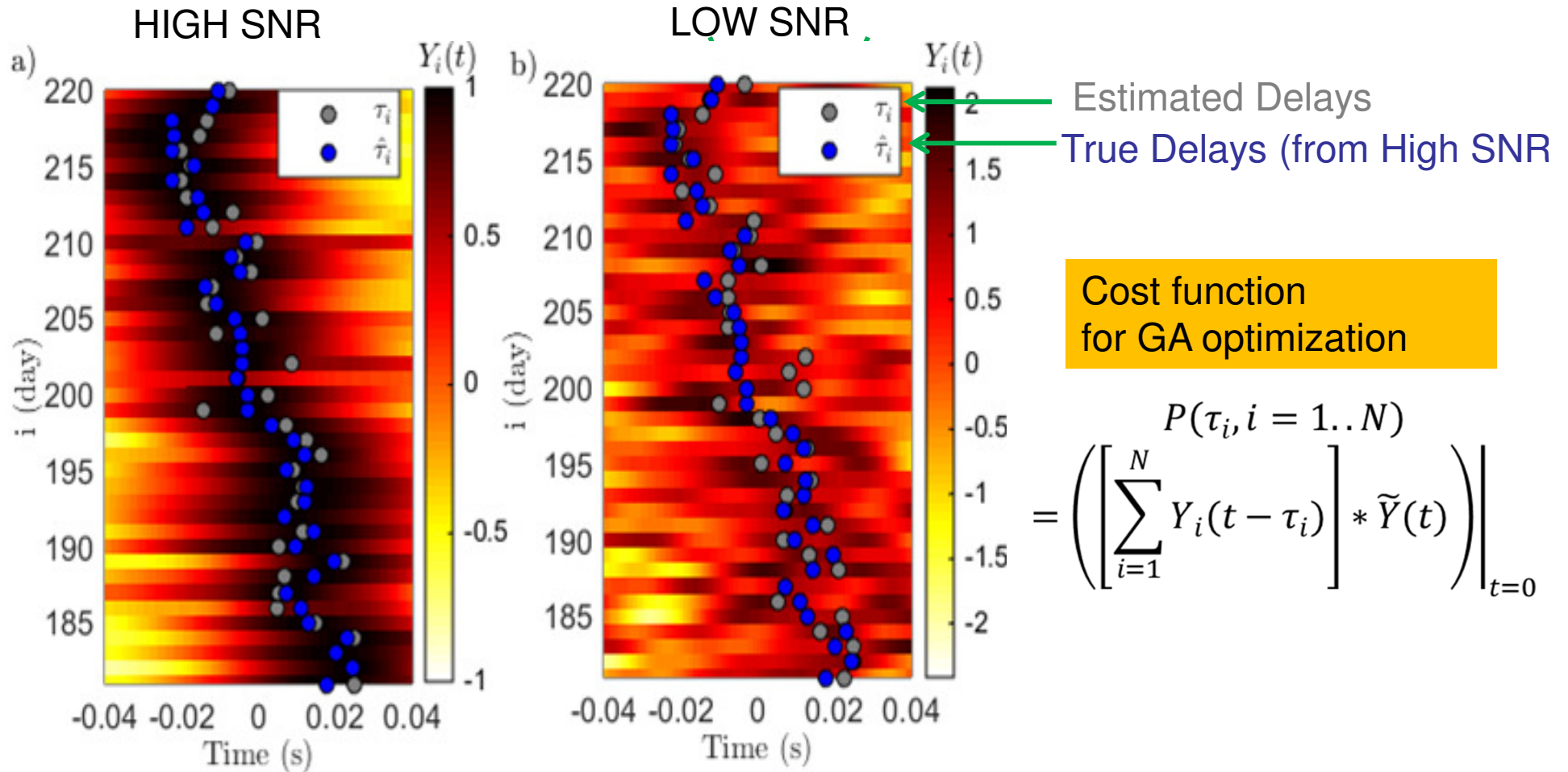
Regular Average

"Reference" waveform

Optimized Average

1Hz < f < 40Hz

# Optimization – "delay search": Experimental results



- Track arrival time of Beamformer Noise correlations (on SOFAR arrivals) over 40days
- 1Hz<f<40Hz

# Conclusions

**Emergence rate of coherent arrivals from ambient noise correlations can be enhanced** using:

- Array beamforming (Increase Spatial diversity -> Reduce Averaging time)
- Optimization/Search Algorithm (“Unravel” medium fluctuations)

**Applications to various domains...**

- “Faster” passive monitoring (on shorter time scales)
- “Selective” monitoring (select spatial regions)
- Mitigate environmental variations (e.g. for passive target detection)

Requires a good understanding of..

- noise sources characteristics /physics
- spatial & temporal scales of medium variations



Questions?

Thank you