

# Modeling and Analysis of Seismogram Envelopes in Short Periods

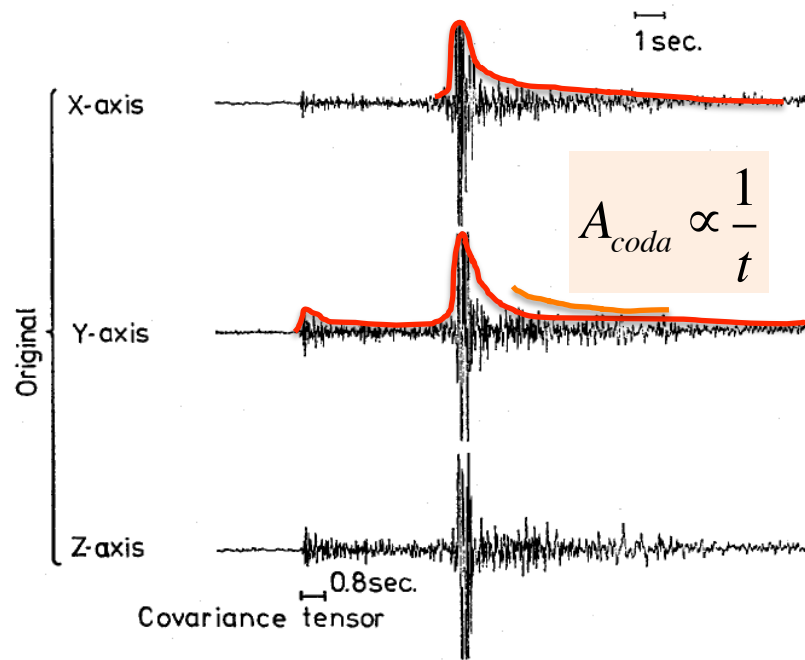
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- **Coda waves** of local earthquakes
- **Random heterogeneities** seen in well log data  
Random fluctuation, ACF/PSDF, Power-law spectra
- **Scalar wave equation in random media**  
Born app., Scattering coefficient
- **Radiative transfer theory (RTT)** in scattering media  
Boltzmann eq.
- **Multiple isotropic scattering model for coda excitation**  
Measurements of total scattering co. , intrinsic abs., and source energy radiation
- **Nonisotropic scattering model**  
Spherical harmonics expansion method, Monte Carlo method
- **Envelope broadening** and Markov approximation method  
Direct derivation of MS envelope based on the parabolic app.
- Model of **whole S-envelope** from onset to coda (ongoing study)  
Use of the Markov propagator and transport scat. co. in the MIS model
- In this talk, we focus on the **scalar wave equation only**.

at NRCDP (NIED) in 1974

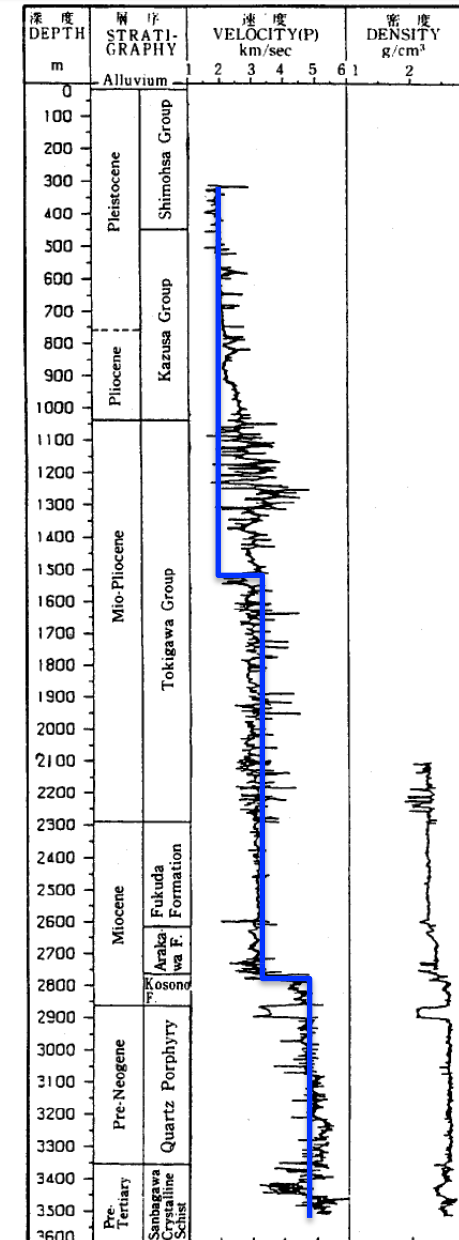
Seismograms registered at a deep well site IWT:  
Direct arrivals+ smoothed envelopes



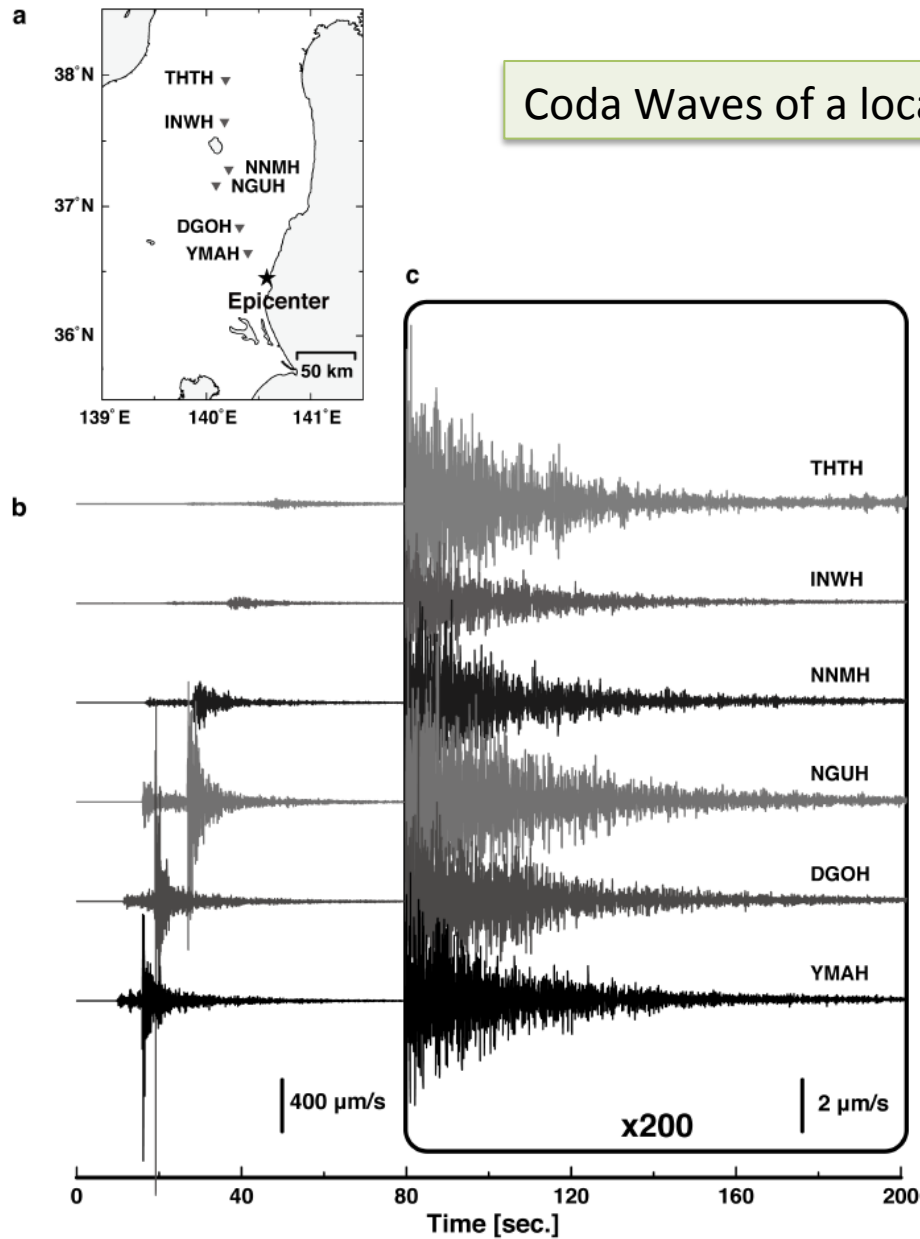
How shall we interpret seismogram envelopes  
in relation with random heterogeneities?

Aki and Chouet, Origin of coda waves, JGR, 1975.

Sonic and density well logs at IWT:  
Step like changes + random fluctuation



-3610m



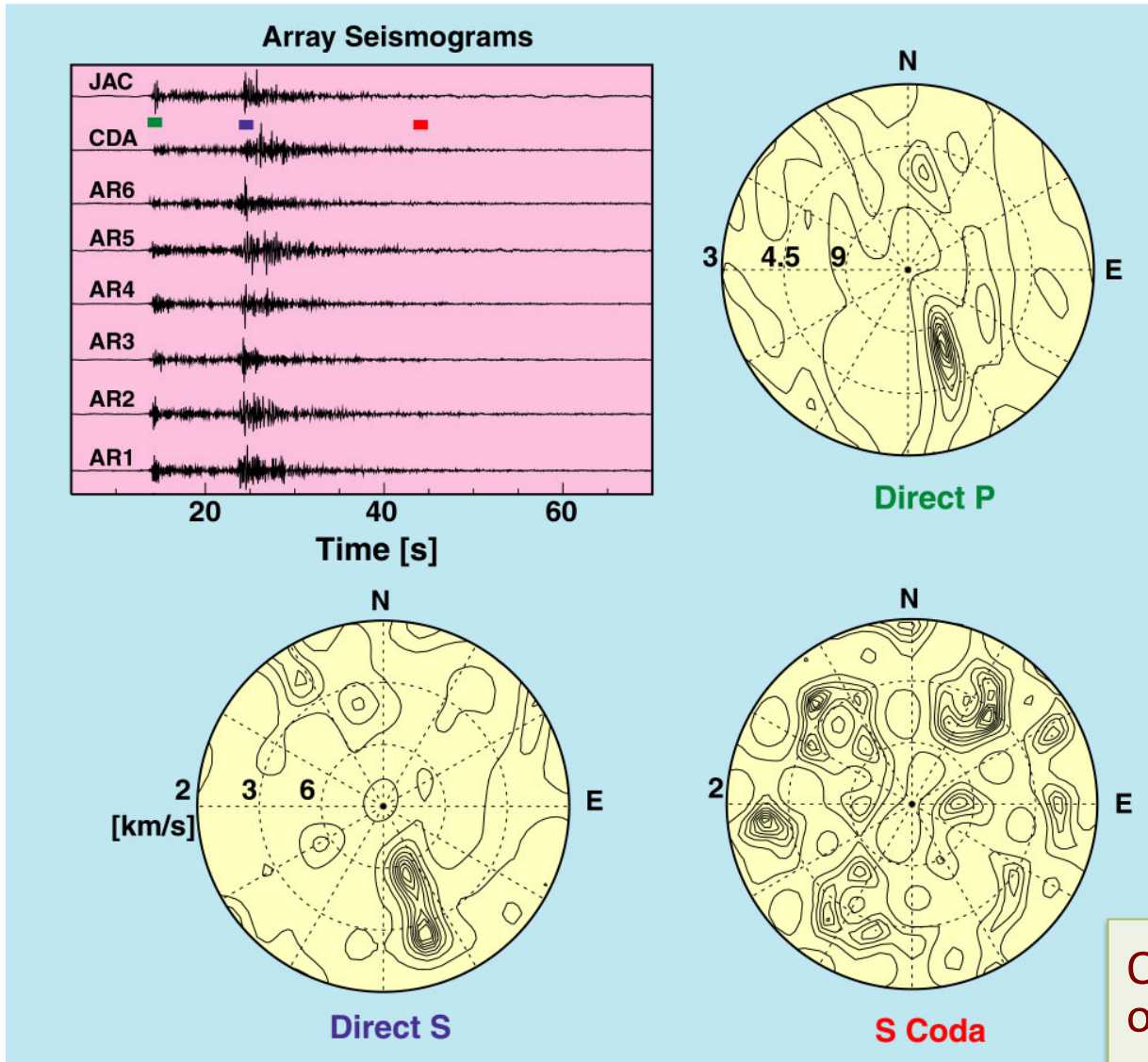
Coda Waves of a local earthquake registered by Hi-net

Common decay curve for coda amplitude envelopes independent of epicentral distances

Fig. 1.1 (a) Epicenter (star) of an  $M_W$  4.8 earthquake with 55.3 km in focal depth and Hi-net stations (reversed triangles) of NIED in Honshu, Japan. (b) Velocity seismograms (horizontal transverse component) arranged from bottom to top by increasing epicentral distance, where the gain is the same for all the traces. (c) Magnification of 200 times.

Courtesy of T. Maeda

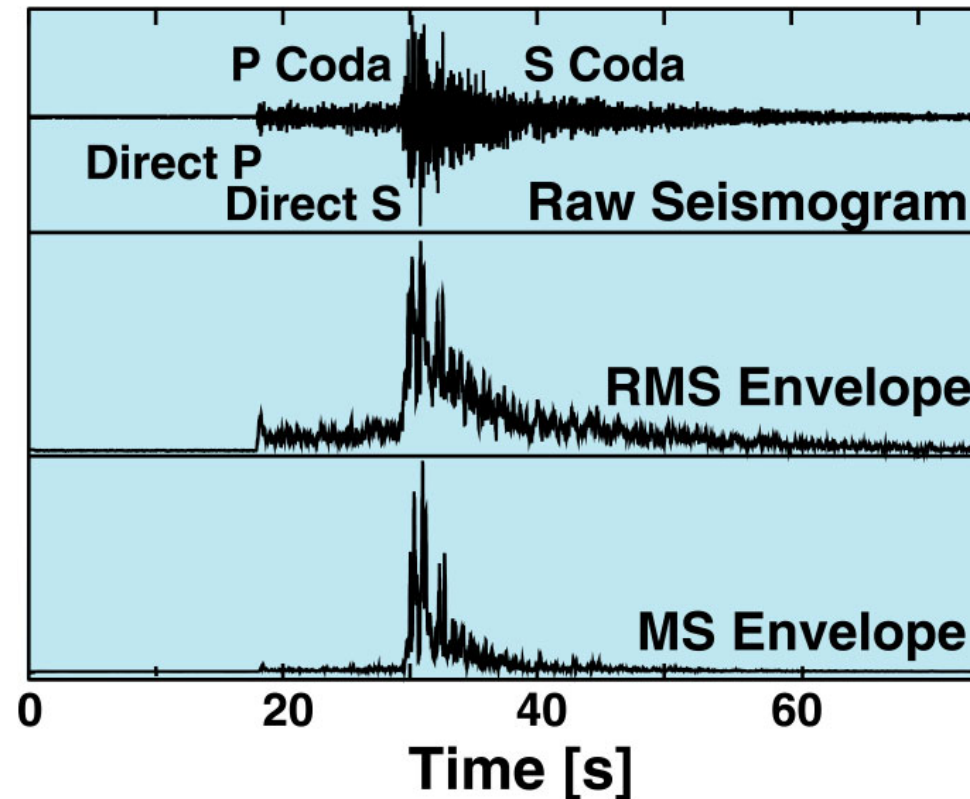
# F-K Analysis



Omni-directional arrivals  
of S-coda waves

superposition of scattered S-waves

- To focus on **amplitude envelope** disregarding complex phases
- To study how seismogram envelopes vary with lapse time, travel distance and ray paths in different frequency bands.

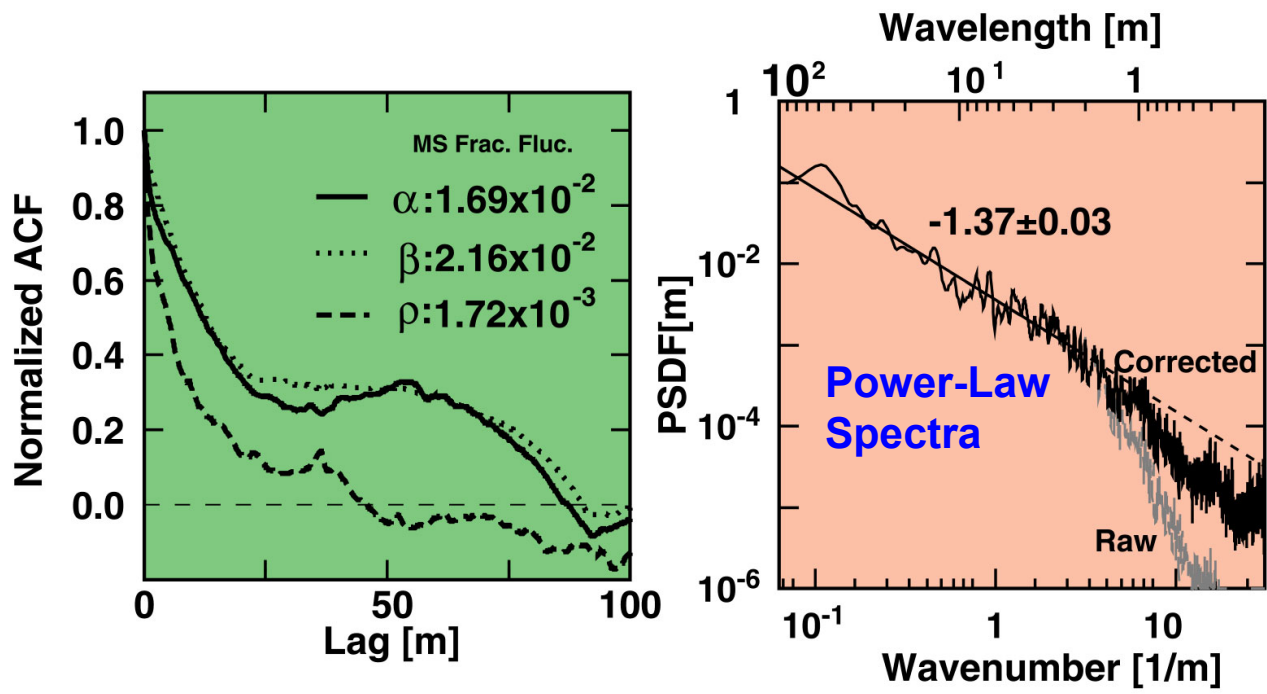


$$E(t, f_c) = \text{MS Amp.}(t, f_c, \Delta f)$$

$\Delta f$  : Octave band width

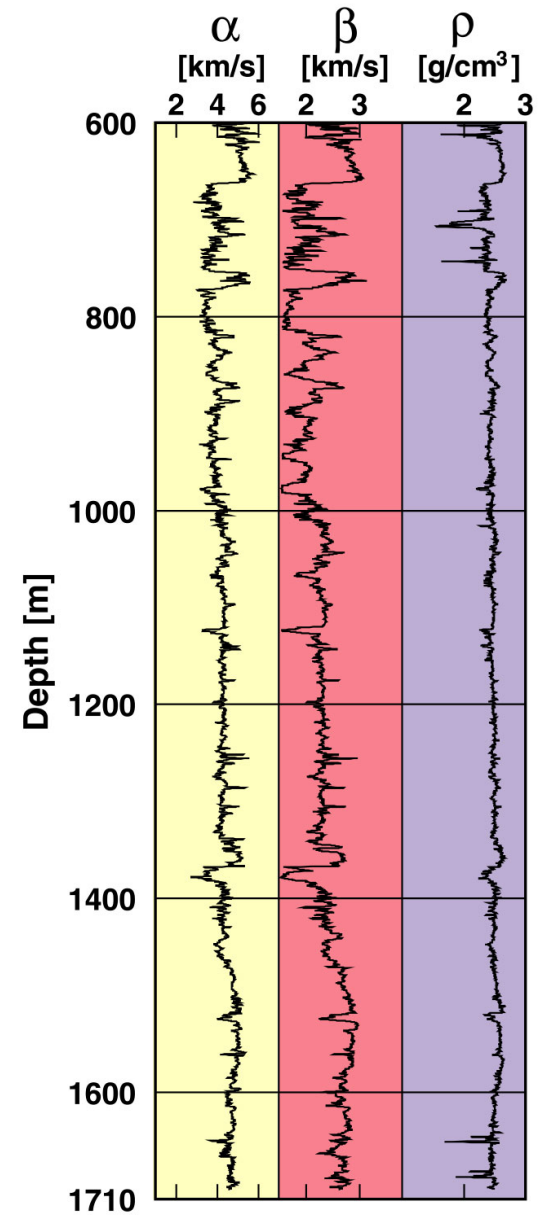
- Random heterogeneities seen in well log data

ACF/PSDF of fractional fluctuation of well log data (Volcanic tuff, Kyushu, Japan)

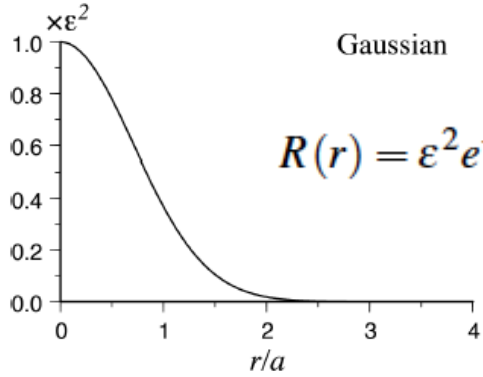


Shiomi et al. [1997]

### YT-2 Well Log Data



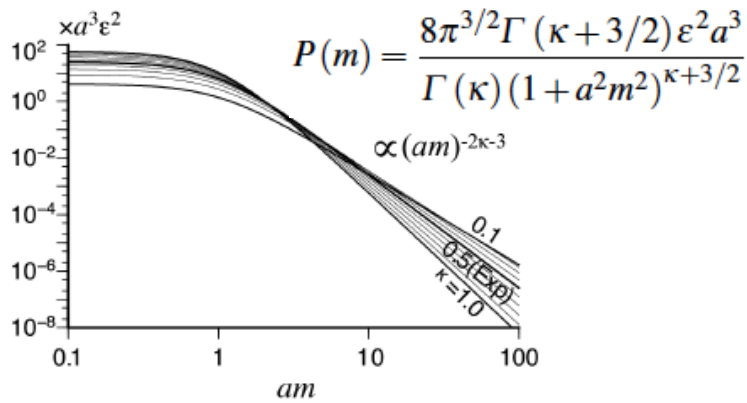
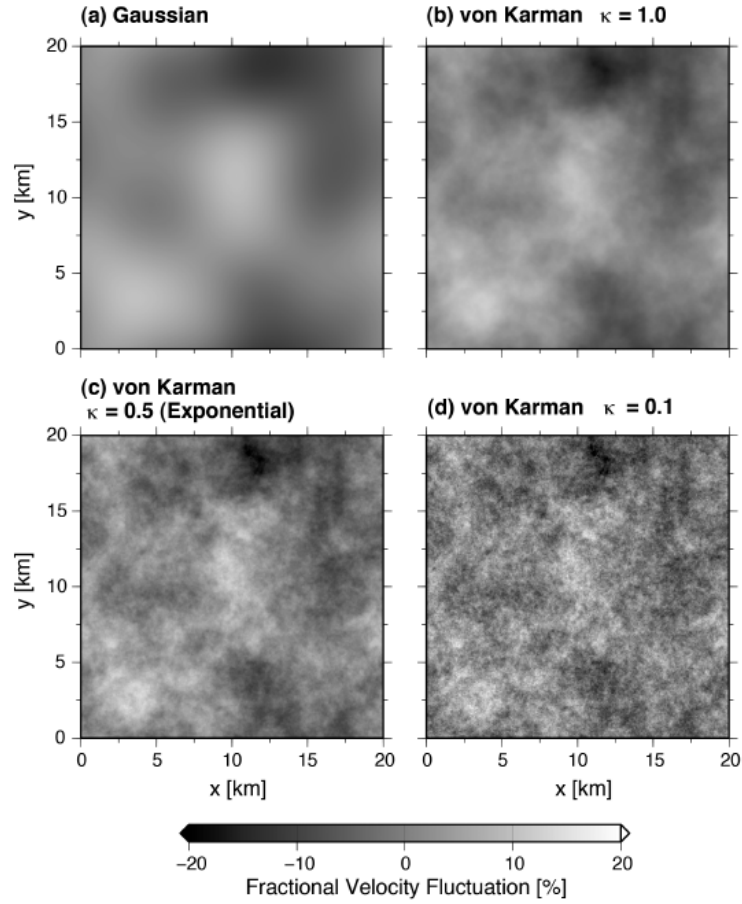
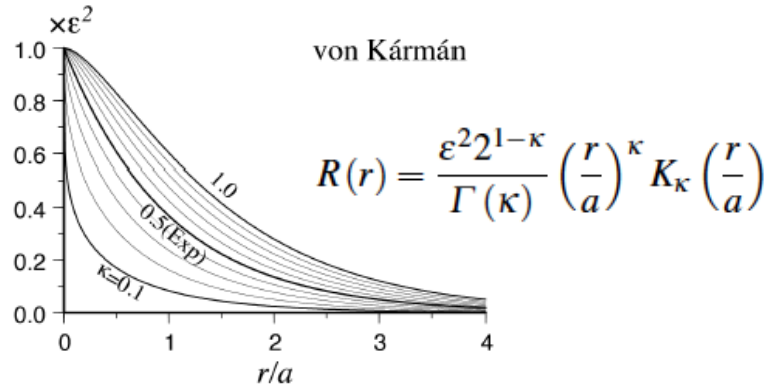
Ensemble of random media



$$V(\mathbf{x}) = V_0 + \delta V(\mathbf{x}) = V_0 [1 + \xi(\mathbf{x})],$$

$$R(\mathbf{x}) \equiv \langle \xi(\mathbf{y}) \xi(\mathbf{y} + \mathbf{x}) \rangle,$$

$$P(\mathbf{m}) = P(m) = \tilde{R}(\mathbf{x}) = \iiint_{-\infty}^{\infty} R(\mathbf{x}) e^{-i\mathbf{m}\mathbf{x}} d\mathbf{x},$$



Parameters: RMS fractional fluctuation  $\varepsilon$  and correlation distance  $a$

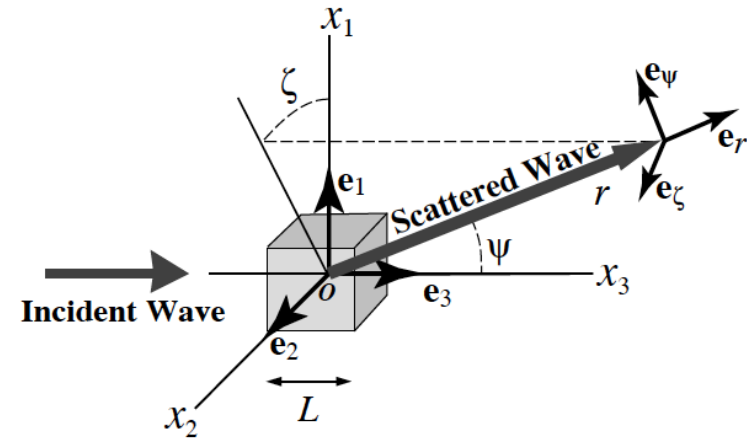
## Scalar wave equation in random media

$$\Delta u - \frac{1}{V_0^2} \partial_t^2 u = -\frac{2}{V_0^2} \xi(\mathbf{x}) \partial_t^2 u.$$

### Born approximation

$$u \approx e^{i(k_0 z - \omega t)} + F \frac{e^{i(k_0 r - \omega t)}}{r}$$

$$F = \frac{-k_0^2}{2\pi} \int \xi(\mathbf{x}') e^{-ik_0(\mathbf{e}_r - \mathbf{e}_z) \cdot \mathbf{x}'} d\mathbf{x}'$$



### Scattering coefficient: Scattering power per unit volume

$$g(\psi; \omega) \equiv 4\pi \frac{1}{L^3} \left\langle \frac{d\sigma}{d\Omega} \right\rangle = 4\pi \frac{1}{L^3} \langle |F|^2 \rangle$$

$$= \frac{k_0^4}{\pi} P(k_0 \mathbf{e}_r - k_0 \mathbf{e}_3) = \frac{k_0^4}{\pi} P\left(2k_0 \sin \frac{\psi}{2}\right)$$

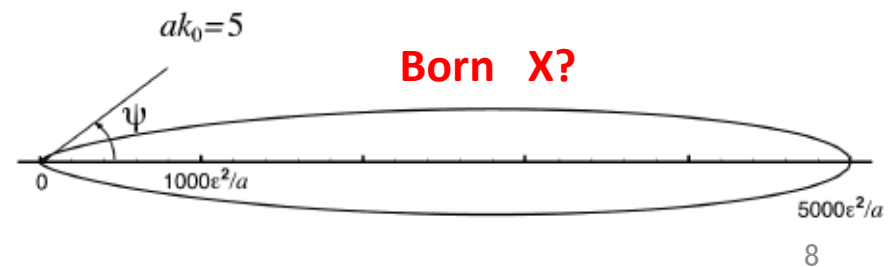
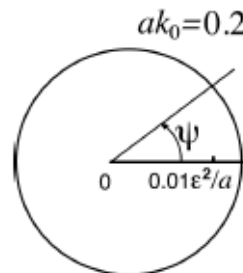
### Ray bending parameter

$$g_0(\omega) = \frac{1}{4\pi} \oint g(\psi, \omega) d\Omega$$

$g_0^{-1}$ : Mean free path

### Exponential ACF:

$$g(\psi; \omega) = \frac{8\epsilon^2 a^3 k_0^4}{(1 + 4a^2 k_0^2 \sin^2 \frac{\psi}{2})^2}$$





Radiative Transfer Equation (Boltzmann equation) in Scattering Media

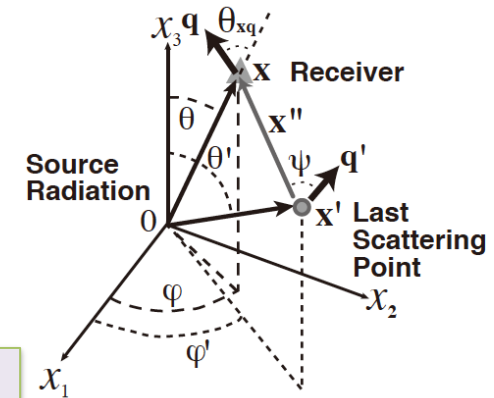
$$\partial_t f(\mathbf{x}, t; \mathbf{q}) + V_0 \mathbf{q} \nabla f(\mathbf{x}, t; \mathbf{q}) = -g_0(k_0) V_0 f(\mathbf{x}, t; \mathbf{q}) + \frac{V_0}{4\pi} \oint g(k_0 \mathbf{q} - k_0 \mathbf{q}') f(\mathbf{x}, t; \mathbf{q}') d\Omega_{q'} + W(\omega) \frac{\Psi(\mathbf{q})}{4\pi} \delta(\mathbf{x}) \delta(t)$$

$$E(\mathbf{x}, t) = \oint f(\mathbf{x}, t, \mathbf{q}) d\Omega_q$$

Ray bending  $g(k_0 \mathbf{q}) = \frac{k_0^4}{\pi} P(k_0 \mathbf{q})$

$$k_0 = \frac{\omega}{V_0}$$

$$g_0(k_0) = \frac{1}{4\pi} \oint g(k_0 \mathbf{q}) d\Omega_q \quad \oint \Psi(\mathbf{q}) d\Omega_q = 4\pi$$



- $f(\mathbf{x}, t, \mathbf{q})$ : Directional distribution of energy density to direction  $\mathbf{q}$  (Specific intensity)
- $E(\mathbf{x}, t)$ : Energy density in a band having center at  $\omega_0$
- $V_0$ : Propagation velocity
- $P(\mathbf{m})$ : Power spectral density of velocity fractional fluctuation
- $g(k_0 \mathbf{q})$ : Scattering coefficient at wavenumber  $k_0$  to direction  $\mathbf{q}$
- $g_0(k_0)$ : Total scattering coefficient at wavenumber  $k_0$
- $W$ : Total radiated energy
- $\Psi(\mathbf{q})$ : Radiation pattern

Phenomenological equation satisfying energy conservation.

## Diffusion Approximation

At a large lapse time, we may expect that spatial distribution of  $f$  is smooth. Boltzmann eq. is well approximated by the diffusion equation:

$$\partial_t E - D\Delta E = W \delta(\mathbf{x}) \delta(t)$$

$$D = V_0 / (3g_m)$$

Transport Scattering Coefficient (TrSC):

$$\begin{aligned} g_m(\omega_0) &= \frac{1}{4\pi} \oint g(k_0, \psi) (1 - \cos \psi) d\Omega \\ &= \frac{1}{4\pi} \int_0^{2k_0} P(m) m^3 dm. \end{aligned}$$

Total Scattering Coefficient :

$$\begin{aligned} g_0(\omega_0) &= \frac{1}{4\pi} \oint g(k_0, \psi) d\Omega \\ &= \frac{k_0^2}{2\pi} \int_0^{2k_0} P(m) m dm \end{aligned}$$

$g_m(k_0)$  works as an effective isotropic scattering coefficient even if  $g(k_0 \mathbf{q})$  is nonisotropic.

$$E(\mathbf{x}, t) = \frac{W}{(4\pi Dt)^{3/2}} e^{-\frac{r^2}{4Dt}} H(t)$$

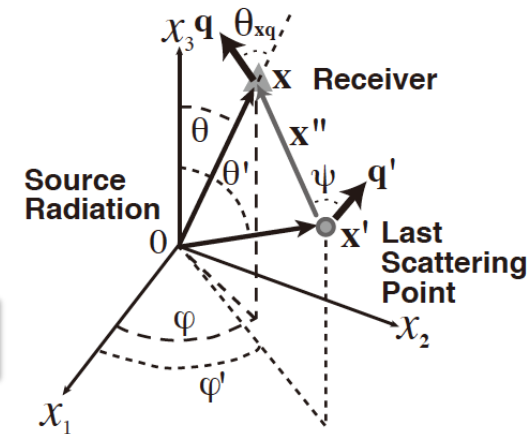
## Radiative Transfer Equation in Integral Form

$$f(\mathbf{x}, t; \mathbf{q}) = W\Psi(\mathbf{q}) G_R(\mathbf{x}, t; \mathbf{q}) + V_0 \int_{-\infty}^{\infty} dt' \iiint_{-\infty}^{\infty} d\mathbf{x}' G_R(\mathbf{x} - \mathbf{x}', t - t'; \mathbf{q}) \oint g(k_0\mathbf{q} - k_0\mathbf{q}') f(\mathbf{x}', t'; \mathbf{q}') d\Omega_{q'}.$$

Propagator:  $G_R(\mathbf{x}, t; \mathbf{q}) = \delta_{\Omega}(\mathbf{x}; \mathbf{q}) G(\mathbf{x}, t)$

$$G(\mathbf{x}, t) = \delta\left(t - \frac{|\mathbf{x}|}{V_0}\right) \frac{e^{-g_0 V_0 t}}{4\pi V_0 |\mathbf{x}|^2} H(t)$$

Causality, Geometrical decay factor, Scattering loss



$\delta_{\Omega}(\mathbf{x}; \mathbf{q})$  : Delta function in solid angle  
Propagation is on the ray direction

## Multiple isotropic scattering model for coda excitation

$$g_0 = g_m$$

Spherical source radiation and isotropic scattering:

$$E(\mathbf{x}, t) = W G(\mathbf{x}, t) + V_0 g_0 \int_{-\infty}^{\infty} G(\mathbf{x} - \mathbf{x}', t - t') E(\mathbf{x}', t') d\mathbf{x}' dt'$$

$$G(\mathbf{x}, t) = \frac{1}{4\pi r^2 V_0} \delta\left(t - \frac{r}{V_0}\right) e^{-g_0 V_0 t} \Rightarrow \frac{1}{4\pi r^2 V_0} \delta\left(t - \frac{r}{V_0}\right) e^{-g_0 V_0 t - Q_I^{-1} \omega_0 t}$$

$Q_I^{-1}$  : Intrinsic absorption at  $\omega_0$

Fourier transform:

$$E(\mathbf{x}, t) = \frac{1}{(2\pi)^4} \iiint \int_{-\infty}^{\infty} \hat{E}(\mathbf{k}, \omega) e^{i(\mathbf{kx} - \omega t)} d\mathbf{k} d\omega$$

Radiative transfer equation in the Fourier space:

$$\hat{E}(\mathbf{k}, \omega) = \frac{W \hat{G}(\mathbf{k}, \omega)}{1 - g_0 V_0 \hat{G}(\mathbf{k}, \omega)}$$

Series expansion [Born approximation]

$$\hat{E}(\mathbf{k}, \omega) \approx W \hat{G} \left[ 1 + g_0 V_0 \hat{G} + \left( g_0 V_0 \hat{G} \right)^2 + \left( g_0 V_0 \hat{G} \right)^3 \dots \right]$$

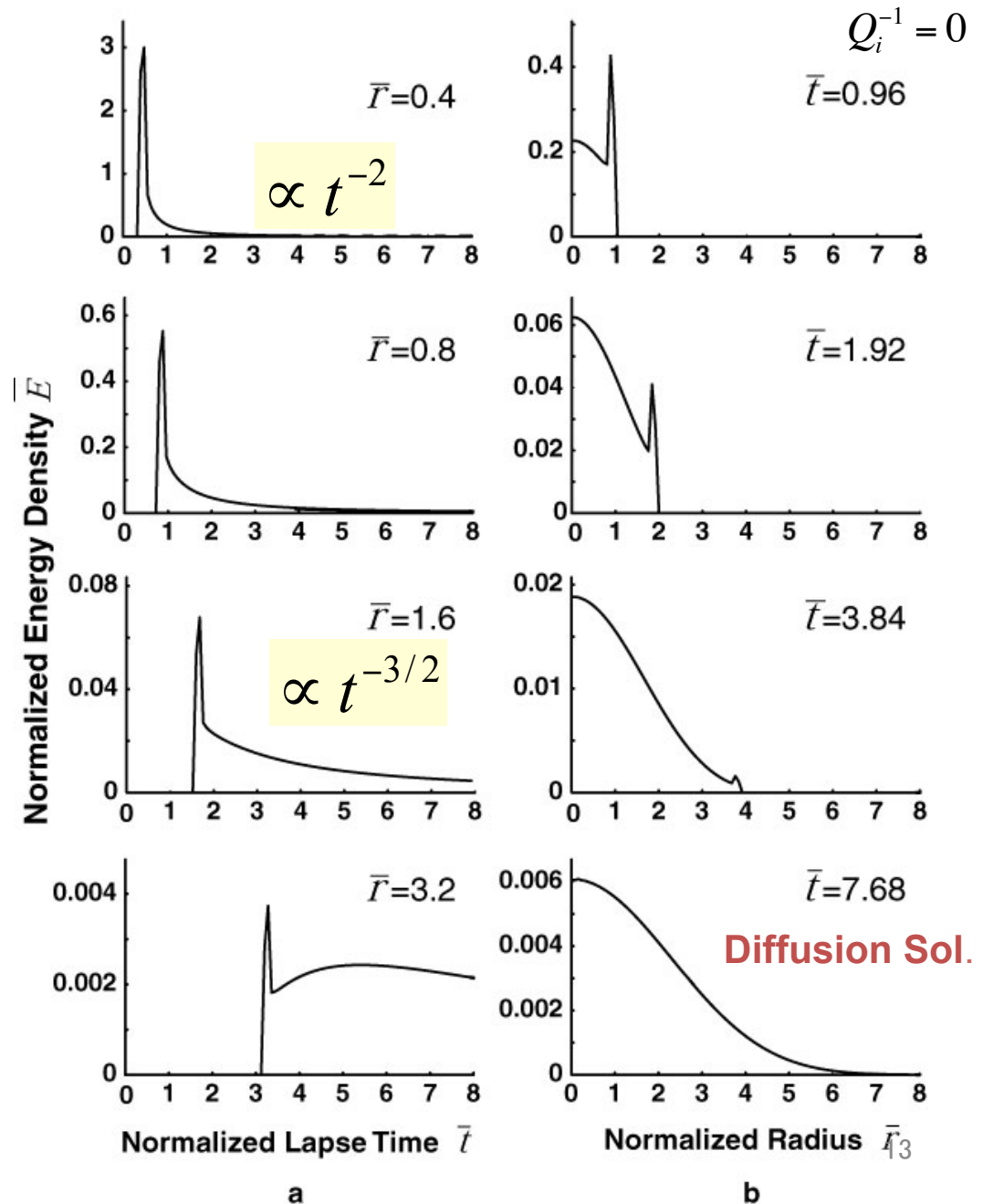
## Temporal Change and Spatial Dist. of Energy Density

Isotropic Scattering Process for Spherical Source Radiation  
[Zeng et al. 1991]

To multiply  $\text{Exp}[-Q_i^{-1} \omega_0 t]$  in the presence of intrinsic absorption

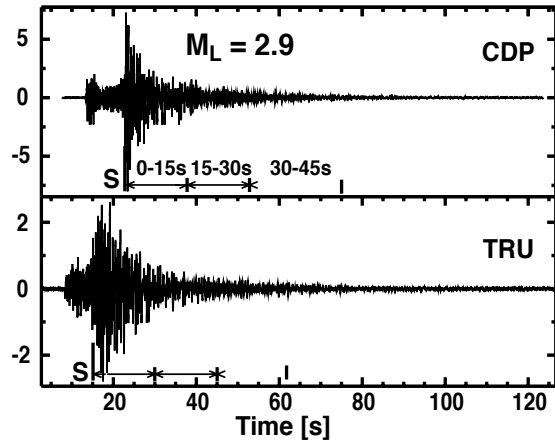
$$\bar{t} = g_0 V_0 t, \quad \bar{r} = g_0 r$$

$$\bar{E} = \frac{E}{W g_0^3}$$



Diffusion Sol.

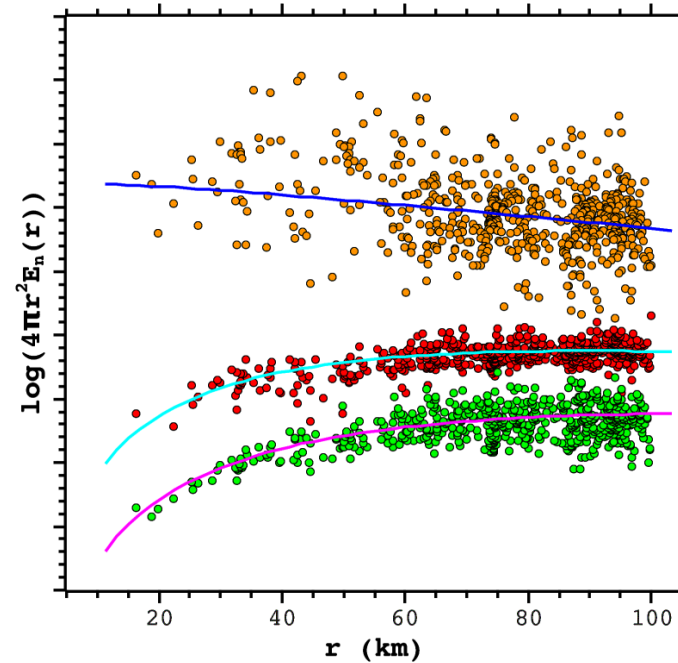
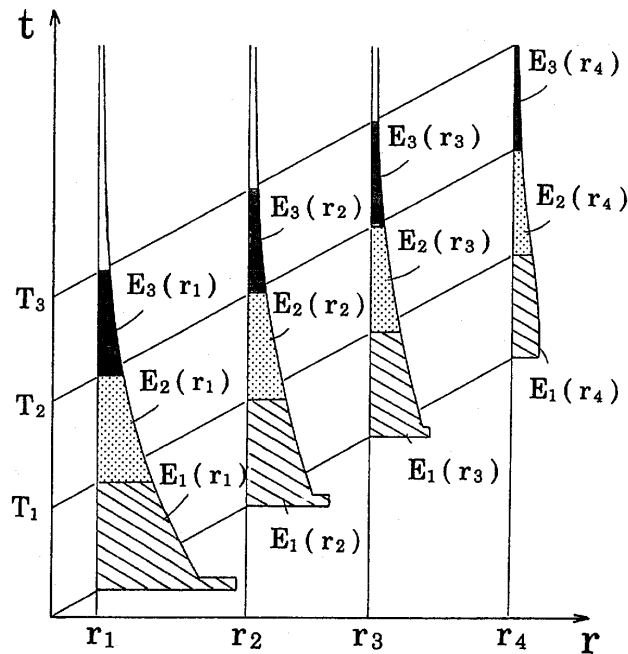
# Multiple Lapse Time Window Analysis [Fehler et al. 1991]



To plot the logarithm of the energy obtained from each window as a function of the hypocentral distance.

To perform a non-linear fit by using a simple model (isotropic scattering and intrinsic absorption):  $Q_S^{-1} = V_0 g_0 / \omega$  and  $Q_i^{-1}$

2-4 Hz

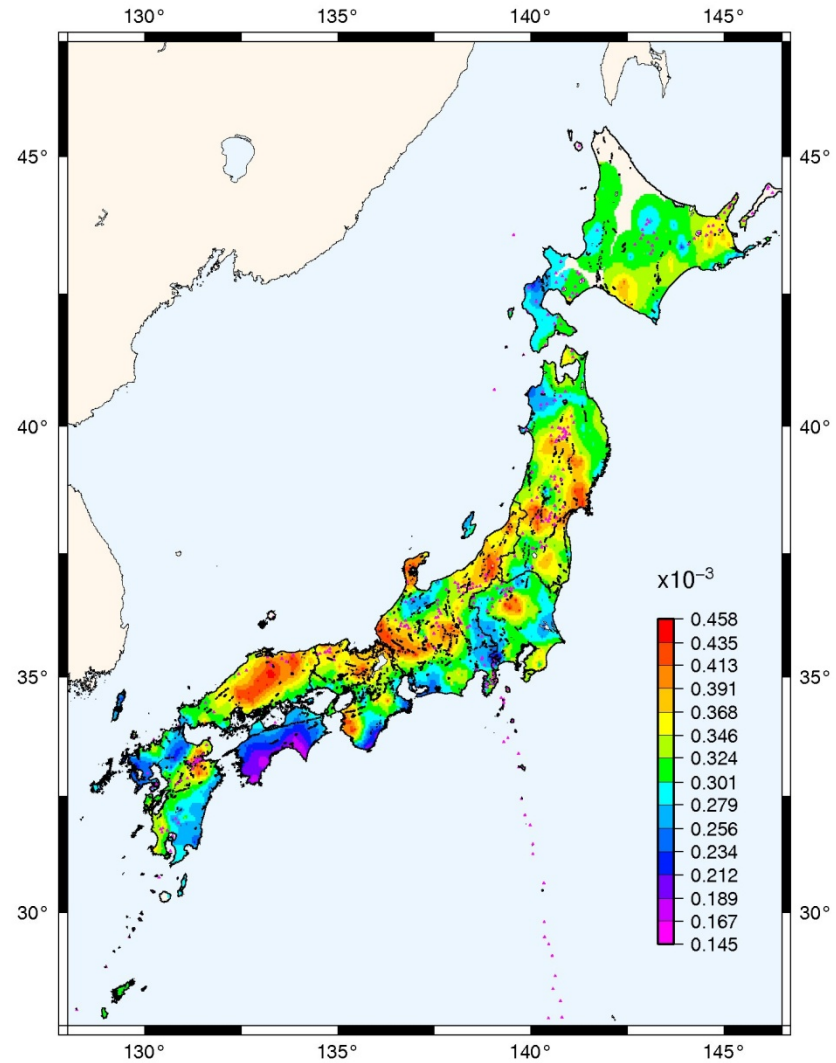
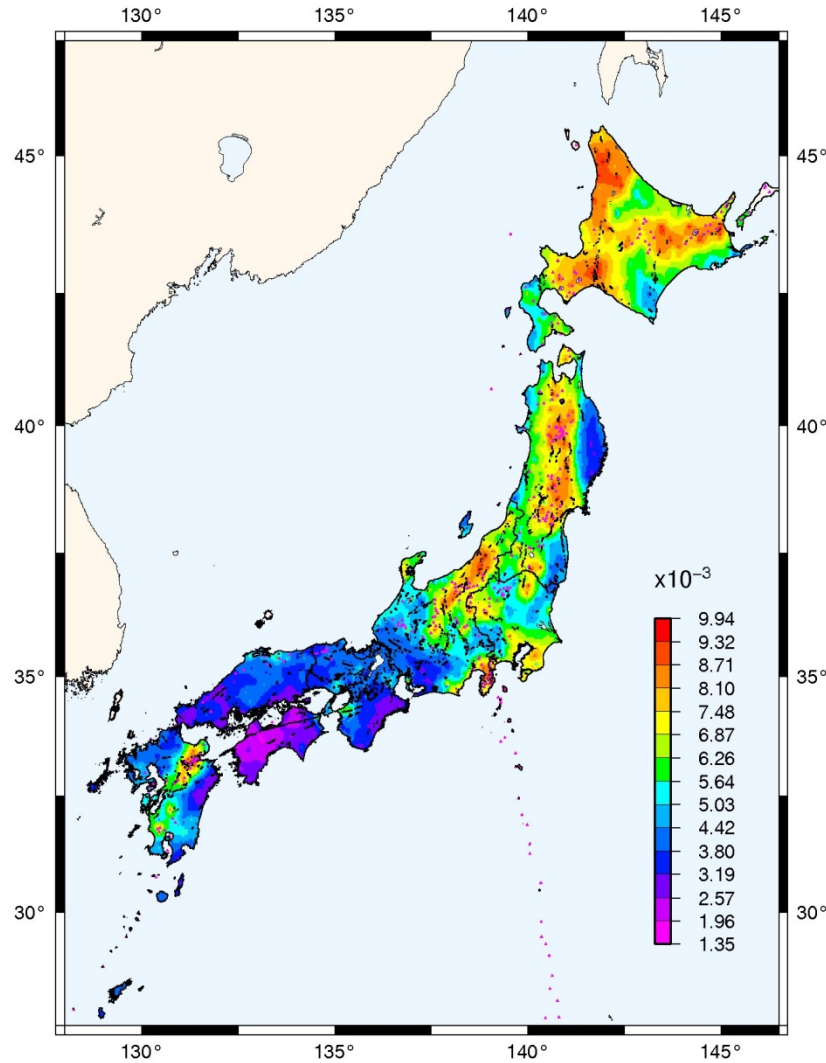


# Scattering Loss (Spatial smoothing)

$$Q_S^{-1} = V_0 g_0 / \omega$$

$Q_S^{-1}$  1-2 Hz

$Q_S^{-1}$  8-16 Hz



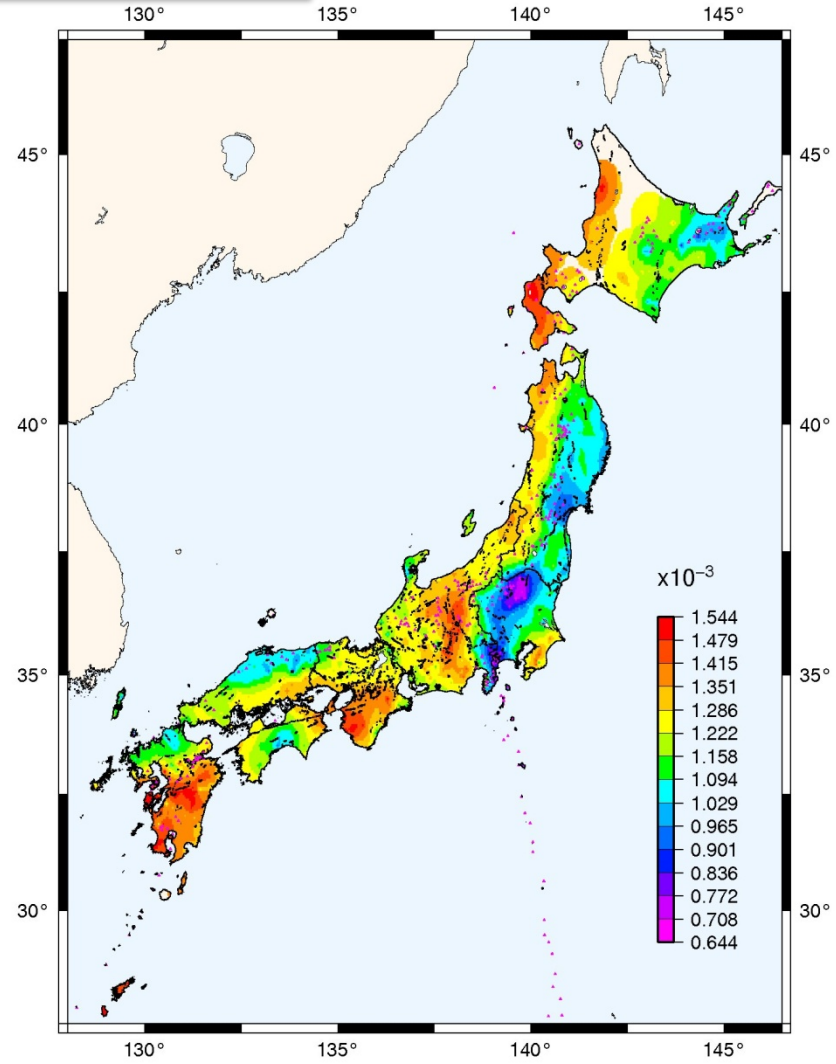
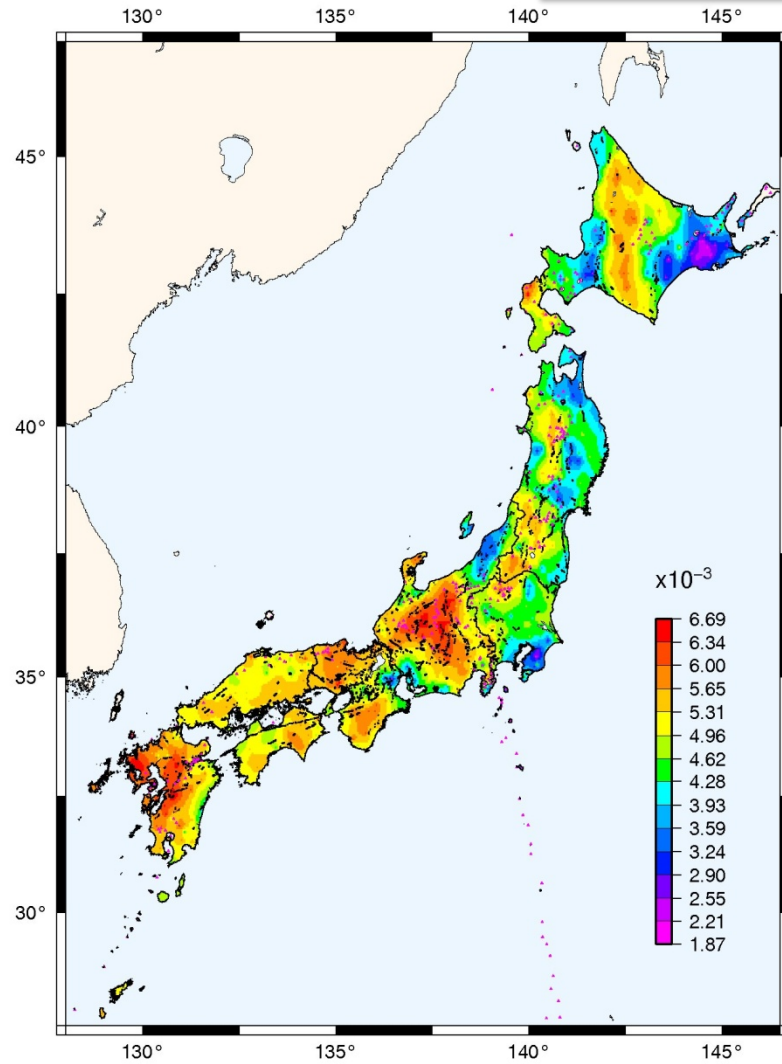
Hi-net: 772 stations,  
135,000 events ( $D < 40\text{km}$ )

Carcole and Sato [2010]

# Intrinsic Absorption (Spatial smoothing)

$Q_i^{-1}$  1–2 Hz

$Q_i^{-1}$  8–16 Hz





## Multiple Isotropic Scattering Model with PS Conversion

Two velocity modes

$$E(\mathbf{x}, t) = E^P(\mathbf{x}, t) + E^S(\mathbf{x}, t)$$

$$E^P(\mathbf{x}, t) = W^P G^P(\mathbf{x}, t) + \iiint \iiint \{E^P(\mathbf{x}', t') \alpha_0 g_0^{PP} + E^S(\mathbf{x}', t') \beta_0 g_0^{SP}\} G^P(\mathbf{x} - \mathbf{x}', t - t') dt' dx'$$

$$E^S(\mathbf{x}, t) = W^S G^S(\mathbf{x}, t) + \iiint \iiint \{E^S(\mathbf{x}', t') \beta_0 g_0^{SS} + E^P(\mathbf{x}', t') \alpha_0 g_0^{PS}\} G^S(\mathbf{x} - \mathbf{x}', t - t') dt' dx'$$

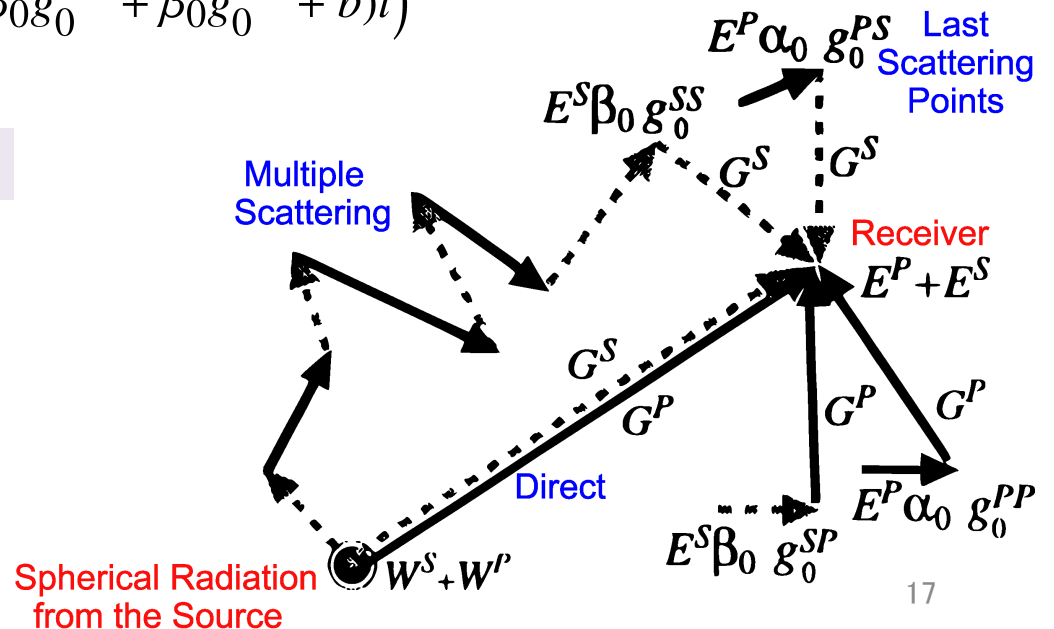
$$G^P(\mathbf{x}, t) = \frac{1}{4\pi r^2 \alpha_0} H(t) \delta\left(t - \frac{r}{\alpha_0}\right) \exp\left(-(\alpha_0 g_0^{PP} + \alpha_0 g_0^{PS} + b)t\right)$$

$$G^S(\mathbf{x}, t) = \frac{1}{4\pi r^2 \beta_0} H(t) \delta\left(t - \frac{r}{\beta_0}\right) \exp\left(-(\beta_0 g_0^{SS} + \beta_0 g_0^{SP} + b)t\right)$$

$g^{**}$  : Total scattering coefficient

Intrinsic attenuation  $b = \omega Q_i^{-1}$

Sato [1994]

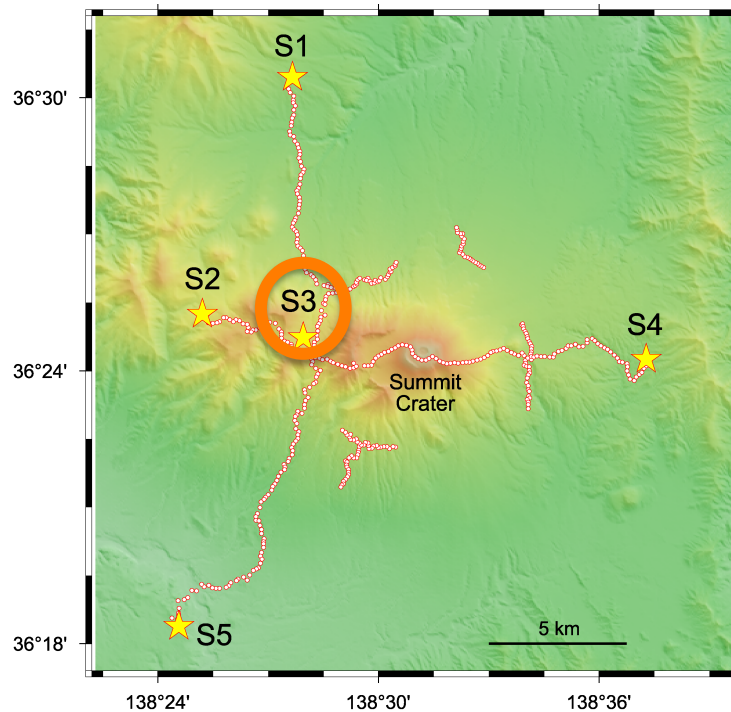


# Structure Study Experiment of Asama Volcano, Japan in 2006

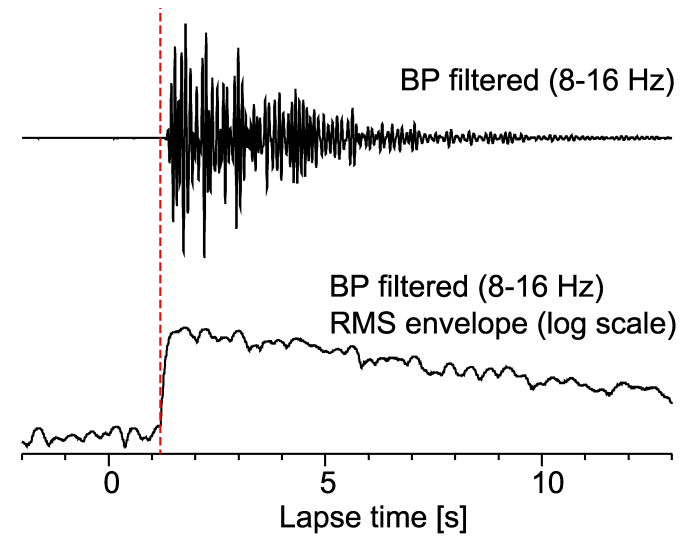


5 Shots, 250-350 kg TNT at 60m in depth.

464 stations (2Hz velocity seismometers) with a 100m separation



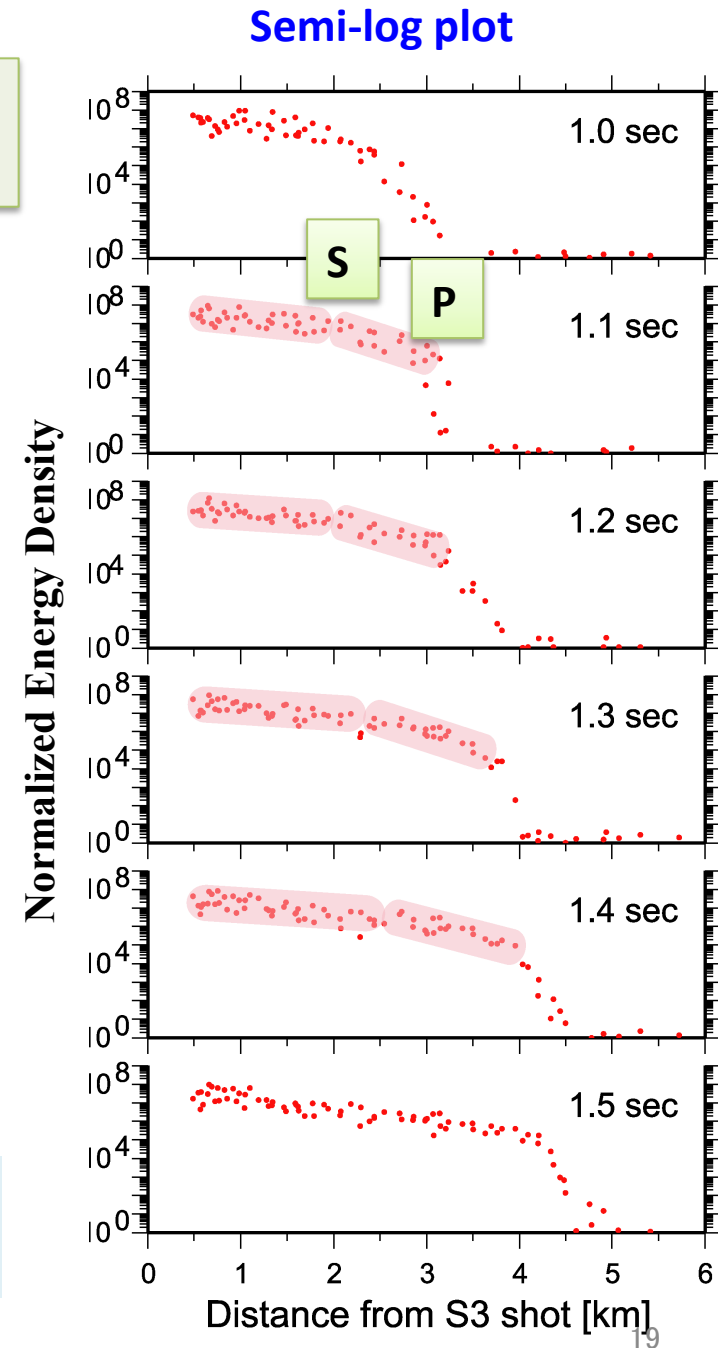
## Explosion...P-Source



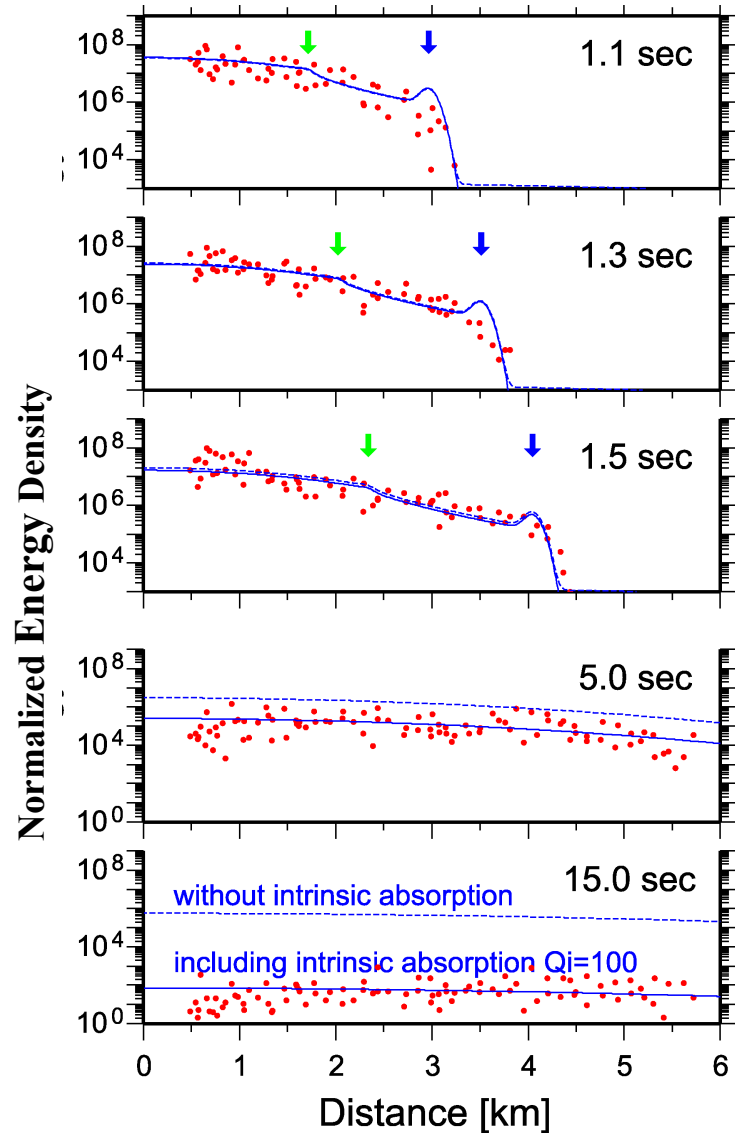
## Temporal Change in the Spatial Distribution of Seismic Energy Density (MS Amplitude)

- P-wave source (Explosion)
- Existence of wave front
- Energy concentration near the source
- Characterized by two different gradient lines
- **Corner propagates with the S wave velocity**

Band-pass filtering 8-16Hz  
Coda normalization at 50s



# Estimation of total scattering coefficients



Yamamoto and Sato [2010]

**Assumption:**

P wave source ( $W_s=0$ )

3-D infinite medium (neglect surface effect)

**Parameters:**

Initial P wave energy  $W_p$

P wave velocity

Isotropic scattering coefficients ( $g^{PP}$ ,  $g^{PS}$ ,  $g^{SS}$ ),

**Constraint:**  $\frac{\alpha_0}{\beta_0} = \sqrt{3}, \quad \frac{g_0^{PS}}{g_0^{SP}} = 2 \frac{\alpha_0^2}{\beta_0^2} = 6$

By using grid search

Data at lapse time 1.0-1.5s



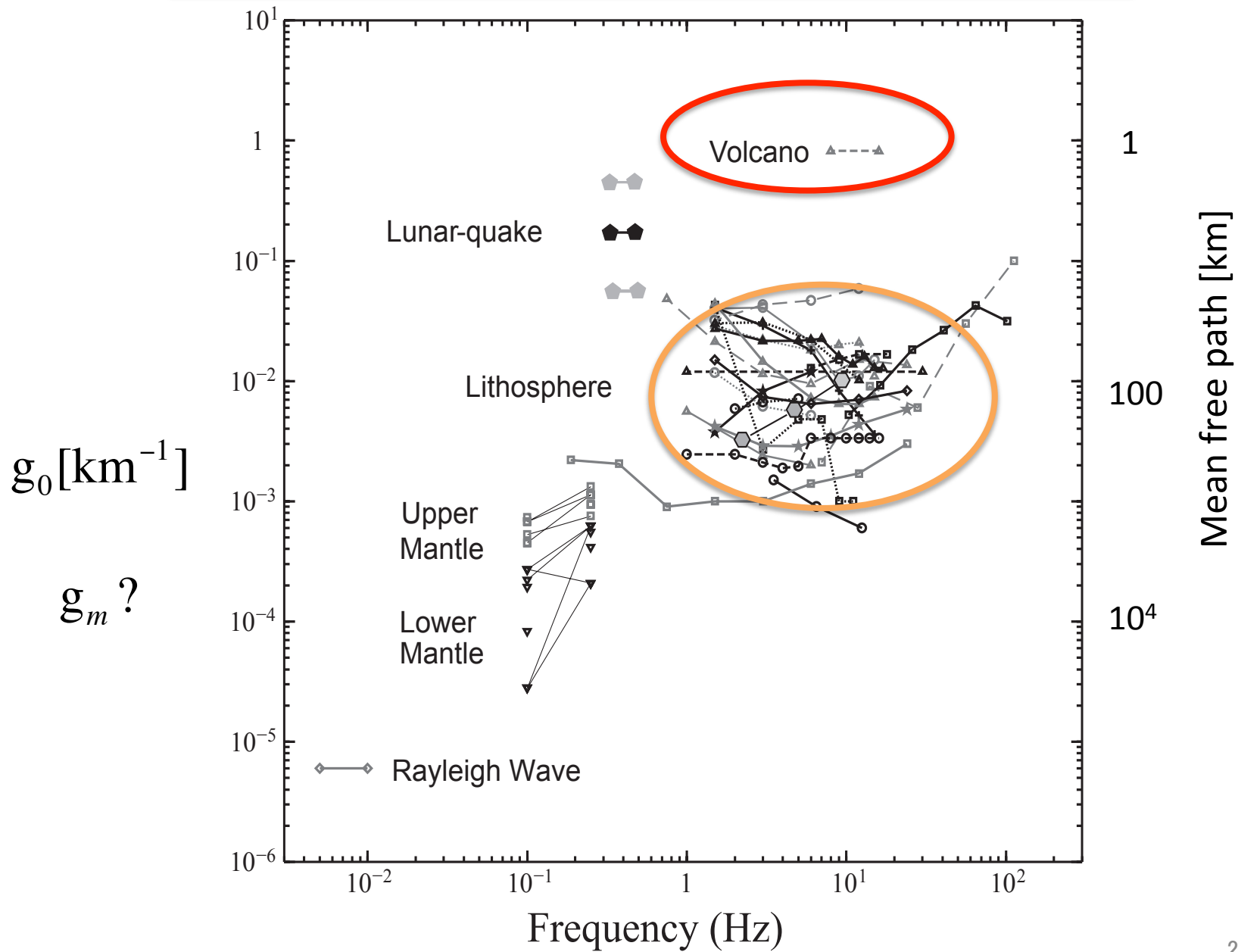
$$\alpha_0 = 2.7 \text{ km/s}, \quad \beta_0 = 1.56 \text{ km/s}$$

$$g_0^{PP} = 0.26 \text{ km}^{-1}, \quad g_0^{PS} = 0.63 \text{ km}^{-1}$$

$$g_0^{SP} = 0.11 \text{ km}^{-1}, \quad g_0^{SS} = 0.79 \text{ km}^{-1}$$

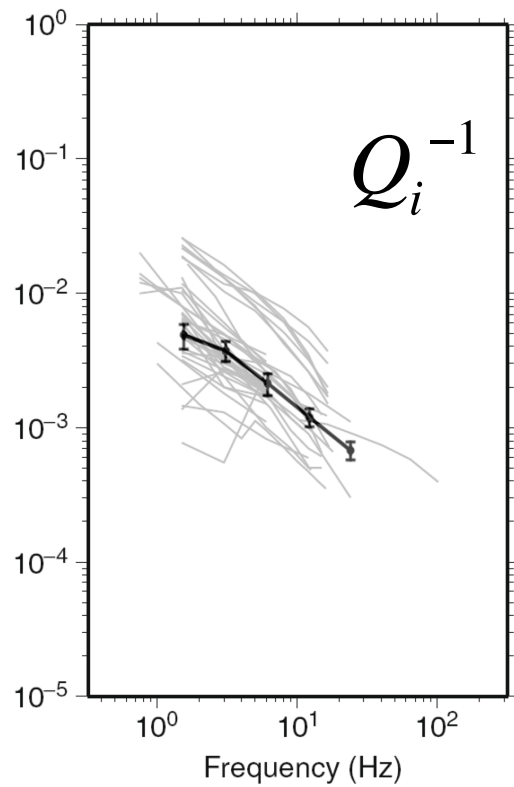
The estimated  $Q_i^{-1} = 0.01$  may contain intrinsic absorption and window effect.

# Total Scattering Coefficient in the Lithosphere

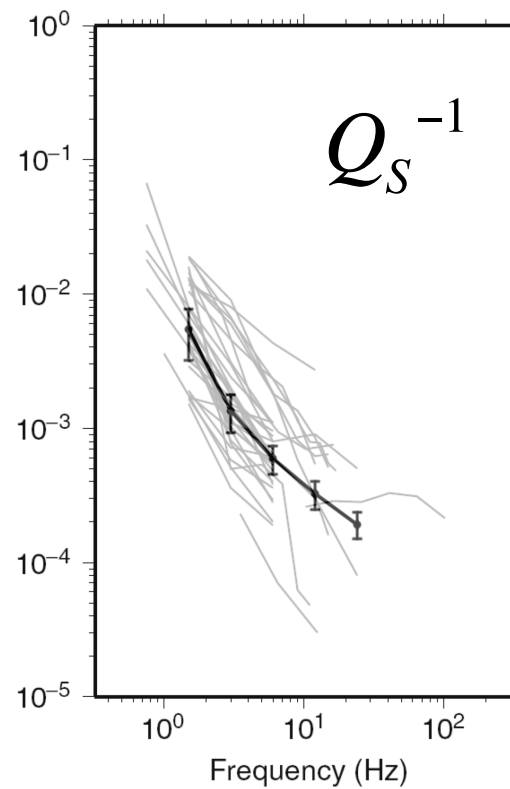


# Frequency Dependence of Attenuation and Scattering in Japan

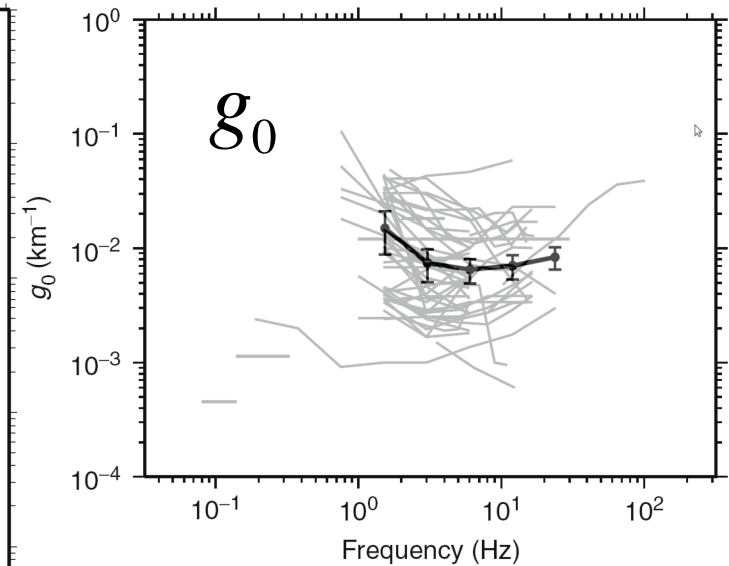
Intrinsic Absorption



Scattering Loss



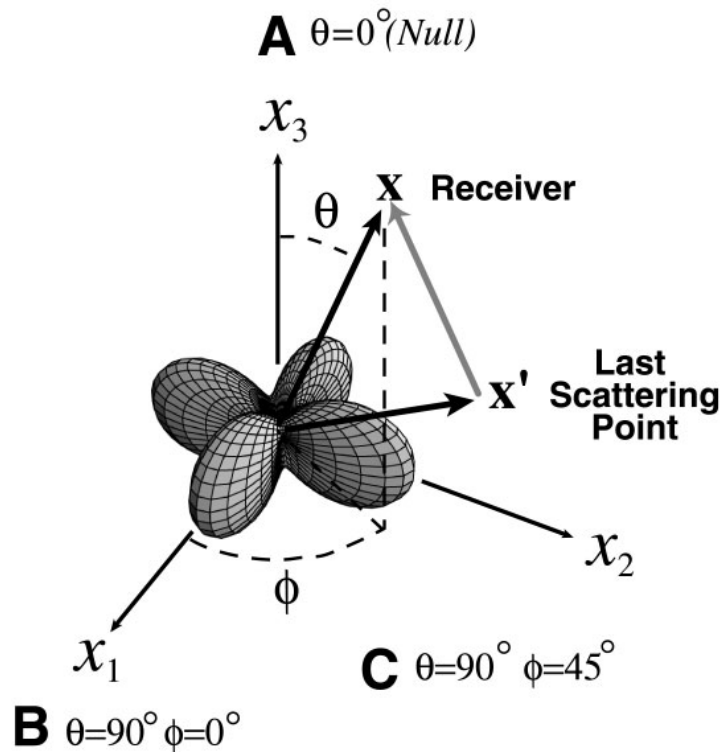
Scattering Coefficient



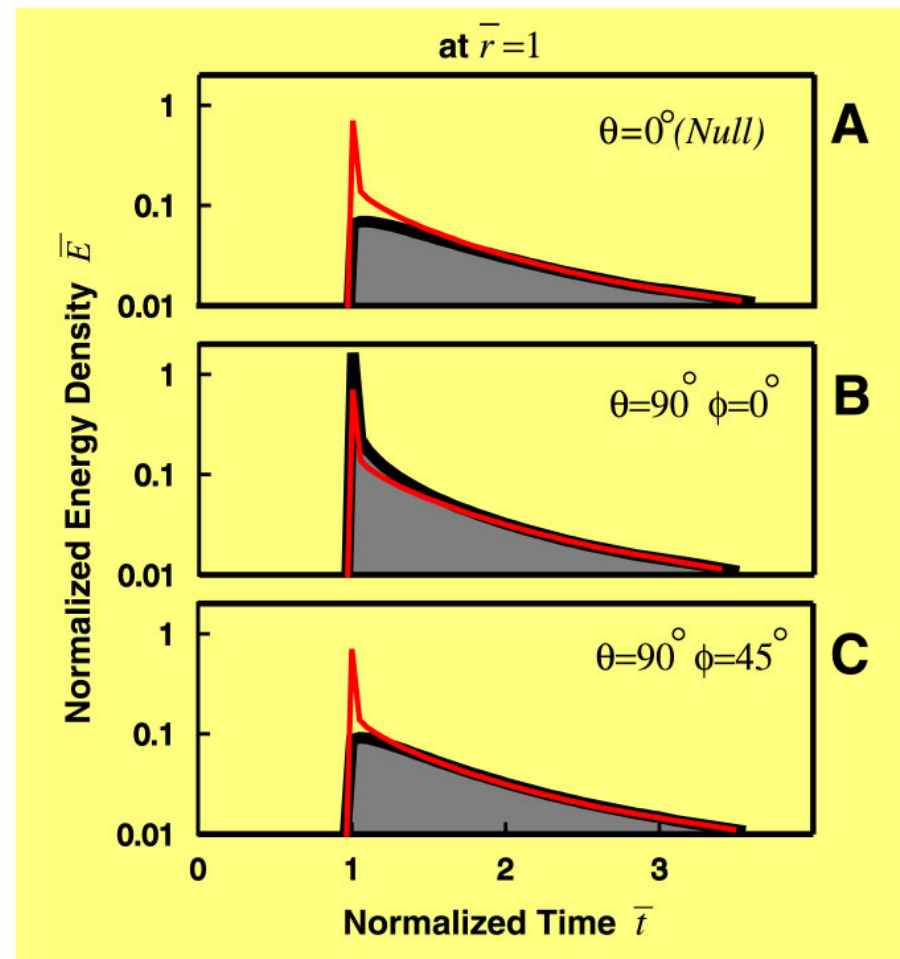
$$E = W\Psi G + V_0 g_0 G * E$$

$$G(\mathbf{x}, t) = \frac{1}{4\pi V_0 r^2} H(t) \delta\left(t - \frac{r}{V_0}\right) e^{-(V_0 g_0 + b)t}$$

$$\oint d\Omega \Psi(\theta, \phi) = 4\pi$$



**Point Shear Dislocation Source**

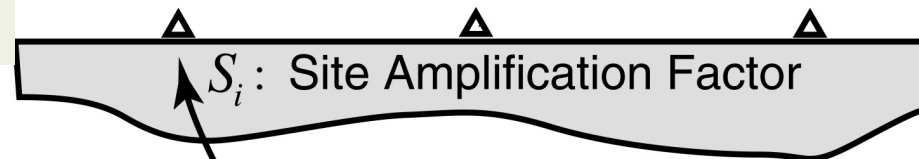


$$\bar{r} = g_0 r \text{ and } \bar{t} = g_0 V_0 t \quad \text{Sato et al. [1997]}$$

# Seismogram Envelope Inversion for the High Frequency Energy Radiation Hisotry

Deconvolution of observed S-wave envelopes by using the envelope Green function

$E_{ij}$ : Energy Density (MS Amp.)  
at the  $i$ -th Station at the  $j$ -th Time

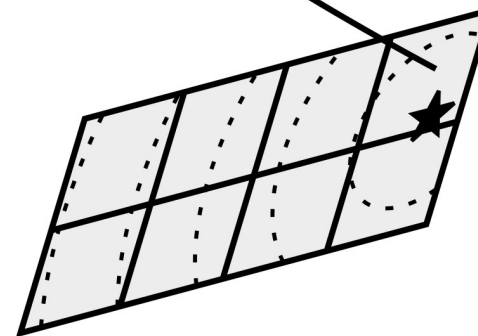


$S_i$ : Site Amplification Factor

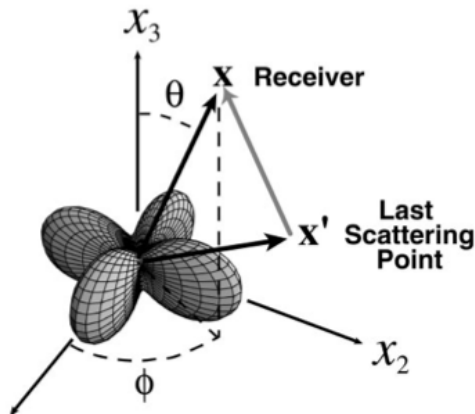
$F_{ijk}$ : Envelope Green Function for  
a Point-Shear Dislocation Source



$W_k$ : Radiated Energy  
from the  $k$ -th Subfault



$$E_{ij} = S_i \sum_k F_{ijk} W_k$$



Point-shear dislocation source

$$\sum_{i,j} \frac{1}{\sigma_i^2} \left| E_{ij} - S_i^{(n-1)} \sum_k F_{ijk} W_k^{(n)} \right|^2 \Rightarrow \text{Min}$$

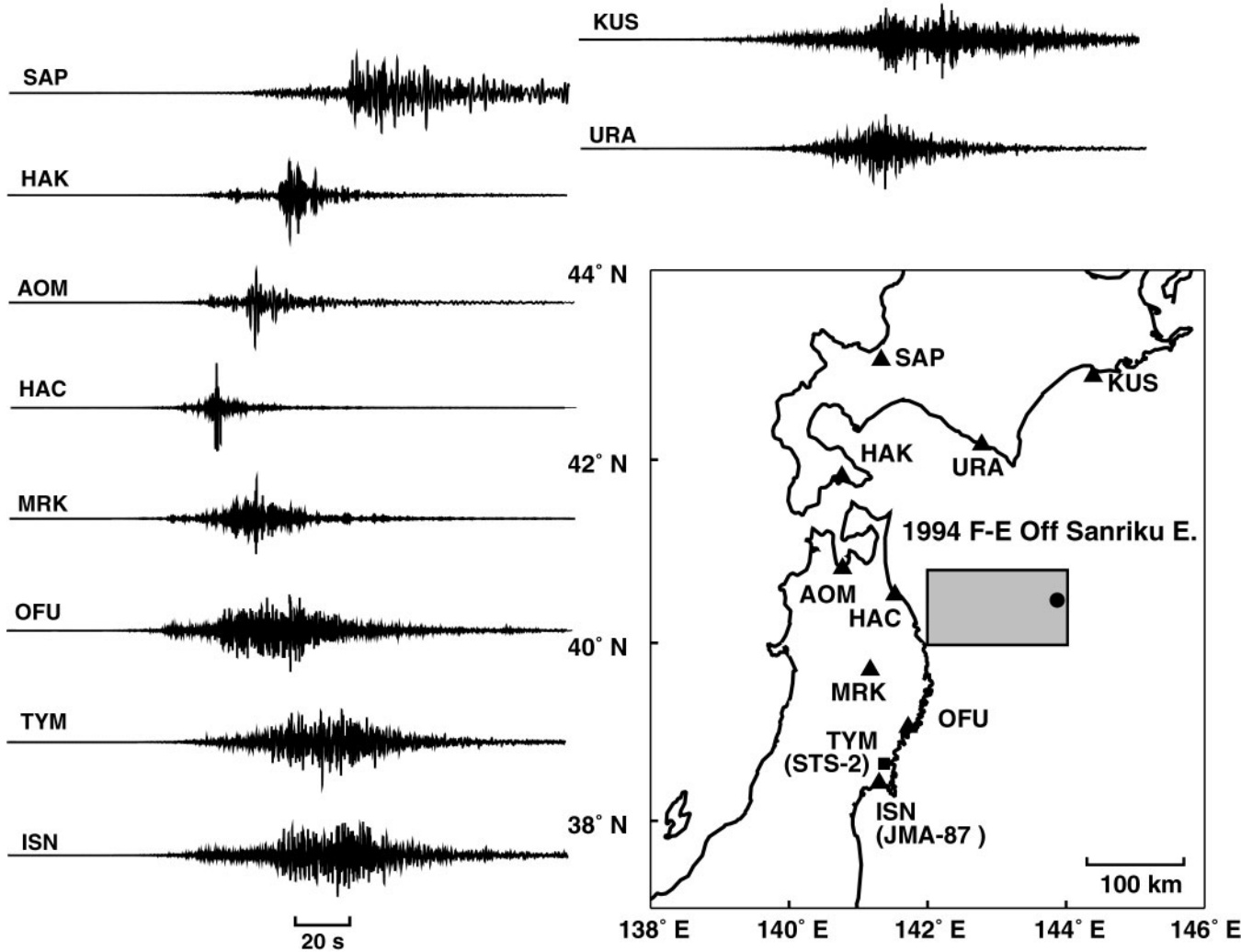
where  $\sigma_i = E_{ij}|_{\text{Max}}$

Positive constraint for  $W_k$

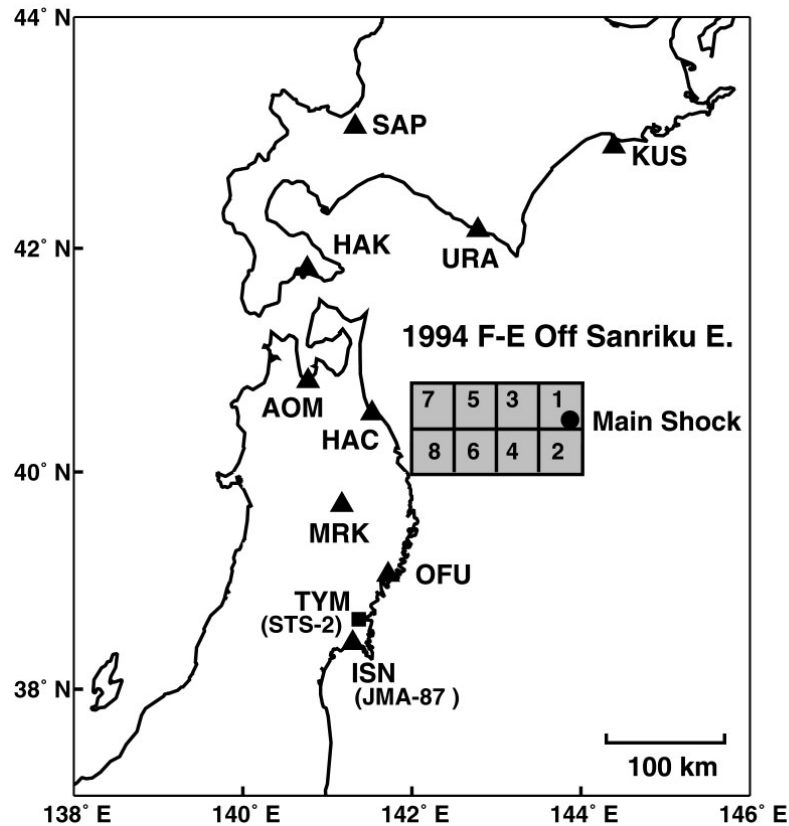


# 1994 Far-East off-Sanriku Earthquake (Mw 7.7)

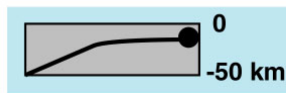
## Velocity Seismograms (EW-Comp. >1 Hz )



# Envelope Inversion Results



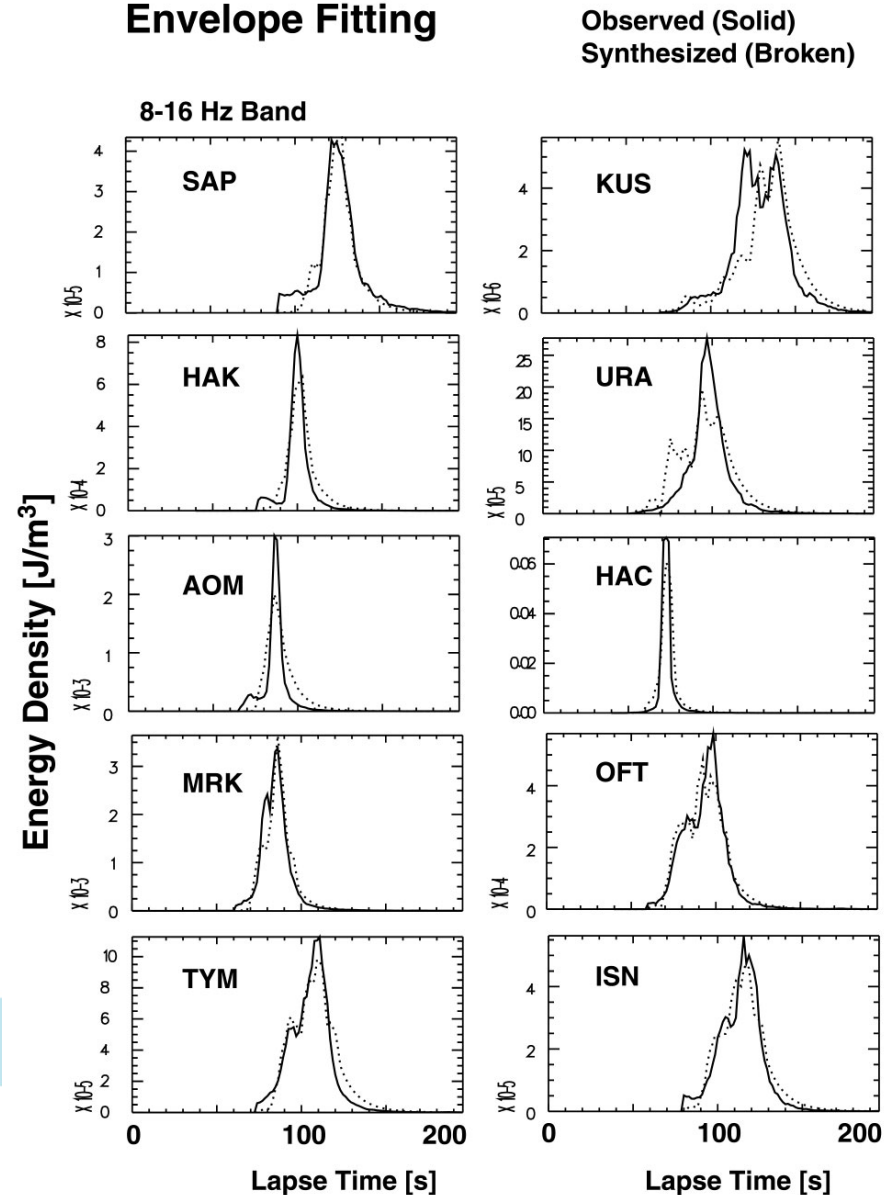
Strike  $184^\circ$ , Rake  $70^\circ$   
 Dip  $0^\circ$  (E)  $\sim$   $30^\circ$  (W)  
 After Harvard CMT  
 $V_{rup}=2.7$  km/s



$V_s=3.9$  km/s,  $g_0=0.005$  km $^{-1}$  for 1-16 Hz  
 $Q_i=400$  (1-2 Hz)  $\sim$  1500 (8-16 Hz)

After Hoshiba[1993] and Sakurai et al. [1995]

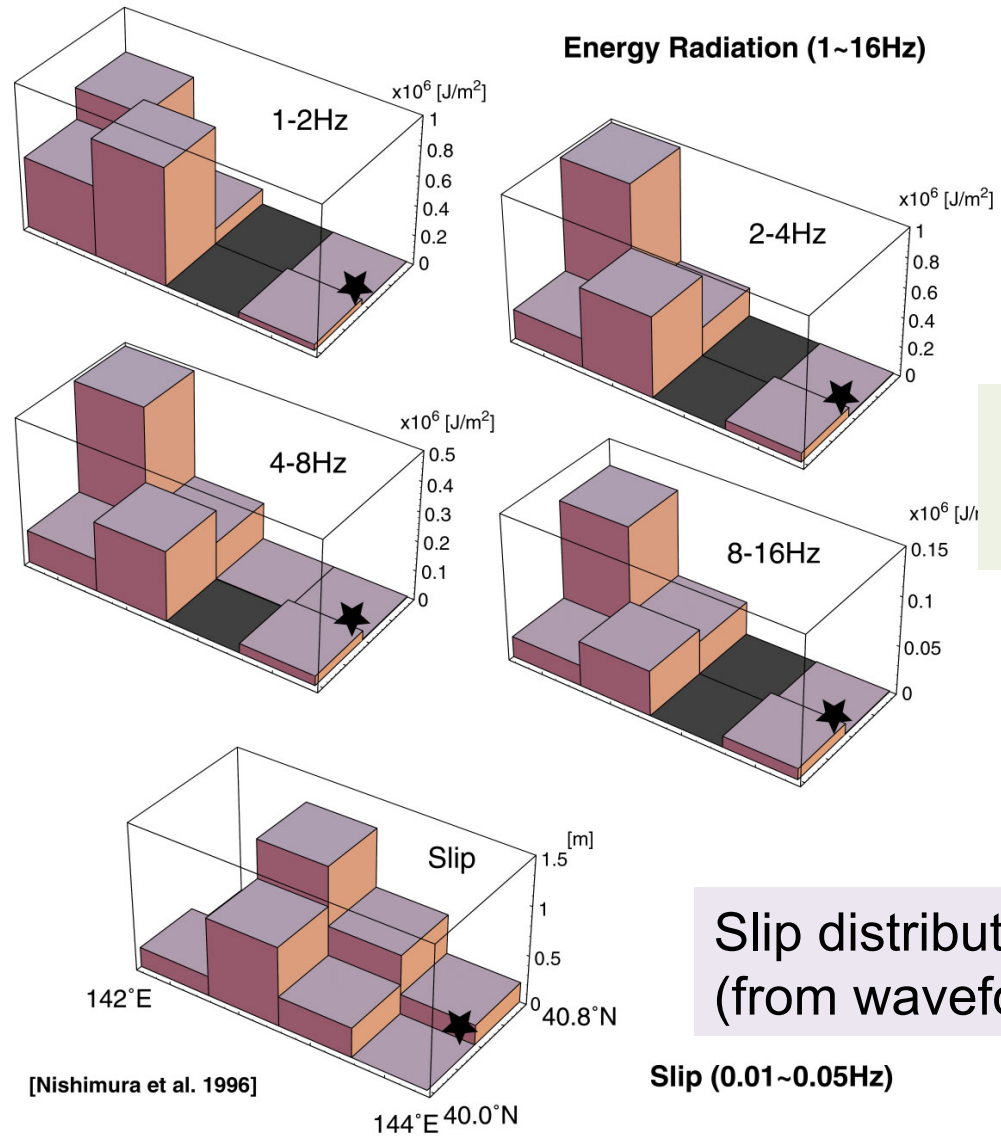
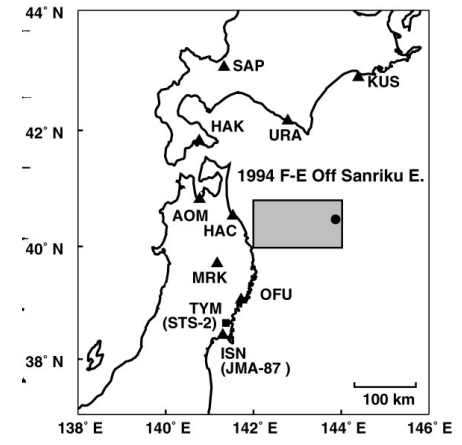
## 1994 Far East Off Sanriku Earthquake Envelope Fitting



Nakahara et al. [1998]

(Free Surface Corrected)

# 1994 Far East Off Sanriku Earthquake (Mw=7.7)



Radiation of high freq. energy (from envelope inversion)

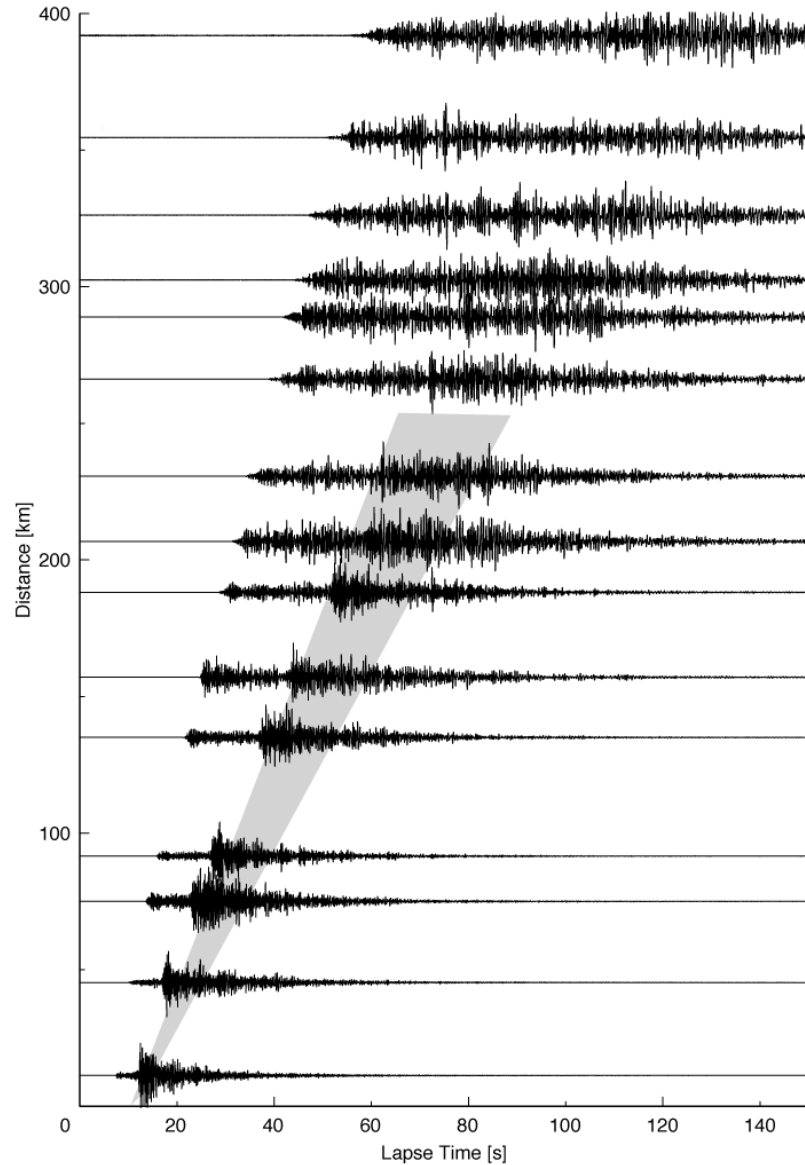
Slip distribution (from waveform inversion)

[Nishimura et al. 1996]

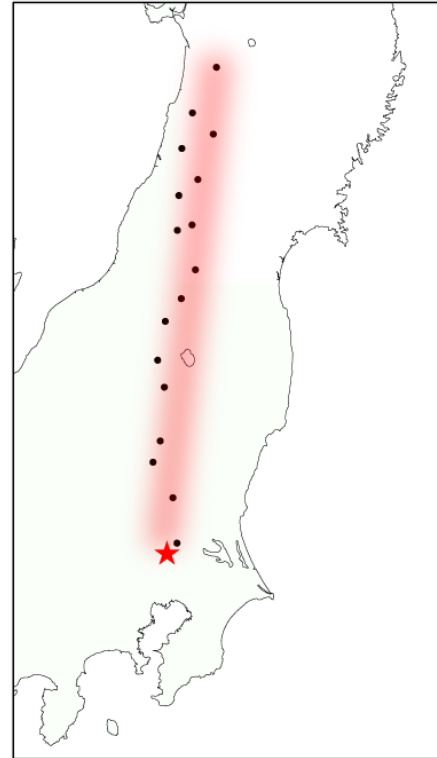
[Nakahara et al. 1998]

# Envelope Broadening with Travel Distance Increasing

Feb. 16, 2005, 04:46:36.21 JST, Southern Ibaraki, Japan  
130.8763E, 36.0360N, Depth 53 km, Mw=5.4  
Hi-net Transverse Component Traces, 2-32 Hz

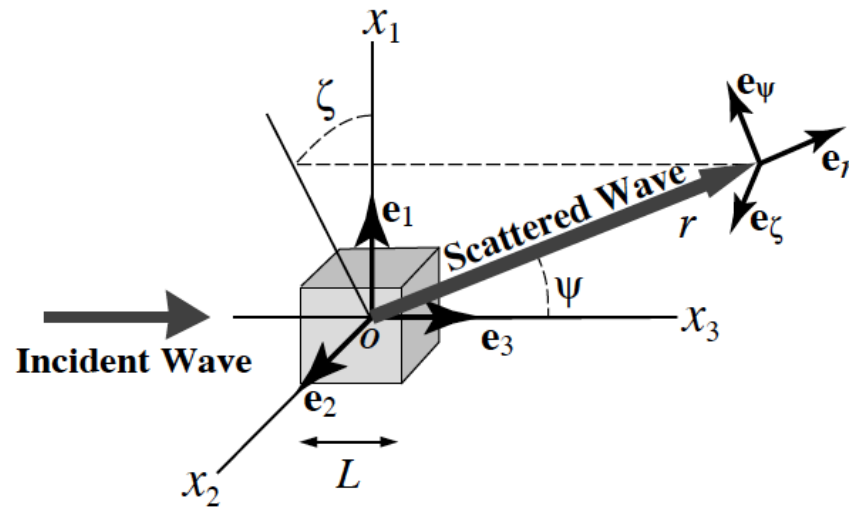


Focus on envelope width and peak delay from the S-onset



Courtesy of T. Maeda

Waves around the direct arrival are forward scattered waves (diffracted waves).



Strong forward scattering (diffraction) for  $\lambda \ll a$

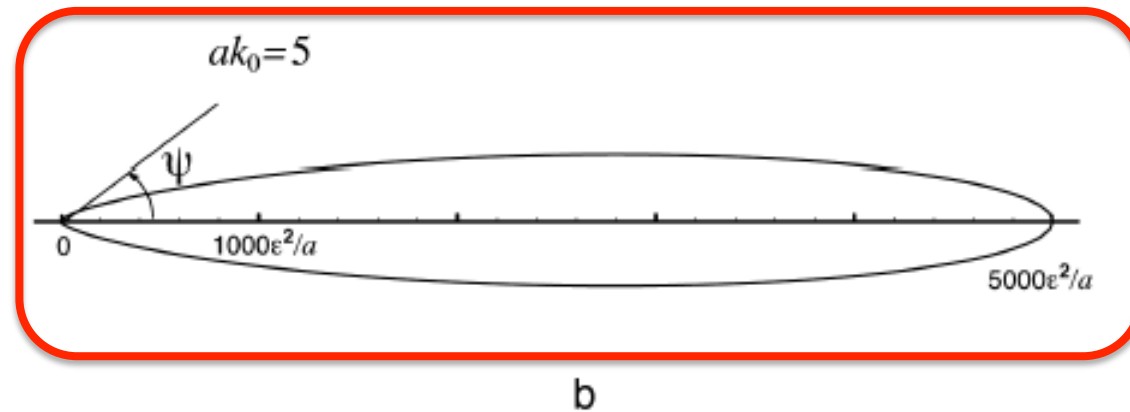
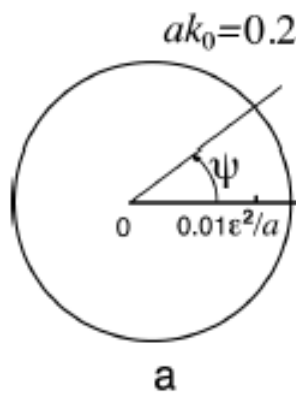


Fig. 4.3 Angular dependence of scattering coefficient (4.29) for scalar waves in 3-D random media characterized by an exponential ACF: (a) Low wavenumber case. (b) high wavenumber case. 29

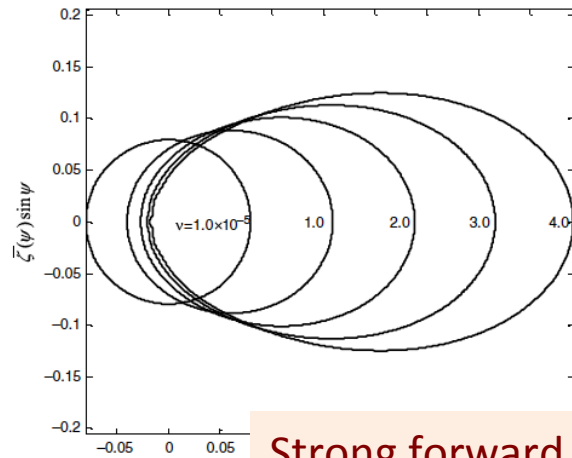
## Radiative Transfer Equation (Boltzmann equation) in Scattering Media

$$\begin{aligned} \partial_t f(\mathbf{x}, t; \mathbf{q}) + V_0 \mathbf{q} \nabla f(\mathbf{x}, t; \mathbf{q}) = & -g_0(k_0) V_0 f(\mathbf{x}, t; \mathbf{q}) \\ & + \frac{V_0}{4\pi} \oint g(k_0 \mathbf{q} - k_0 \mathbf{q}') f(\mathbf{x}, t; \mathbf{q}') d\Omega_{q'} + W(\omega) \frac{\Psi(\mathbf{q})}{4\pi} \delta(\mathbf{x}) \delta(t), \end{aligned}$$

$$E(\mathbf{x}, t) = \oint f(\mathbf{x}, t, \mathbf{q}) d\Omega_q \qquad g_0(k_0) = \frac{1}{4\pi} \oint g(k_0 \mathbf{q}) d\Omega_q$$

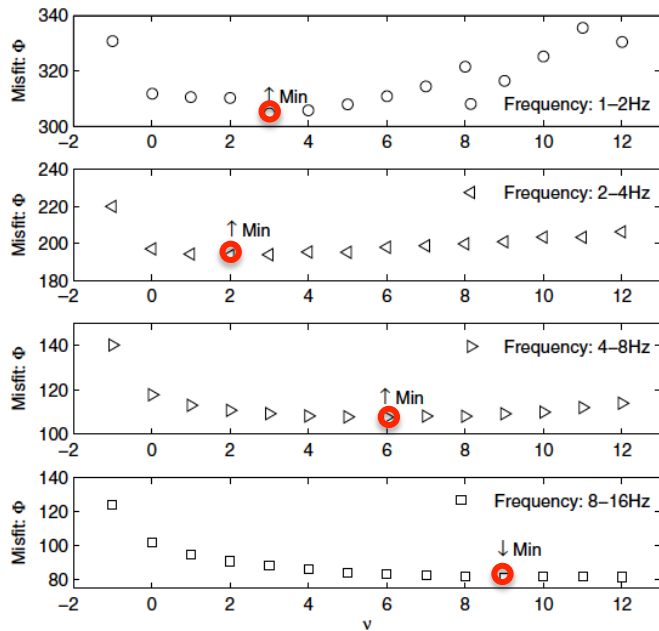
Using the Fourier transform of  $f$  in space and time, and spherical harmonic expansion of  $g$  for scattering angle, we can write down the above equation as **linear equations using Wigner 3- $j$  symbols** [Sato, 1995]

Inversion of scattering pattern from the envelope analysis of the 2008 Wells Earthquakes in Nevada, USA



$$g(\psi) \propto \frac{1 + \nu}{\left(1 + \nu \sin^2 \frac{\psi}{2}\right)^2}$$

Strong forward scattering (diffraction)=Envelope broadening



RMS envelope in 8-16 Hz (Event 4)

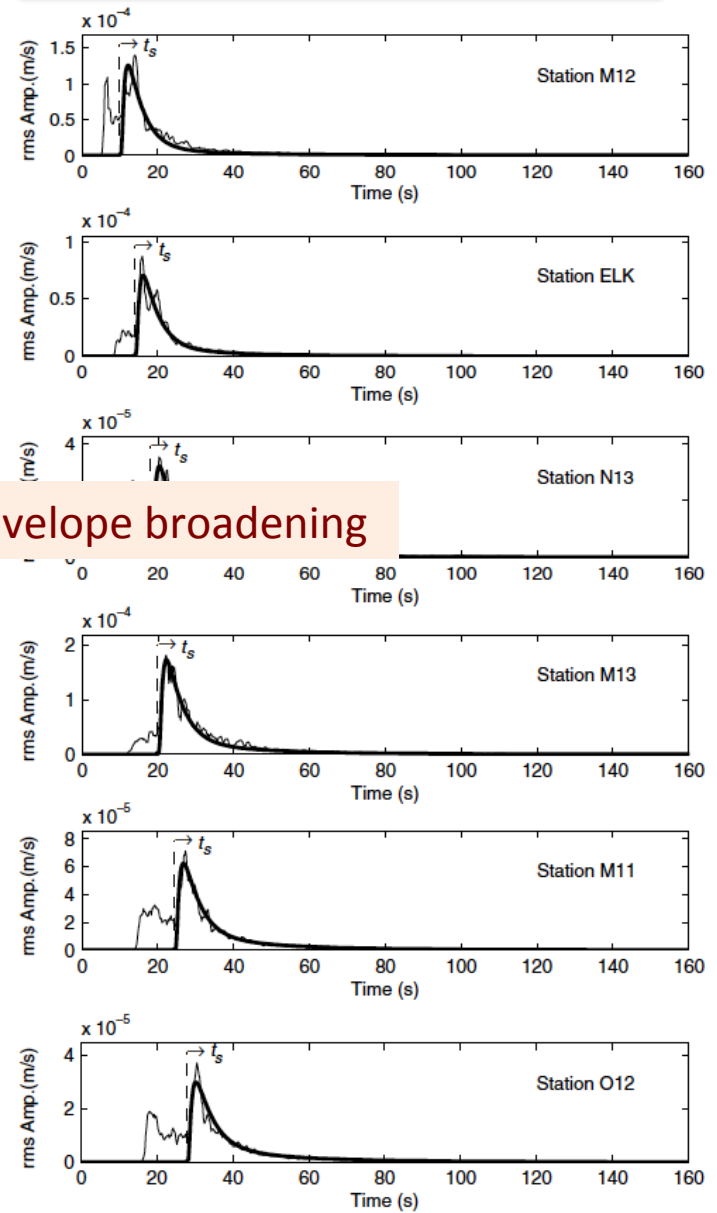


Figure 6. The change of misfit of event 4 with different scattering patterns. Arrows denote the minimum of misfits.

## Markov approximation: stochastic extension of the phase screen method

Spherically outgoing wavelet:  $u(r, \theta, \varphi, \omega) = \frac{1}{2\pi r} \int_{-\infty}^{\infty} U(\mathbf{x}_{\perp}, r, \omega) e^{ik_0 r - i\omega t} d\omega,$

When  $a k_c \gg 1$ , we may apply the **parabolic approximation**:

$$2ik_c \partial_r U + \Delta_{\perp} U - 2k_c^2 \xi U = 0$$

We define the **two-freq. mutual coherence func. (TFMCF)** on the transverse plane:

$$\Gamma_2(\mathbf{x}_{\perp c}, \mathbf{x}_{\perp d}, r, \omega', \omega'') \equiv \langle U(\mathbf{x}'_{\perp}, r, \omega') U(\mathbf{x}''_{\perp}, r, \omega'')^* \rangle$$

Its Fourier transform gives the intensity spectral density:

$$\widehat{I}(r, t; \omega_c) = \frac{1}{2\pi r^2} \int_{-\infty}^{\infty} d\omega_d e^{-i\omega_d(t-r/V_0)} \Gamma_2(\mathbf{x}_{\perp c}, \mathbf{x}_{\perp d} = 0, r, \omega_c, \omega_d)$$

Longitudinal integral of  $R$ :  $A(r_{\perp d}) = 2 \int_0^{\infty} R(r_{\perp d}, z) dz$  ← Delta correlation  
 $R(\mathbf{x}) = A(\mathbf{x}_{\perp}) \delta(z)$

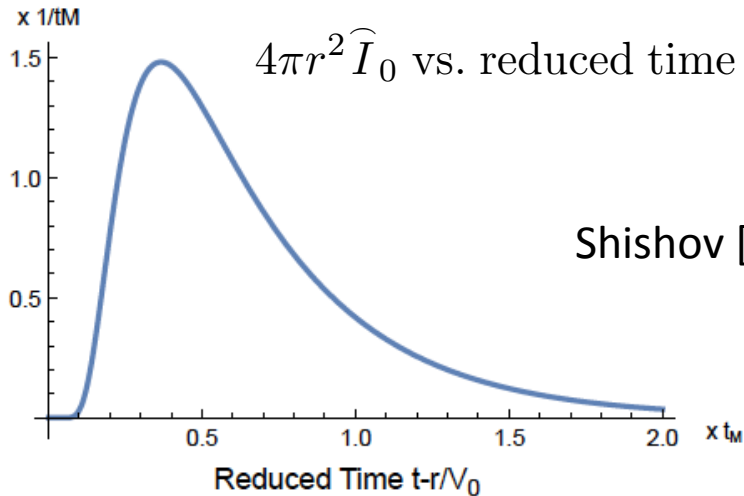
To solve the master equation for a quasi-monochromatic wavelet around  $\omega_c$ :

$$\partial_r \Gamma_2 + \underbrace{i \frac{k_d}{2k_c^2} \Delta_{\perp d} \Gamma_2 + k_c^2 (A(0) - A(r_{\perp d})) \Gamma_2}_{\text{Diffraction effect}} + \underbrace{\frac{k_d^2}{2} A(0) \Gamma_2}_{\text{Wandering effect}} = 0.$$

Initial condition:  $\Gamma_2(r \sim 0, t) = \frac{1}{4\pi}$        $\widehat{I}(r \sim 0, t) = \frac{1}{4\pi r^2} \delta\left(t - \frac{r}{V_0}\right)$



# Synthesized spherical envelope in 3-D random media based on the Markov approximation



Gaussian ACF:  $R(\mathbf{x}) = \varepsilon^2 \exp[-r^2 / a^2]$

Characteristic time of envelope width:

$$t_M = \frac{\sqrt{\pi}\varepsilon^2}{2V_0 a} r^2 \propto r^2$$

$$\widehat{I}_0(r, t, \omega_c) = \frac{1}{4\pi r^2} \frac{\pi^2}{16 t_M(r)} \vartheta_4'' \left( 0, e^{-\frac{\pi^2}{4} \frac{(t-r/V_0)}{t_M(r)}} \right)$$

Elliptic theta func.  $\vartheta_4(v, q) = 1 + 2 \sum_{n=1}^{\infty} (-1)^n q^{n^2} \cos 2nv$

$$\widehat{I} = \widehat{I}_0 * w$$

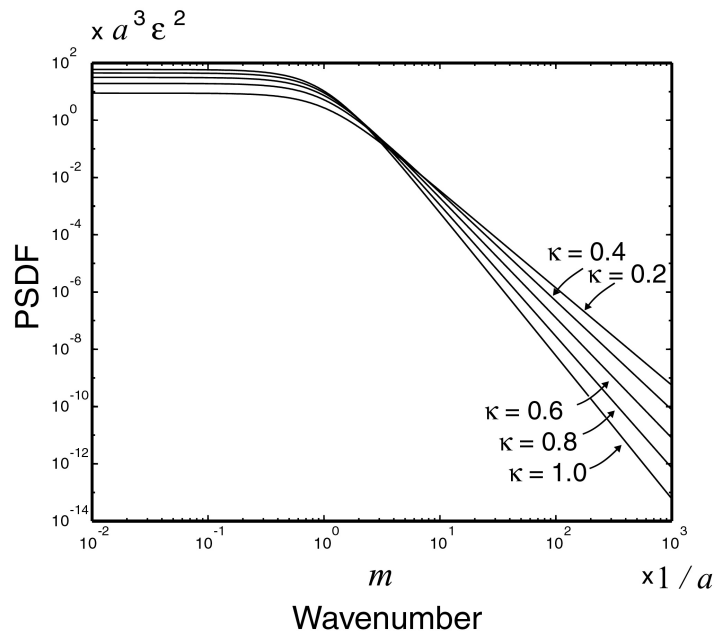
Freq. independent

Diffraction effect    Wandering effect

$$w(r, t) \propto e^{-\frac{t^2}{t_w(r)^2}}$$

$$t_w = \sqrt{\frac{2\sqrt{\pi}\varepsilon^2 ar}{V_0^2}} \propto \sqrt{r}$$

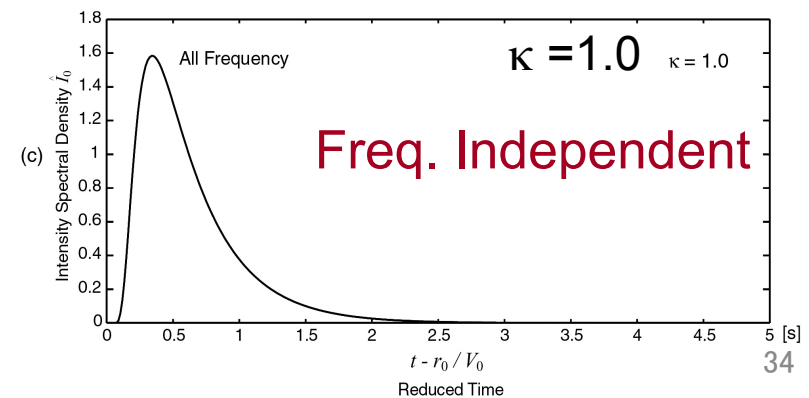
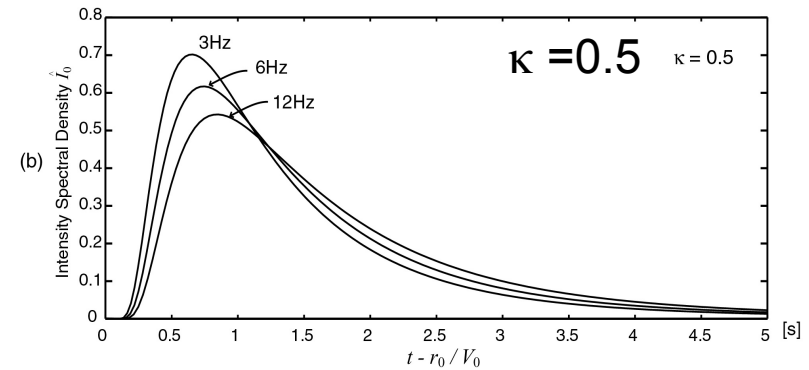
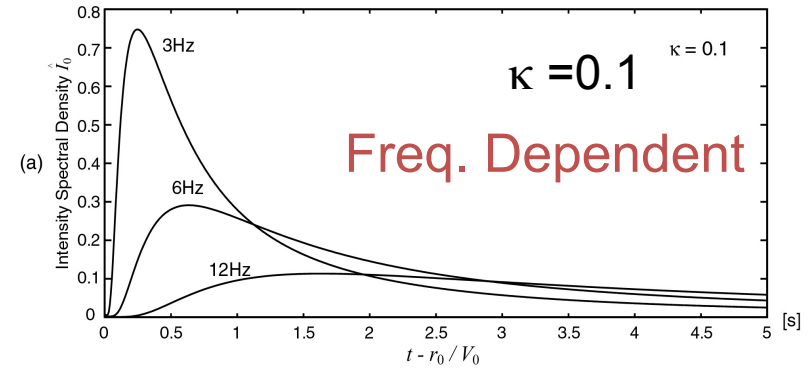
# MS Envelopes of spherical wavelet in von Karman-type Random Media



$$a k_c \gg 1$$

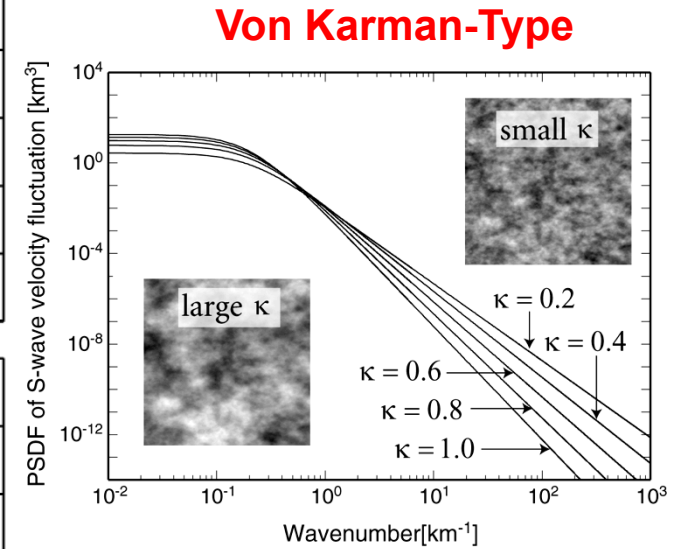
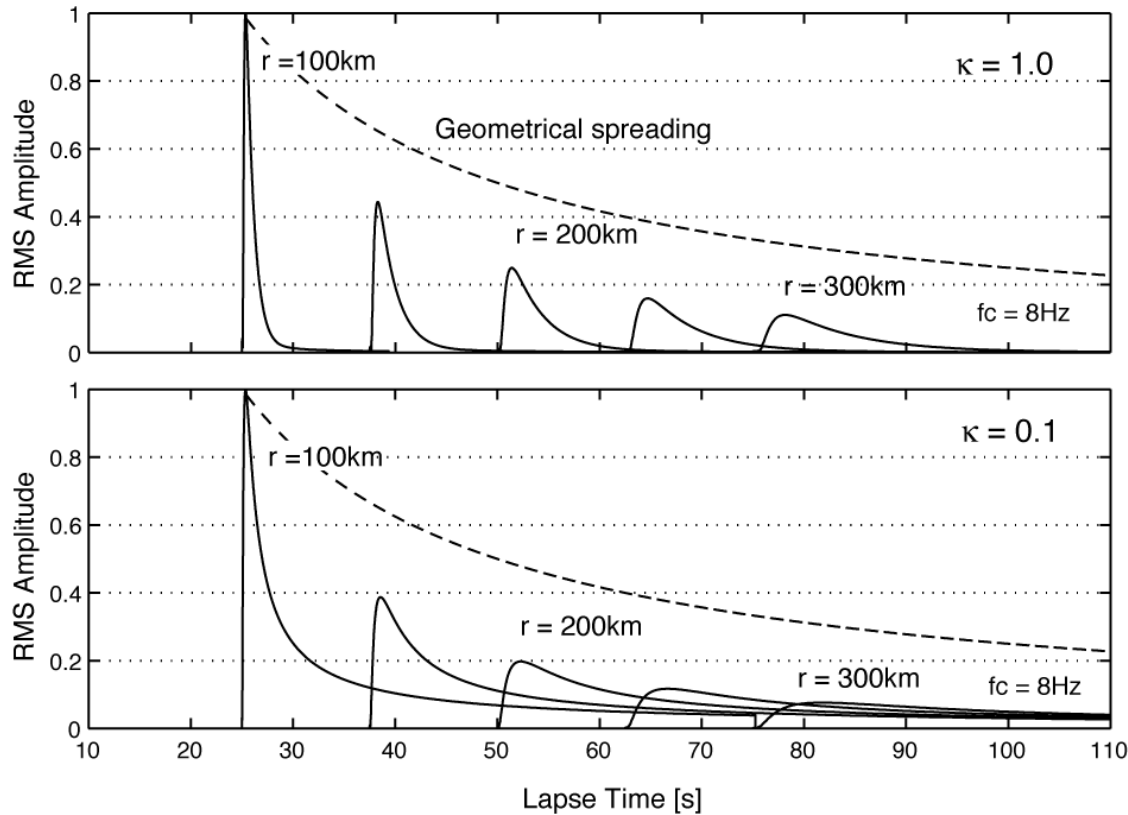
Saito et al. [2001]

$\widehat{I}_0$  at 100km,  $\epsilon=0.05$ ,  $a=5$  km



# Envelope Broadening of Spherical Scalar Wavelet in 3-D

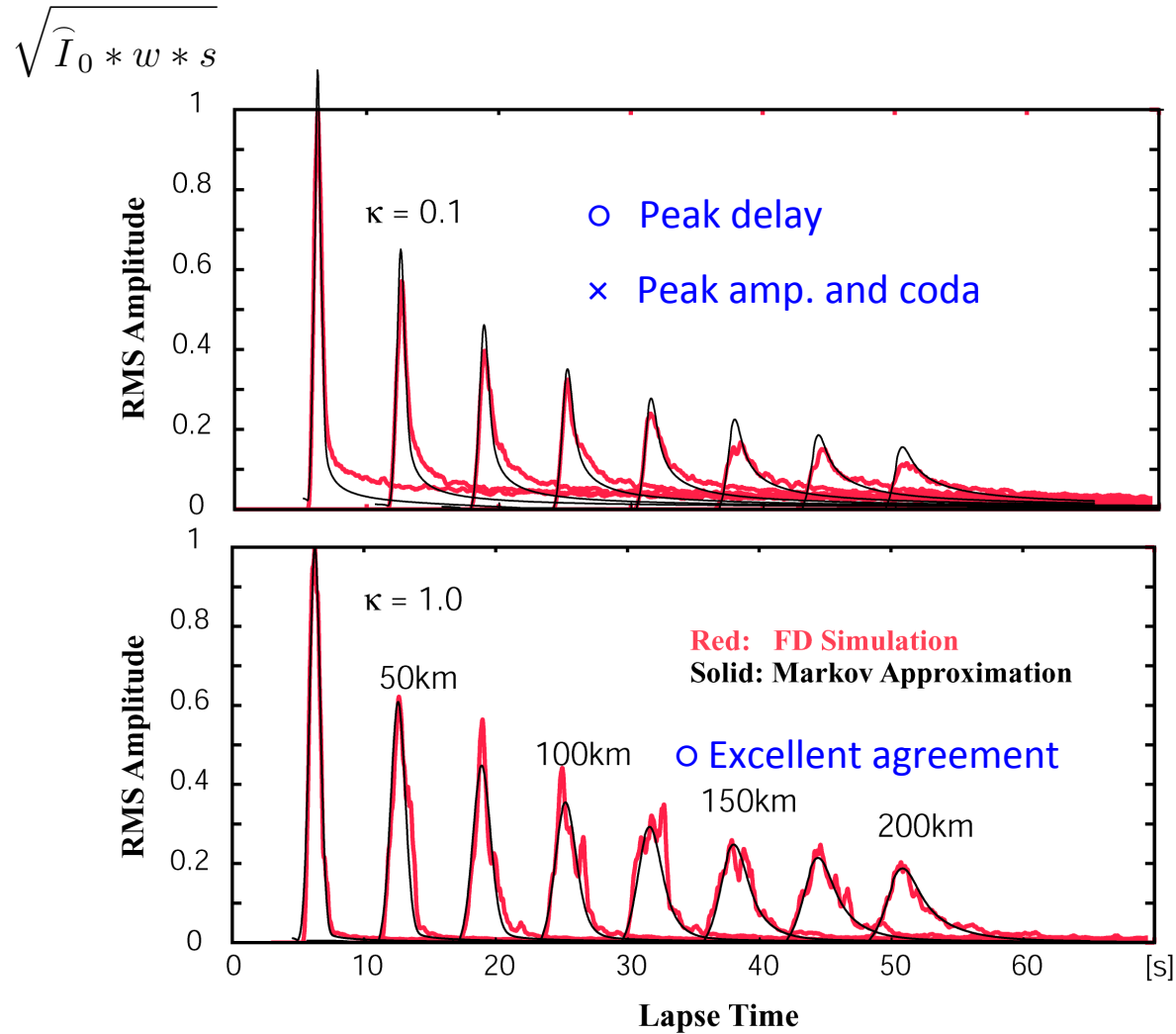
$$\sqrt{\widehat{I}_0}$$



$$P(m) = \frac{8\pi^{3/2} \varepsilon^2 a^3 \Gamma(\kappa + 3/2)}{\Gamma(\kappa) (1 + a^2 m^2)^{\kappa + 3/2}}$$

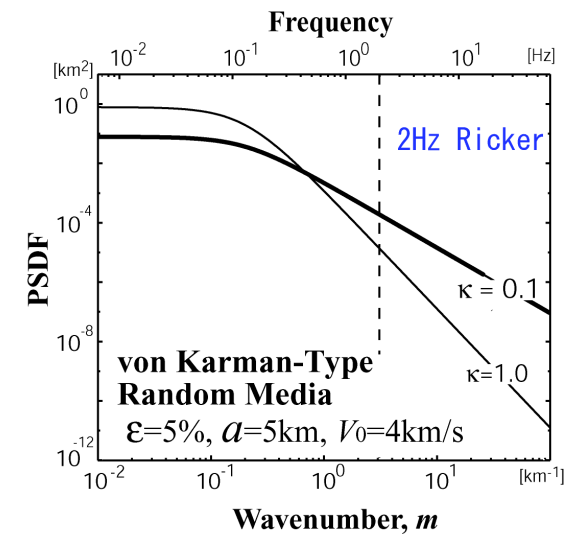
For an impulsive source radiation, the peak delay from the onset and the envelope width increase with travel distance.

# Comparison of Markov- and FD-Envelopes in 2-D

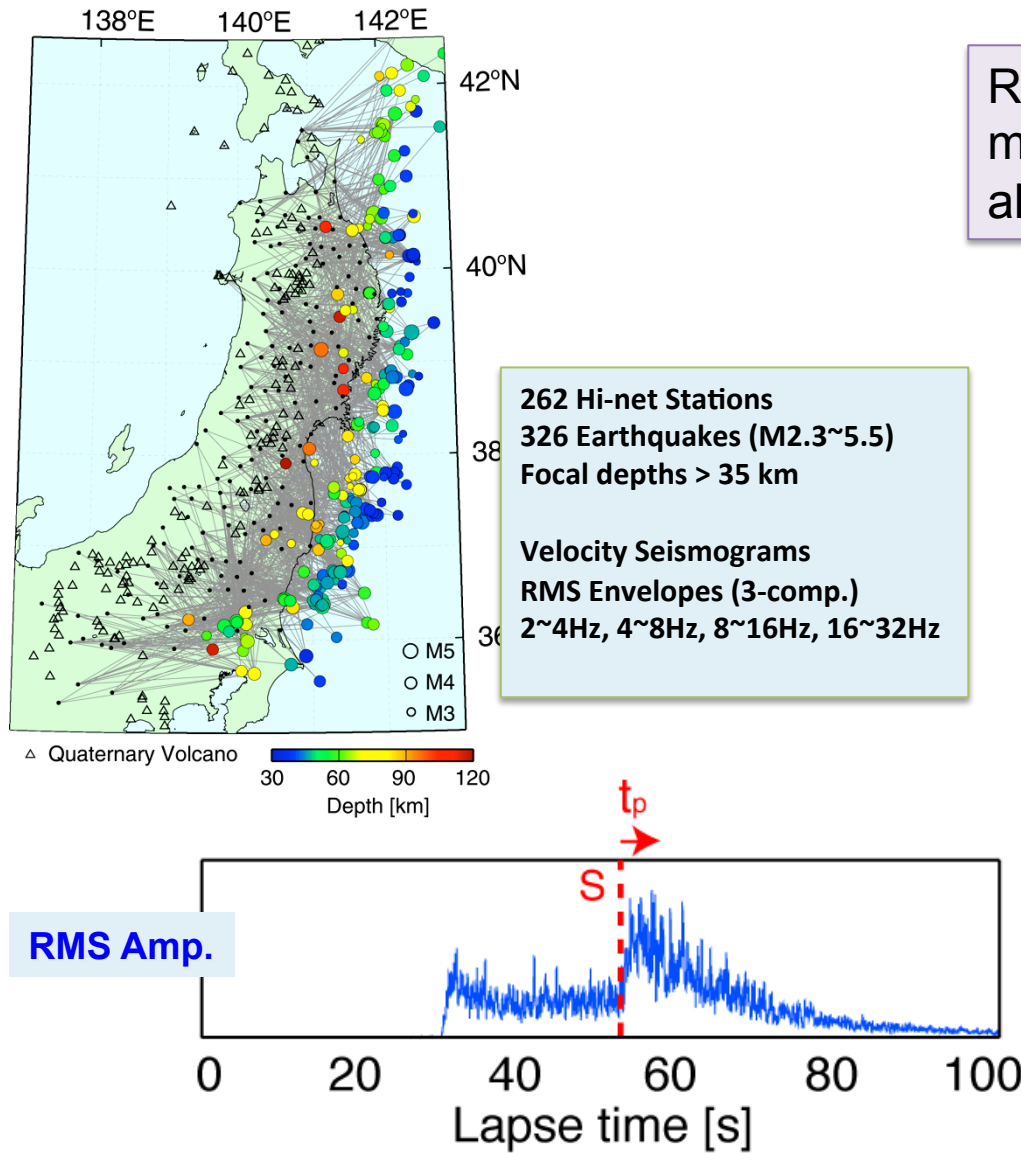


## Envelopes of 2Hz Ricker Wavelet in von Karman-Type Random Media

$\epsilon=0.05, a=5\text{km}$

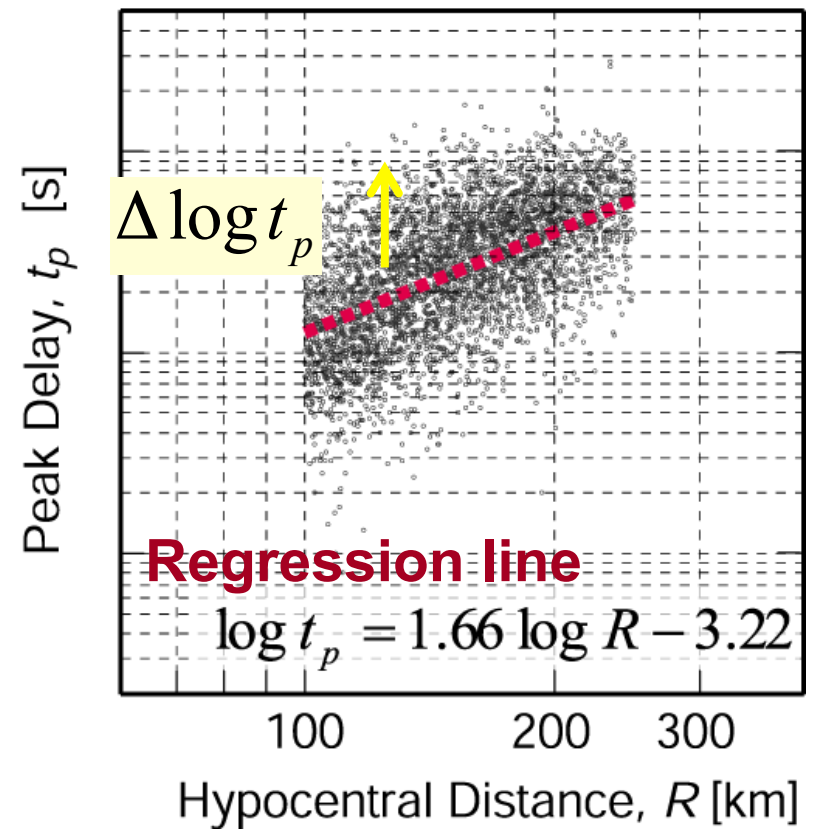


# Peak delay ( $t_p$ ) from the S-onset vs. travel distance

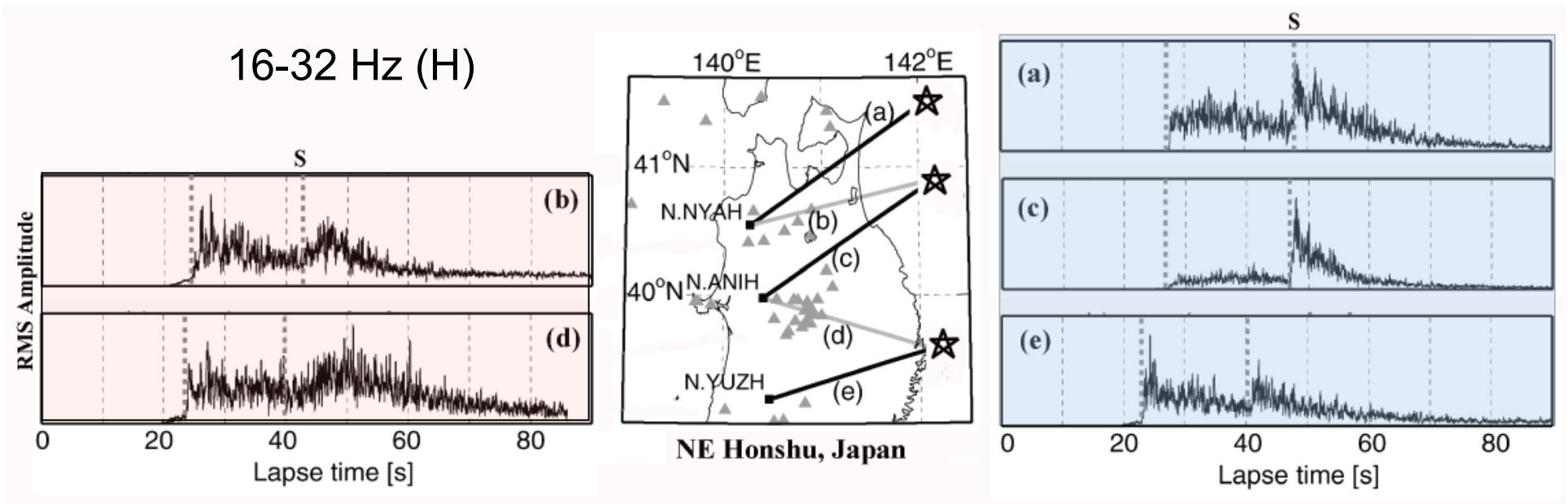


Residual from the regression line means relative strength of scattering along the ray

S-wave (16~32Hz)



**Peak delay is large for a ray traveling beneath Quaternary volcanoes**



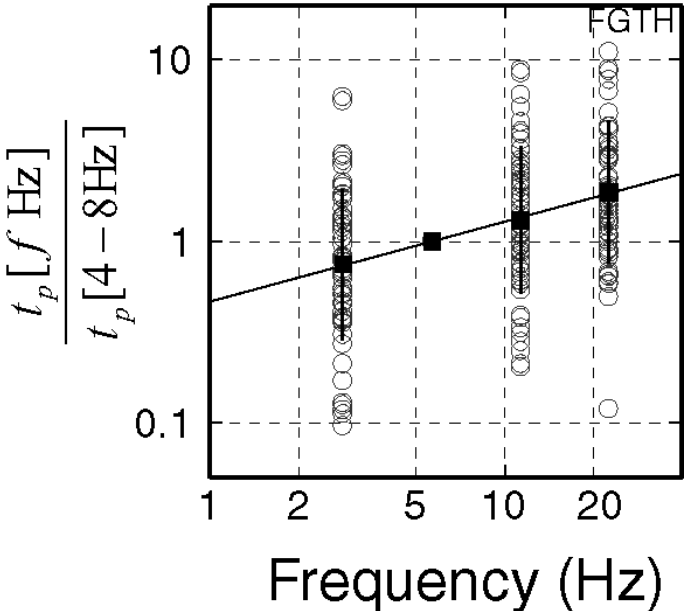
Paths beneath Q. volcanoes

Paths between Q. volcanoes

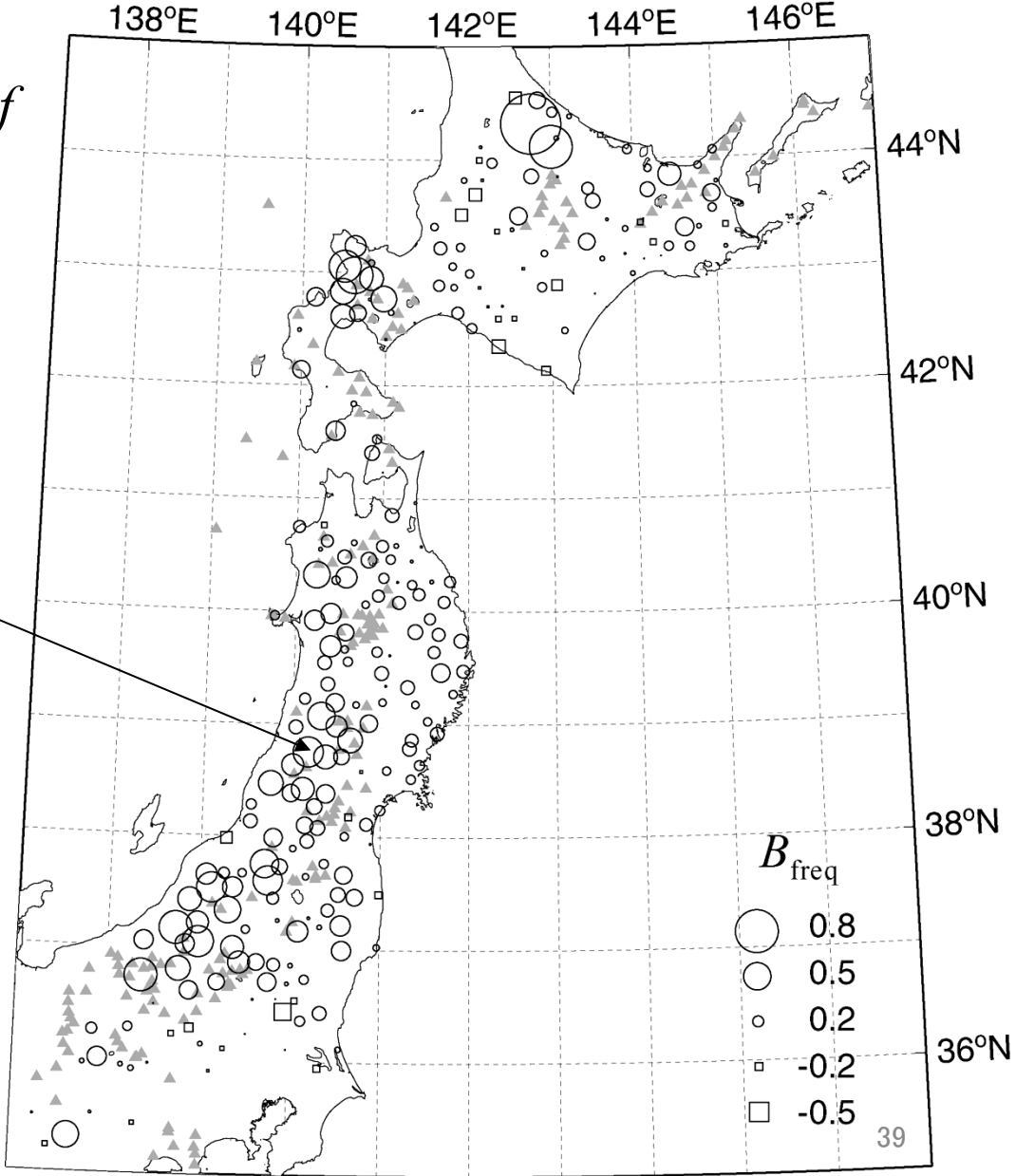
**Strong inhomogeneity beneath Quaternary volcanoes**

# Frequency dependence of peak delay in Tohoku, Japan

$$\log\left(\frac{t_p[f]}{t_p[4-8\text{Hz}]}\right) = A_{\text{freq}} + B_{\text{freq}} \log f$$

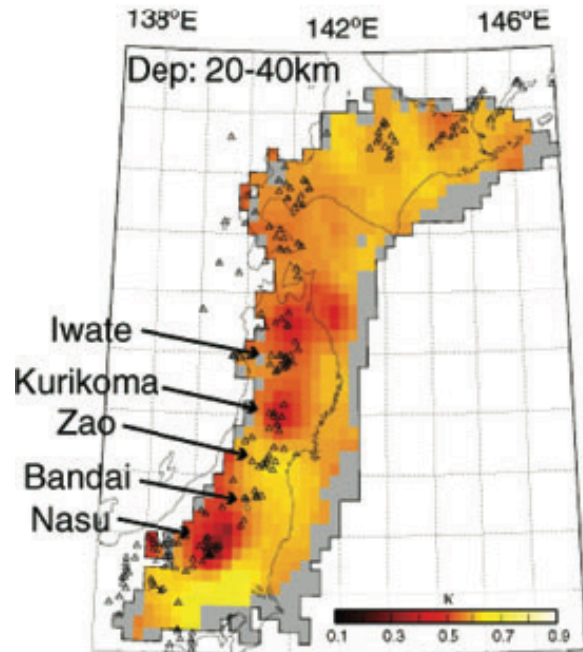


Strong freq. dependence at the back-arc side of the Quaternary Volcanoes.



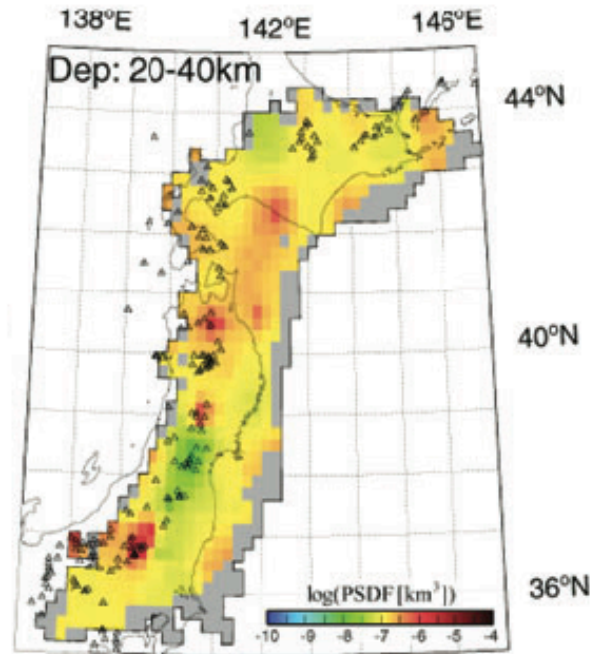
# Inversion of PSDF of Velocity Frac. Fluc.

$\kappa$  Spectrum Role-off



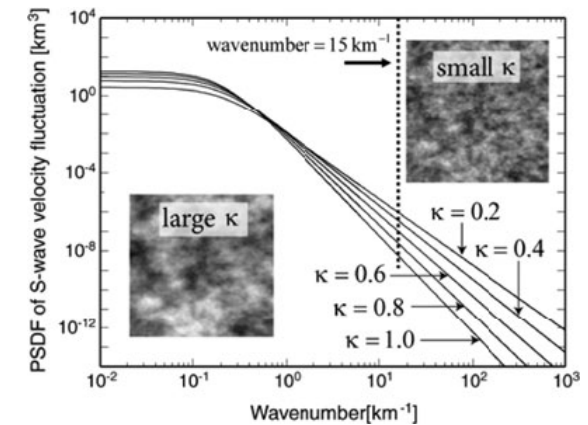
PSDF ( $m=15\text{km}^{-1}$ )

S-wave 10Hz



Δ Quaternary Volcano

von Karman-type random media

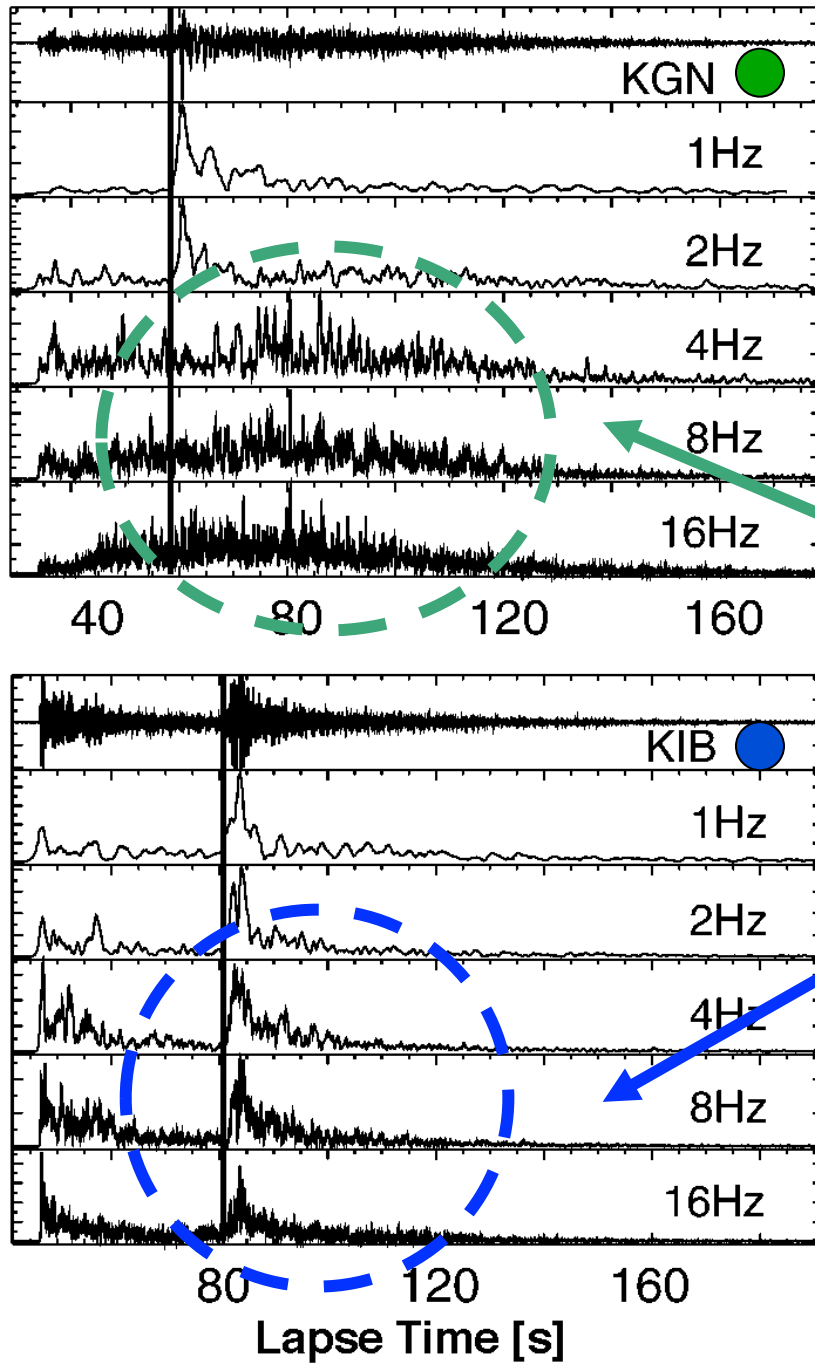


Spectral role-off is small and random inhomogeneity is strong especially beneath Quaternary volcanoes.

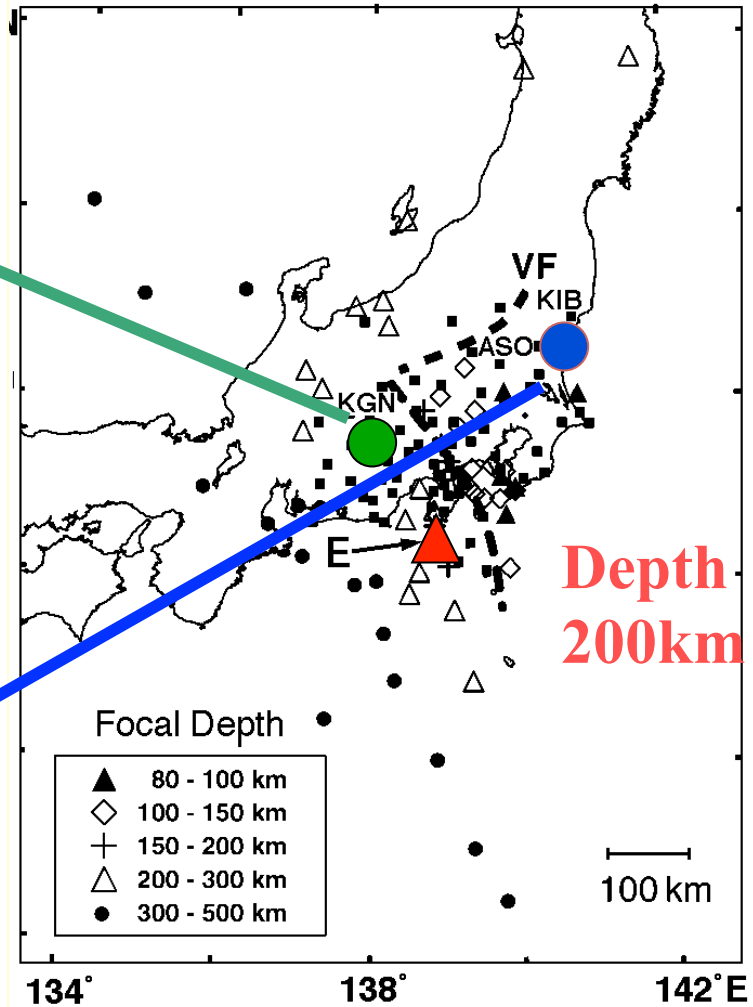
Takahashi et al. [2009]



### S-Arrival

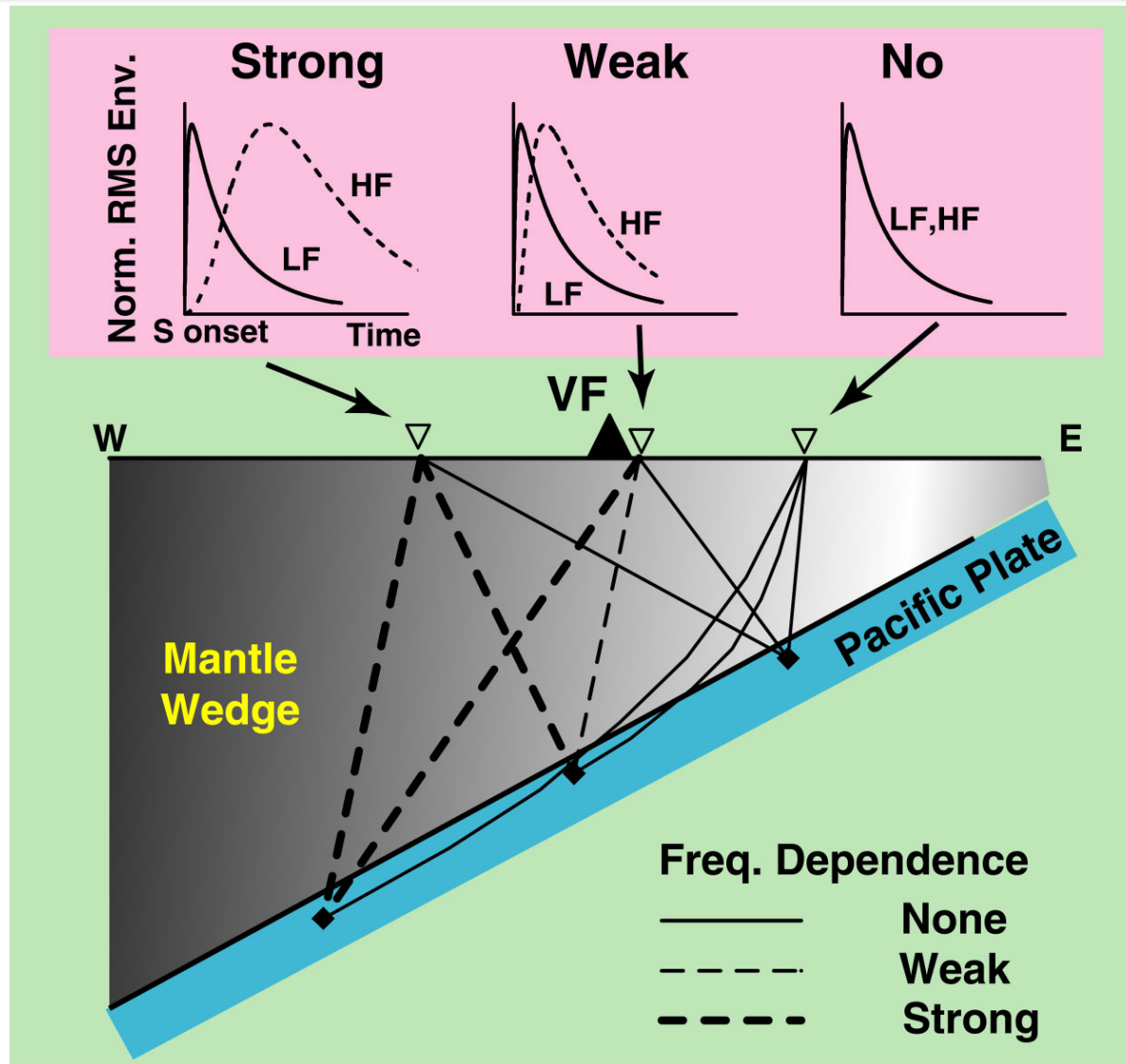


### Envelope broadening in Kanto



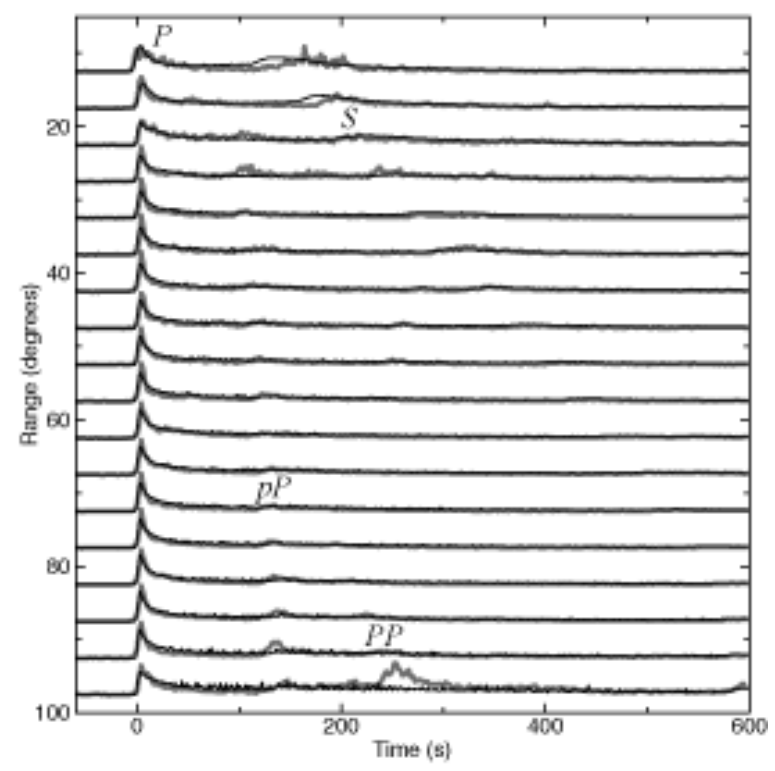
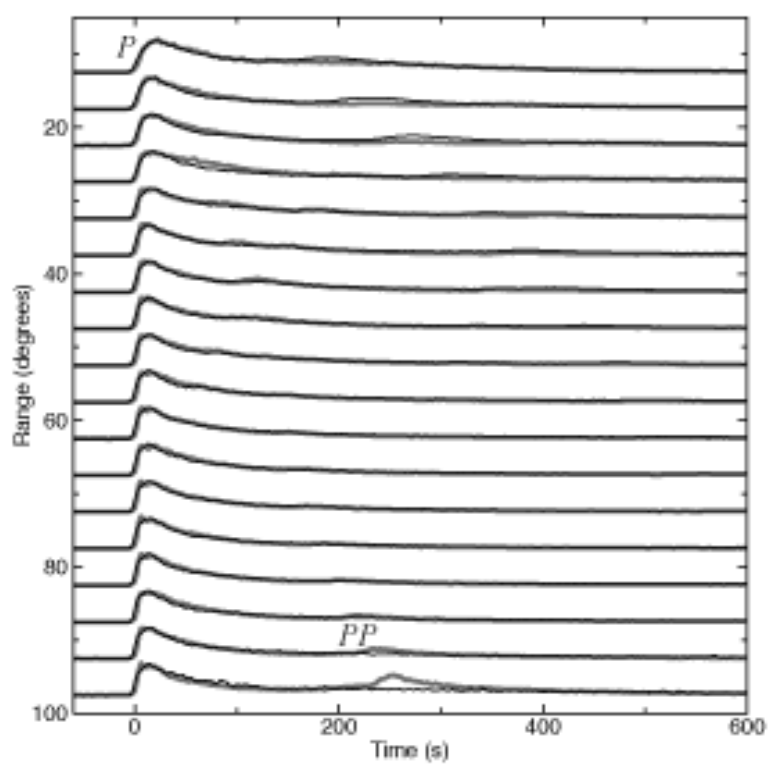
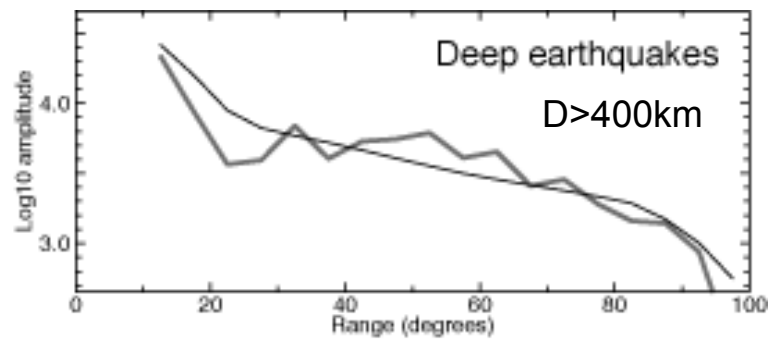
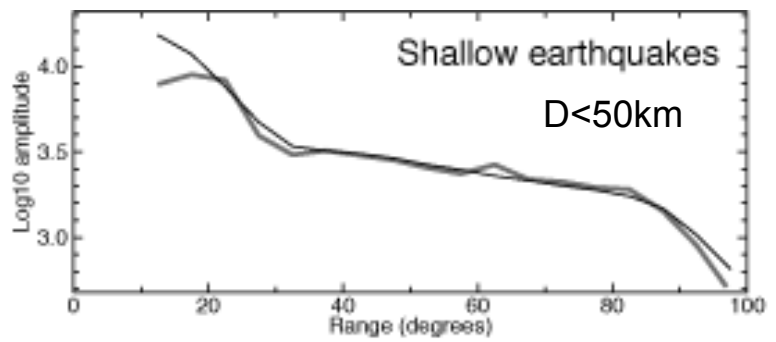
Obara and Sato [1995]

# Inhomogeneity beneath the Kanto-Tokai Area, Japan



Obara and Sato [1995]

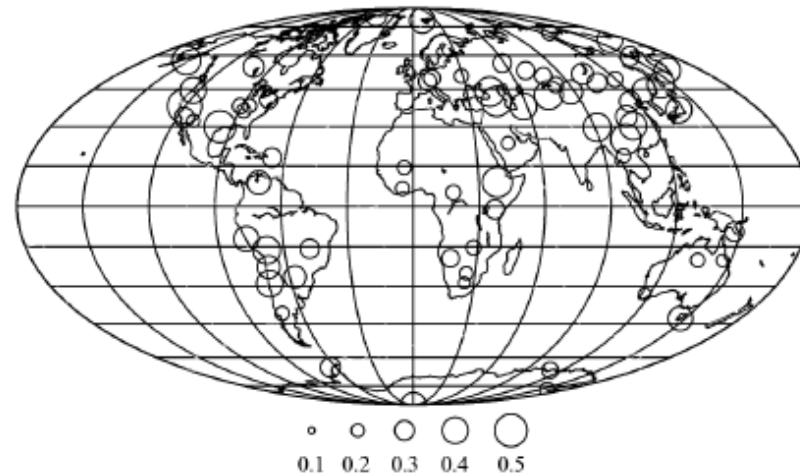
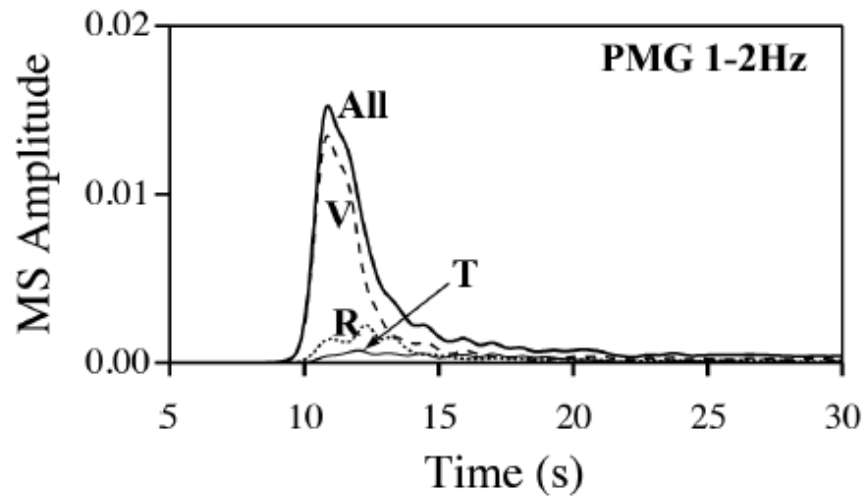
# P-Wave Envelopes of Teleseismic Events



Estimate the randomness from RTT using Born Scattering Amplitudes ( $f_c \sim 1\text{Hz}$ )

Shearer and Earle [2004]

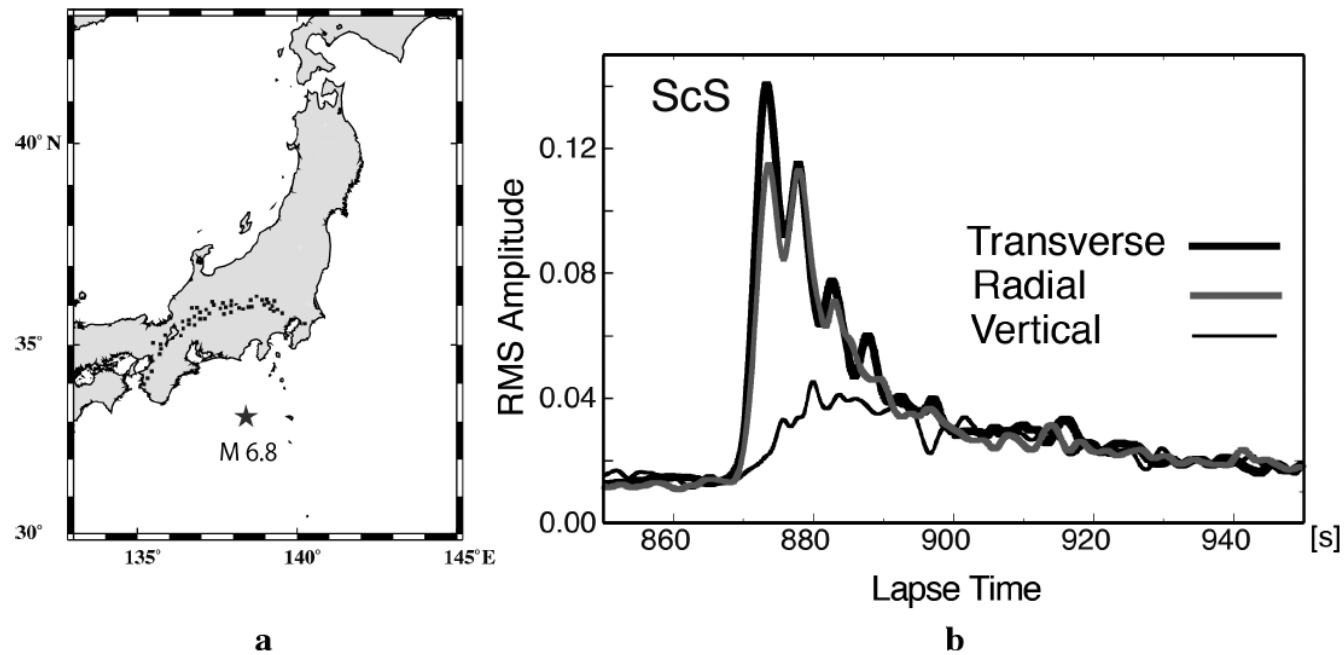
Energy partition of teleseismic P-wavelet into the transverse component as a measure of lithospheric heterogeneity



$$\sqrt{\frac{T}{V + R + T}}$$

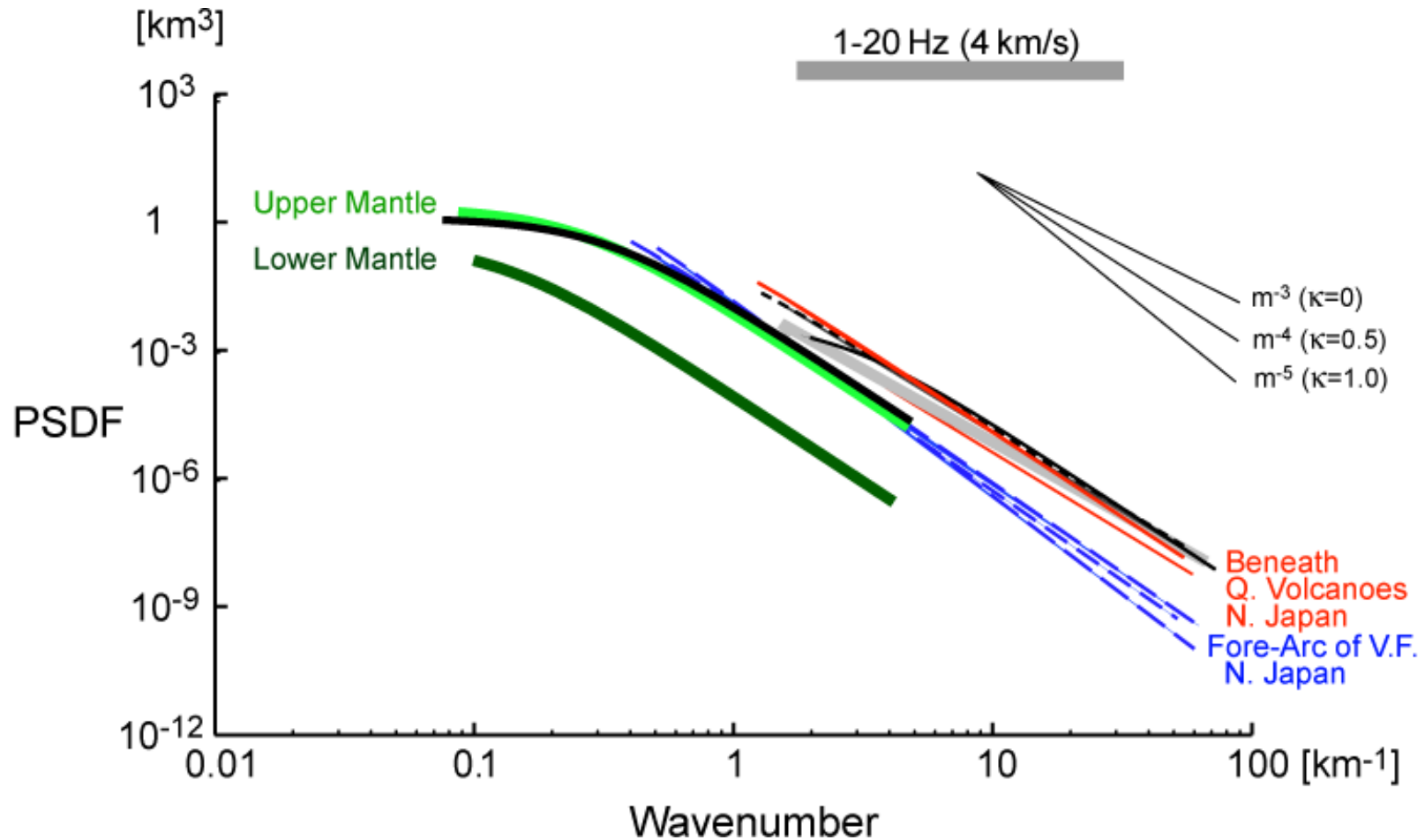
Kubanza et al. [2006]

## Excitation of vertical-component in stacked ScS wavelet envelope



**Fig. 9.41** (a) Epicenter (star) of an M 6.8 earthquake with a focal depth of 333 km and 53 Hi-net stations (dots) located within a range of 300 - 350 km in epicentral distance. (b) RMS envelopes around the ScS phase for 0.33 - 0.67 Hz. [Courtesy of K. Emoto]

## Summary of Reported PSDF of Velocity Fractional Fluctuation



$$V(\mathbf{x}) = V_0 [1 + \xi(\mathbf{x})] \quad \xi(\mathbf{x}) \Rightarrow P(m)$$

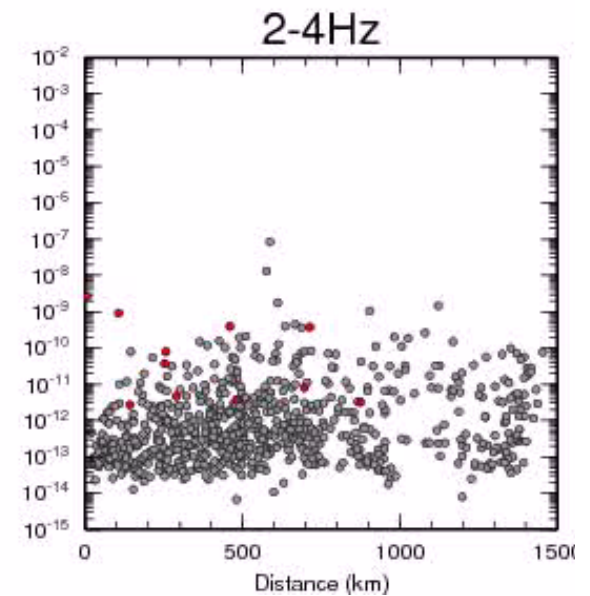
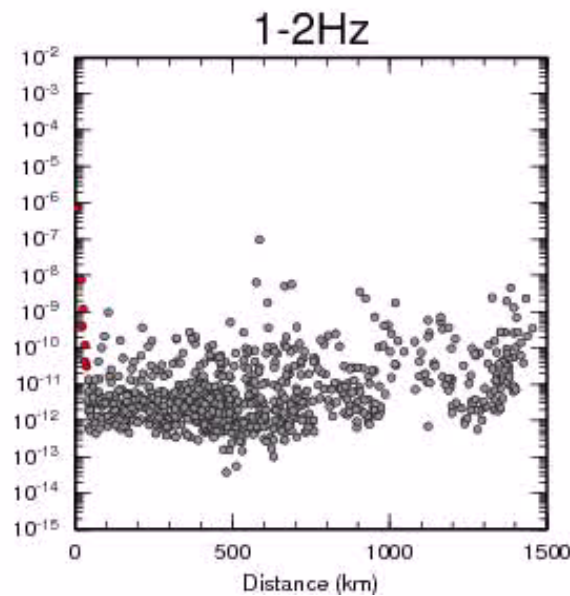
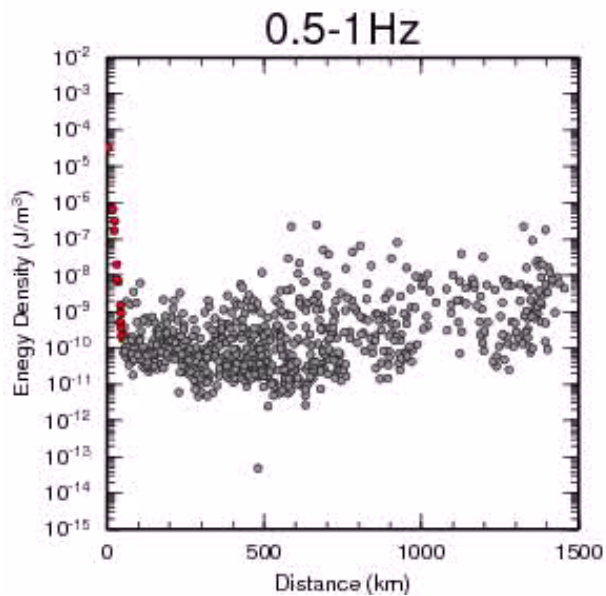
Time lapse movie of  
the space distribution of  
seismic energy density  
of a local earthquake

Hi-net data  
Courtesy of K. Emoto

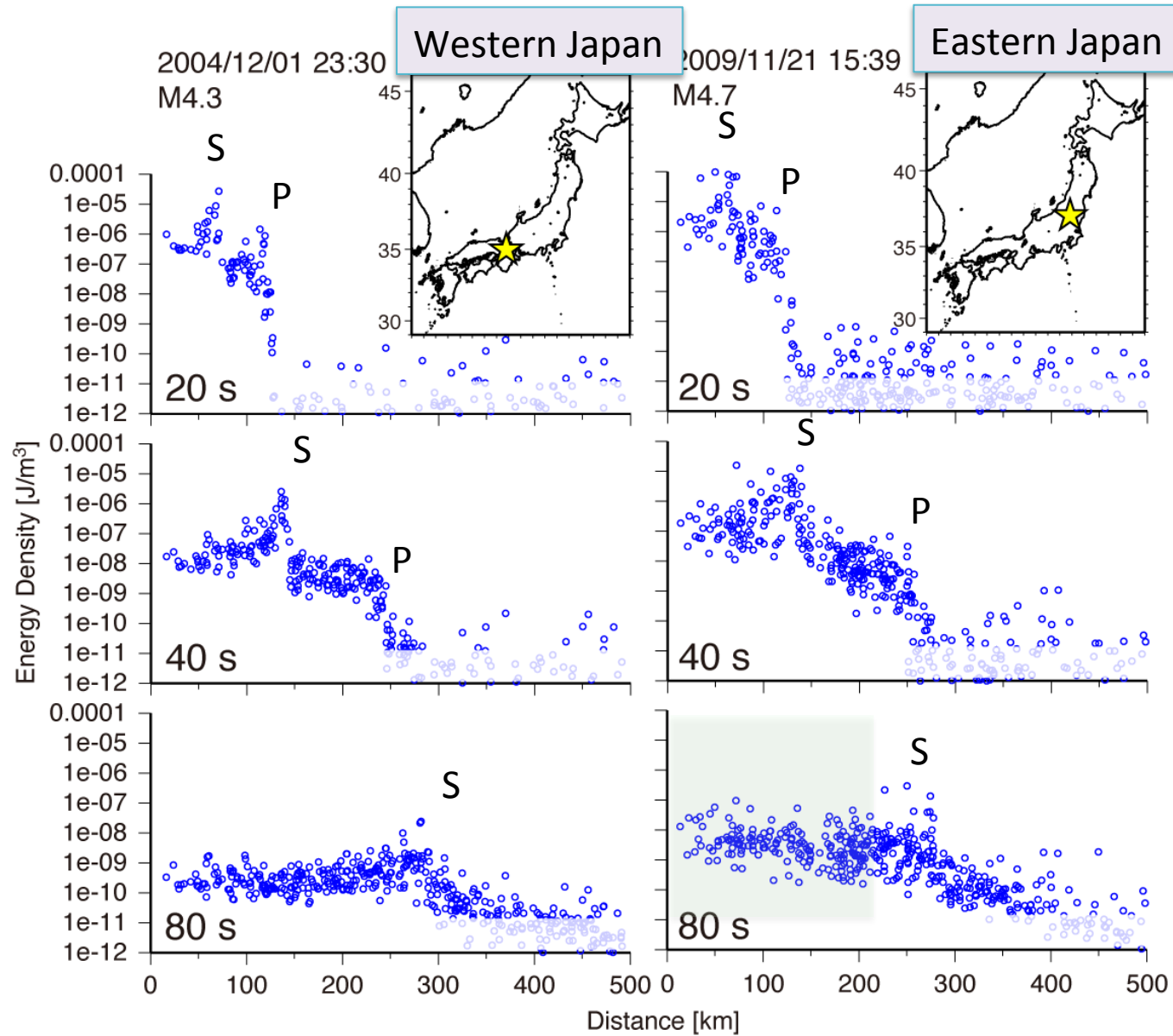


W. Japan  
2011/11/21, M 5.2

Semi-log plot of Energy density vs. distance



Snap shot of the space distribution of seismic energy (4-8 Hz)  
 [Asano and Saito, 2011]



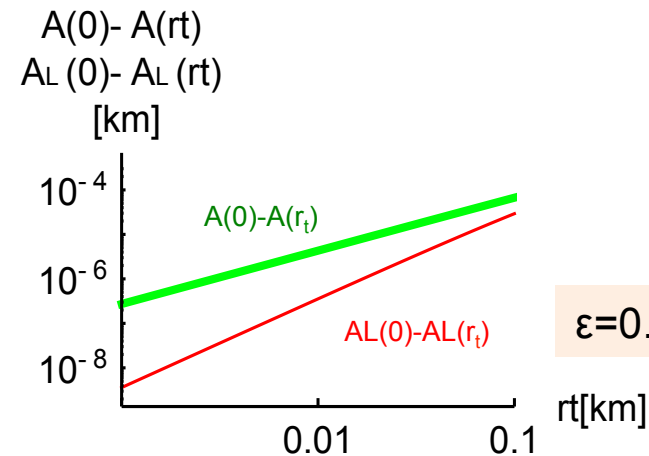
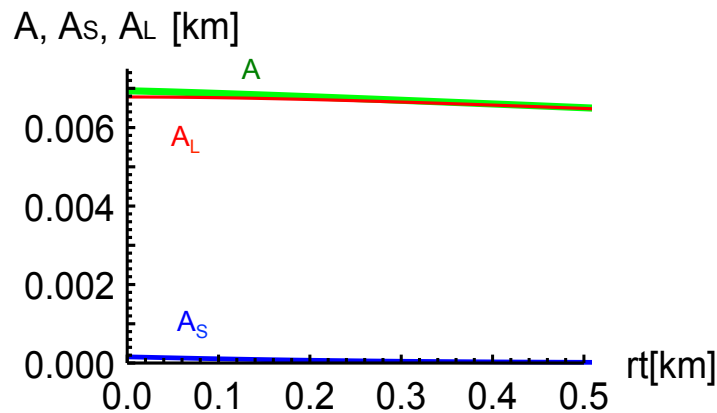
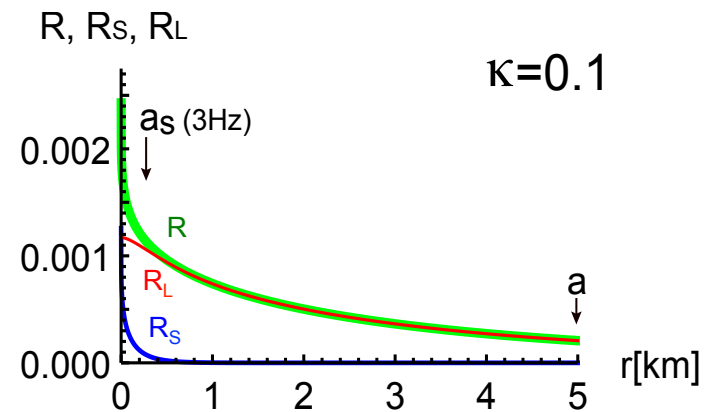
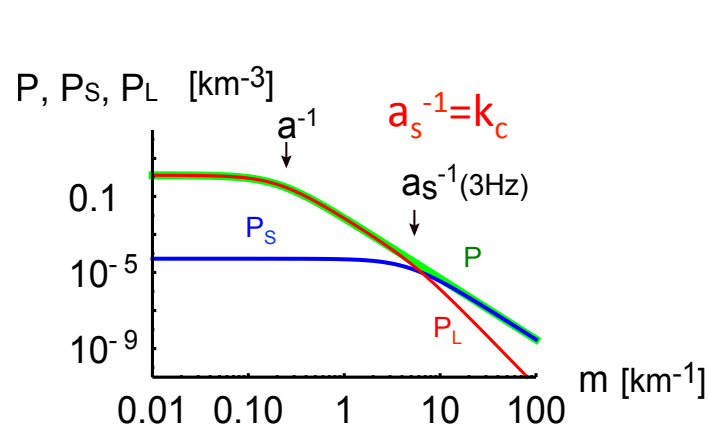


For von-Karman-type random media,  $k_c \gg a^{-1}$  means the power-law range, where  $k_c^{-1}$  is the unique scale of length.

To apply strictly the parabolic approximation for slowly varying fluctuations  $\xi_L$  :

$$2ik_c \partial_r U + \Delta_{\perp} U - 2k_c^2 \xi_L(k_c, \mathbf{x}) U = 0$$

$$\partial_r \Gamma_2 + i \frac{k_d}{2k_c^2} \Delta_{\perp d} \Gamma_2 + k_c^2 (A_L(0) - A_L(r_{\perp d})) \Gamma_2 + \frac{k_d^2}{2} A_L(0) \Gamma_2 = 0, \quad A_L(r_{\perp}) \equiv 2 \int_0^{\infty} dz R_L(\mathbf{x}_{\perp}, z)$$

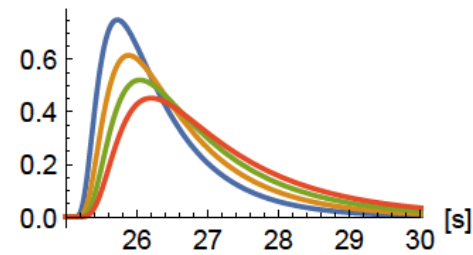
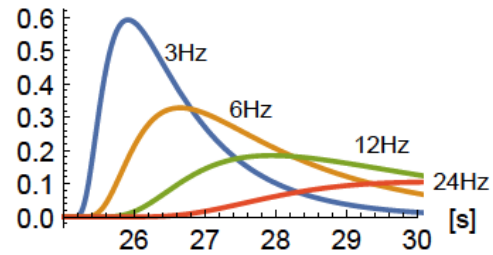


$$G_{M0}(\mathbf{x}, t, \omega_c) = \frac{1}{4\pi r^2 V_0} \frac{\pi^2}{16 t_M} \vartheta_4'' \left( 0, e^{-\frac{\pi^2}{4} \frac{(t-r/V_0)}{t_M}} \right)$$

vK type\_eps0.05\_a5km\_v4km/s\_eta1. at 100.km  
 4 Pi r^2 V0 GM0

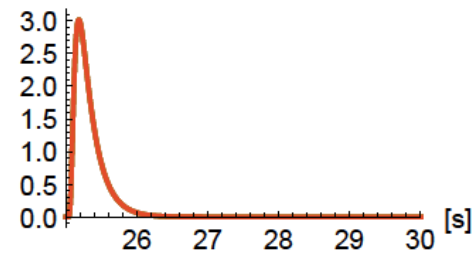
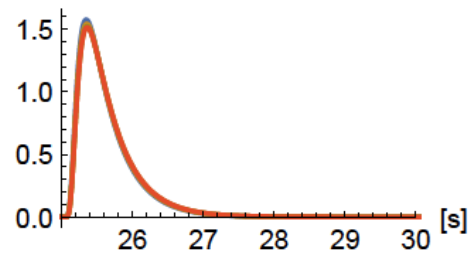
kappa=0.1

kappa=0.5

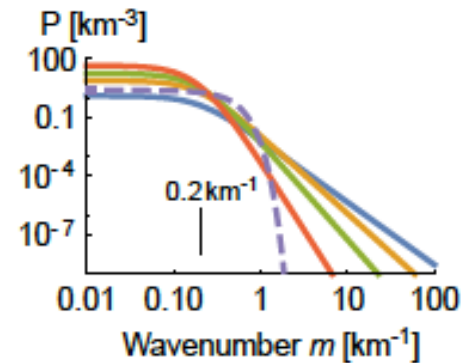
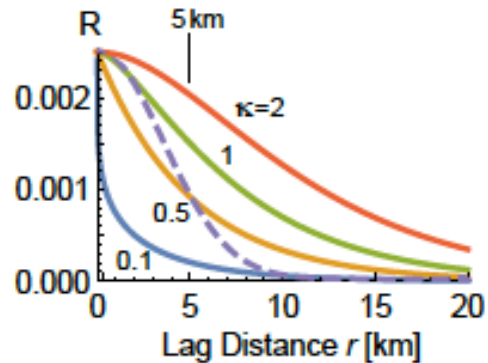


kappa=1

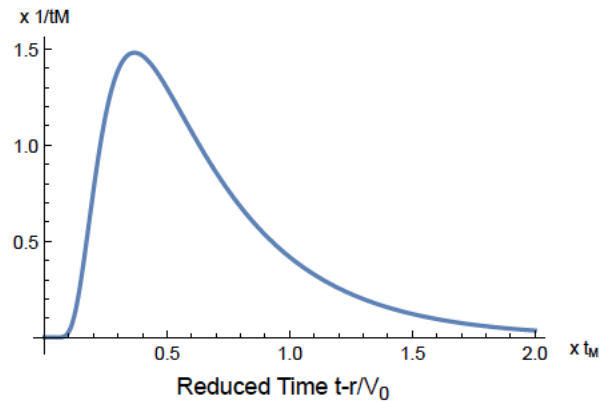
kappa=2



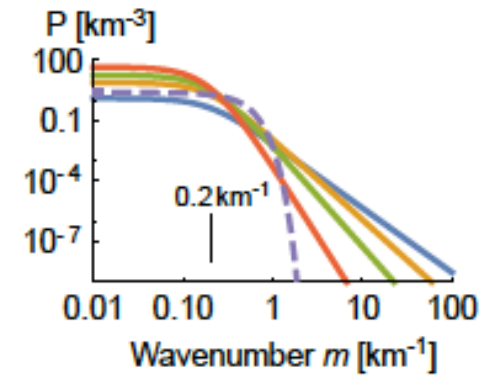
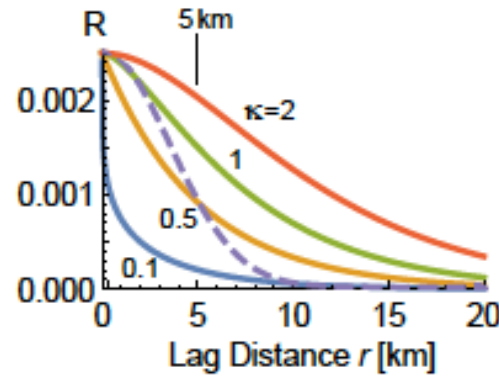
von Karman-type random media ( $\epsilon=0.05$ ,  $a=5\text{km}$ ,  $V_0=4\text{km/s}$ )



$4\pi r^2 V_0 G_{M0}$  vs. reduced time



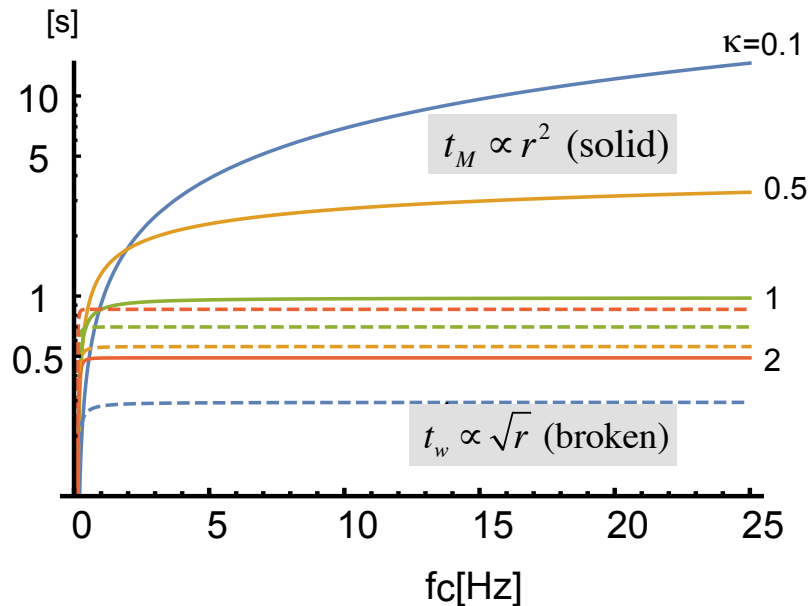
von Karman-type random media ( $\epsilon=0.05$ ,  $a=5\text{km}$ ,  $V_0=4\text{km/s}$ )



Broadening and wandering

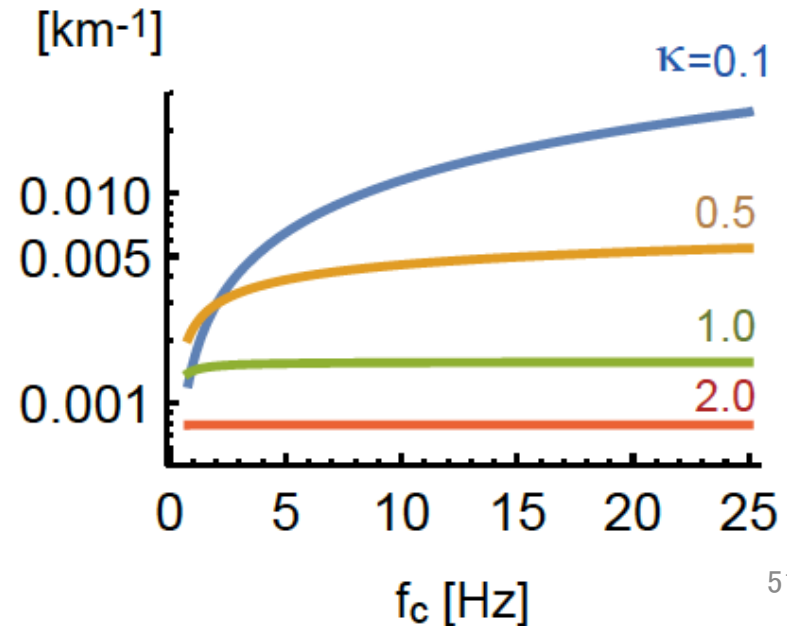
$t_M(\kappa, r, f_c)$  and  $t_W(\kappa, r, f_c)$

at 100km



TrSC

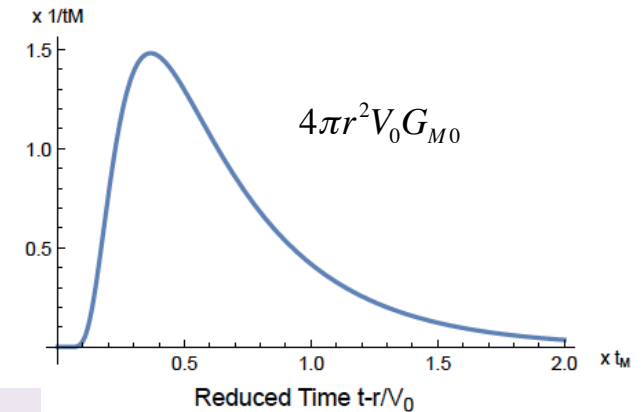
$$g_m(\kappa, f_c) = \frac{1}{4\pi} \int_0^{2\kappa} P(m) m^3 dm$$



## Multiple isotropic scattering model using the Markov propagator and TrSC

$$G_{M0}(\mathbf{x}, t, \omega_c) = \frac{1}{4\pi r^2 V_0} \frac{\pi^2}{16 t_M} \vartheta_4'' \left( 0, e^{-\frac{\pi^2}{4} \frac{(t-r/V_0)}{t_M}} \right)$$

$$t_M(\kappa, r, \omega_c) \text{ and } t_W(\kappa, r, \omega_c)$$



Markov propagator = diffraction + wandering + scattering loss

$$G_M = G_{M0} * w$$

$$G_M^S(\mathbf{x}, t; \omega_c) = G_M(\mathbf{x}, t; \omega_c) e^{-V_0 g_m(\omega_c) t}$$

$$E(\mathbf{x}, t; \omega_c) = G_M^S(\mathbf{x}, t; \omega_c) + V_0 g_m(\omega_c) \int_{-\infty}^{\infty} G_M^S(\mathbf{x} - \mathbf{x}', t - t'; \omega_c) E(\mathbf{x}', t'; \omega_c) dt' d\mathbf{x}'$$

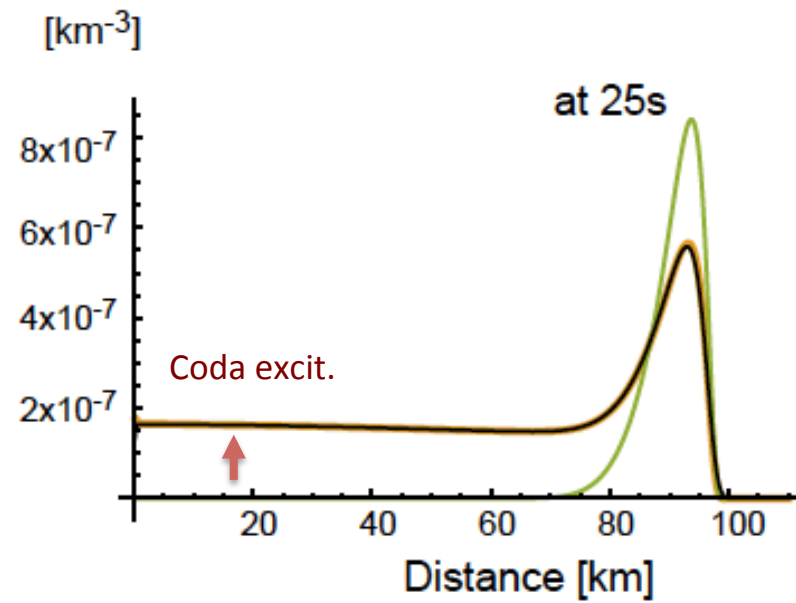
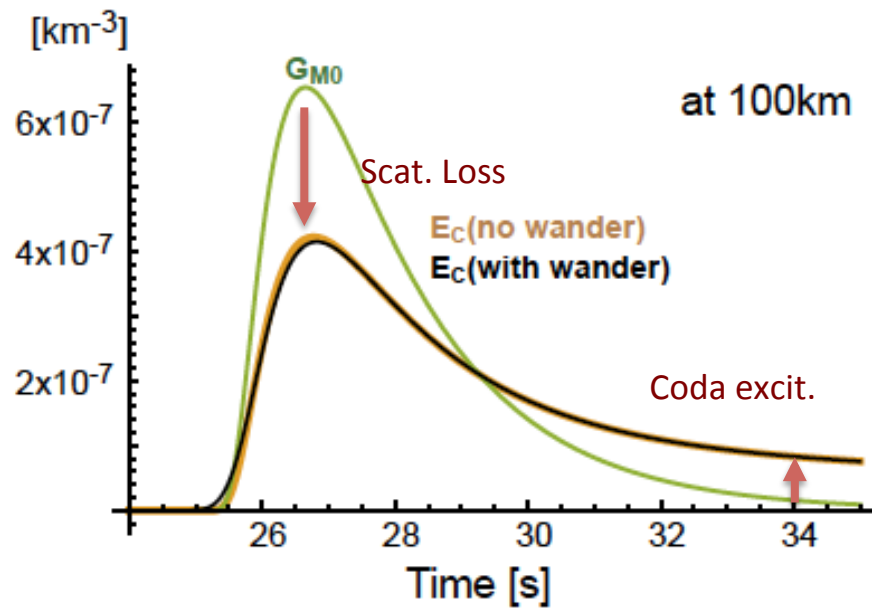
TrSC (Effective isotropic scat. co.):  $g_m(\kappa, \omega_c) = \frac{1}{4\pi} \int_0^{2\kappa_c} P(m) m^3 dm$

Correct to satisfy the energy conservation:  $E_c(\mathbf{x}, t; \omega_c) = \frac{E(\mathbf{x}, t; \omega_c)}{\int E(\mathbf{x}, t; \omega_c) d\mathbf{x}}$

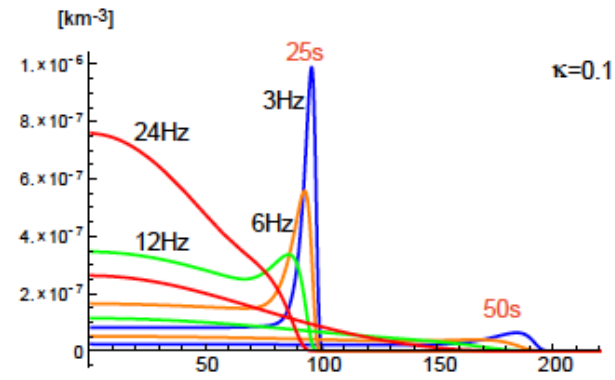
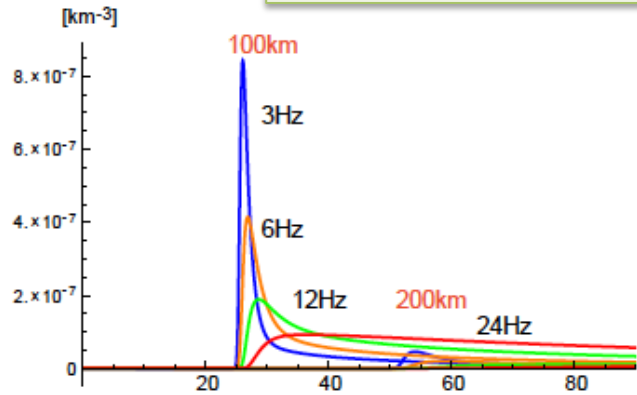
$$4\pi r^2 V_0 \int G_{M0}(\mathbf{x}, t, \omega_c) dt = 1, \quad \int E_c(\mathbf{x}, t, \omega_c) d\mathbf{x} = 1$$

$G_{M0}$  and  $E_c$

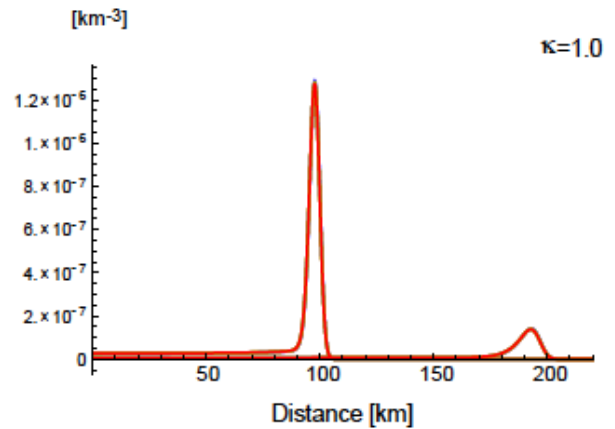
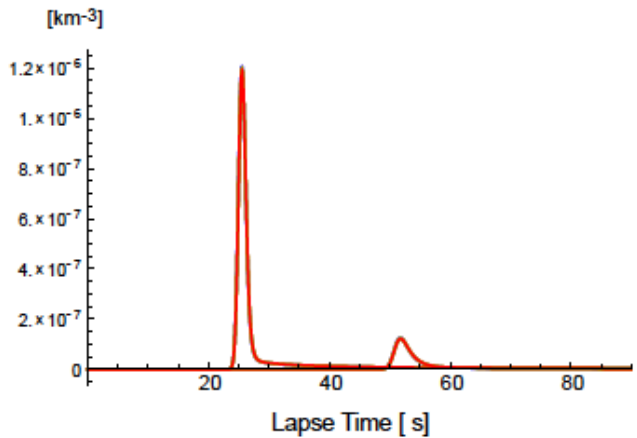
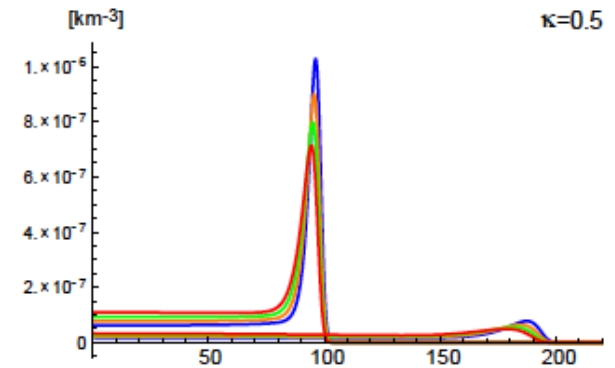
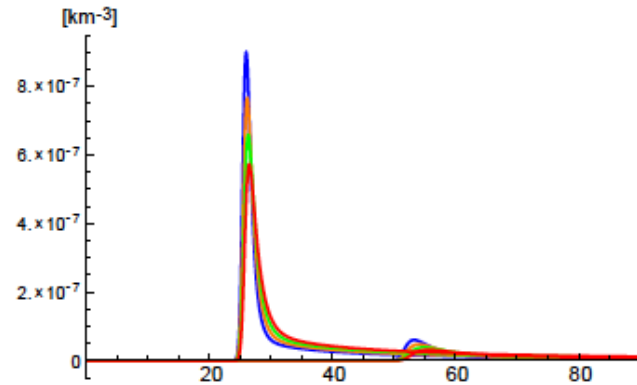
$\kappa=0.1$       $t_M=4.5s, t_W=0.29s$  at 100km,  $g_m=0.0076 \text{ km}^{-1}$



Linear plot of the time- and space-distribution of  $E_c(r,t)$

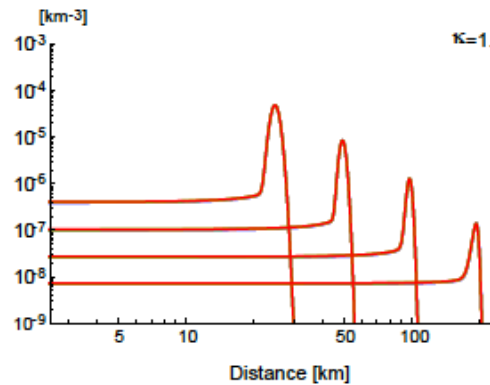
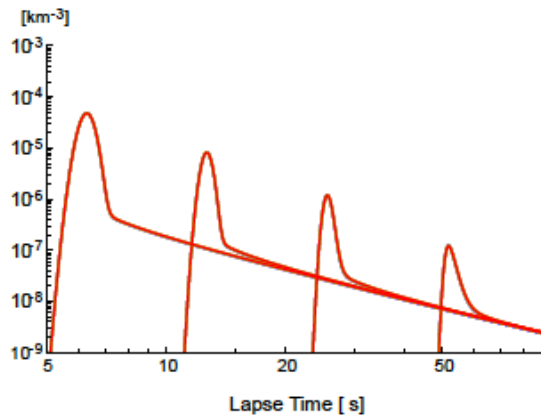
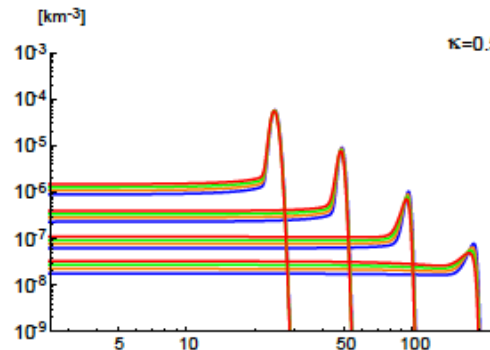
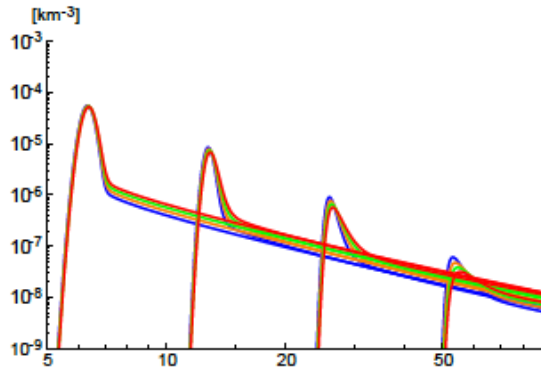
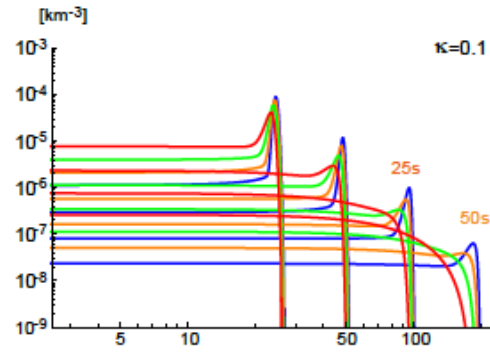
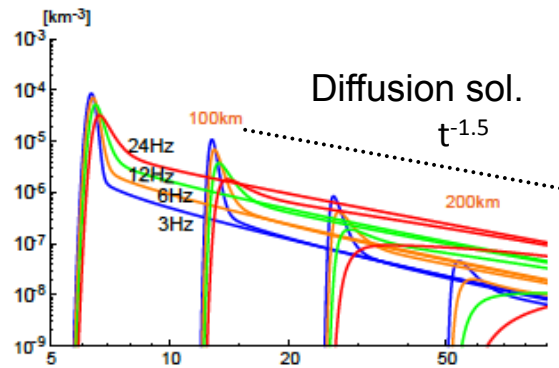


Freq. dependent  
Large coda  
Strong scattering loss



Freq. independent  
Weak coda

Log-log plot of the time- and space-distribution of  $E_c(r,t)$



Freq. dependent  
Large coda  
Strong scattering loss

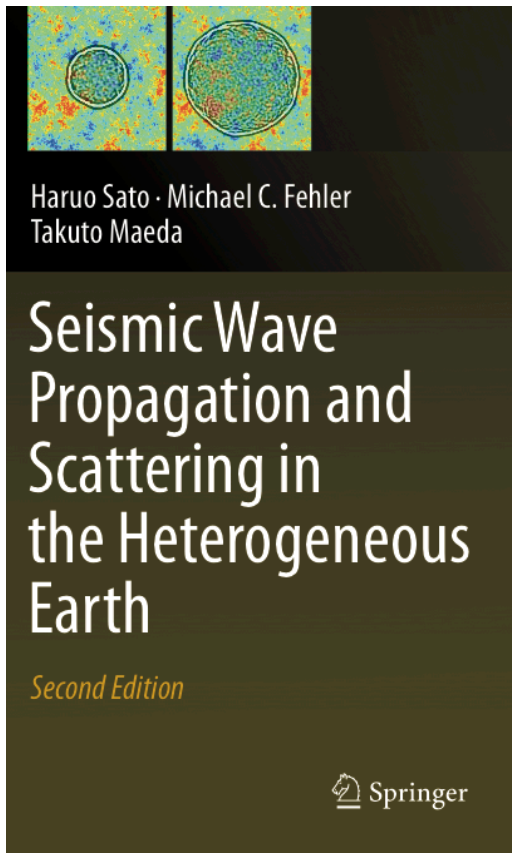
Freq. independent  
Weak coda

## Future Studies

- Observation
  - Analysis of whole envelope from P-onset to S-coda
  - Regional differences in scattering characteristics in relation with tectonic settings
  - Measurements of the PSDF of random velocity fluctuation
- Theory
  - Comparison of the RTT solution with FD wave-simulations in 3-D random media having power law spectra
  - Introduction of depth dependence and regional differences in the envelope synthesis
  - Extension of RTT for elastic vector-wave envelopes



## References



- Bridge between Boltzmann eq. and Stochastic wave theory in random media: Howe [1974], Ryzhik et al. [1996], Margerin [2005].
- RTT of elastic waves through the distribution of spherical obstacles: Margerin et al. [2000].
- Monte Carlo simulation of the Boltzmann eq.: Gusev and Abubakirov [1996], Sens-Schonfelder et al. [2009].
- RTT vs. FD simulation of waves in random media: Frankel and Clayton [1986], Przybilla and Korn [2008], Przybilla et al [2006].
- Markov app. v.s. FD simulation of waves in random media: Fehler et al. [2000], Emoto et al. [2010, 2011].
- Use of RTT solution as a kernel for in the time lapse change measurement of medium parameters: Pacheco and Snieder [2005], Obermann et al. [2013].
- Use of RTT solution for the inversion of source energy radiation: Nakahara et al. [1998], Nakahara [2008]
- Fractal scattering model for the power law decay of amplitude: Sato and Fukushima [2013].

Sato, Fehler and Maeda, “Seismic wave propagation and scattering in the heterogeneous earth: the 2<sup>nd</sup> edition”, Springer, 2012.