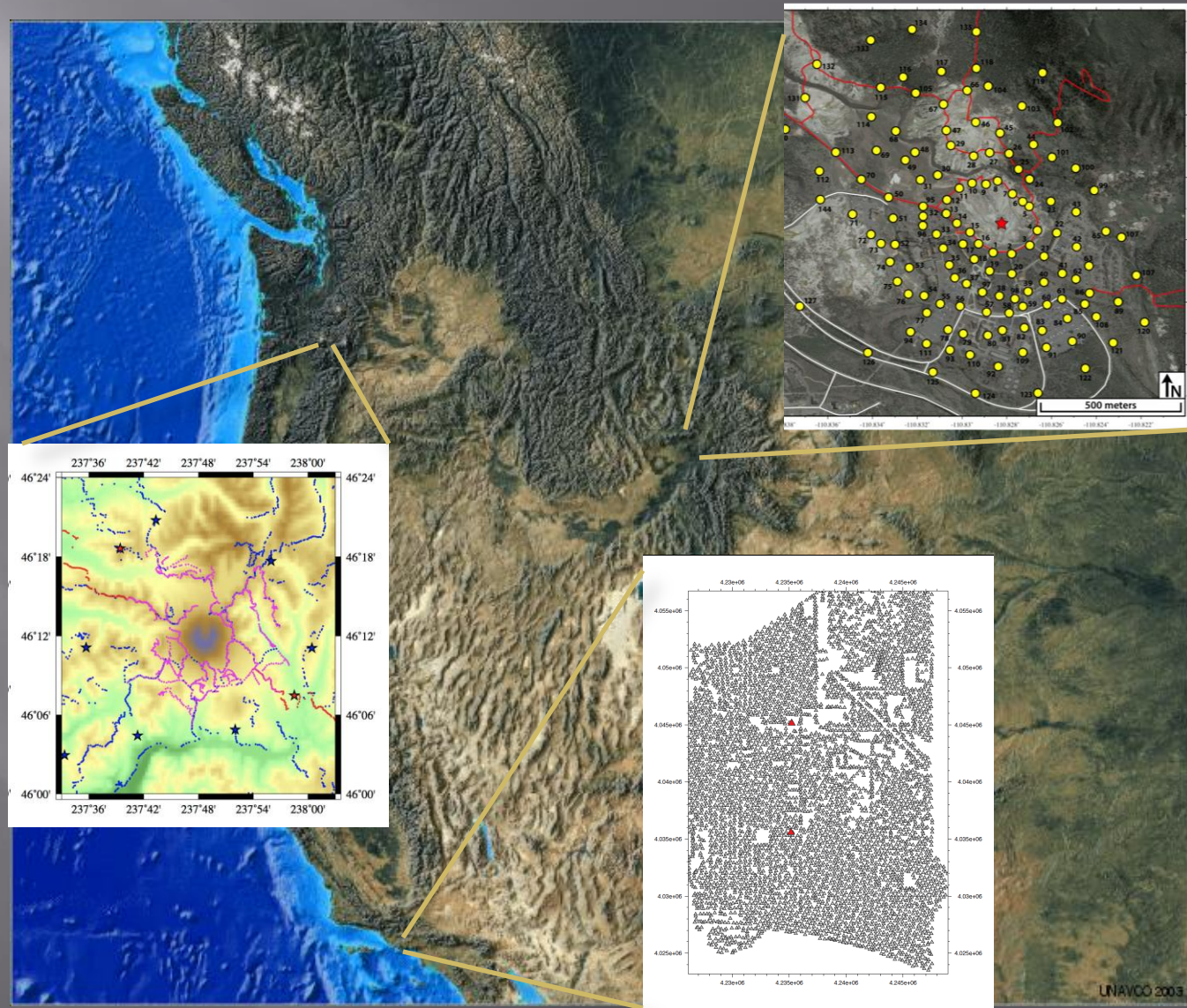


Passive seismic imaging based on dense geophone arrays

Fan-Chi Lin

University of Utah

June 6, 2017
at Cargèse France



Acknowledgement

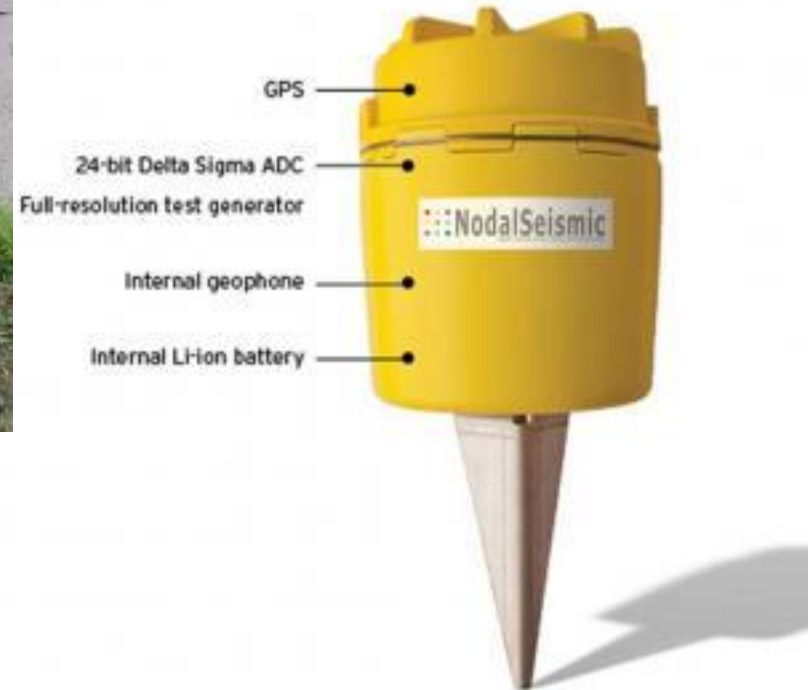
- ▣ Yadong Wang (Utah)
- ▣ Sin-Mei Wu (Utah)
- ▣ Jamie Farrell (Utah)
- ▣ Kevin Ward (Utah)
- ▣ Brandon Schmandt (U. New Mexico)



Cableless geophone system

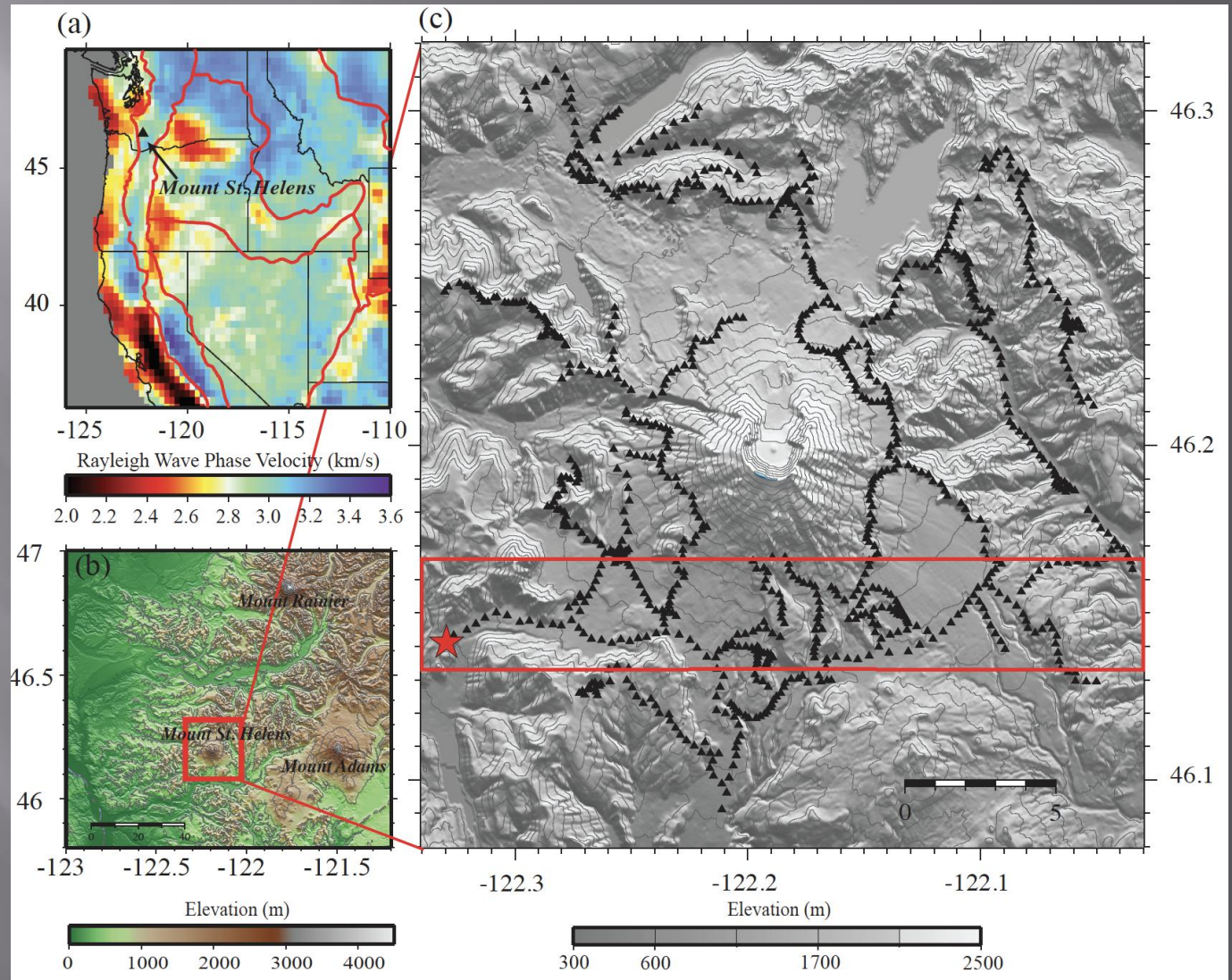


Autonomous
Cheap and easy to deploy
5 or 10 Hz corner frequency
14~30 days battery life
1 or 3 components

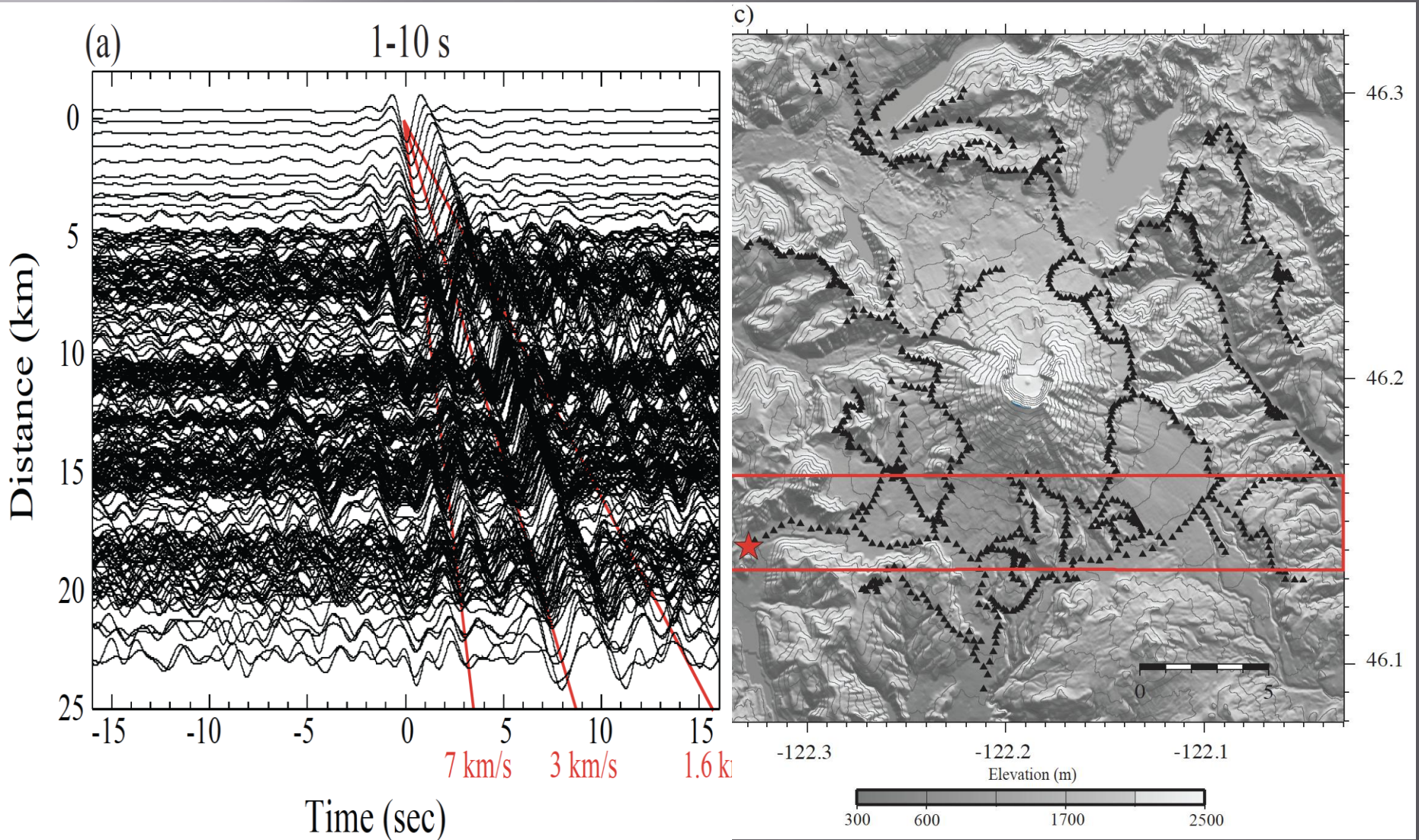


Mt St Helens

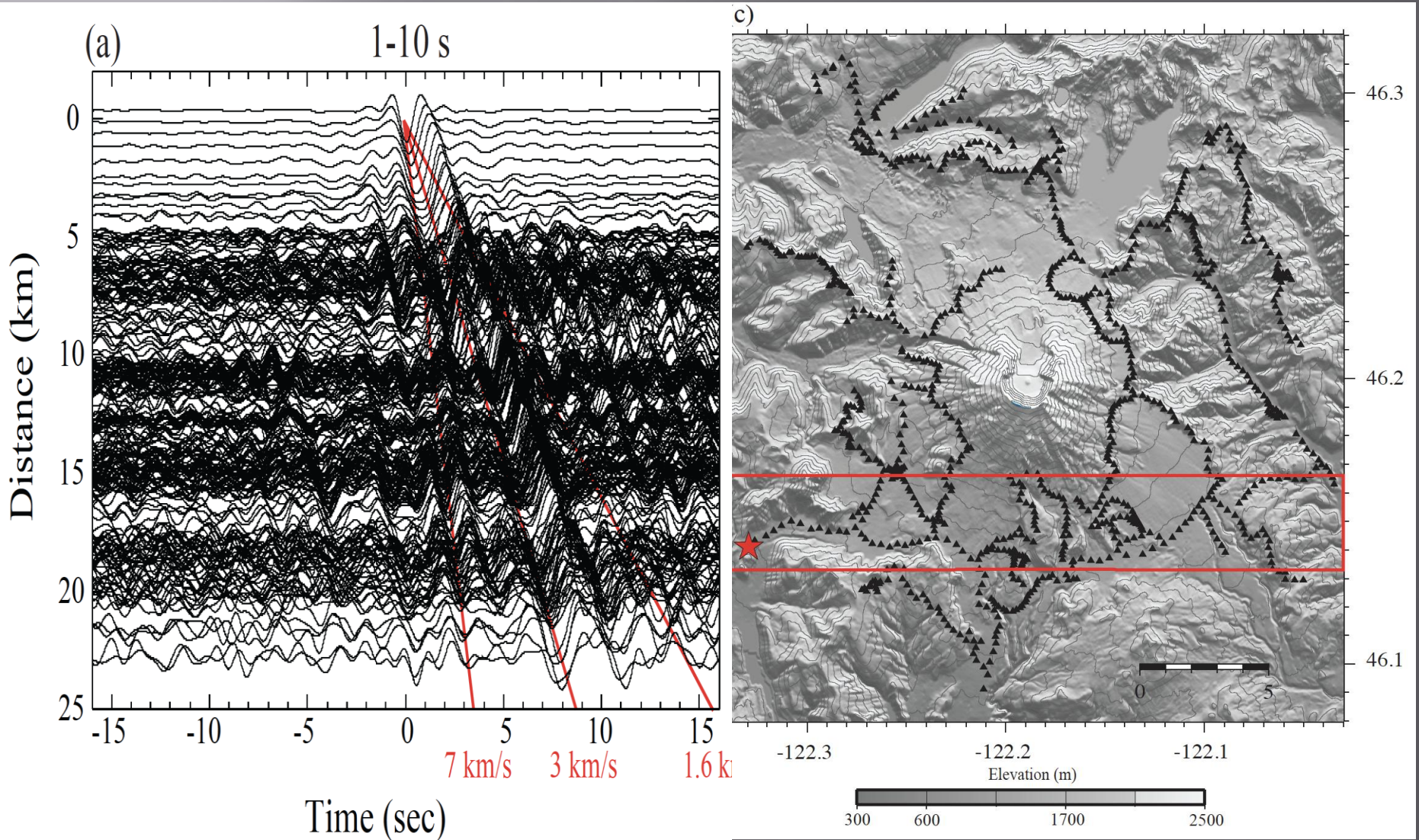
904 nodes deployed
between 18 July 2014
and 5 August 2014



Mt St Helens

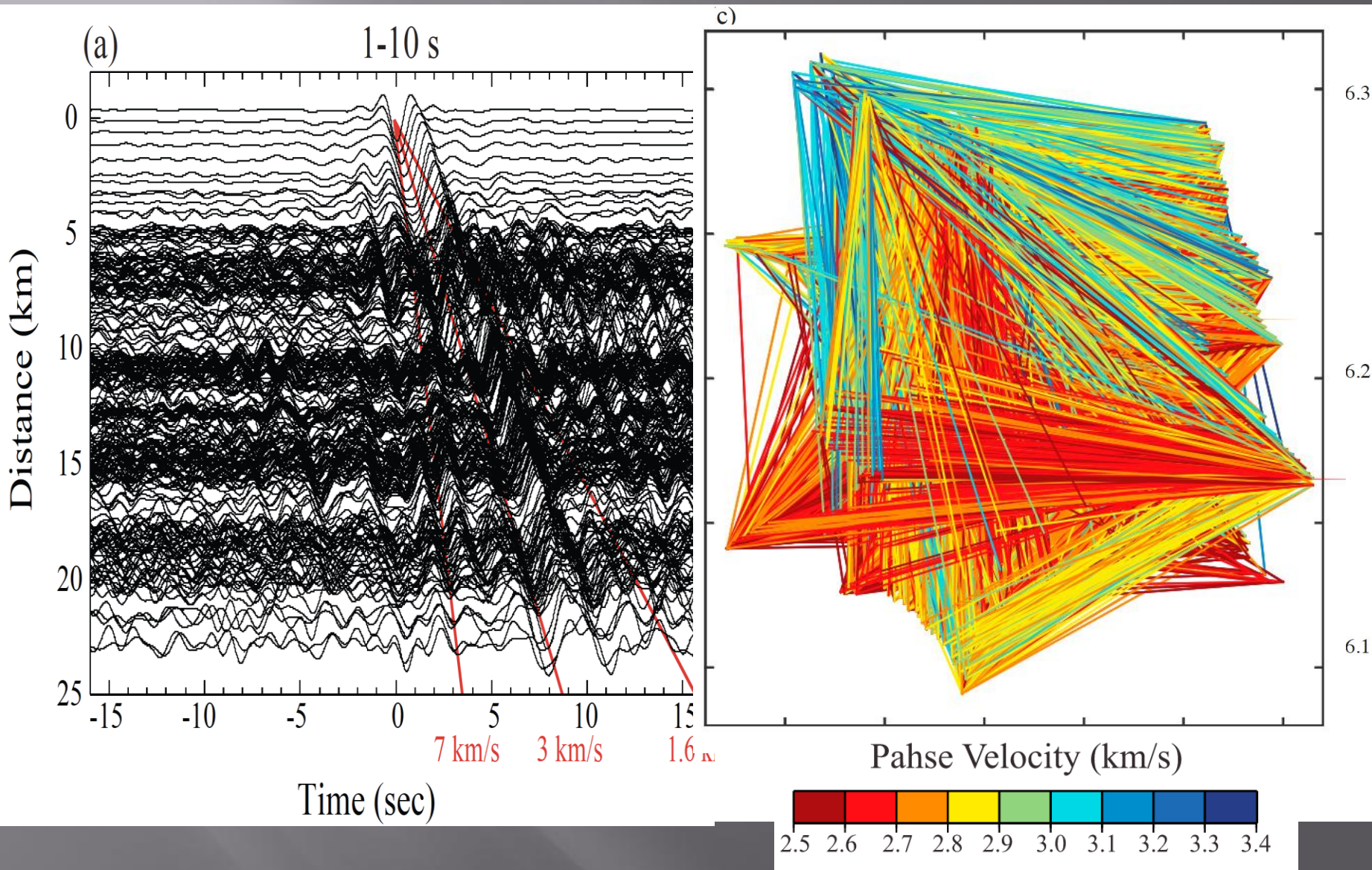


Mt St Helens

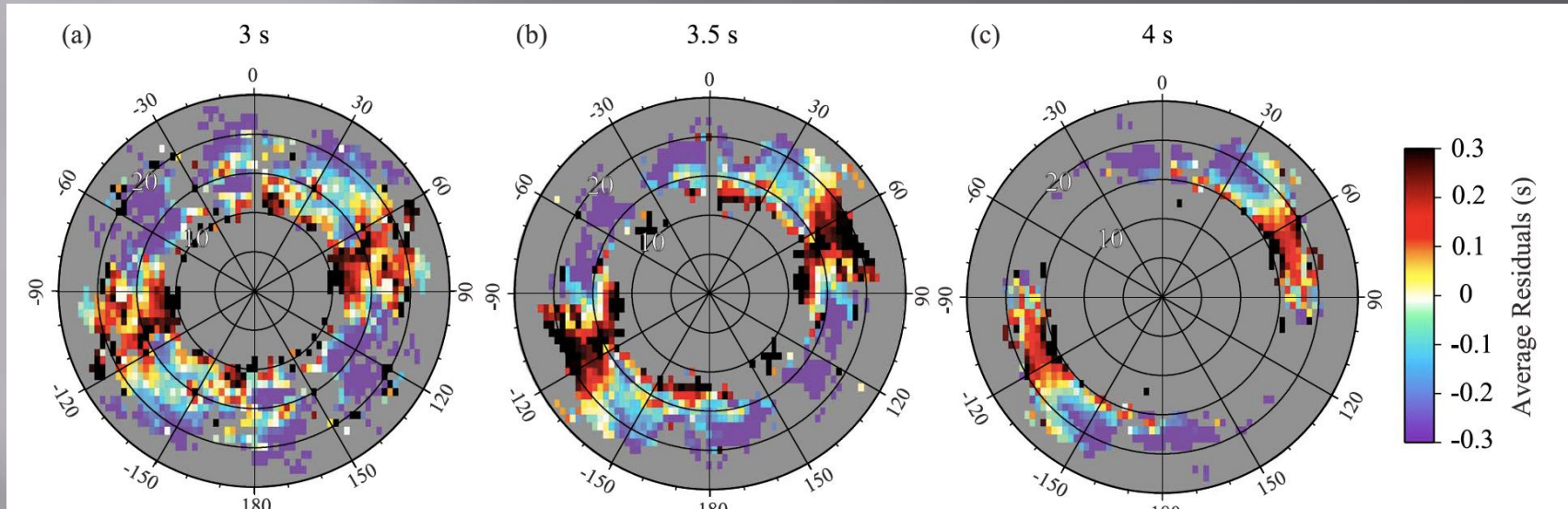


Mt St Helens

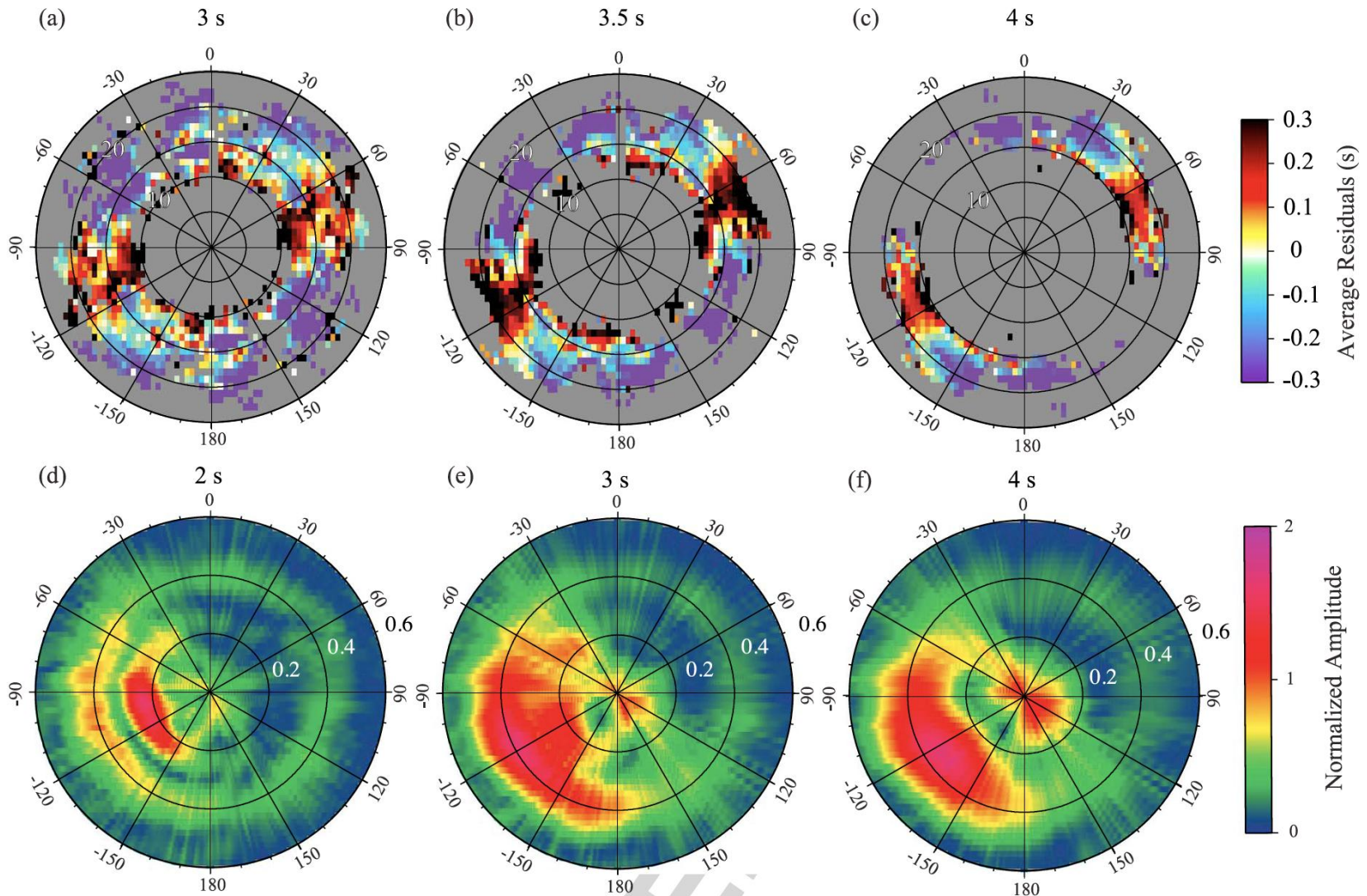
3s



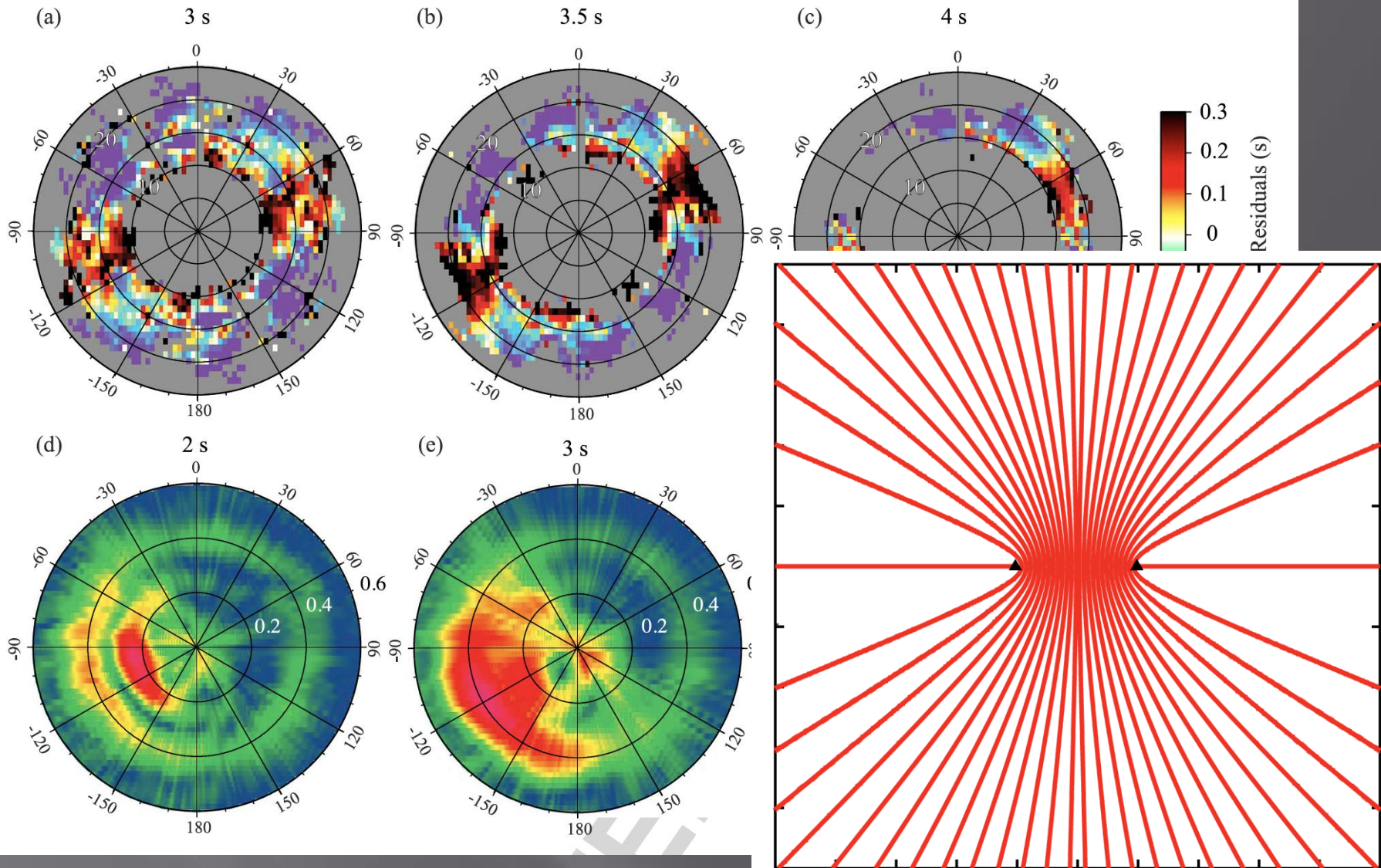
Systematic Traveltime Biases



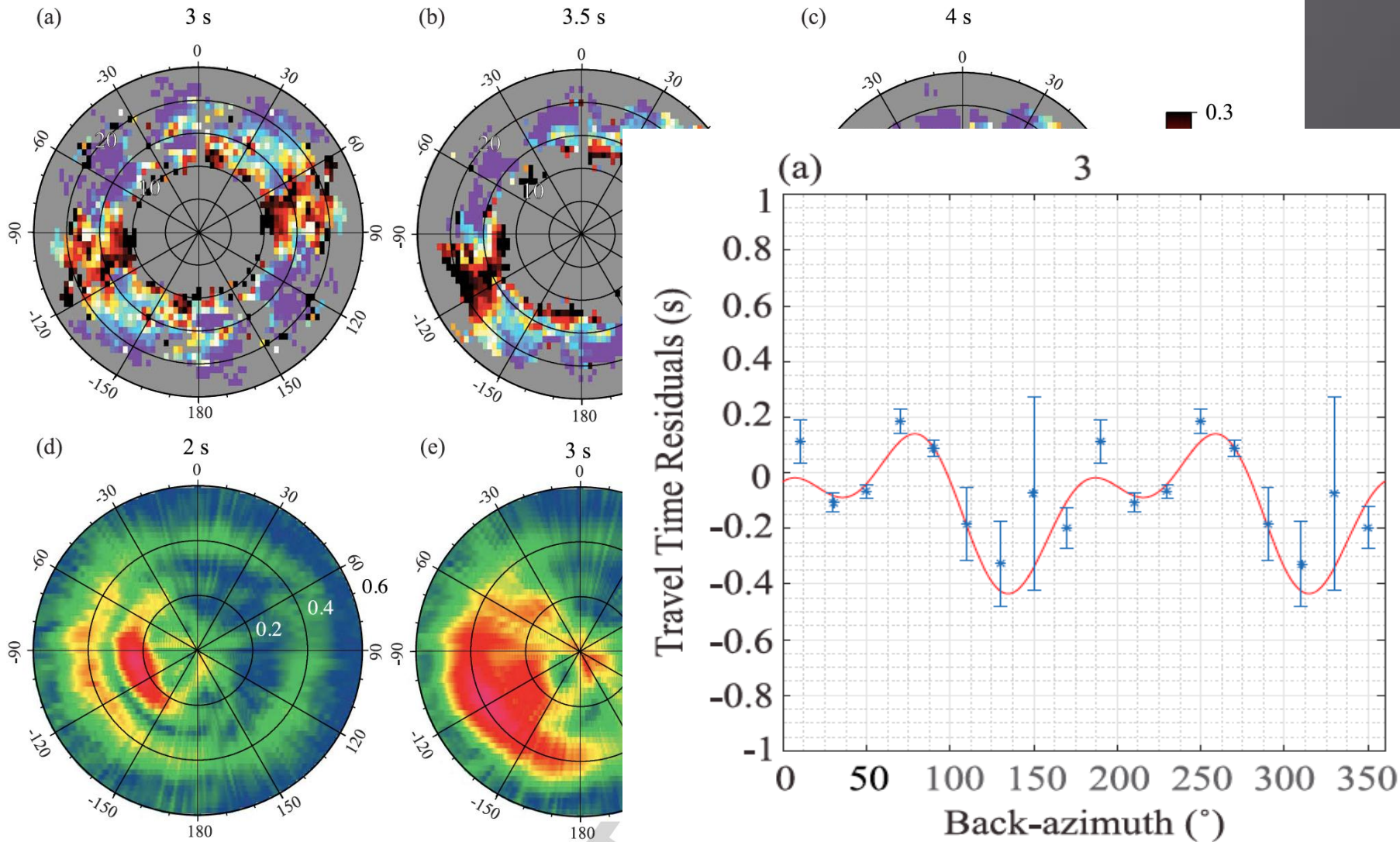
Systematic Traveltime Biases



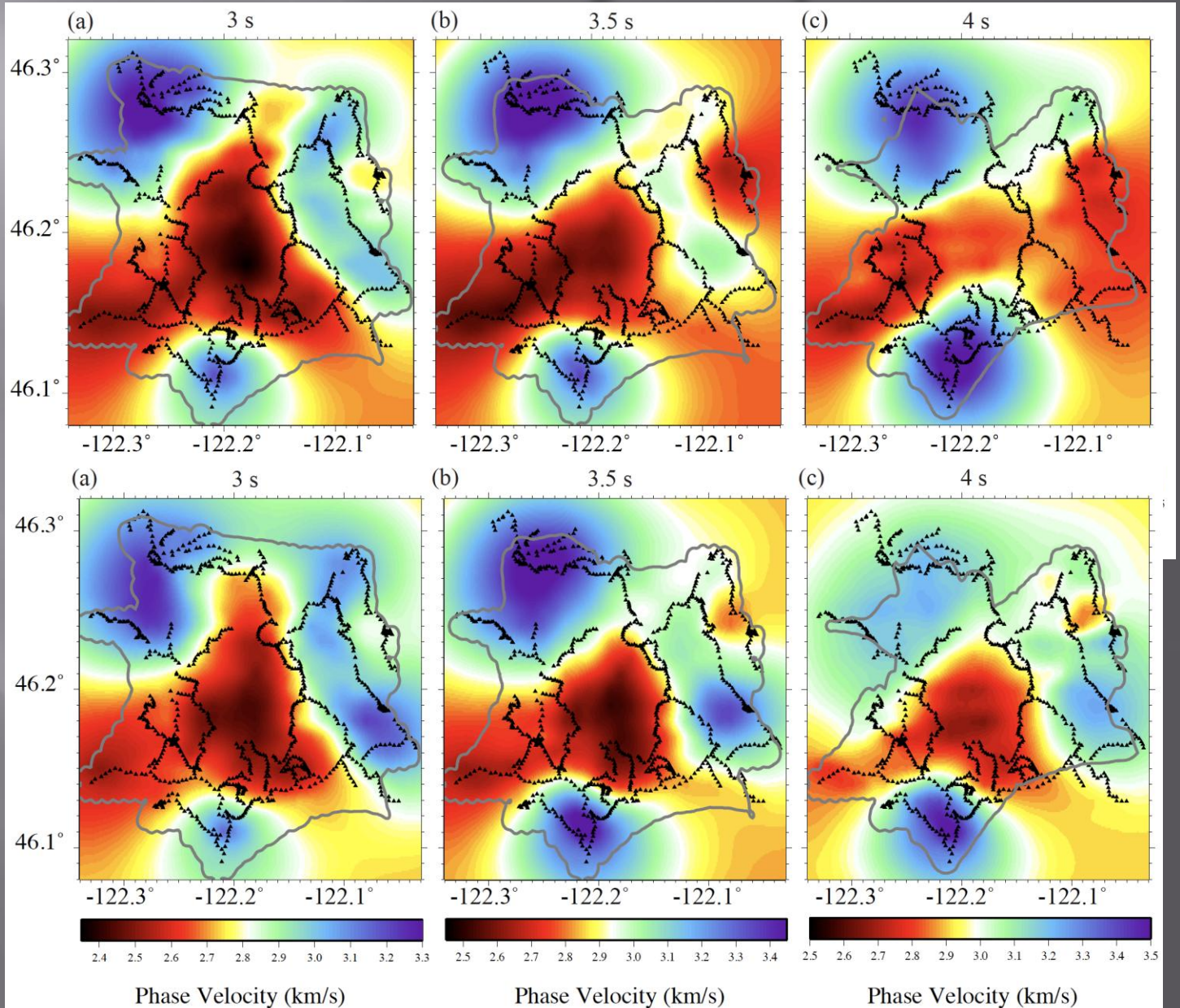
Systematic Traveltime Biases



Systematic Traveltime Biases



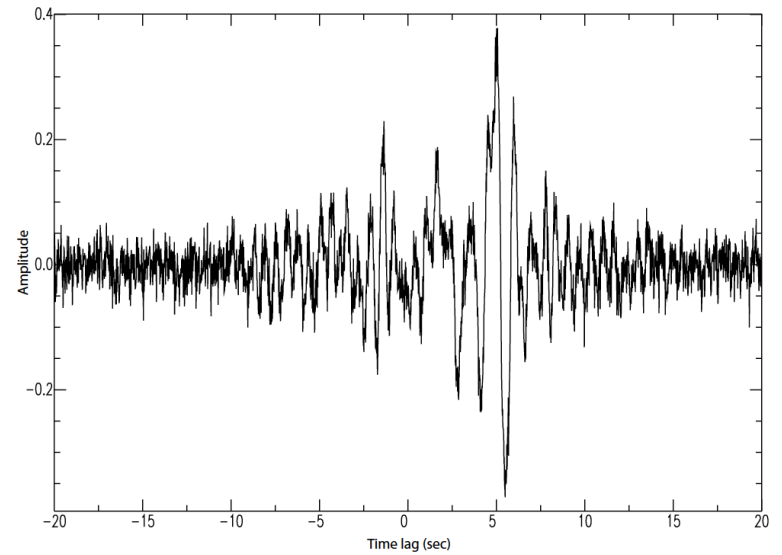
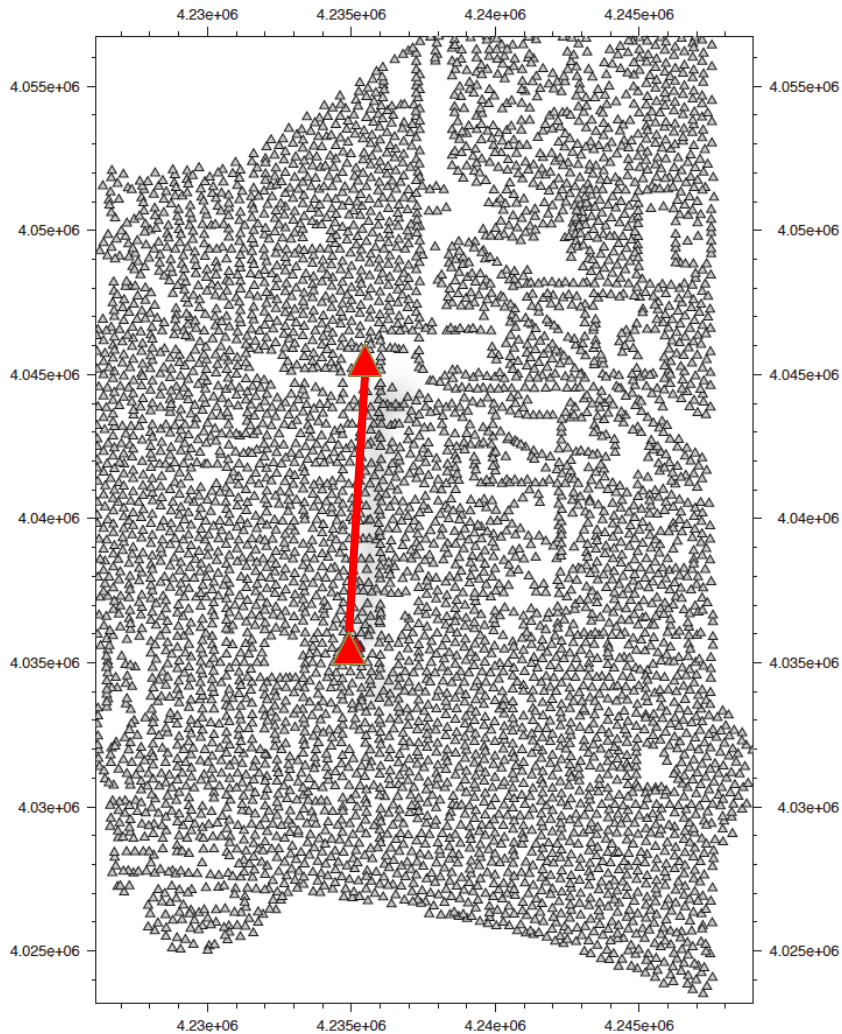
Phase velocity maps



Before Correction

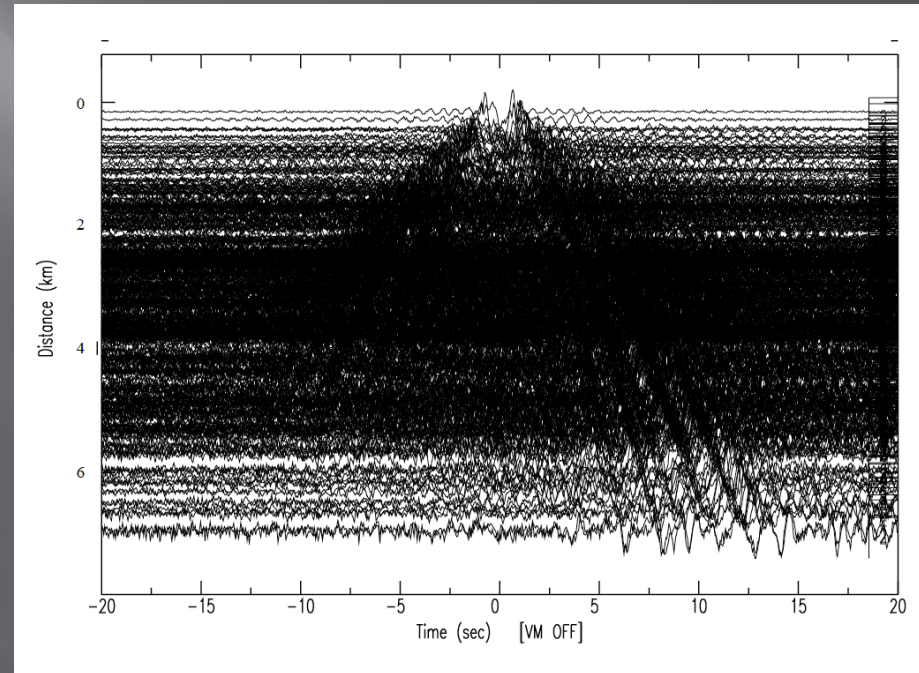
After Correction

Long Beach Dense Array



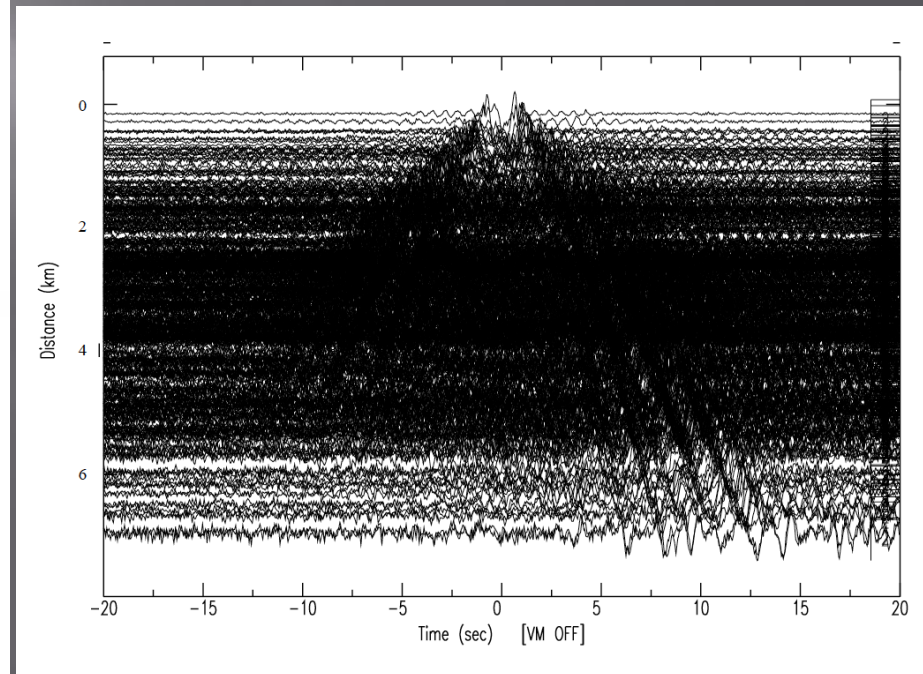
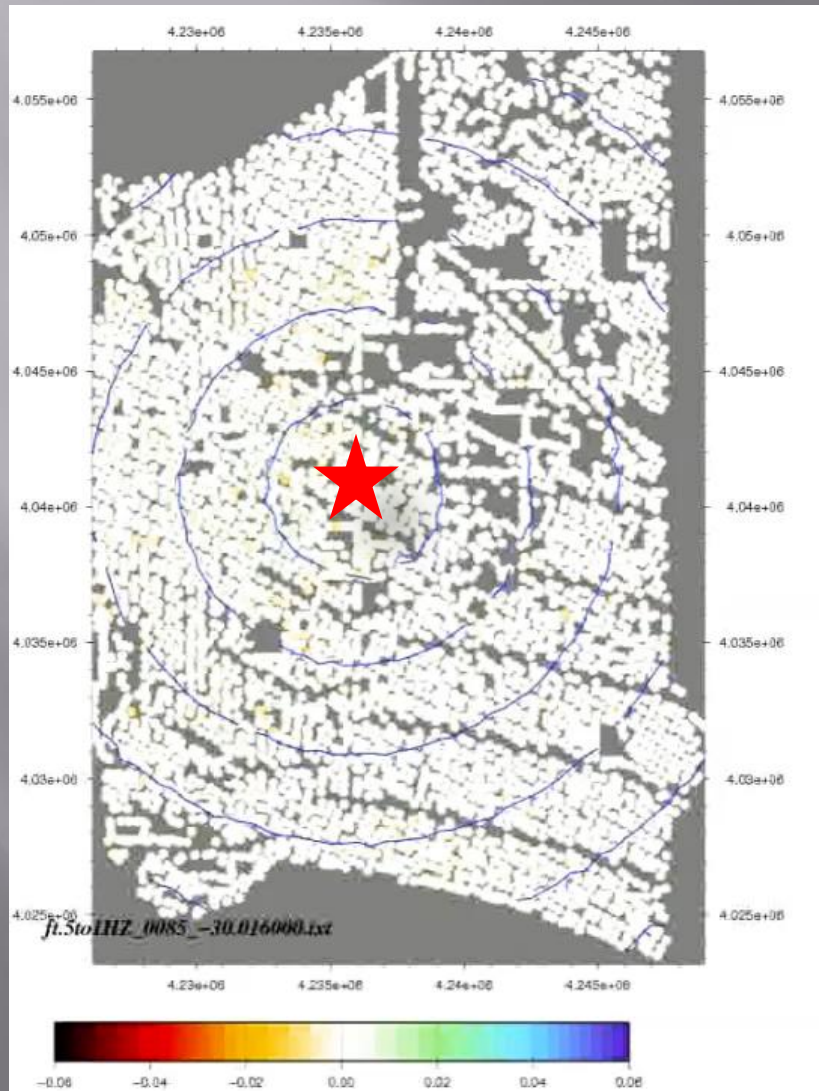
3 week of noise

Long Beach Dense Array

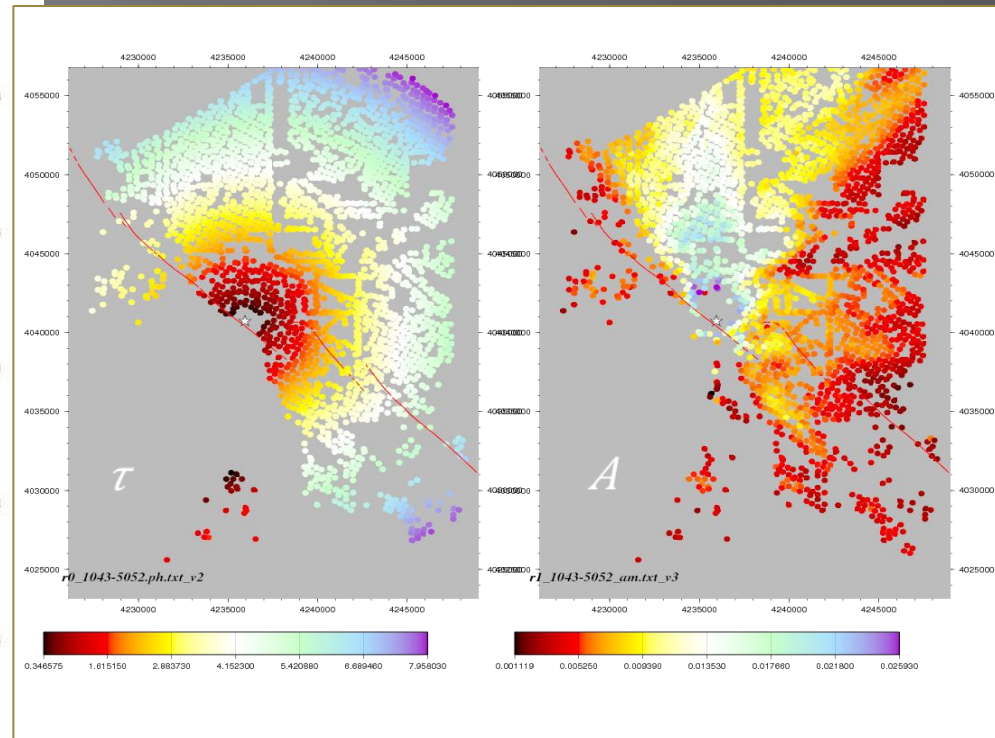
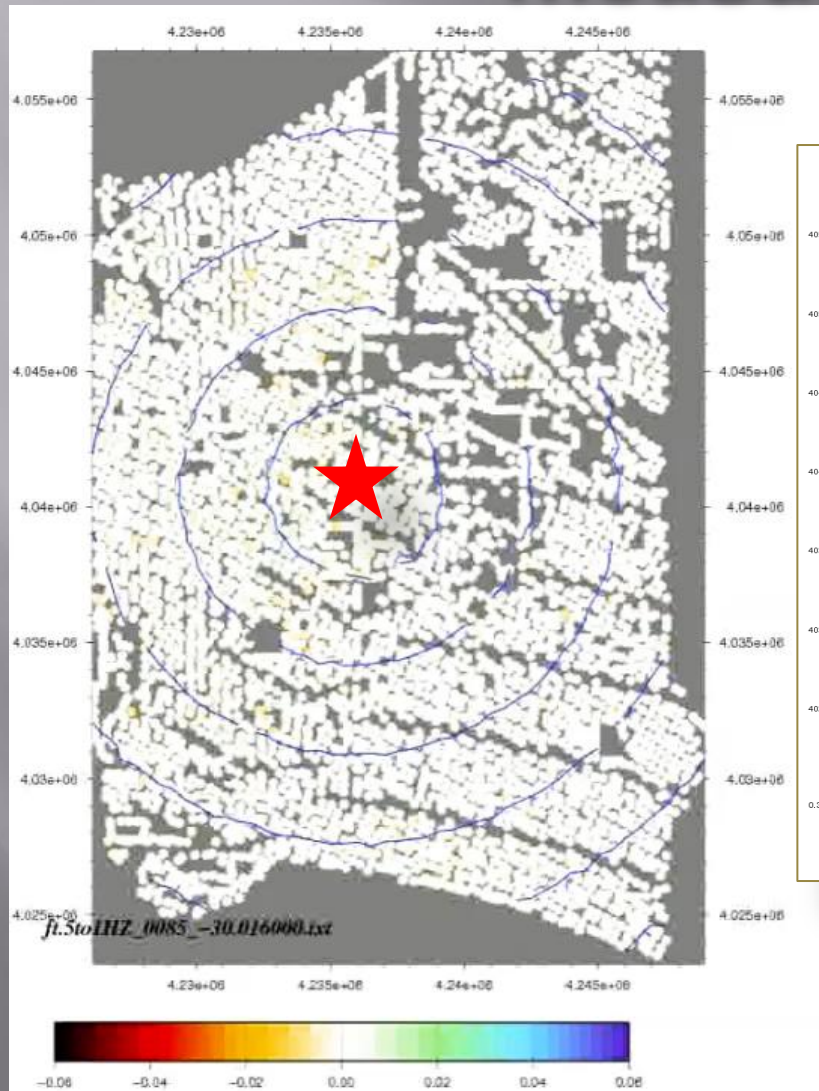


3 week of noise

Long Beach Dense Array



Traveltime and amplitude measurements



The cross-correlation wavefield

- Assume the noise wavefield $u(\mathbf{x}, t)$ satisfies the wave equation L :

$$Lu(\mathbf{x}, t) = 0$$

- The cross-correlation wavefield between a source station A and all other locations is:

$$C(\mathbf{x}_A, \mathbf{x}, t) = \int u^*(\mathbf{x}_A, t) u(\mathbf{x}, t + t) dt$$

- The cross-correlation wavefield satisfies the wave equation:

$$LC(\mathbf{x}_A, \mathbf{x}, t) = \int u^*(\mathbf{x}_A, t) Lu(\mathbf{x}, t + t) dt = 0$$

Solution of the 2D wave equation

- Real part:

$$\frac{1}{\alpha(\mathbf{r})^2} = \nabla t \cdot \nabla t - \frac{\nabla^2(A/b)}{w^2(A/b)}$$

Helmholtz tomography (Lin & Ritzwoller 2011, GJI)

- High-frequency approximation (ray-theory approach)

$$\frac{\hat{\mathbf{k}}}{c} @ \nabla t$$

Eikonal tomography (Lin et al. 2009, GJI)

- Imaginary part:

$$\frac{-2a}{c} + \frac{2\nabla b \cdot \nabla t}{b} = \frac{2\nabla A \cdot \nabla t}{A} + \nabla^2 t$$

Corrected amplitude decay
Apparent amplitude decay
Focusing/defocusing

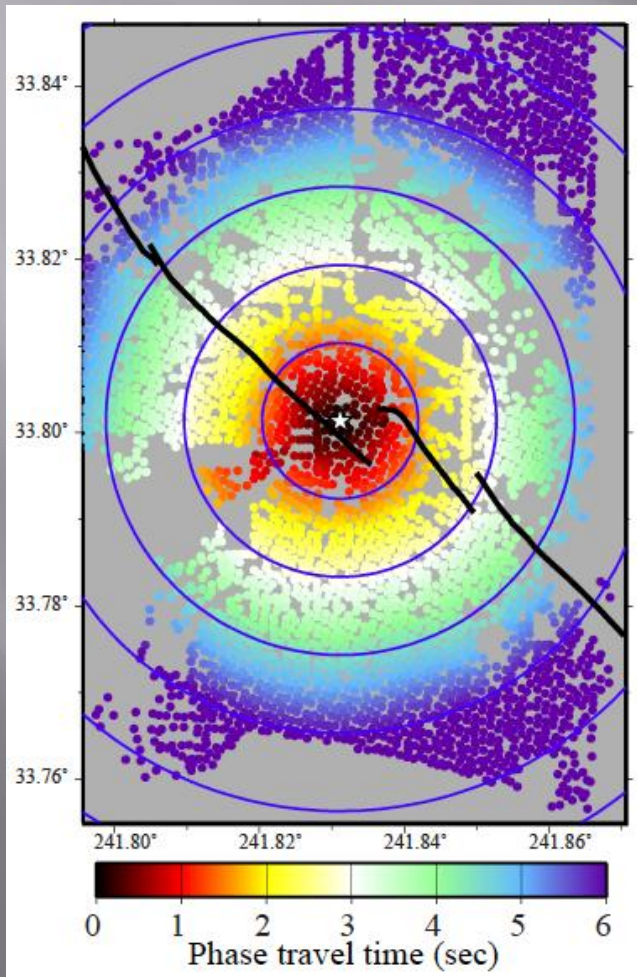
A: observed amplitude

τ : observed phase travel time

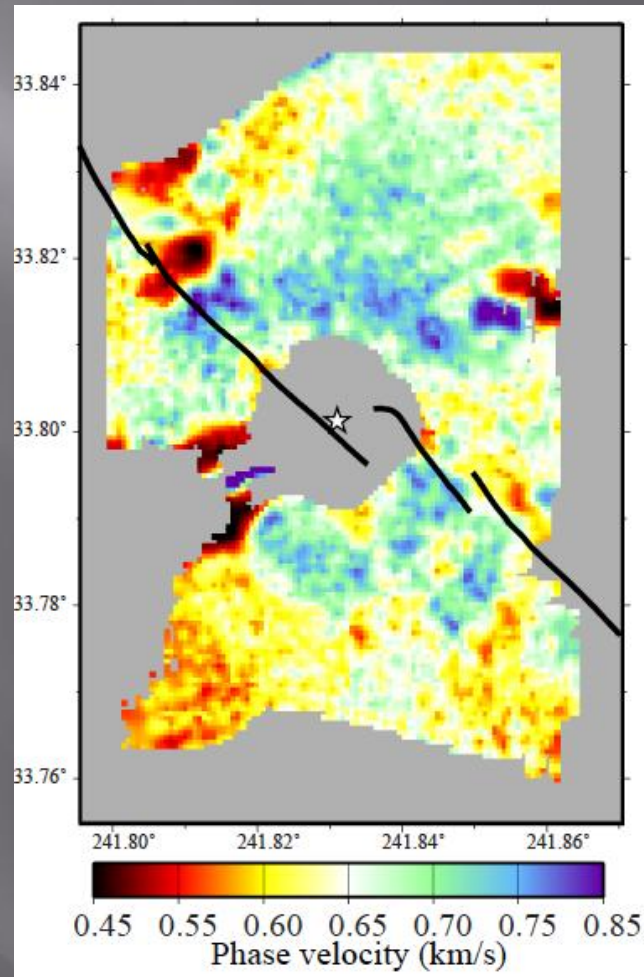
β : local amplification factor

Eikonal tomography

Travel time



Phase velocity



$$\frac{\hat{k}}{c} @ \nabla t$$

Eikonal Tomography

Eikonal tomography

N=20

N=100

N=520

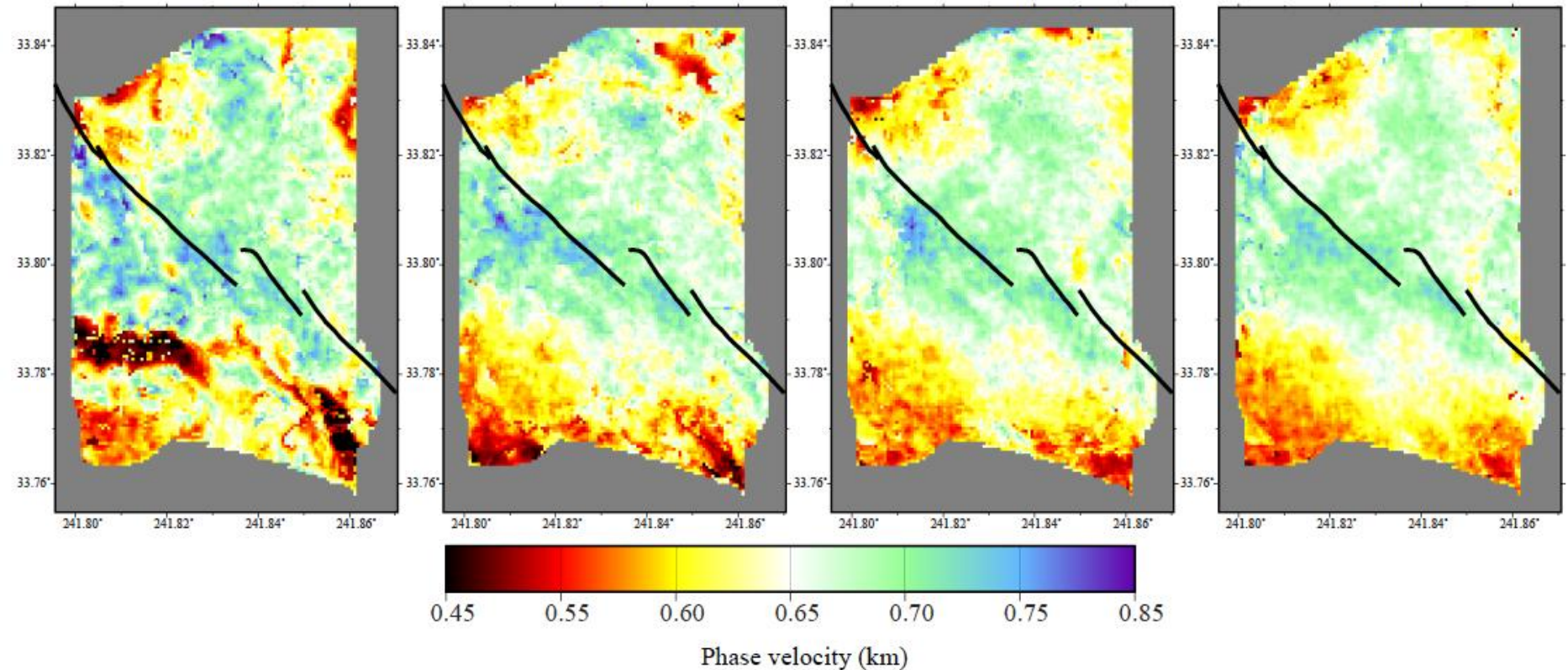
N=5204

N=20

N=100

N=520

N=5204



Azimuthal anisotropy

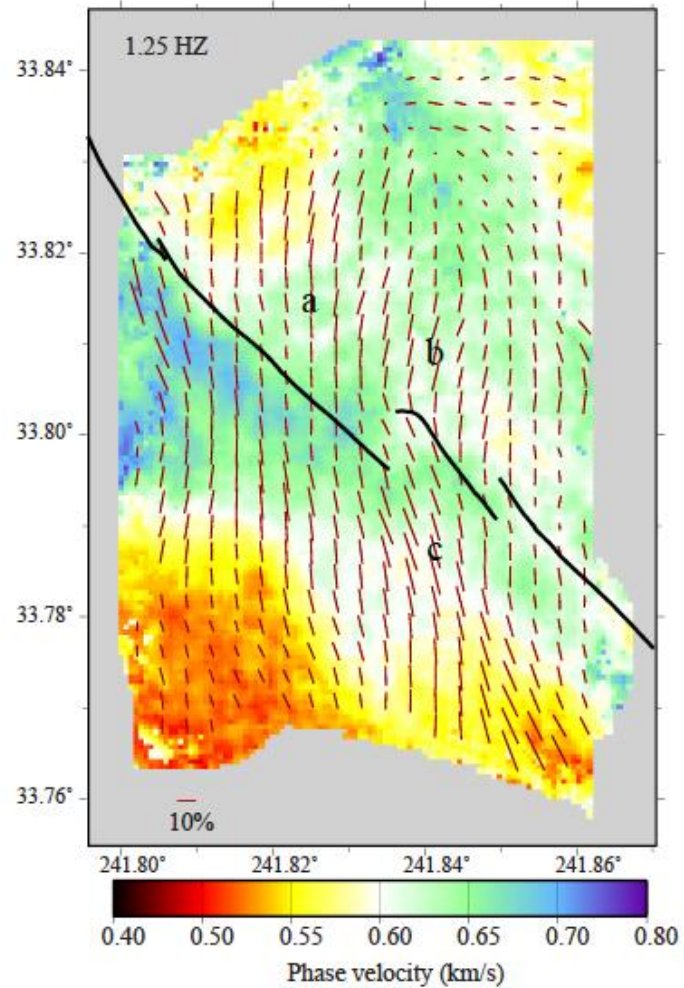
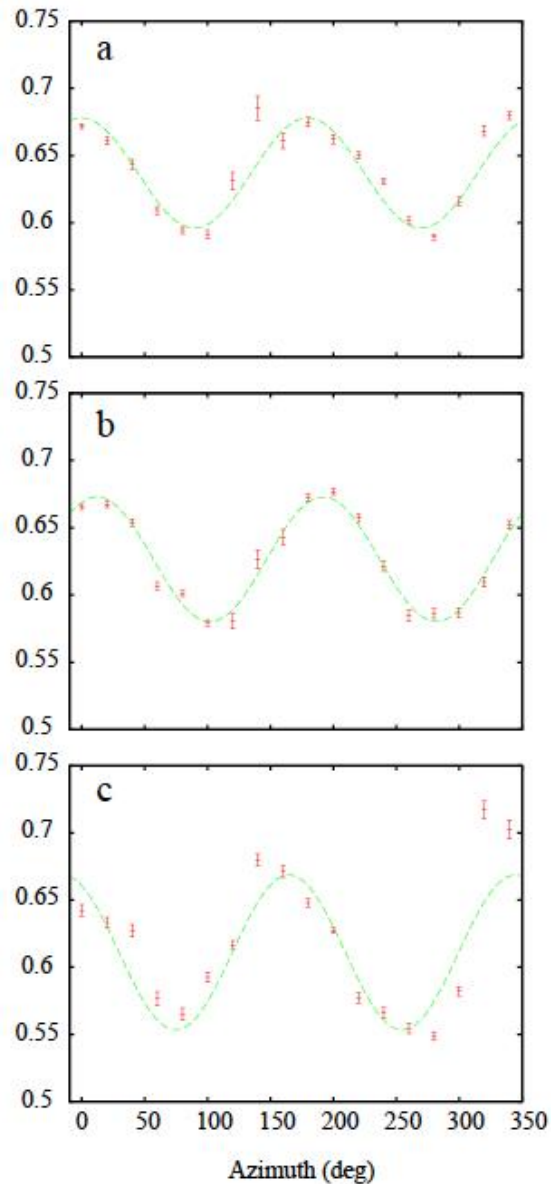
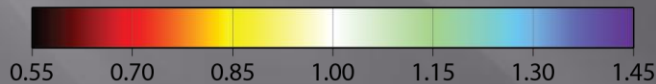
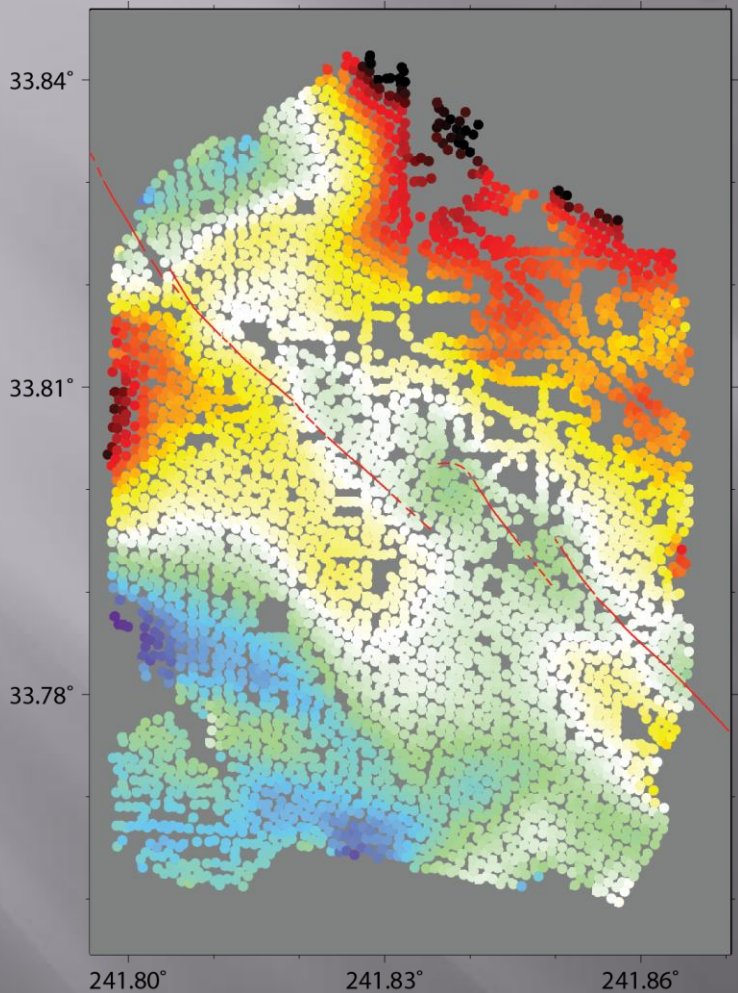


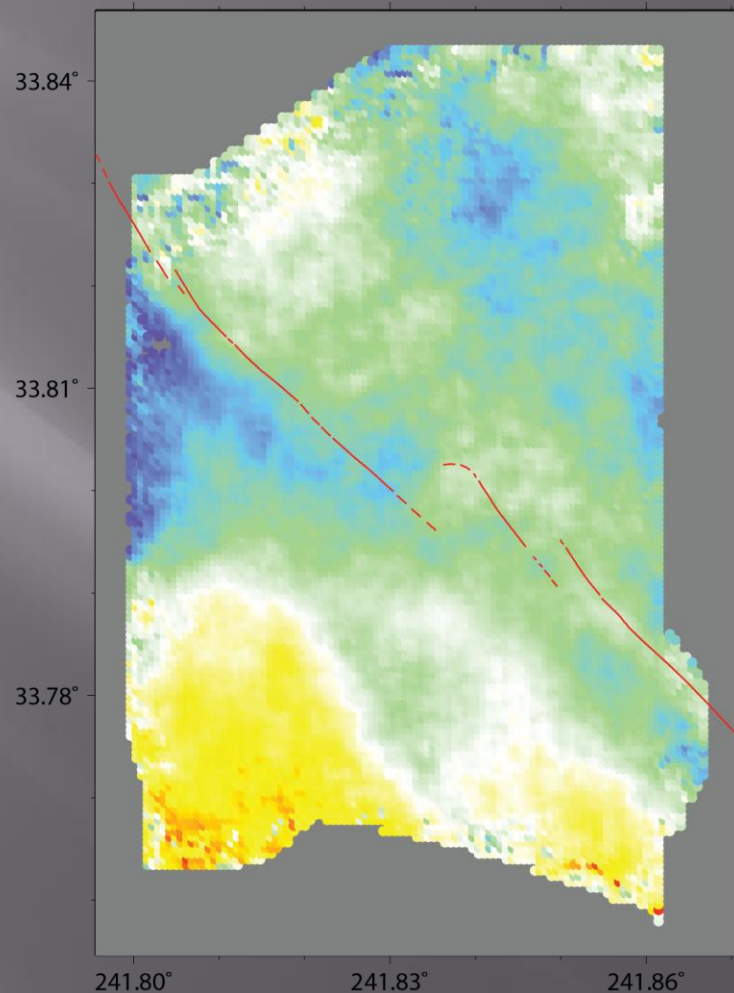
Figure 12

Amplification



1 Hz Amplification Factor

Bowden et al., 2015



1.4 Hz Phase Velocity (km/s)

Lin et al., 2013

THE UPPER GEYSER BASIN IMAGING EXPERIMENT



Photo by Stephen Wilkes, National Geographic Magazine

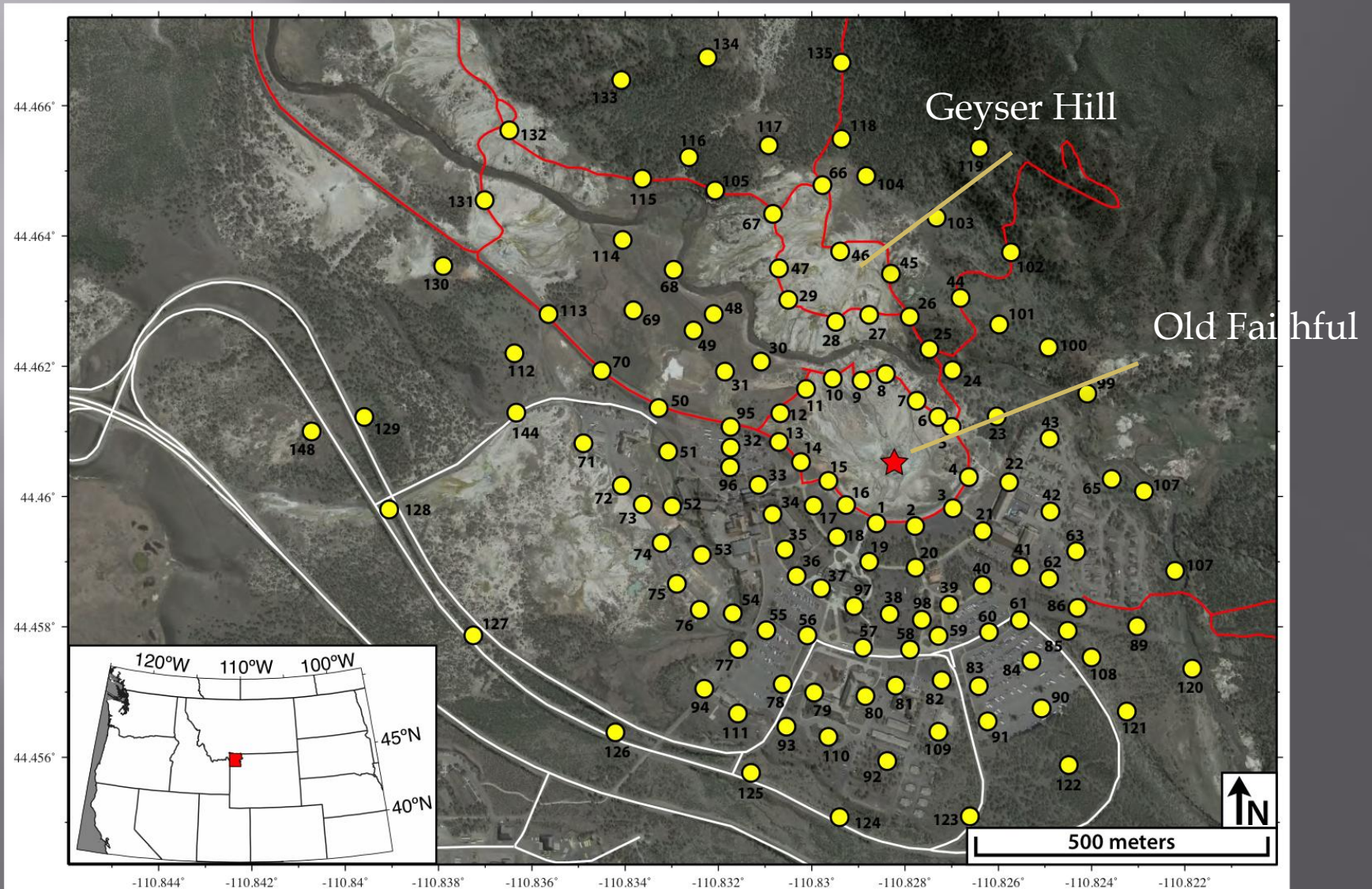
3C geophones



Cableless 3C geophone system



The Old Faithful Array

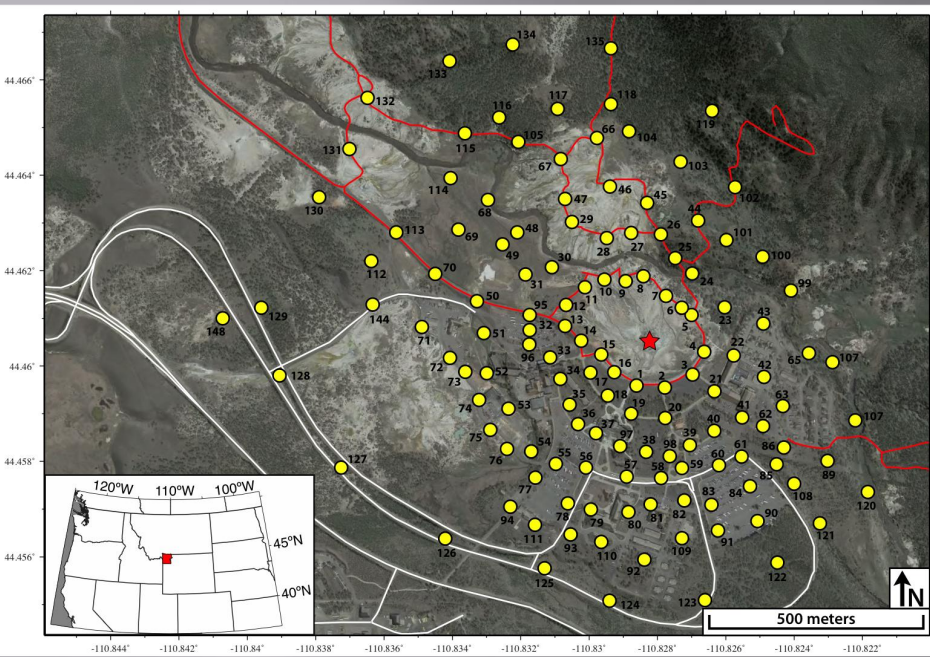


Old Faithful deployment

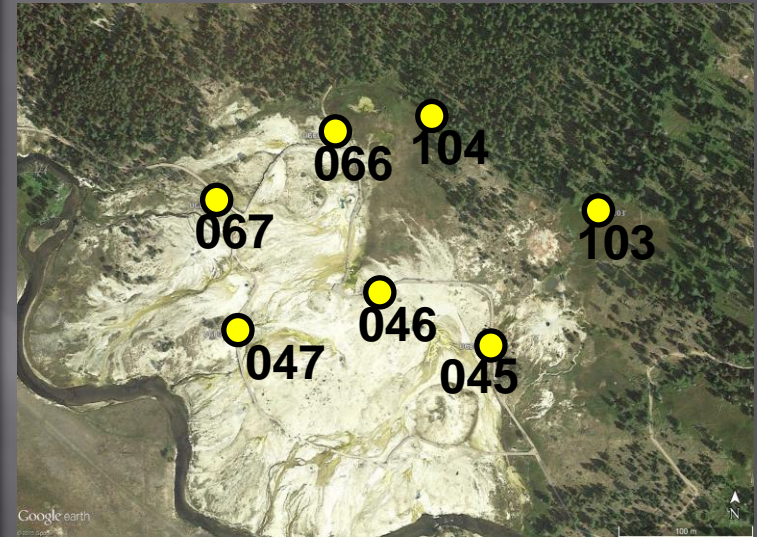
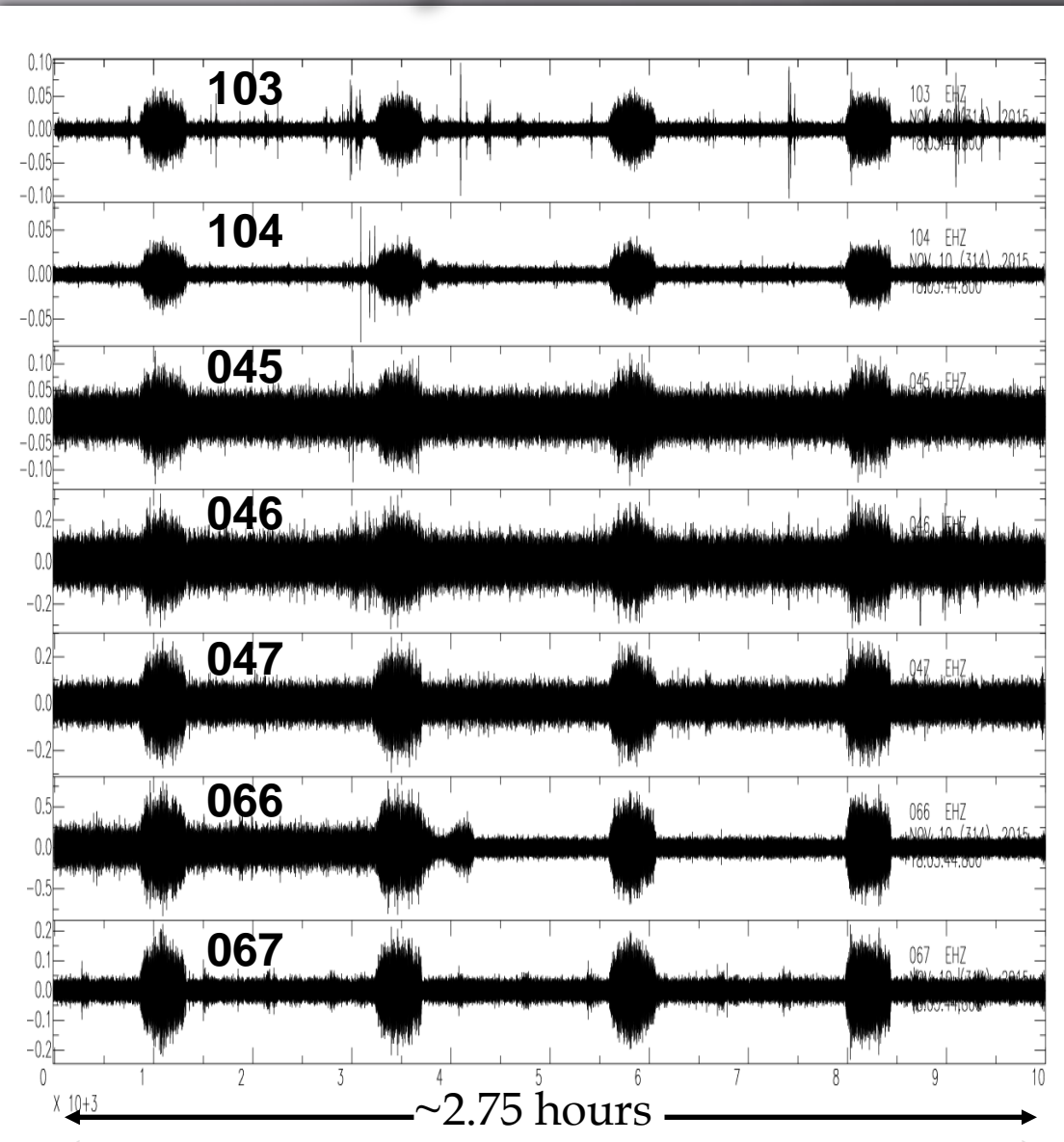




Temporal Noise Variation

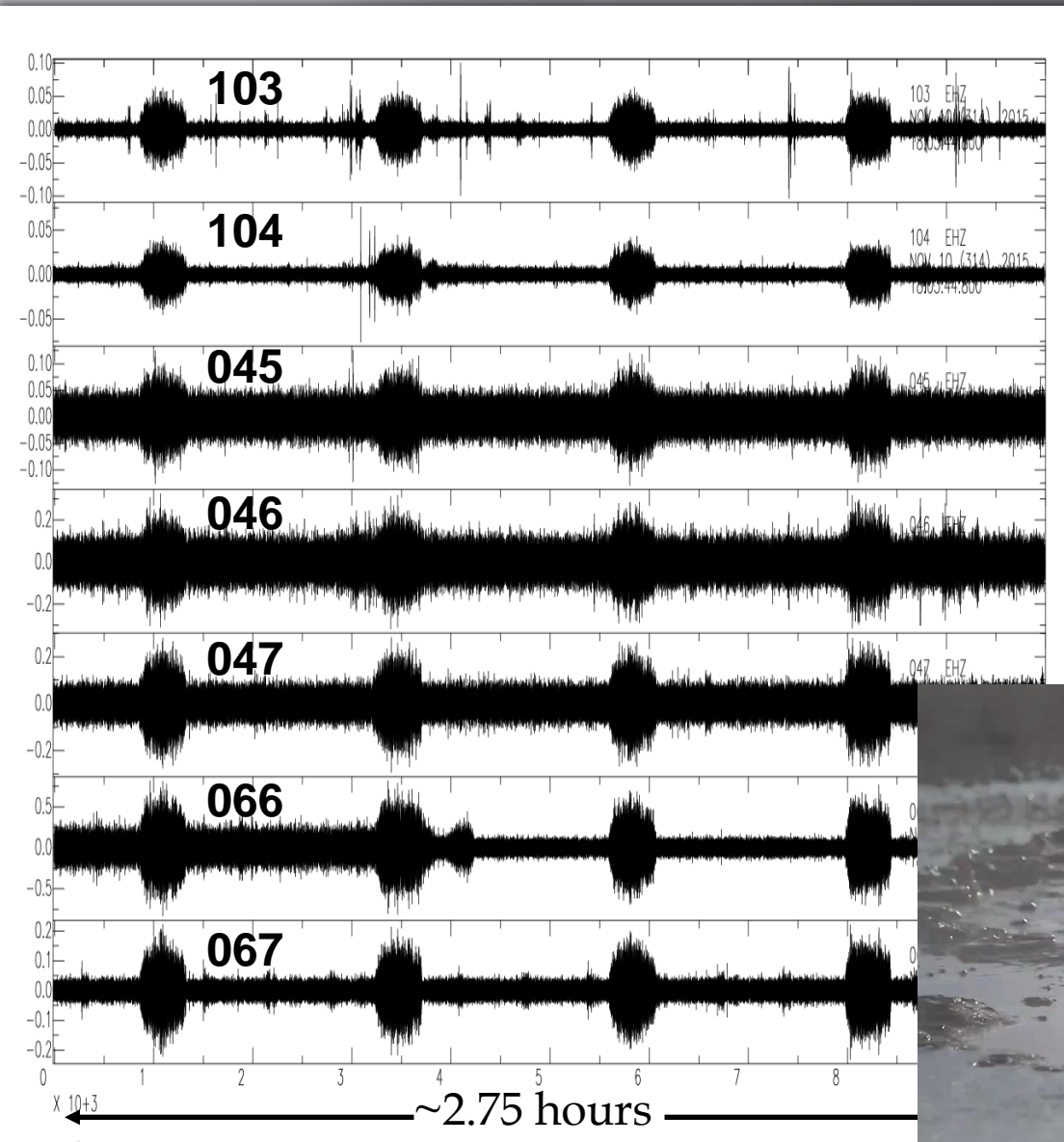


Hydrothermal Tremor

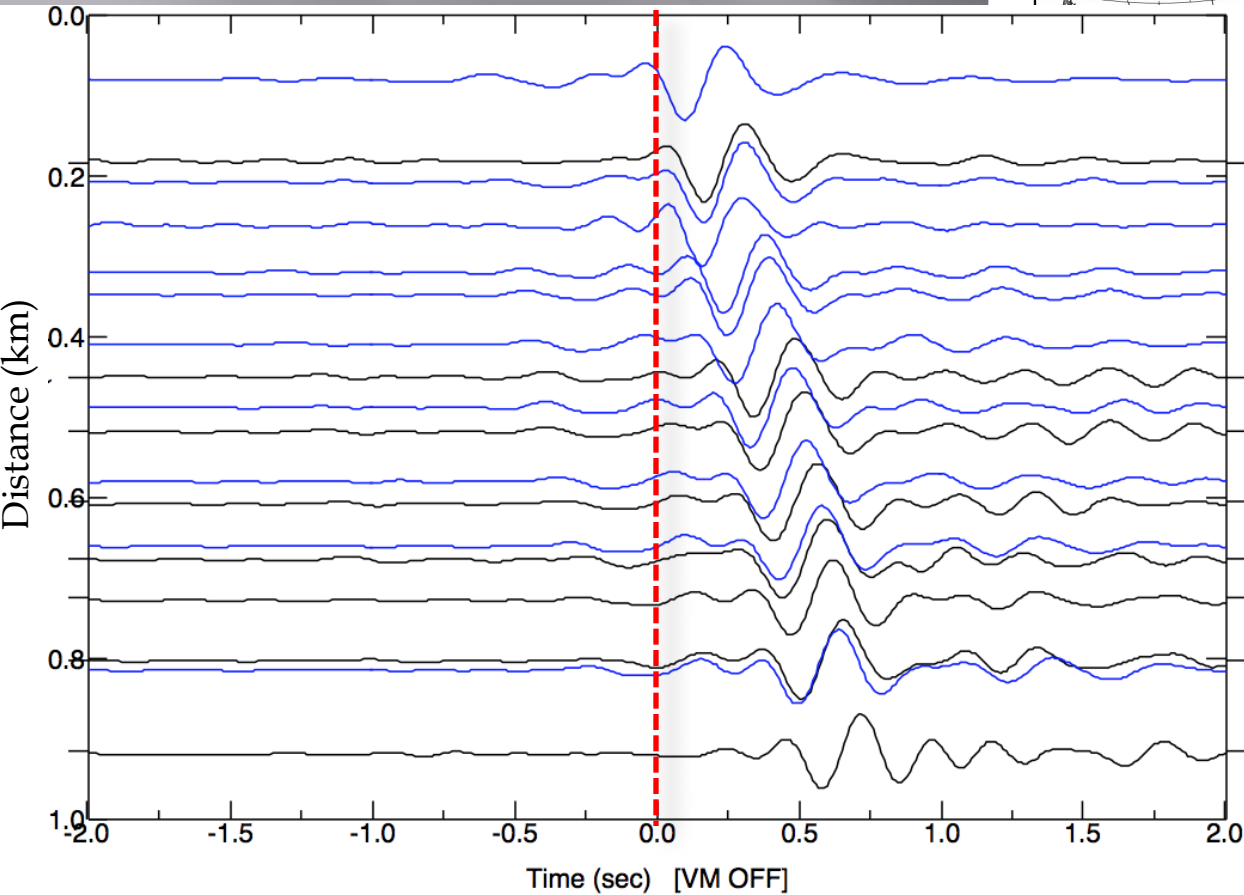
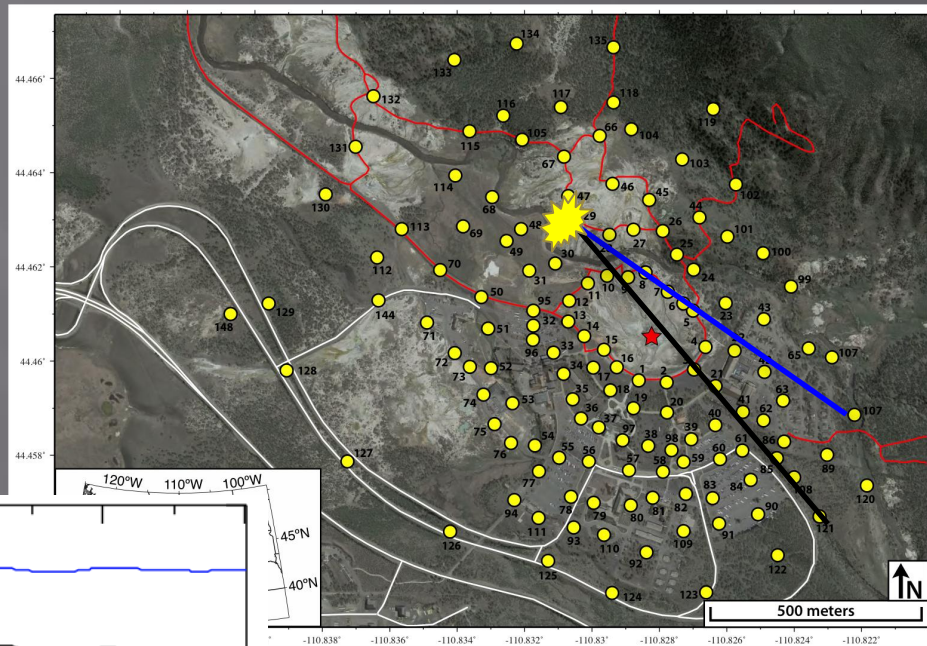


Raw seismic waveforms

Doublet Pool

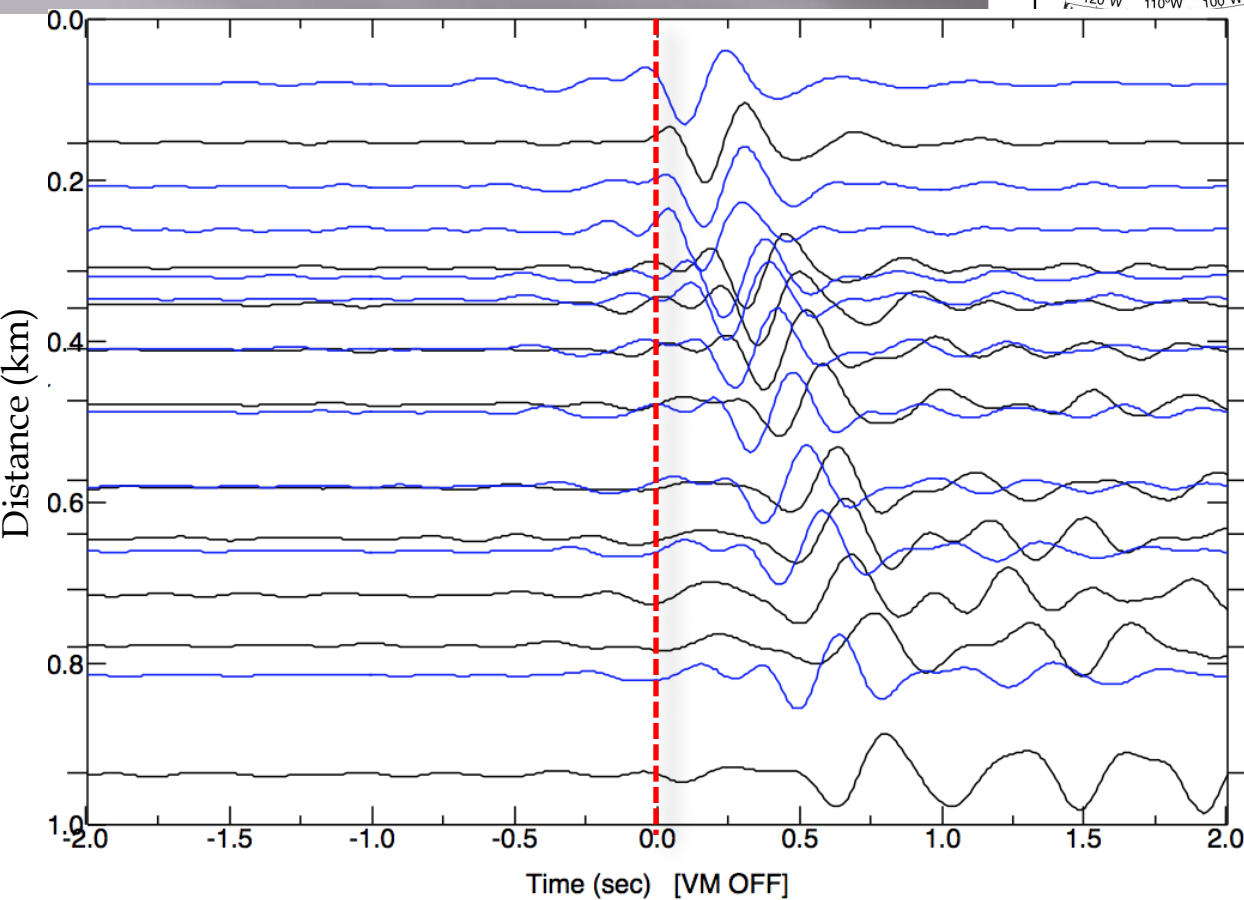
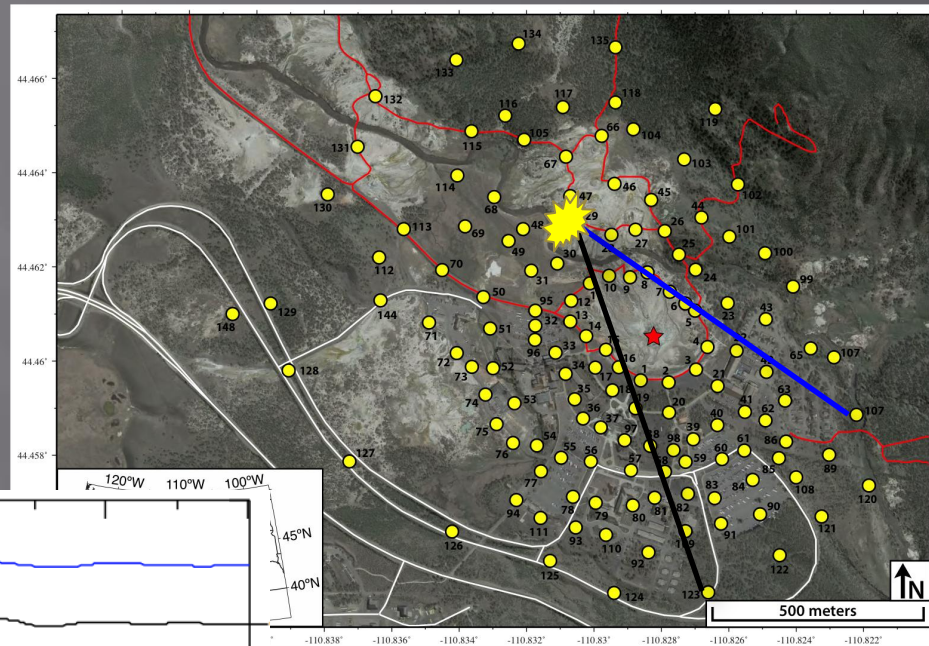


Virtual Record Section



1-5 Hz

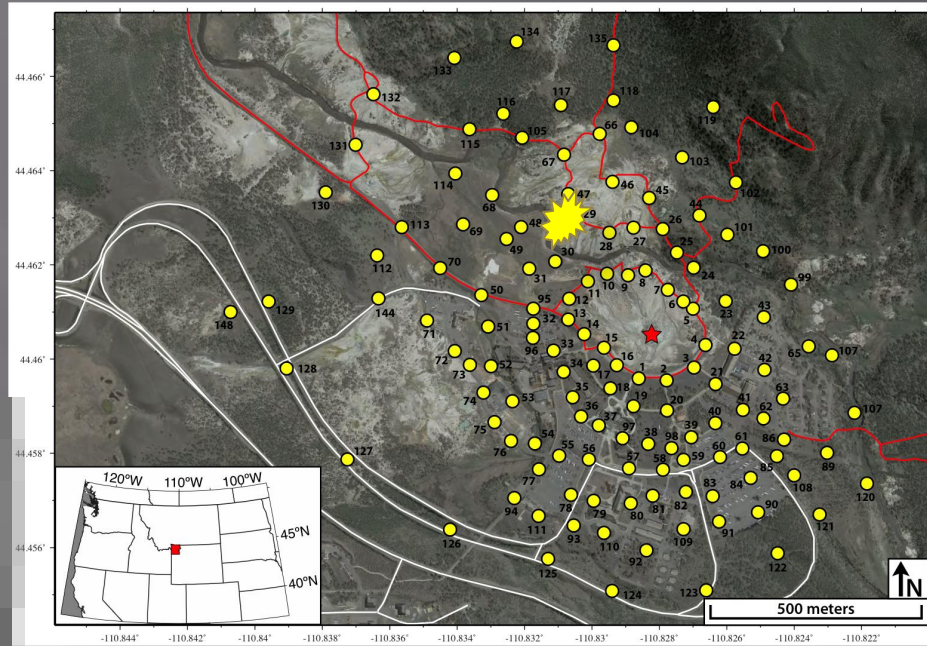
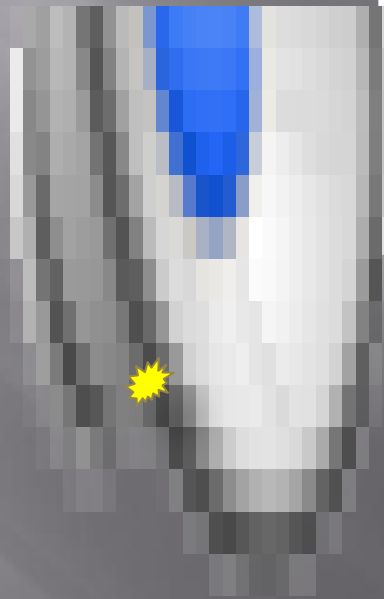
Virtual Record Section



1-5 Hz

Cross-correlation wavefield

0.5 - 3 Hz



Conclusions

- ▣ The availability of low cost autonomous geophones has promoted the development of passive shallow imaging with dense geophone arrays.
- ▣ The noise source distribution can be highly heterogeneous at short periods and can affect the retrieval of the Green's function based on the noise cross-correlation method.
- ▣ 2D array analysis method can be used to mitigate the source effect and study both surface wave velocity and amplification structures.
- ▣ Seismic imaging in Yellowstone has now revealed new information on shallow structure associated with the Old Faithful hydrothermal system.

Questions?

