





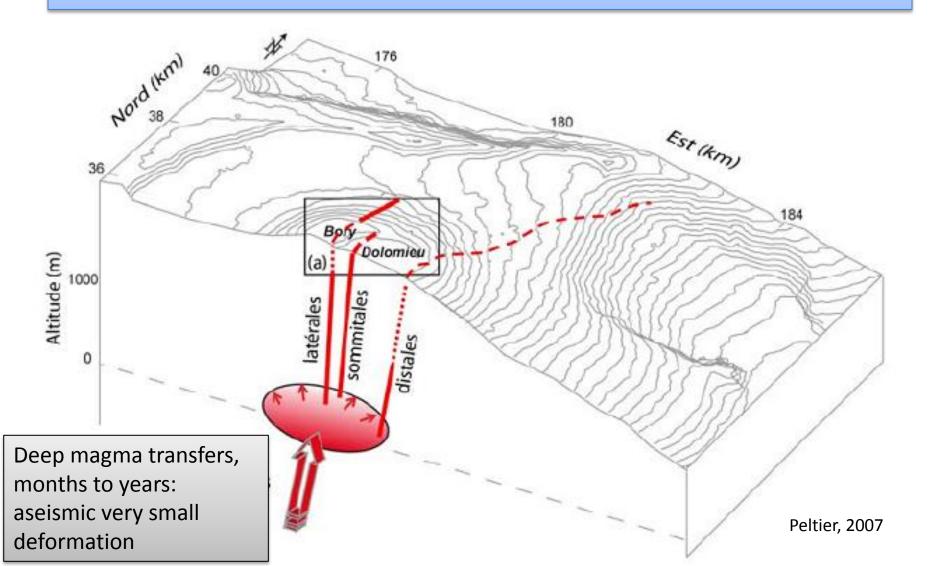
What can we learn about volcanoes using ambient seismic noise correlations?

Florent Brenguier, D. Rivet, A. Obermann, D. Clarke, M. Campillo, N. Shapiro, E. Larose

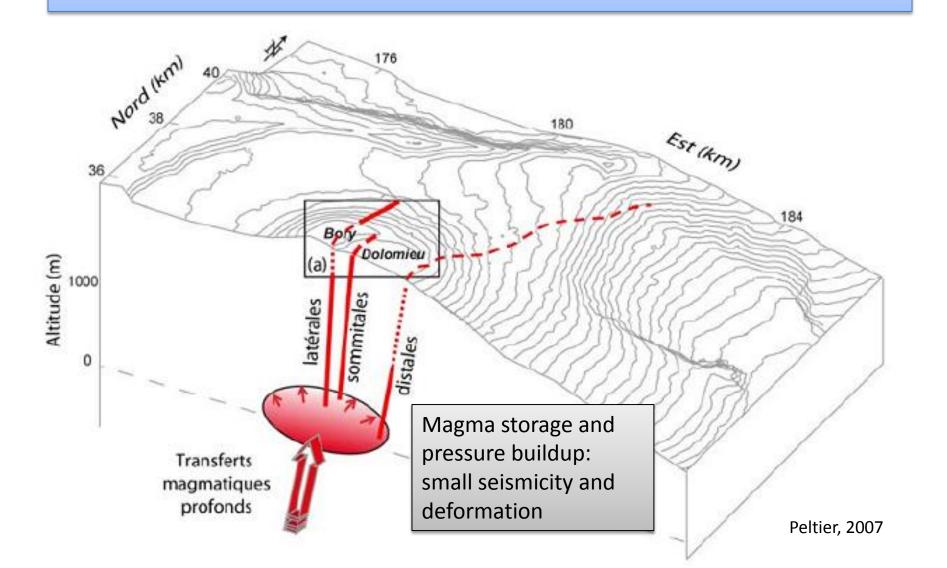




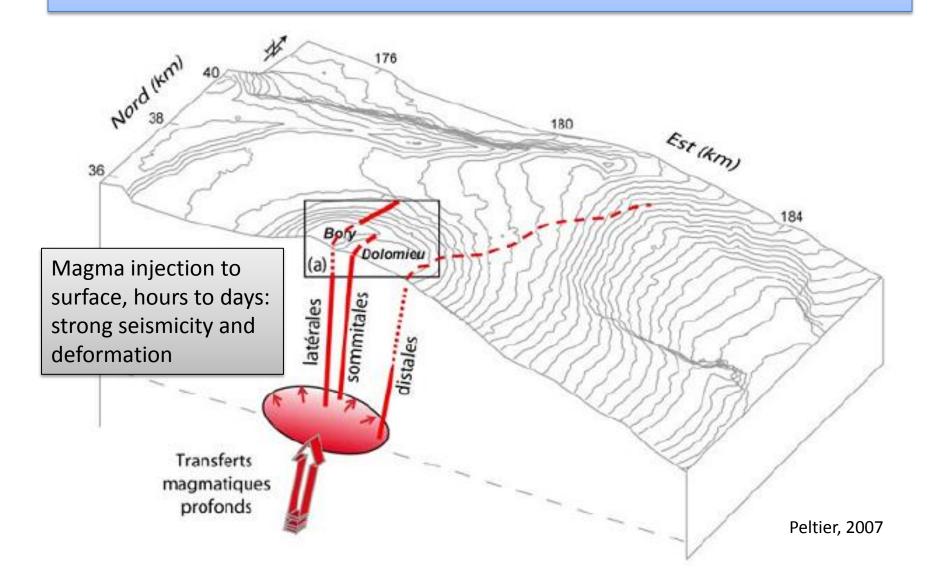
Volcanoes are among the most active parts of the Earth



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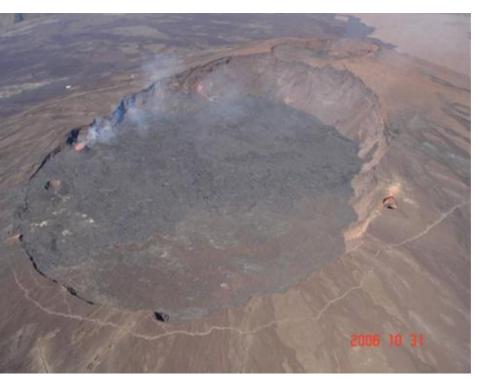
Volcanoes are among the most active parts of the Earth



Magma eruption



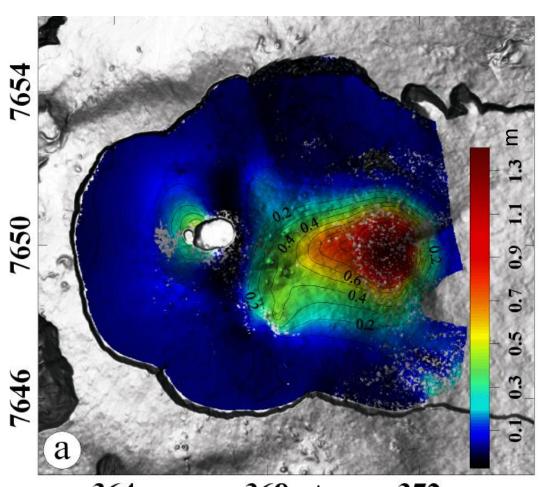
Magma withdrawal may generate caldera collapse





Piton de la Fournaise, Before and after the April 2007 eruption and caledra collapse

Strong stress perturbations may trigger slope instability



364 368 km **372** Courtesy of J.-L. Froger

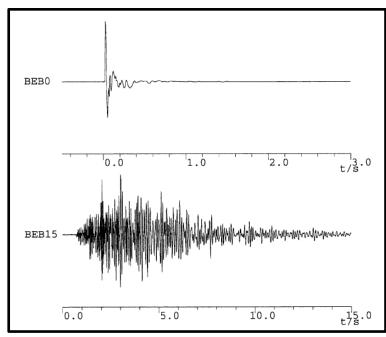


Mt St Helens, 1982. Courtesy of Tom Casadevall

Picture insar

Volcanoes are strongly scattering medium



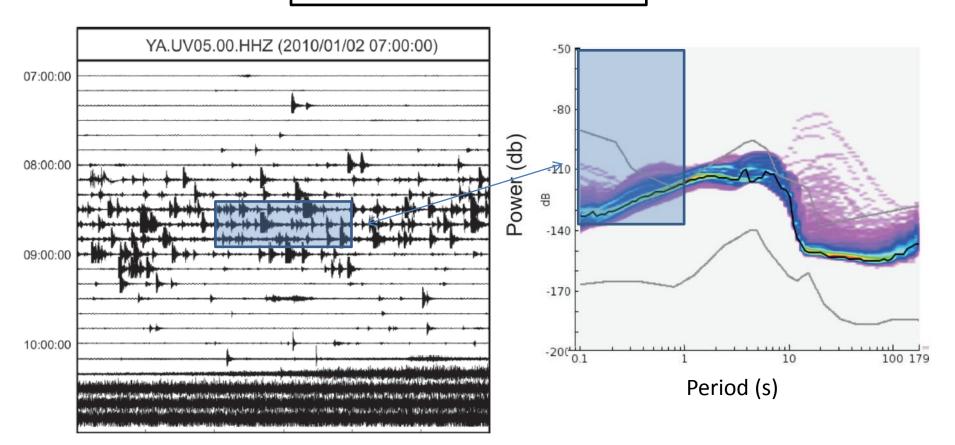


Very small mean free path

$$(\eta_{\rm s}^{-1} \approx 100 \text{ m})$$

A variety of seismic signals on volcanoes!

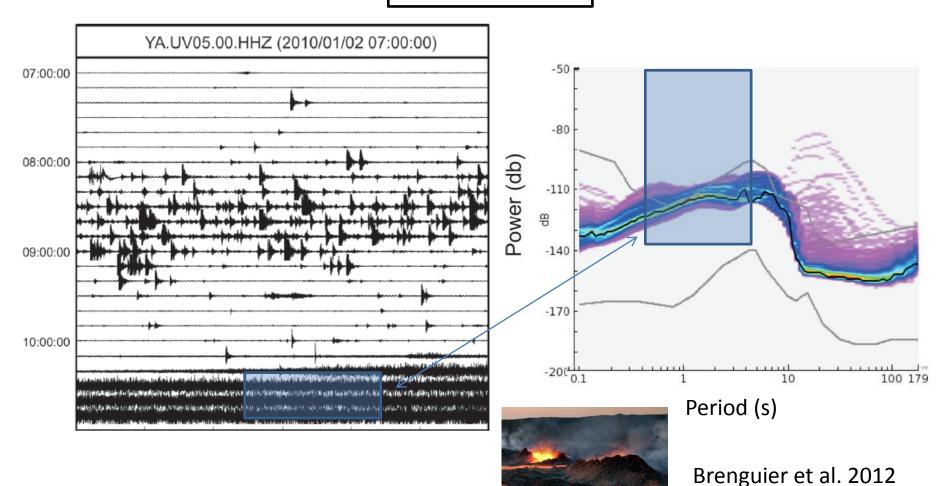
Volcano-tectonic earthquakes



Brenguier et al. 2012

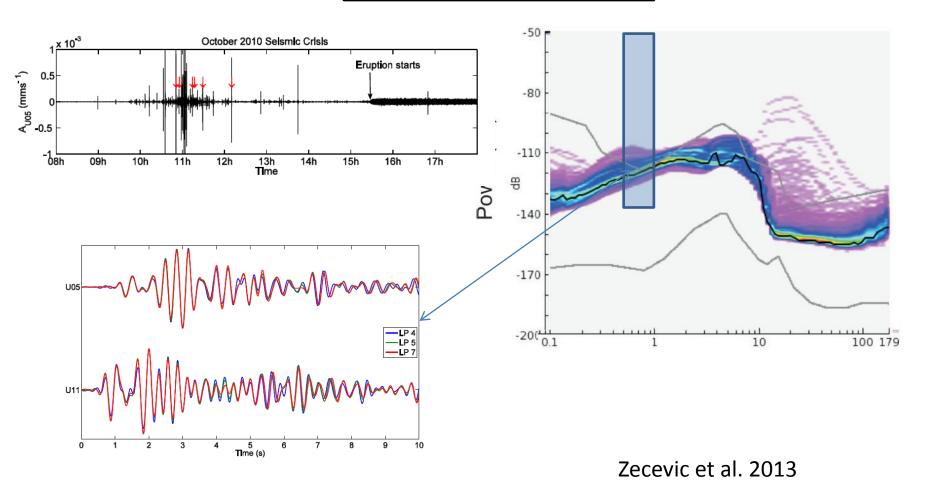
A variety of seismic signals on volcanoes!

Eruption tremor



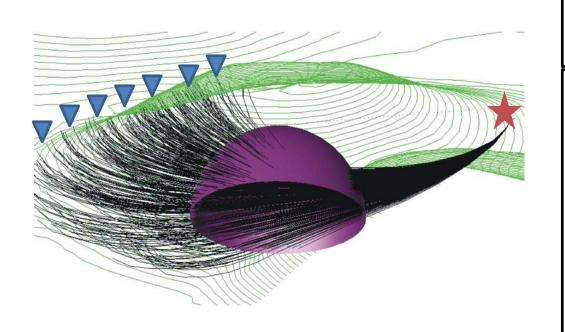
A variety of seismic signals on volcanoes!

Long-period earthquakes



Our interest: Image and monitor volcanoes

Seismic waves probe the rock mass at depth and their properties are affected by volcanic activity



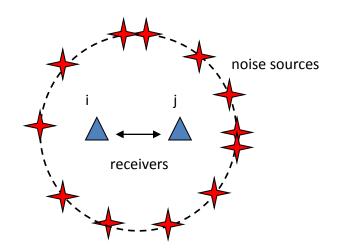
Velocity contrasts in volcanoes are linked to:

- Contrasts between intrusive and effusive material

Changes of the volcano interiors affect:

- Seismic velocities
- Seismic intrinsic attenuation
- Scattering of seismic waves

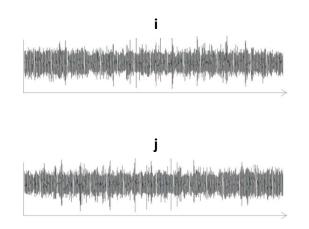
Extraction of Green's functions from correlations of seismic noise

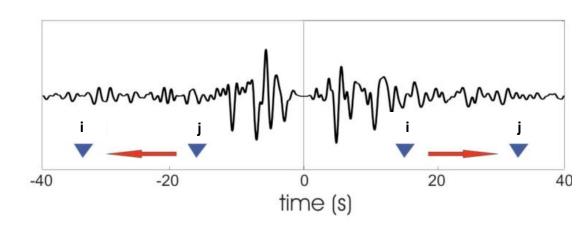


$$\frac{d}{d\tau} C(\tau, \vec{r}_A, \vec{r}_B) = \frac{-\sigma^2}{4 a} (G_a(\tau, \vec{r}_A, \vec{r}_B) - G_a(-\tau, \vec{r}_A, \vec{r}_B))$$

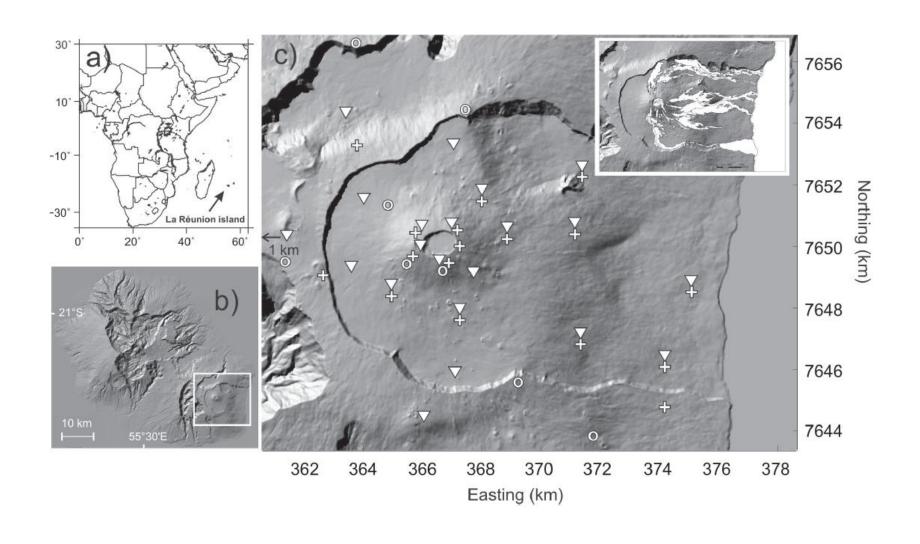
Campillo. 2006

$$\mathsf{G}_{\mathsf{i},\mathsf{j}} = \left(\begin{array}{ccc} Z_i Z_j & Z_i R_j & Z_i T_j \\ R_i Z_j & R_i R_j & R_i T_j \\ T_i Z_j & T_i R_j & T_i T_j \end{array} \right)$$

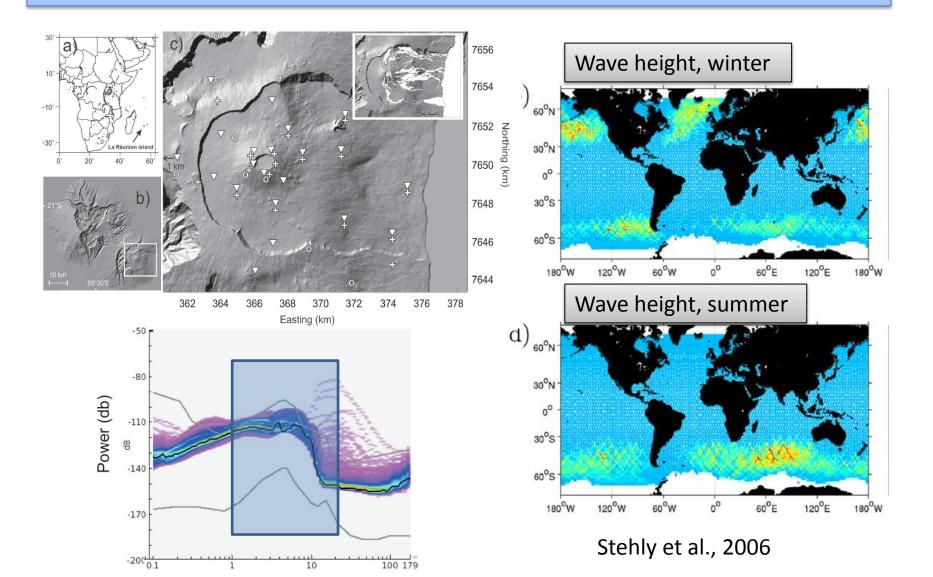




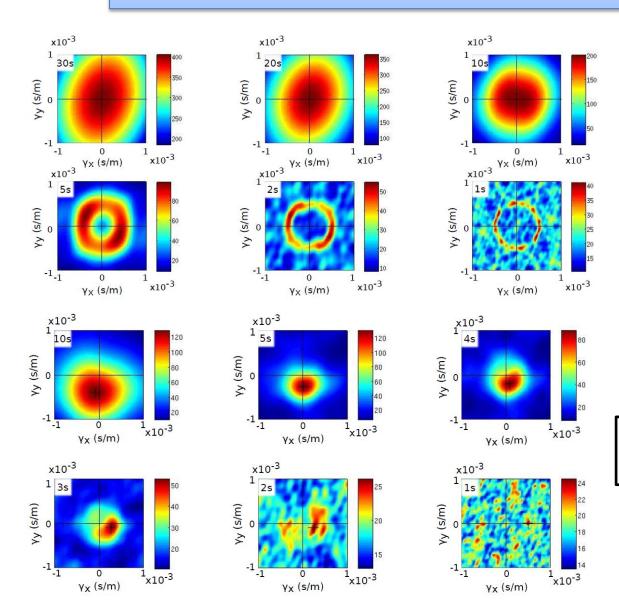
Piton de la Fournaise Volcano



Ambient seismic noise sources properties at Piton de la Fournaise Volcano, La Réunion



Properties of seismic noise at PdF

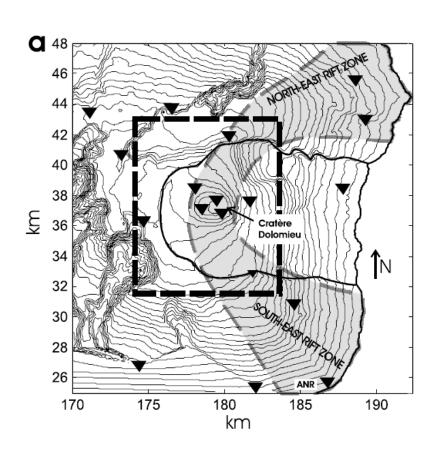


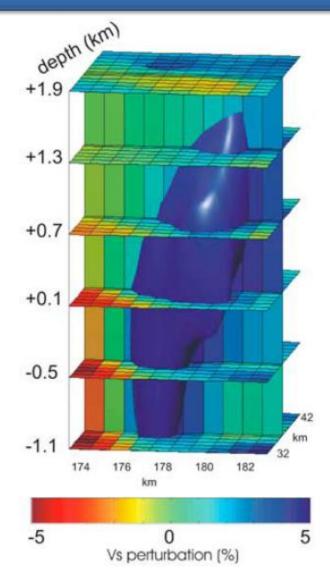
Theoretical array response

Beamforming of 1 day of noise, again, scattering helps

Tsang-Hin-Sun, 2012

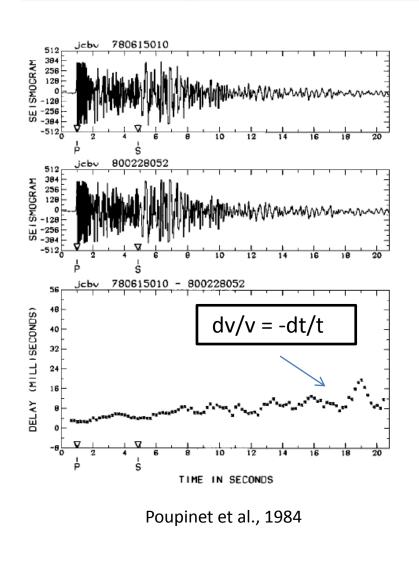
3D tomography using ambient noise

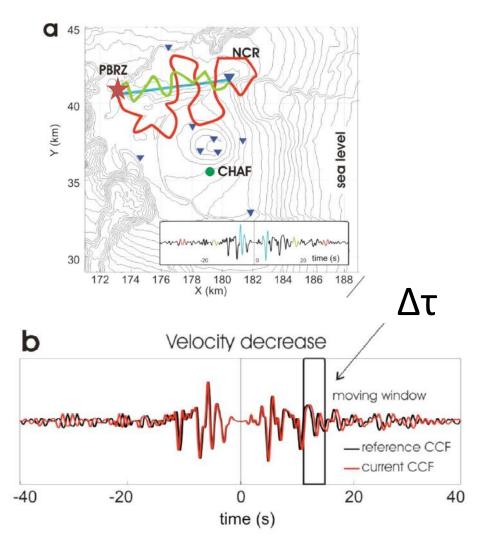




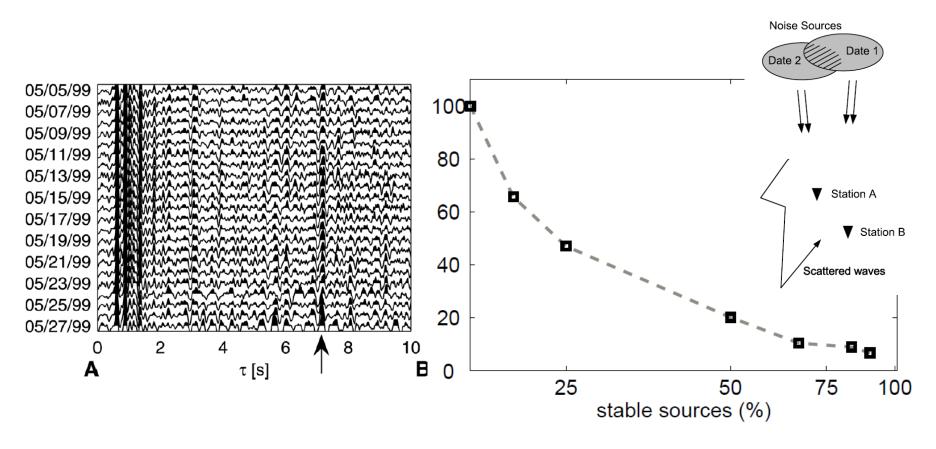
Brenguier et al., 2007

Measurements of seismic velocity changes: coda wave interferometry





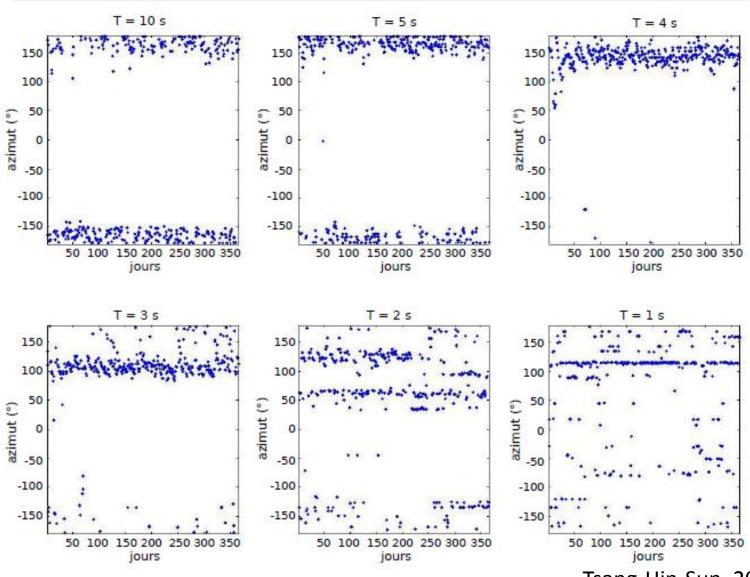
May noise correlations be used for the reconstruction of virtual repeating seismic sources?



Sens-Schönfelder and Wegler, 2006

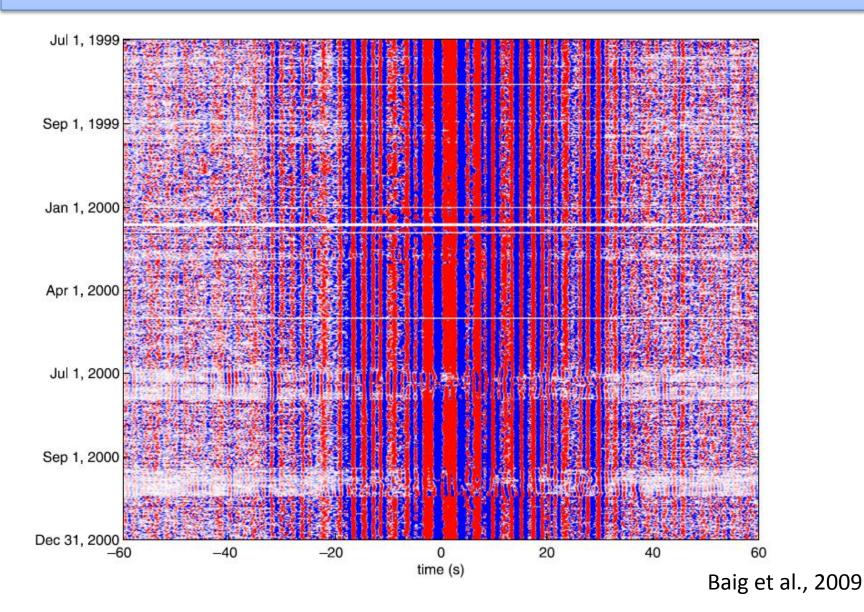
Hadziioannou et al. 2009

Temporal evolution of seismic noise azimuth at PdF

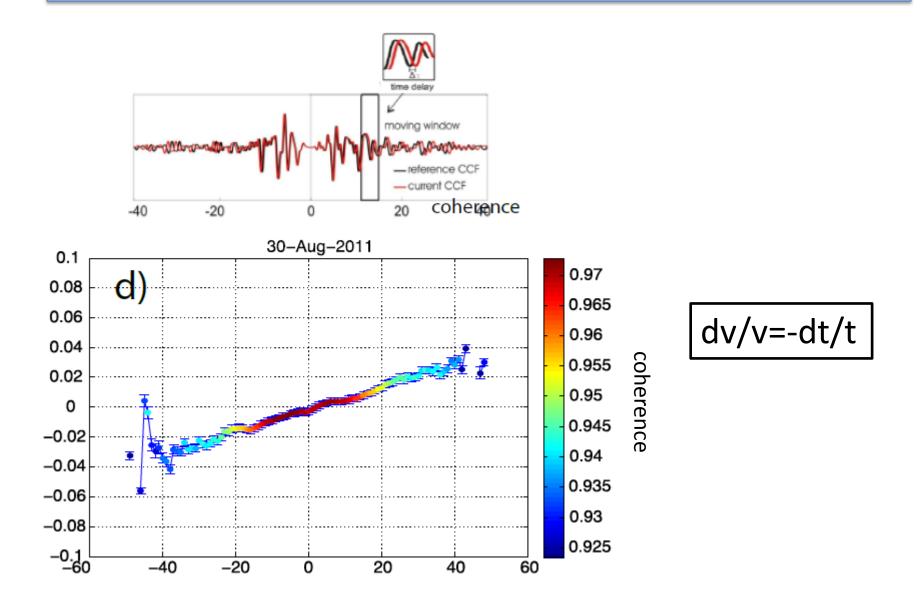


Tsang-Hin-Sun, 2012

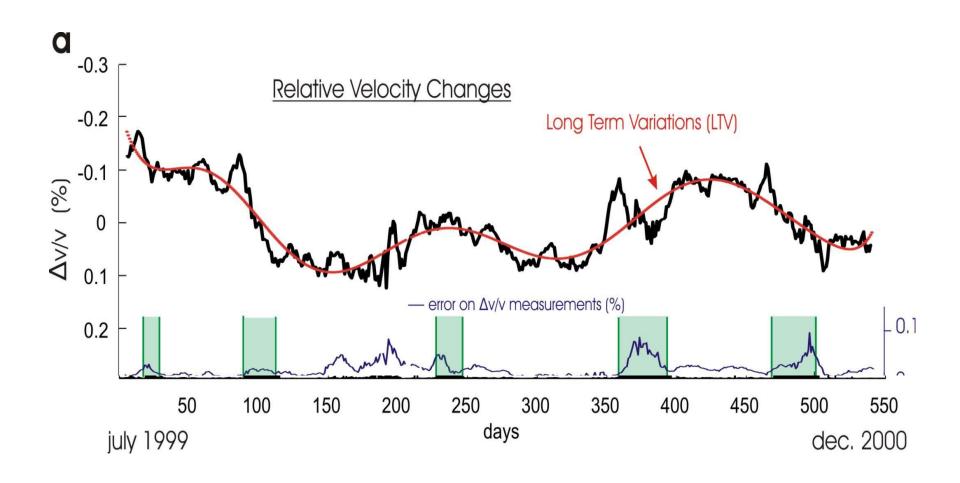
Measurements of velocity changes from the coda of noise correlations at PdF Volcano



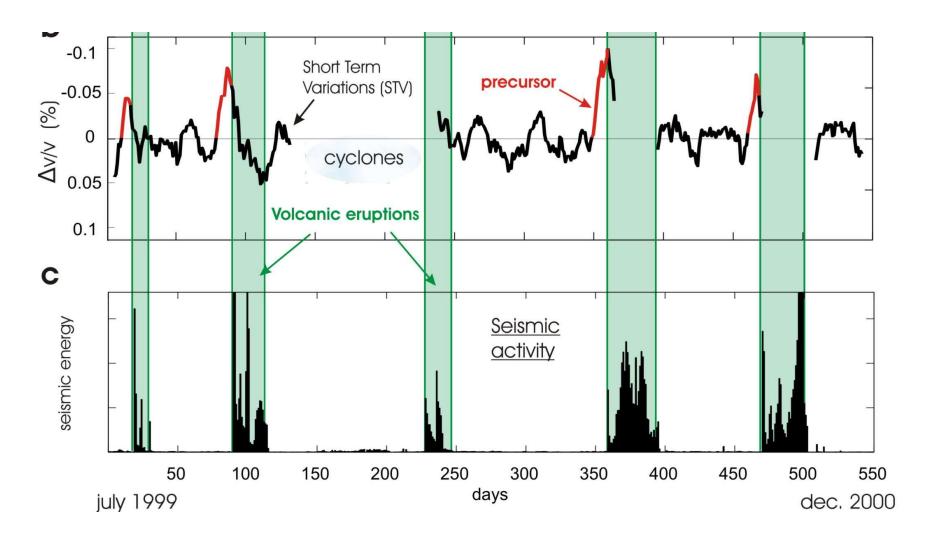
Measurements of velocity changes from the coda of noise correlations at PdF Volcano



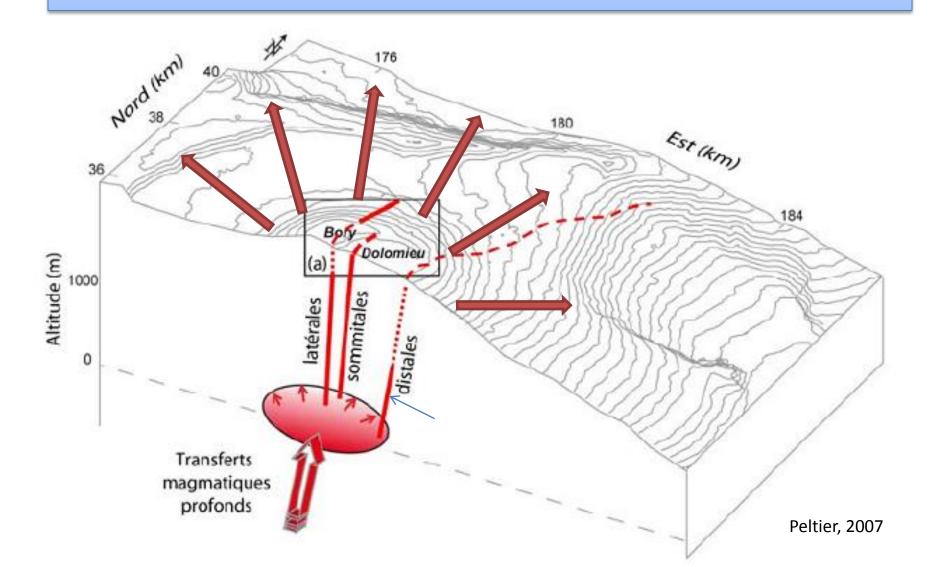
Velocity change measurements at PdF volcano



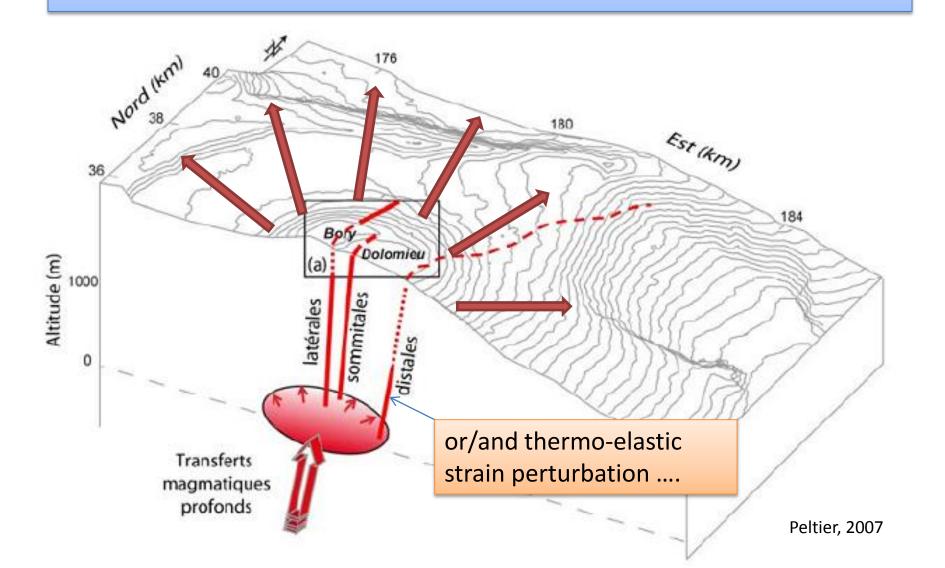
Velocity change measurements at PdF volcano



Stress build-up within the reservoir "dilates" the edifice



Stress build-up within the reservoir "dilates" the edifice



Improving microseism noise records: UnderVolc project broad-band stations

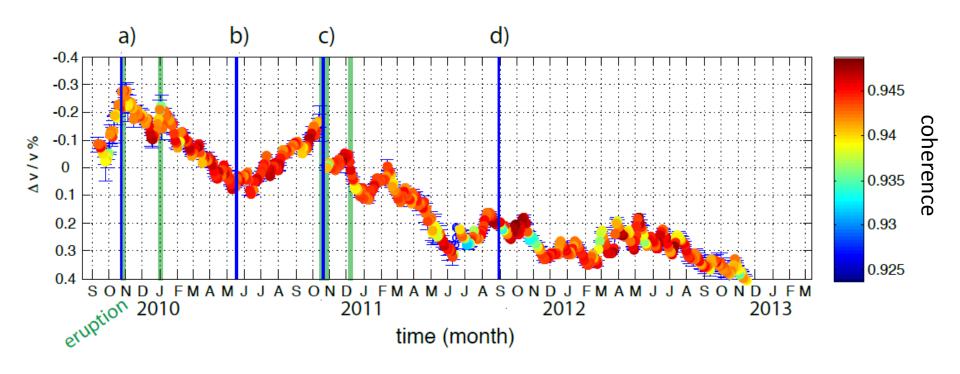




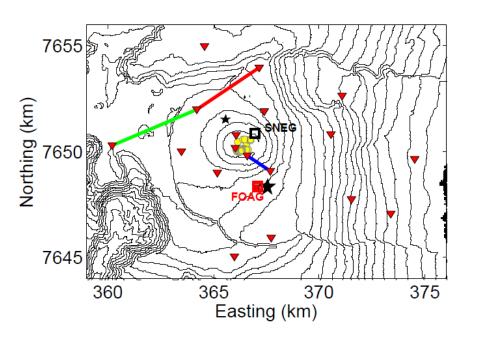


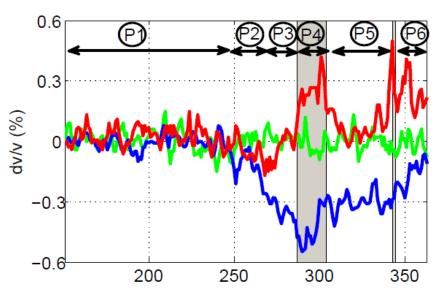


Recent velocity change measurements at PdF volcano



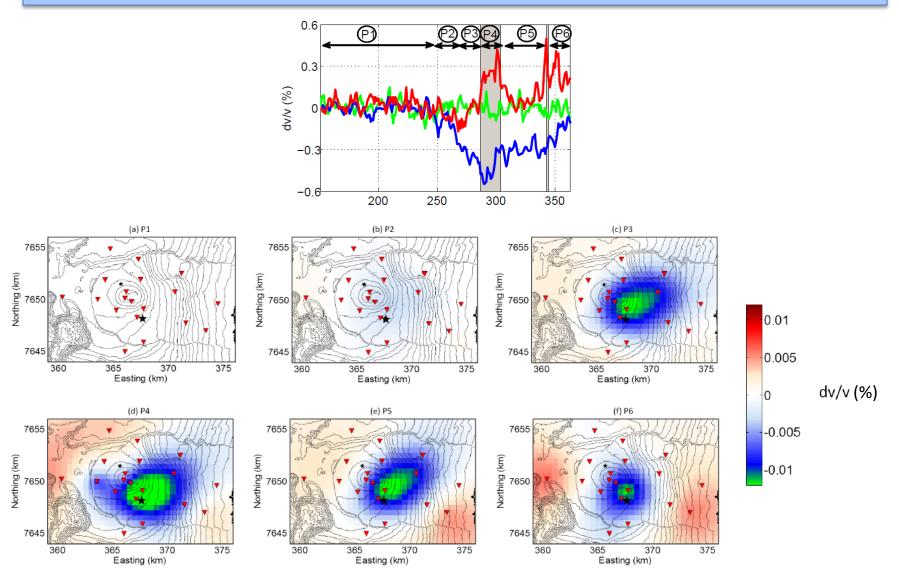
Localization of pre-eruptive seismic velocity changes (1)



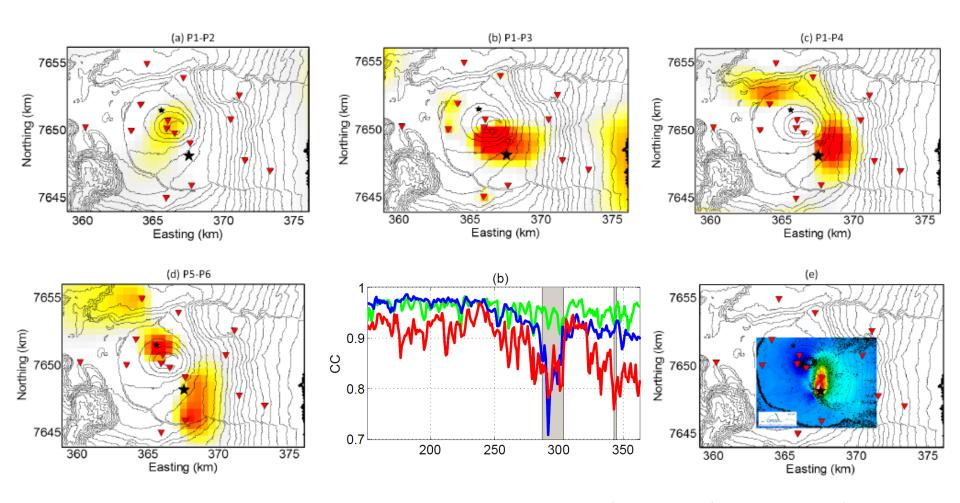


Obermann et al. 2013

Localization of pre-eruptive seismic velocity changes (2)

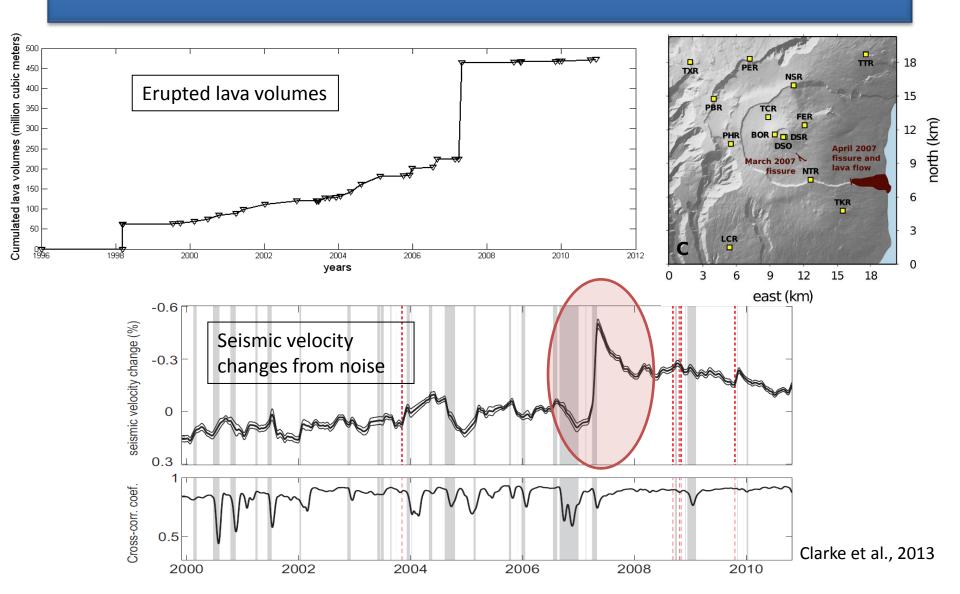


Localization of pre-eruptive seismic velocity changes (3)



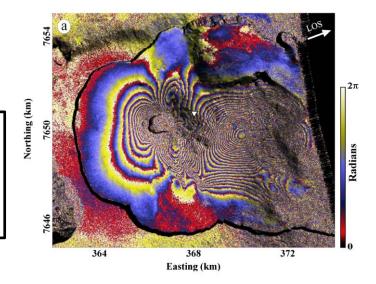
Larose et al., 2010, Obermann et al. 2013

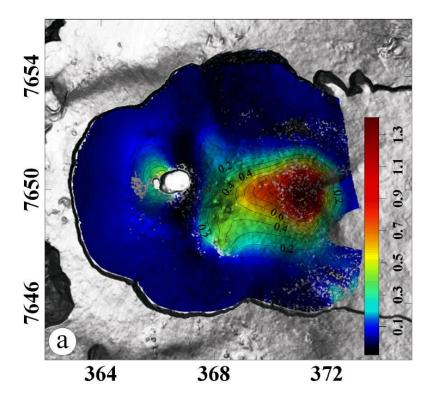
The case of the April 2007 eruption

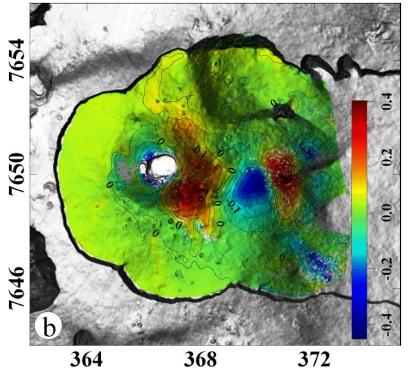


Deformation from InSAR

- Widespread eastward flank movement (1.4 m max)
- Volumetric deformation as strong as 10^-4
- PB : <u>no temporal resolution : When the movement initiated?</u>



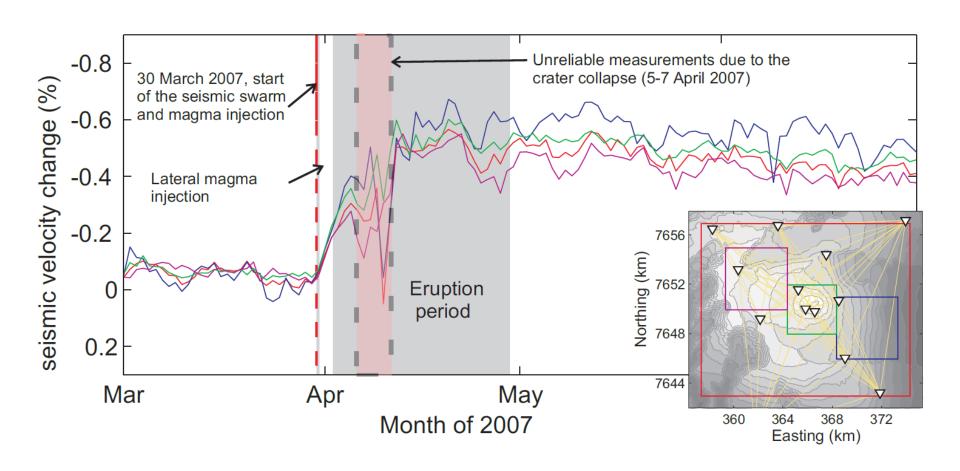




Courtesy of J.L. Froger

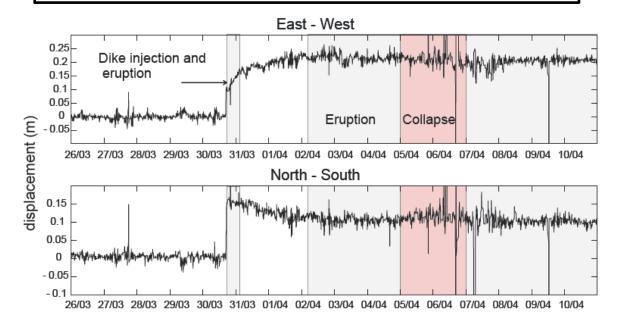
Short-term velocity variations

- Large velocity drop starts on March 30
- Seems to localize on the eastern part of the volcano
- Velocity drop lasts about 15 days



GPS measurements

An eastern flank movement starts on March 30

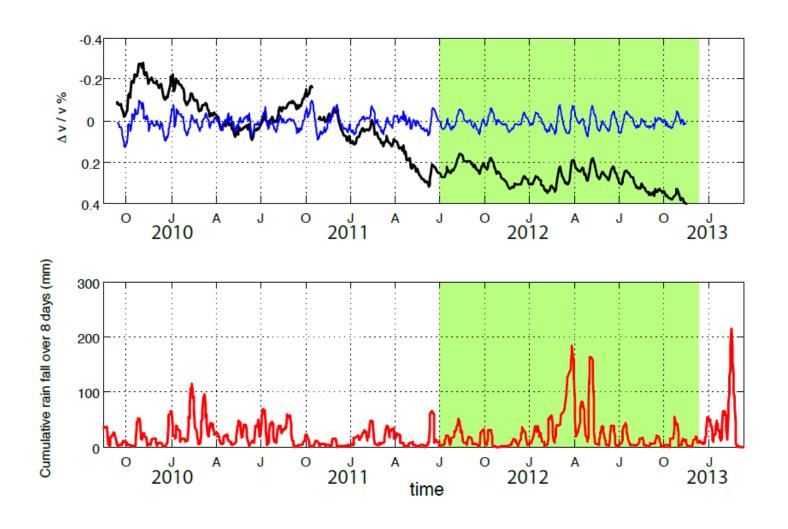


Courtesy of A. Peltier

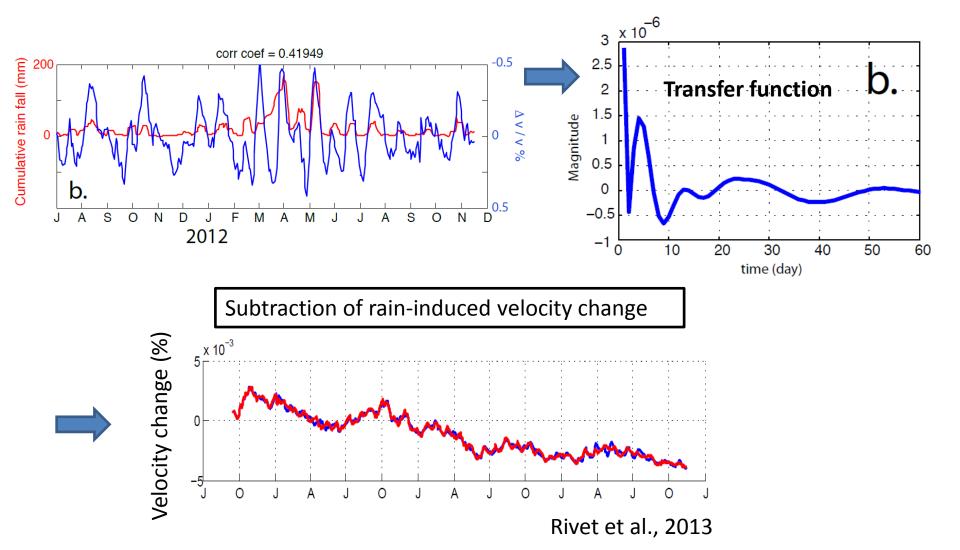
CONCLUSIONS:

- A large eastern flank movement occurred during the March-April volcanic episode
- GPS measurements as well as seismic velocity changes indicate that this movement started at the time of the March 30 magma injection and eruption
- Might thus probably have been triggered by the increase of horizontal stress due to the March 30 eruption and may have favored the horizontal migration of magma and the large April 2007 eruption.

Prospect: effect of non-volcanic perturbations on velocity change measurements (tides, air pressure change, rainfall, T°)



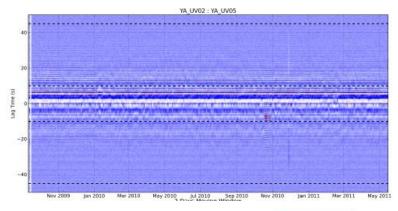
Effect of rainfall on velocity change measurements

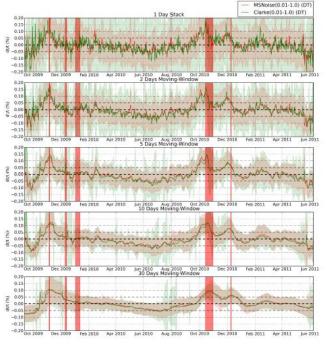


Conclusions

- Continuous monitoring of seismic velocities using ambient seismic noise allows a resolution of less than 0.01 %, ~ a stress detection threshold of ~1 kPa, and a temporal resolution of less than 1 day.
- On volcanoes, such monitoring allows to capture eruption precursory velocity changes and markers for flank movements
- Such a precise observations are sensitive to many phenomena acting on the rock mass (rainfall, tides, thermo-elastic strain, barometric pressure change, ...). These must be known precisely in order to extract the volcanic contribution to the observed velocity changes

Providing an operational tool for real-time noise-based monitoring (MSNoise)





- Written in Python (no requirement for Matlab)
- Automatic processing of real-time data and archives
- Ideally suited for volcano monitoring

T. Lecocq et al., 2013

