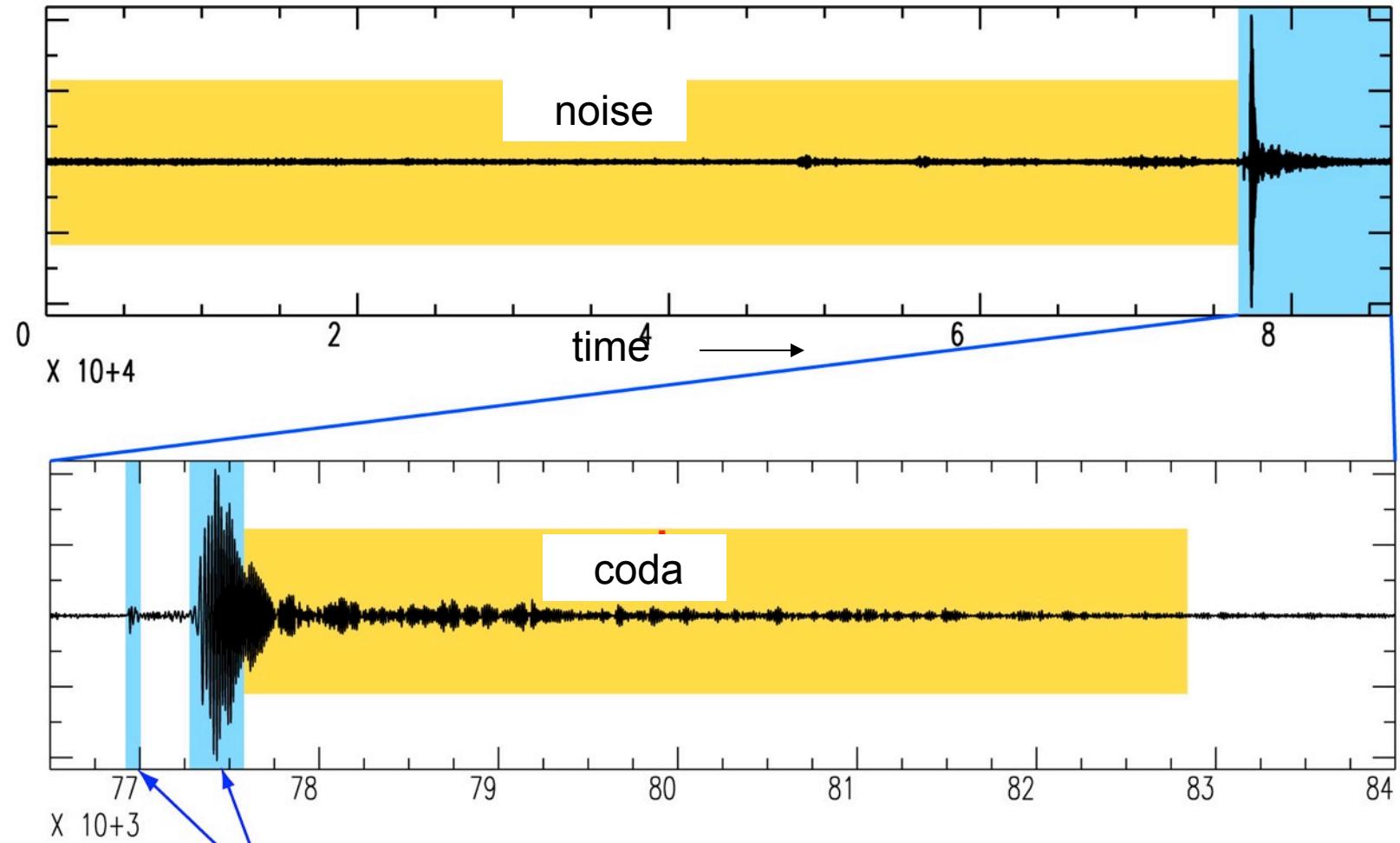


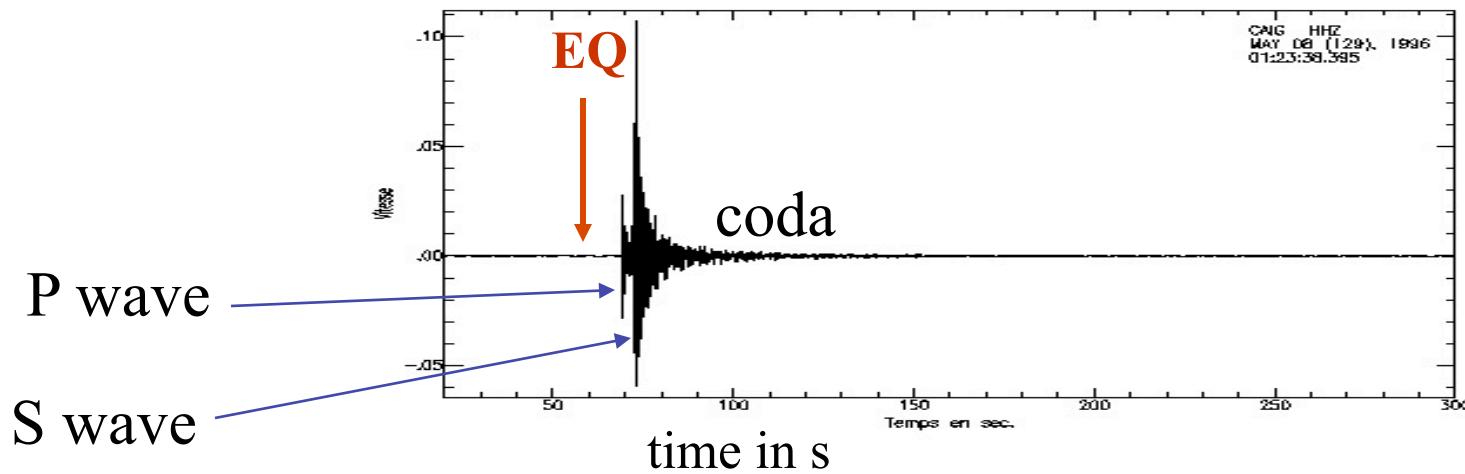
Ground displacement



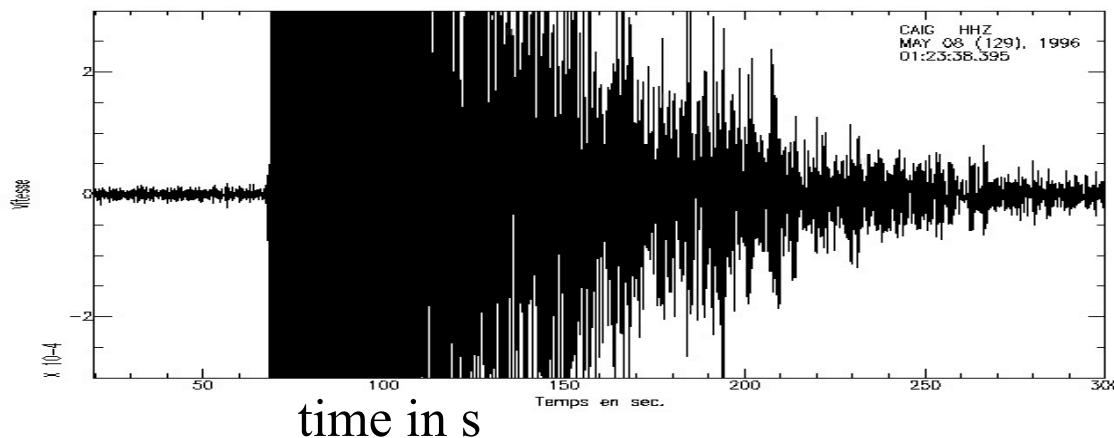
Direct waves

Ballistic waves used for imaging

Example of a record of a local earthquake in the band .5-20Hz



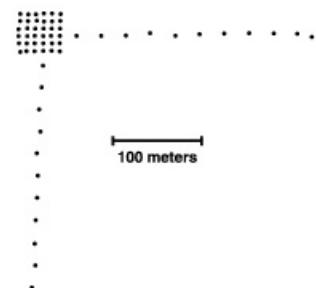
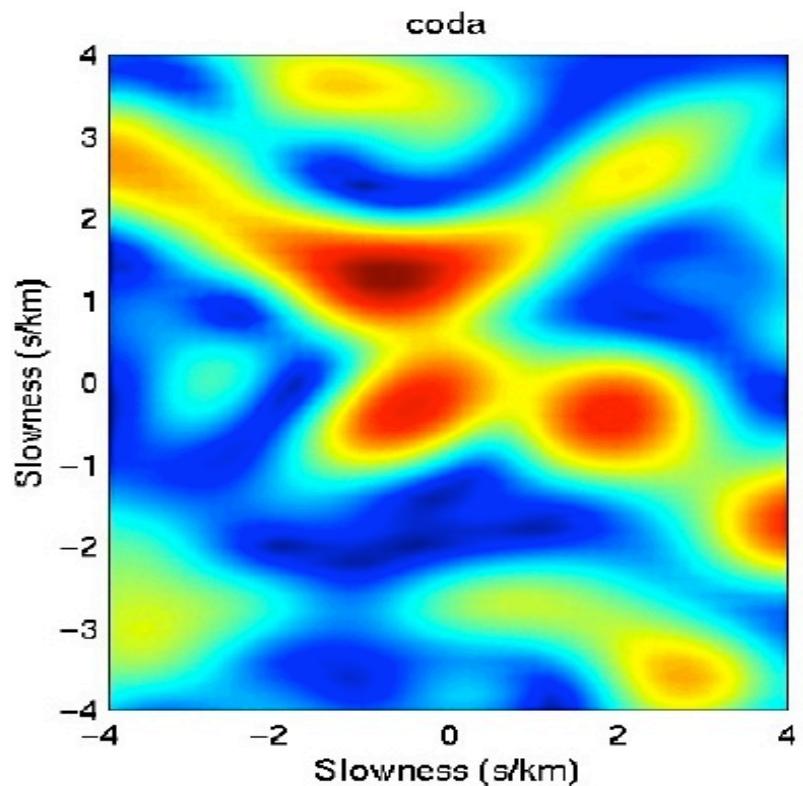
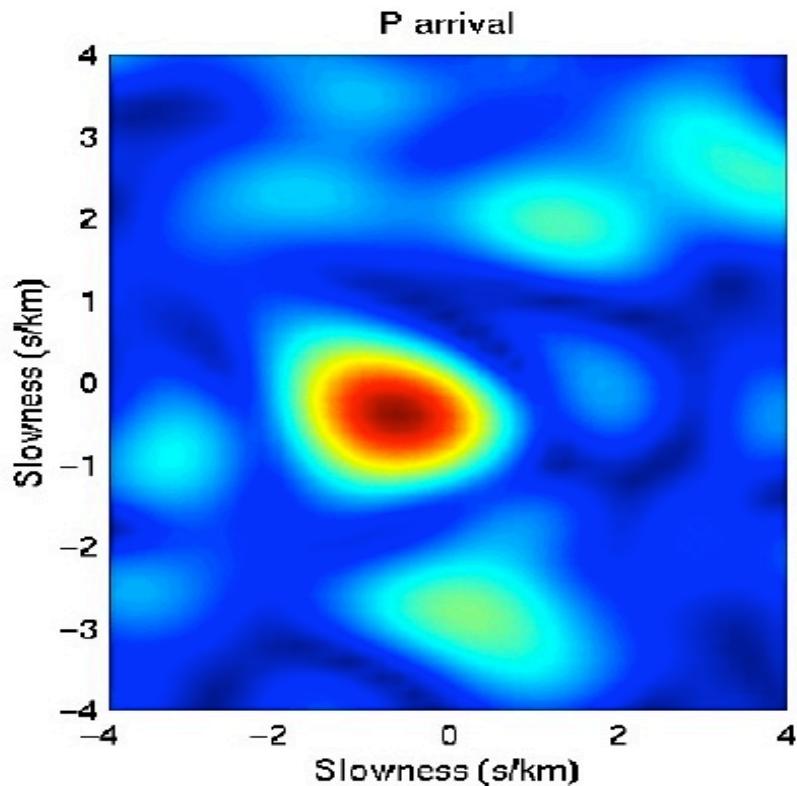
scale x 1000:



# Frequency-wavenumber analysis

(Pinon Flat Seismometer Array)

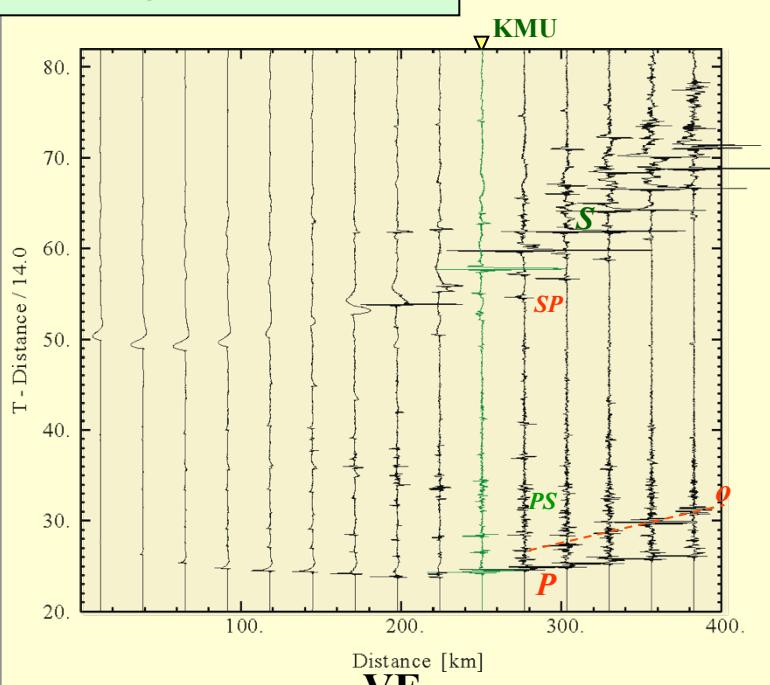
$$u(x,y) \rightarrow u(k_x, k_y)$$



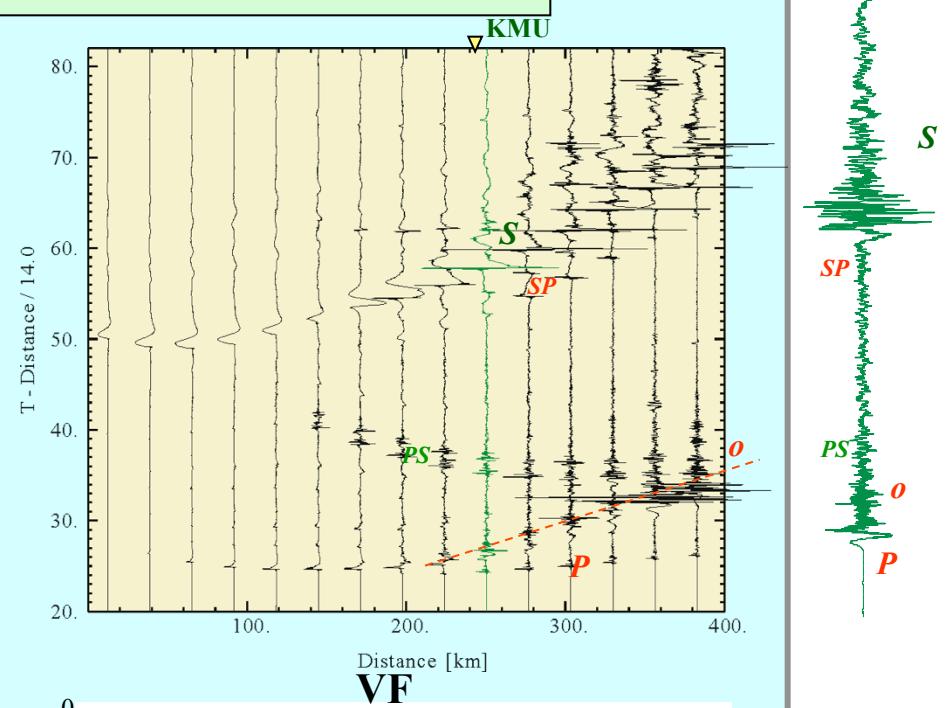
# Computer Simulation (1): Homogeneous Plate (*iasp81*)

Observation  
(KMU)

(a) Homogeneous Plate



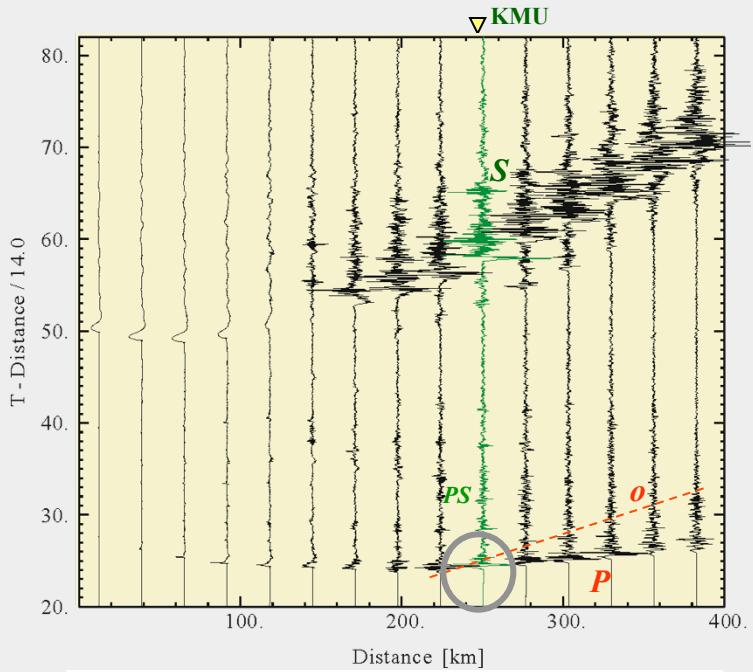
(b) Extended Oceanic Crust



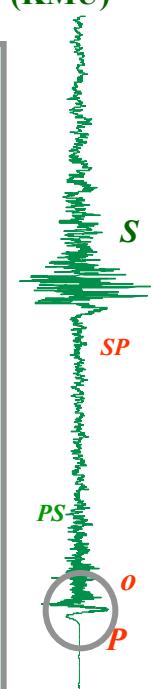
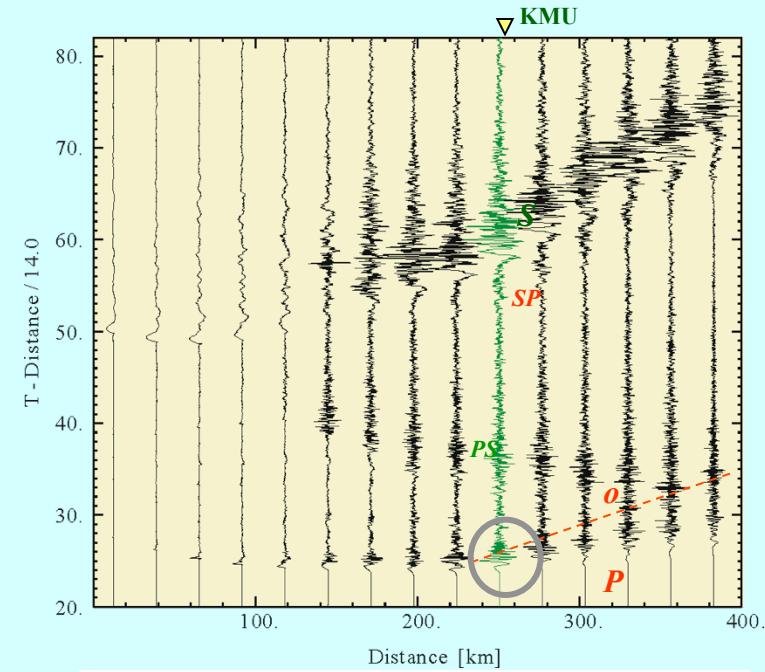
# Heterogeneous Plate Model

Observation  
(KMU)

(c) Isotropic Scatter



(d) Anisotropic Scatter



## Propagation in a weakly heterogeneous medium (*Ref. Aki et Richards 1980*)

Reference medium  $(\lambda_0, \mu_0, \rho_0)$

$\vec{u}$  displacement

$$\rho_0 \ddot{u}_i = (\lambda_0 \vec{\nabla} \cdot \vec{u})_{,i} + [\mu_0 (u_{i,j} + u_{j,i})]_{,j}$$

$$(\text{same as: } \rho_0 \frac{\partial^2 \vec{u}}{\partial t^2} = (\lambda_0 + \mu_0) \vec{\text{grad}} (\text{div } \vec{u}) + \mu_0 \vec{\Delta} \cdot \vec{u})$$

Perturbed medium

$$\rho = \rho_0 + \delta\rho (\vec{r}); \lambda = \lambda_0 + \delta\lambda (\vec{r}); \mu = \mu_0 + \delta\mu (\vec{r})$$

$$\delta\rho, \delta\lambda, \delta\mu \ll \rho, \lambda, \mu$$

Equation of motion:

$$\rho \ddot{u}_i = (\lambda \vec{\nabla} \cdot \vec{u})_{,i} + [\mu (u_{i,j} + u_{j,i})]_{,j}$$

$$(\rho_0 + \delta\rho) \ddot{u}_i = ((\lambda_0 + \delta\lambda) \vec{\nabla} \cdot \vec{u})_{,i} + [(\mu_0 + \delta\mu)(u_{i,j} + u_{j,i})]_{,j}$$

$$\Rightarrow \rho_0 \ddot{u}_i - \lambda_0 (\vec{\nabla} \cdot \vec{u})_{,i} - \mu_0 (u_{i,j} + u_{j,i})_{,j} = -\delta\rho \ddot{u}_i + \delta\lambda (\vec{\nabla} \cdot \vec{u})_{,i}$$

$$+ \delta\lambda_{,i} \vec{\nabla} \cdot \vec{u} + \delta\mu (u_{i,j} + u_{j,i})_{,j} + (\delta\mu)_{,j} (u_{i,j} + u_{j,i})$$

$$\text{with } (u_{i,j} + u_{j,i})_{,j} = \nabla^2 u_i + (\vec{\nabla} \cdot \vec{u})_i$$

$$\begin{aligned} \Rightarrow \rho_0 \ddot{u}_i - (\lambda_0 + \mu_0) (\vec{\nabla} \cdot \vec{u})_{,i} - \mu_0 \nabla^2 u_i &= -\delta\rho \ddot{u}_i + (\delta\lambda + \delta\mu) (\vec{\nabla} \cdot \vec{u})_{,i} \\ &+ \delta\mu \nabla^2 u_i + (\delta\lambda)_i \vec{\nabla} \cdot \vec{u} + (\delta\mu)_{,j} (u_{i,j} + u_{j,i}) \end{aligned}$$

$$\Rightarrow \rho_0 \ddot{u}_i - (\lambda_0 + \mu_0)(\vec{\nabla} \cdot \vec{u}),_i - \mu_0 \nabla^2 u_i = -\delta\rho \ddot{u}_i + (\delta\lambda + \delta\mu)(\vec{\nabla} \cdot \vec{u}),_i \\ + \delta\mu \nabla^2 u_i + (\delta\lambda)_i \vec{\nabla} \cdot \vec{u} + (\delta\mu)_{,j} (u_{i,j} + u_{j,i})$$

$$u = u^0 + u^d$$

→  $u^0$  satisfies elastodynamic equation for  $(\rho_0, \lambda_0, \mu_0)$

Hypothesis: weak perturbations

→ neglect terms like  $\delta\mu \times (u^d)'$  ( $\delta\mu \ll \mu_0$  ;  $u^d \ll u_0$ )

(First order Born approximation)

$$\rho_0 \ddot{u}_i^d - (\lambda_0 + \mu_0)(\nabla u^d),_i - \mu_0 \nabla^2 u_i^d = Q_i$$

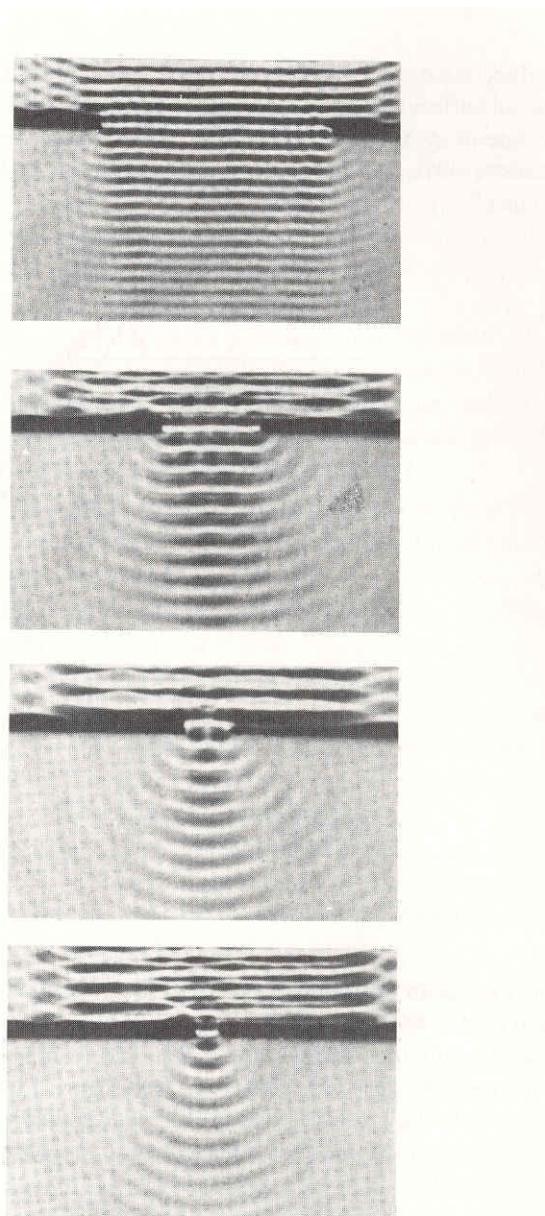
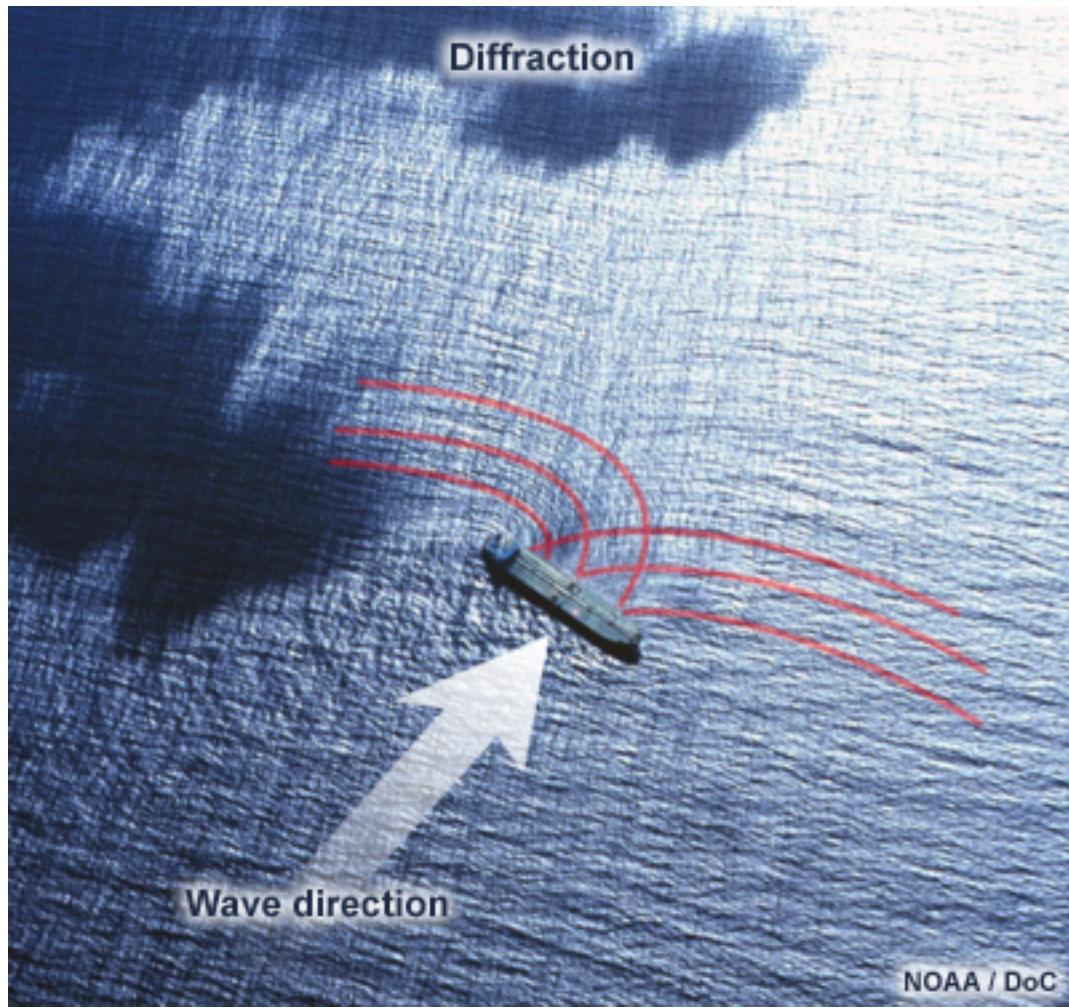
with :

$$Q_i = -\delta\rho \ddot{u}_i^o + (\delta\lambda + \delta\mu)(\vec{\nabla} \cdot \vec{u}^o),_i + \delta\mu \nabla^2 u_i^o + (\delta\lambda)_i \vec{\nabla} \cdot \vec{u}^o + (\delta\mu)_{,j} (u_{i,j}^o + u_{j,i}^o)$$

⇒ extra source terms in the reference model

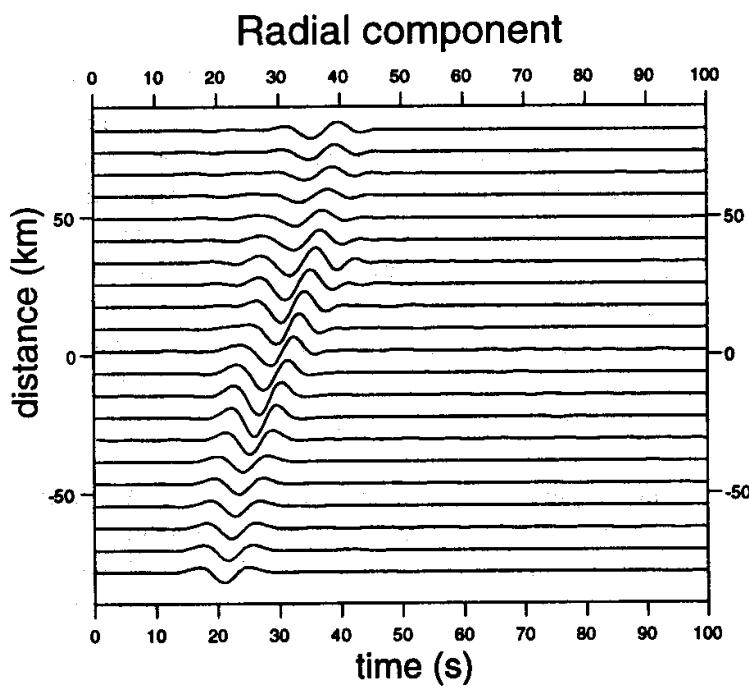
→ diffraction = virtual sources

→ base formula for linearized inverse problem

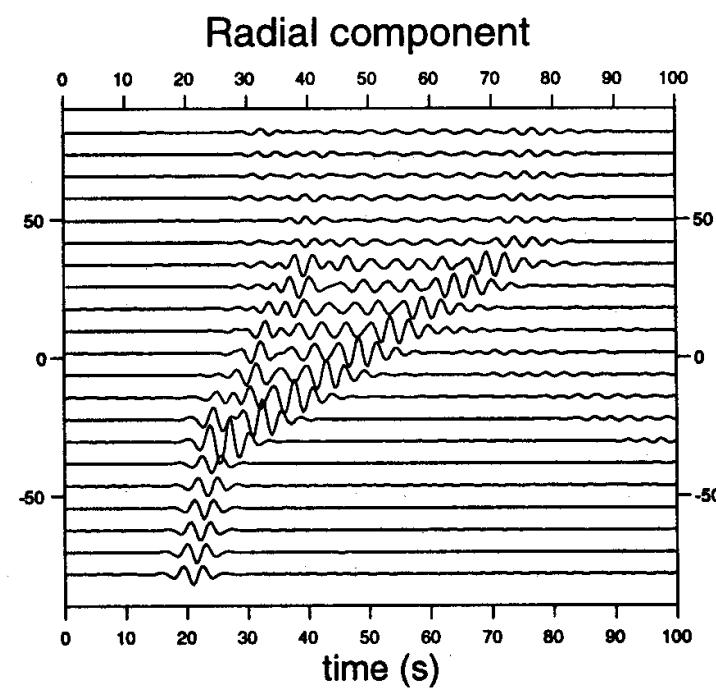


**Figure 8.1** Ripples diffracted into quiet water by an opening.

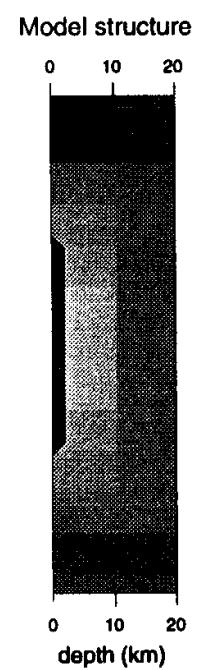
## Diffraction of Rayleigh waves by a basin



<0.15 Hz



0.2-1 Hz



## DIFFRACTION OF P WAVES :

Primary wave propagating in direction  $x_1$  :

$$u_i^o = \delta_{1i} \exp\left(-i \omega (t - \frac{x_1}{\alpha_0})\right)$$

$$\alpha_0 = \left( \frac{\lambda_0 + 2\mu_0}{\rho_0} \right)^{1/2}$$

$\Rightarrow$

$$Q_1 = \left\{ \delta\rho \omega^2 - \frac{(\delta\lambda + 2\delta\mu)\omega^2}{\alpha_0^2} + i \frac{\omega}{\alpha_0} (\delta\lambda)_{,1} + 2i \frac{\omega}{\alpha_0} (\delta\mu)_{,1} \right\} \exp\left(-i\omega(t - \frac{x_1}{\alpha_0})\right)$$

$$Q_2 = i \frac{\omega}{\alpha_0} (\delta\lambda)_{,2} \exp\left(-i\omega(t - \frac{x_1}{\alpha_0})\right)$$

$$Q_3 = i \frac{\omega}{\alpha_0} (\delta\lambda)_{,3} \exp\left(-i\omega(t - \frac{x_1}{\alpha_0})\right)$$

Identification :

- velocity fluctuation :

$$\alpha = \frac{(\lambda + 2\mu)^{1/2}}{\rho^{1/2}} \quad \Rightarrow \delta\alpha = \frac{\partial\alpha}{\partial\lambda}\delta\lambda + \frac{\partial\alpha}{\partial\mu}\delta\mu + \frac{\partial\alpha}{\partial\rho}\delta\rho$$

$$\delta\alpha = \frac{1}{2} \frac{1}{\rho} \frac{1}{\alpha} \delta\lambda + \frac{1}{2} \frac{2}{\rho} \frac{1}{\alpha} \delta\mu - \frac{(\lambda + 2\mu)}{\rho^2} \frac{1}{2} \frac{1}{\alpha} \delta\rho$$

$$\frac{\delta\alpha}{\alpha} = \frac{1}{2} \left( \frac{\delta\lambda + 2\delta\mu}{\lambda + 2\mu} - \frac{\delta\rho}{\rho} \right)$$

→ For  $Q_1$  :

$$\delta \rho \omega^2 - \frac{(\delta\lambda + 2\delta\mu)}{\alpha_0} \omega^2 = -\omega^2 \rho_0 \left( 2 \frac{\delta\alpha}{\delta_0} \right)$$

→ force  $q_1$  proportional to the velocity perturbation:

$$q_1 = -2 \omega^2 \rho_0 \frac{\delta\alpha}{\alpha_0} \exp(-\omega(t - \frac{x_1}{\alpha_0}))$$

(simple force dans la direction de propagation)

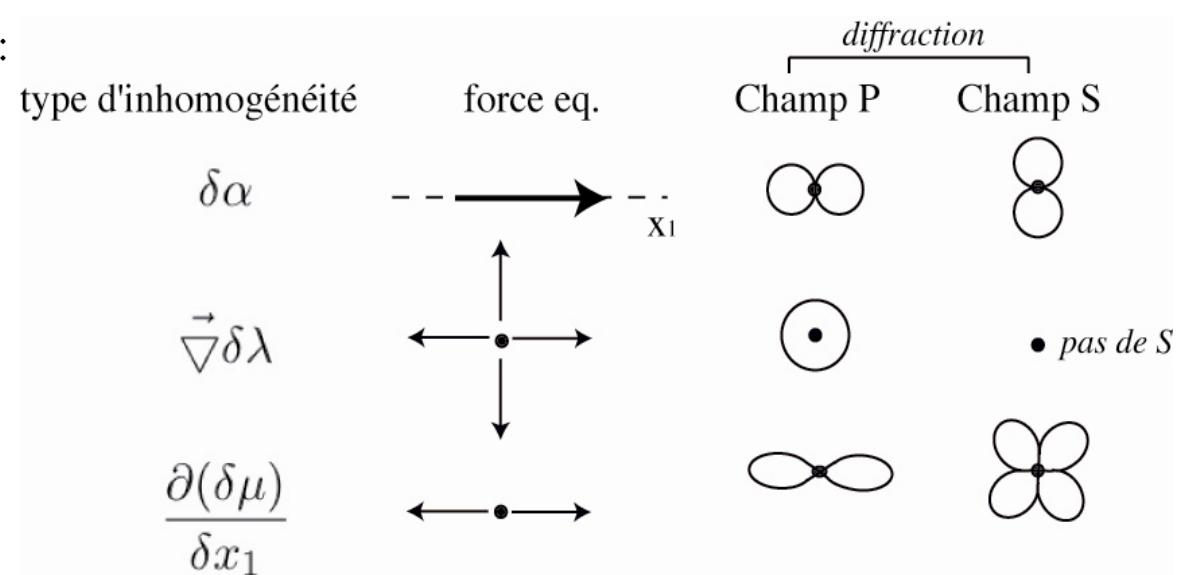
- term proportional to the gradient of  $\lambda$

$$q' = i \frac{\omega}{\alpha_0} \vec{\nabla} \cdot \delta \lambda$$

- terme proportional to the spatial variation of  $\mu$

$$q''_1 = 2i \frac{\omega}{\alpha_0} (\delta\mu)_{,1}$$

For a point perturbation :



- No scattered S wave in the forward direction
- scattered P wave maximum along  $x_1 \rightarrow$  backscattering

DIFFRACTION OF S WAVE:  $u_i^o = \delta_{2i} \exp\left(-i\omega(t - \frac{x_1}{\beta_o})\right)$

$\Rightarrow$

- term proportional to the velocity perturbation:

$$q_2 = -2\omega^2 \rho_o \frac{\delta\beta}{\beta_o} \exp\left(-i\omega(t - \frac{x_1}{\beta})\right)$$

- term proportional to the variation of  $\delta\mu$ :

$$q'_1 = i \frac{\omega}{\beta_o} (\delta\mu)_{,2} \exp\left(-i\omega(t - \frac{x_1}{\beta})\right)$$

$$q'_2 = i \frac{\omega}{\beta_o} (\delta\mu)_{,1} \exp\left(-i\omega(t - \frac{x_1}{\beta})\right)$$

Point perturbation:

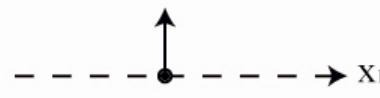
Type d'inhomogénéité

force

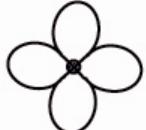
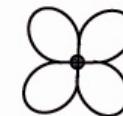
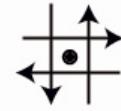
onde P

onde S

$\delta\beta$



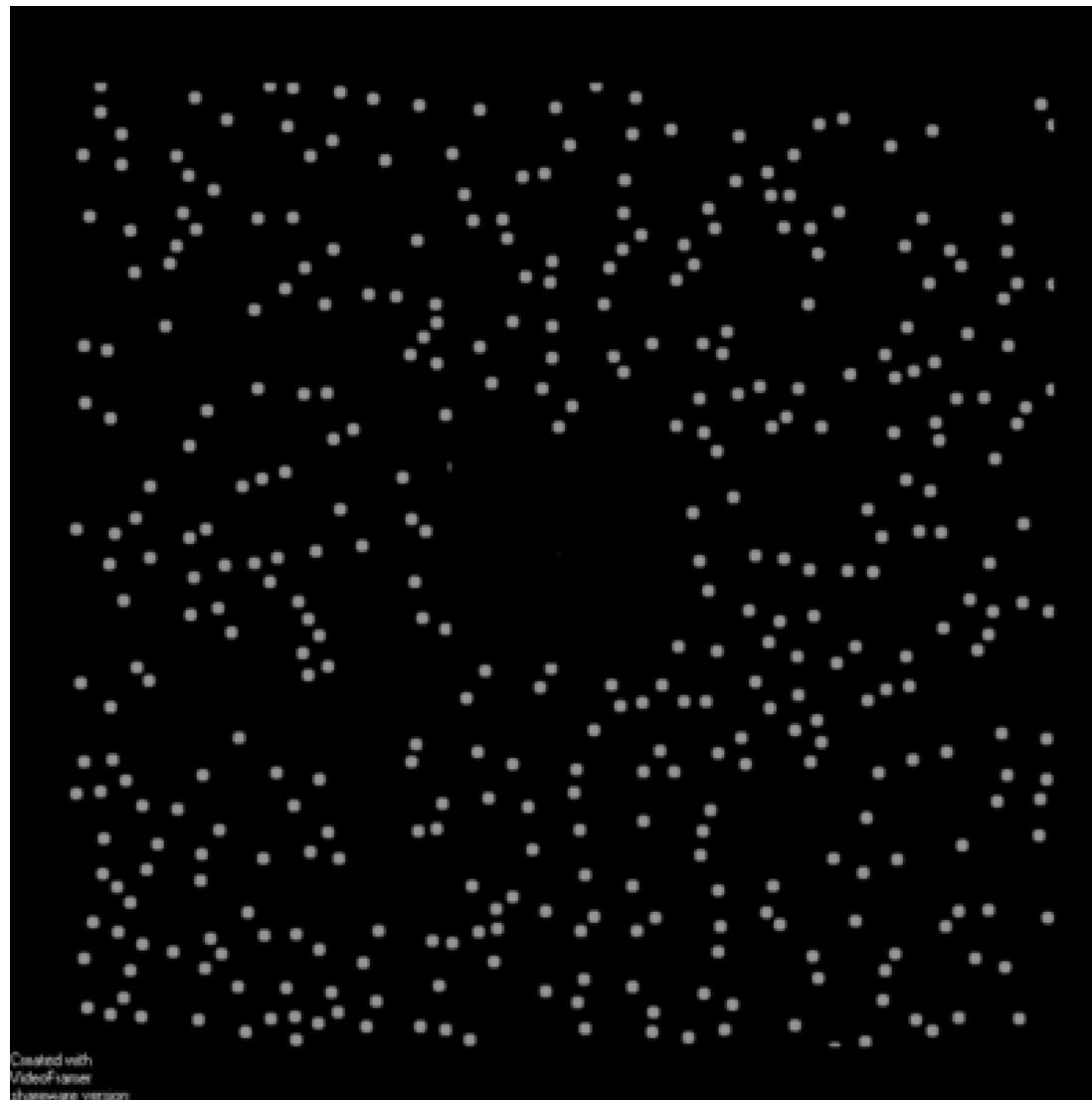
$\frac{\partial(\delta\mu)}{\partial x_1}; \frac{\partial(\delta\mu)}{\partial x_2}$



- No scattered P along  $x_1$

- P to S coupling via multiple scattering

Transient signals in a complex medium.....



Created with  
VideoFramer  
shareware version

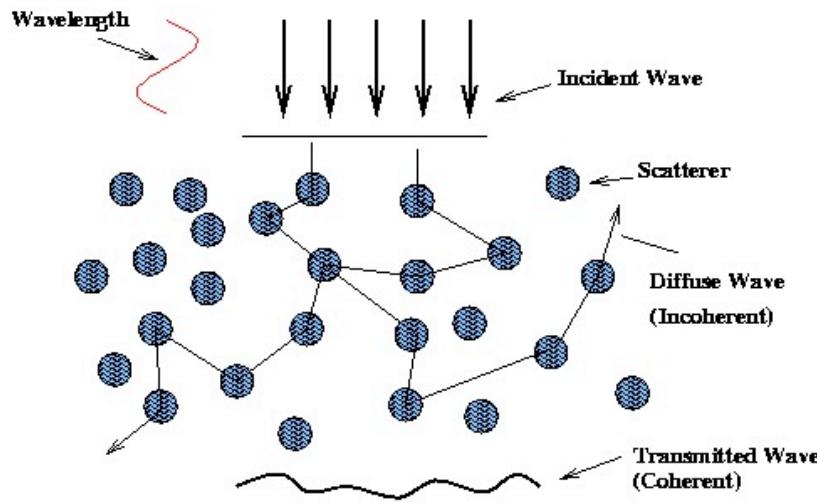
What is the regime relevant to coda waves:

single scattering?

multiple scattering?

diffusion?

## Wave Propagation through Random Media



Length Scales:  $\lambda$ , Correlation Length, Propagation Distance

Question: Ensemble Average Response?

Precise Definition of Coherent and Incoherent Waves

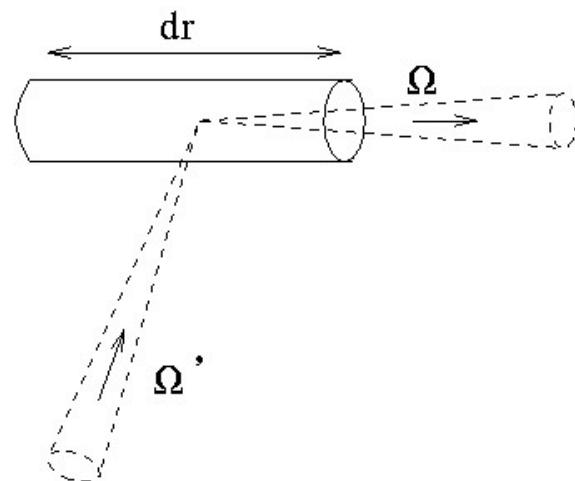
quality factor:

$$Q = \frac{2\pi}{\lambda} l$$

# Radiative transfert equation

## Heuristic View

Energy balance of a beam of energy propagating a distance  $dr$  in the scattering medium



Variation of Intensity

=

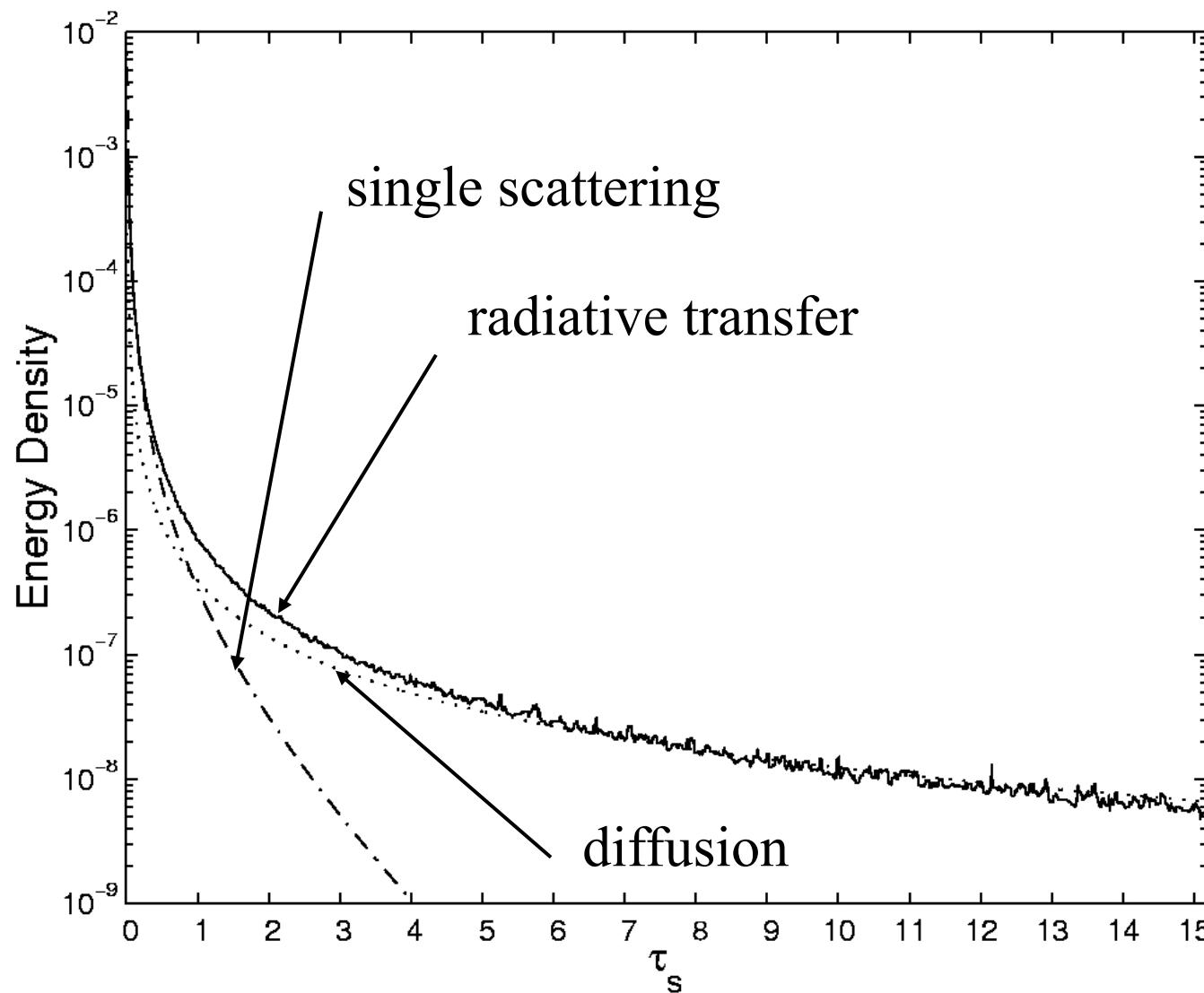
Loss due to scattering into all space  
directions

+

Gain due to scattering from direction  
 $\vec{\Omega}'$  to direction  $\vec{\Omega}$

Elasticity: polarisation + P/S coupling

# The approximations of the Radiative Transfer Equation (RTE)



With the diffusion approximation

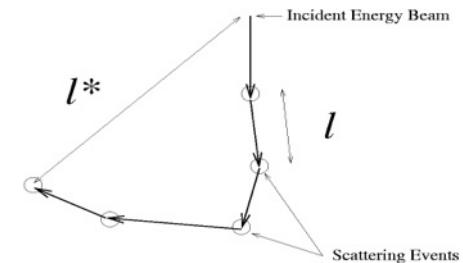
$$\partial_t \rho(\mathbf{R}, t) - D \nabla^2 \rho(\mathbf{R}, t) = \delta(t) \delta(\mathbf{R})$$

$\rho$  is the energy density and  $\mathbf{J}$  the energy current vector

$$\mathbf{J}(\mathbf{R}, t) = -D \nabla \rho(\mathbf{R}, t)$$

$D$  is the diffusion constant of the waves and is related to the transport mean free time  $\tau^*$  as follows:

$$D = v^2 \tau^* / 3$$

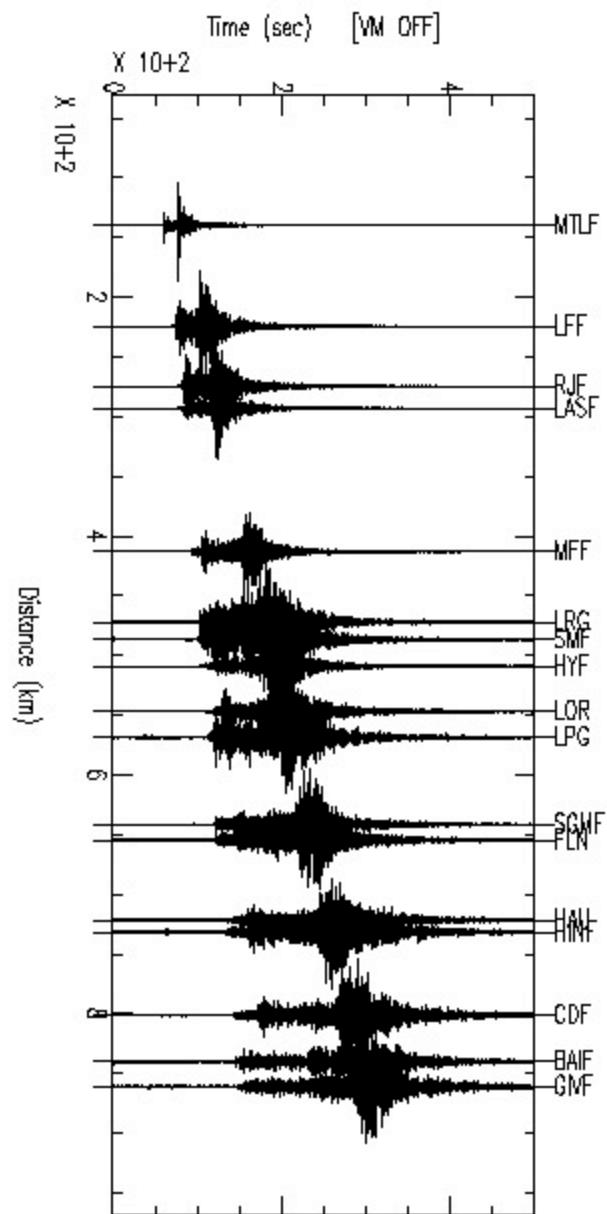
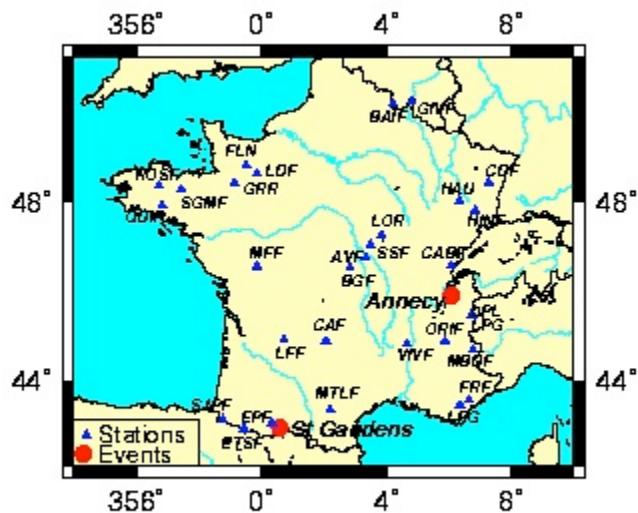


$$\rho(R, t) \approx \exp\left(-\frac{R^2}{4Dt}\right) \frac{1}{t^{3/2}}$$

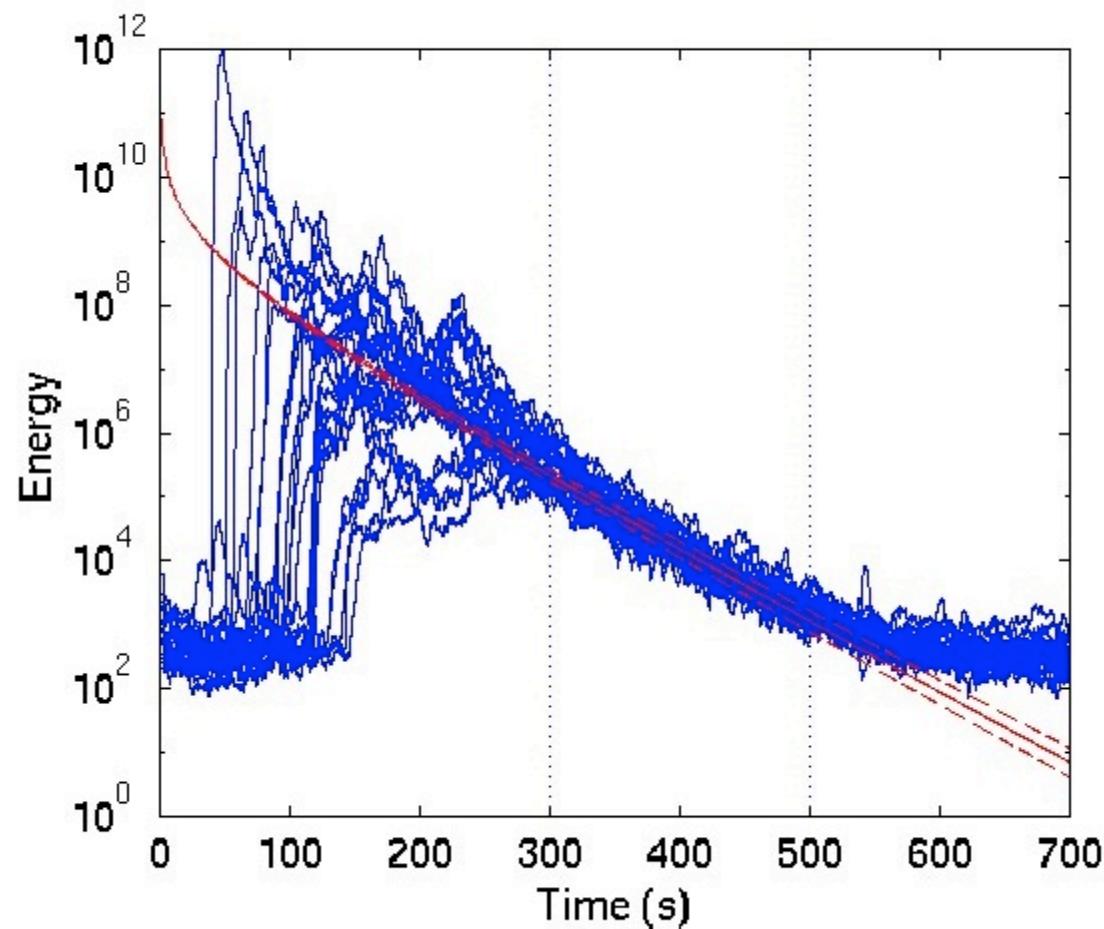
for large lapse times:

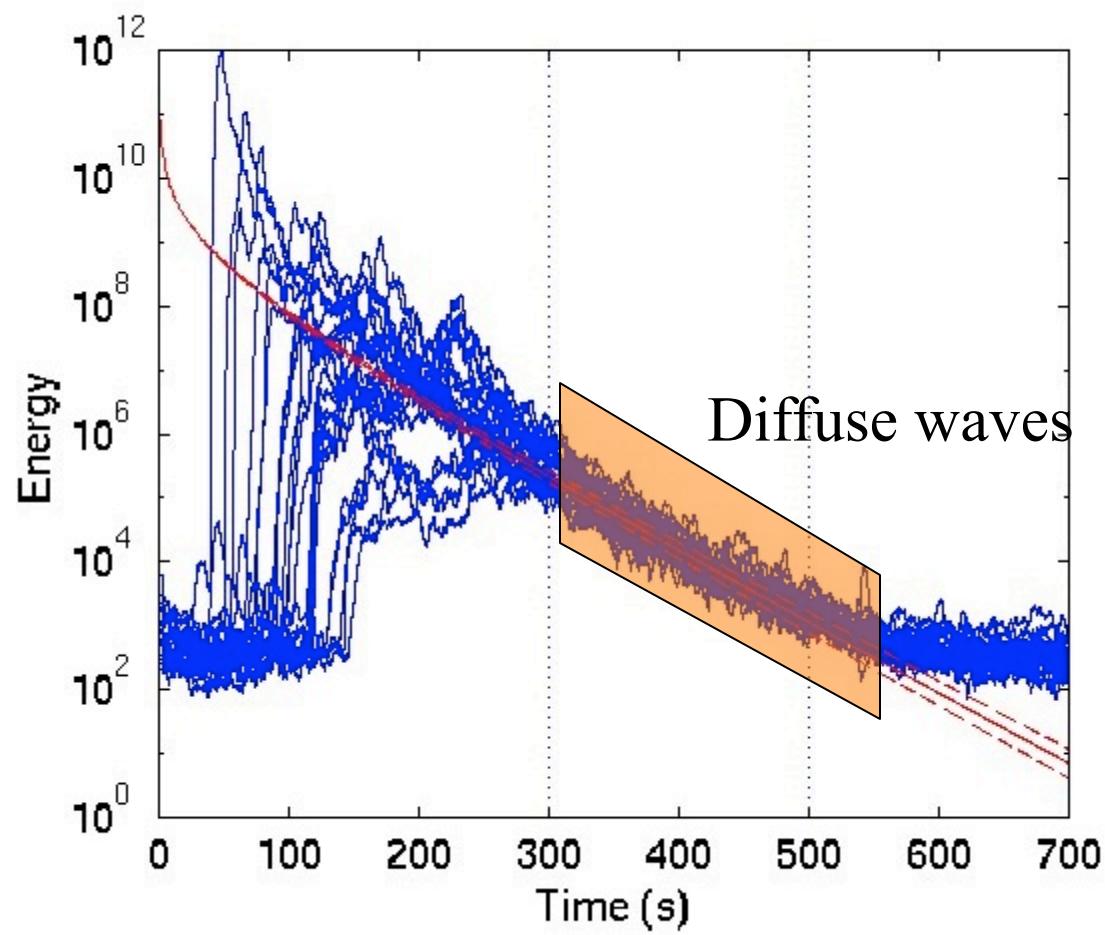
$$\rho(t, \vec{r}) \sim \frac{1}{(Dt)^{3/2}}$$

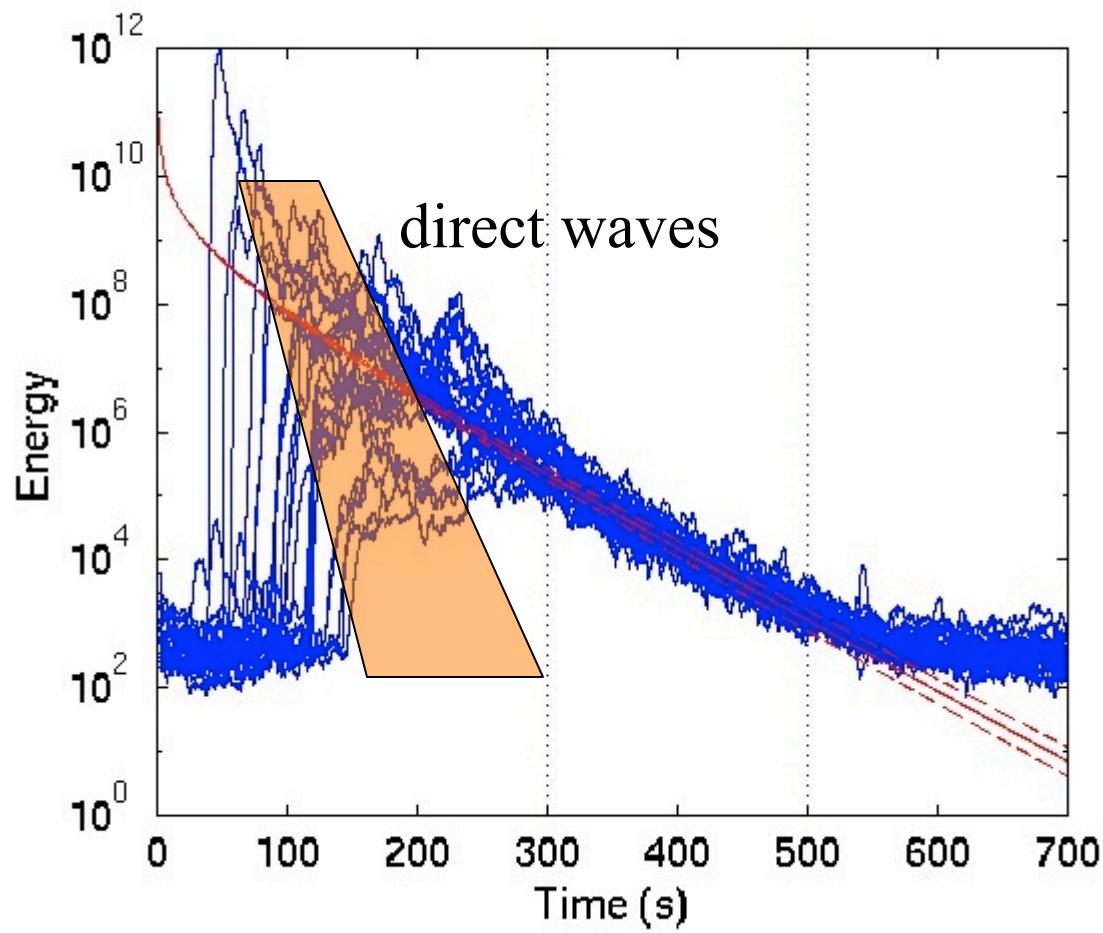
# Coda of regional seismograms and the separation of scattering and absorption



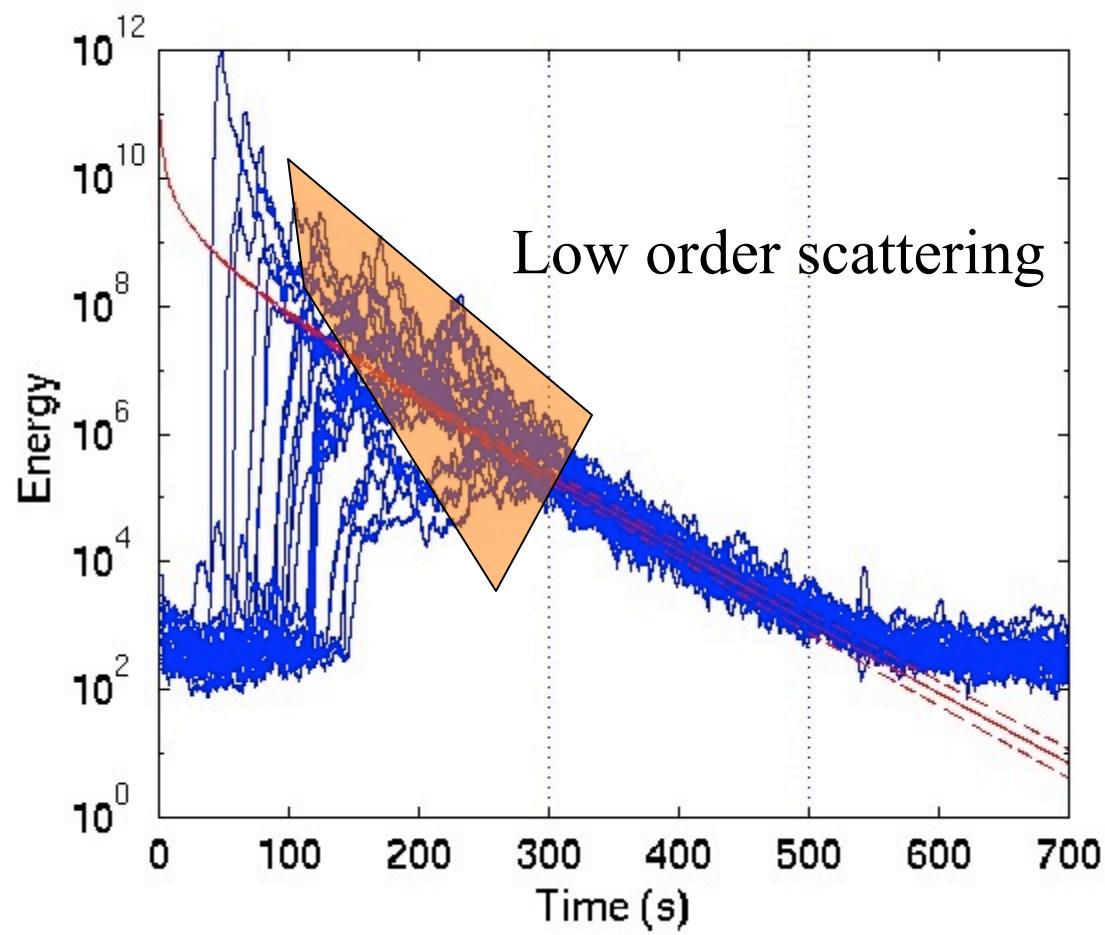
Observations at distances between 150 and 800 km!!



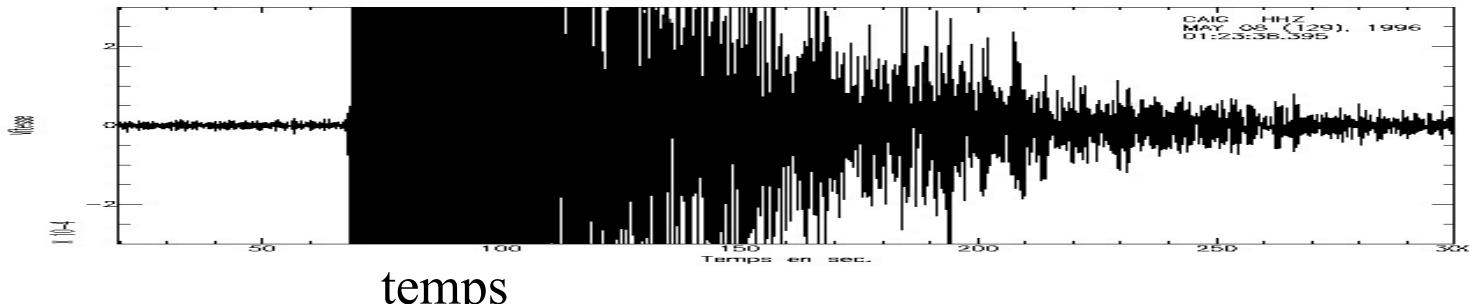
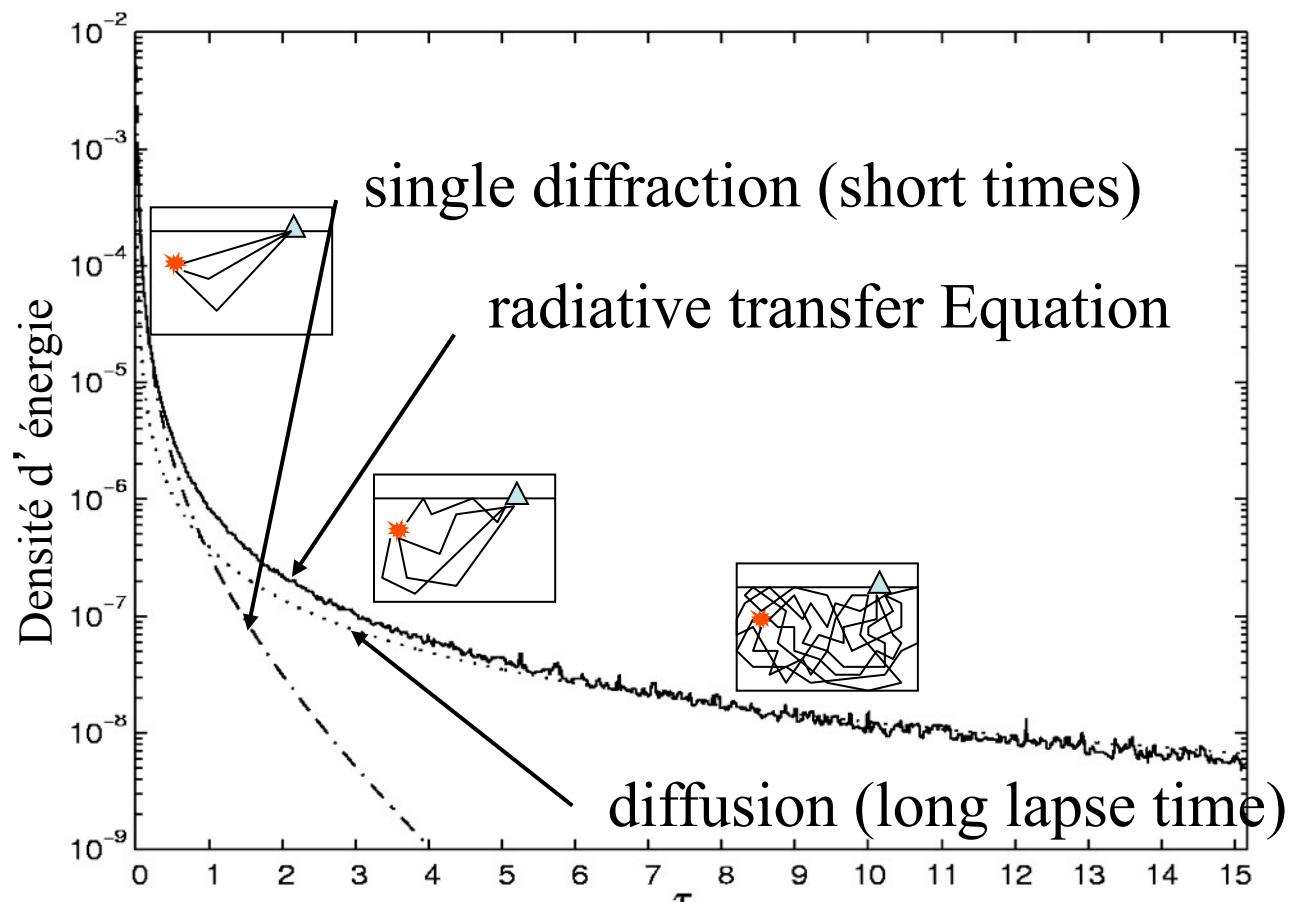




Rapid decay with distance



# Propagation regimes and energy description



## **Propagation Regimes**

**Multiple scattering, diffusion and equipartition**

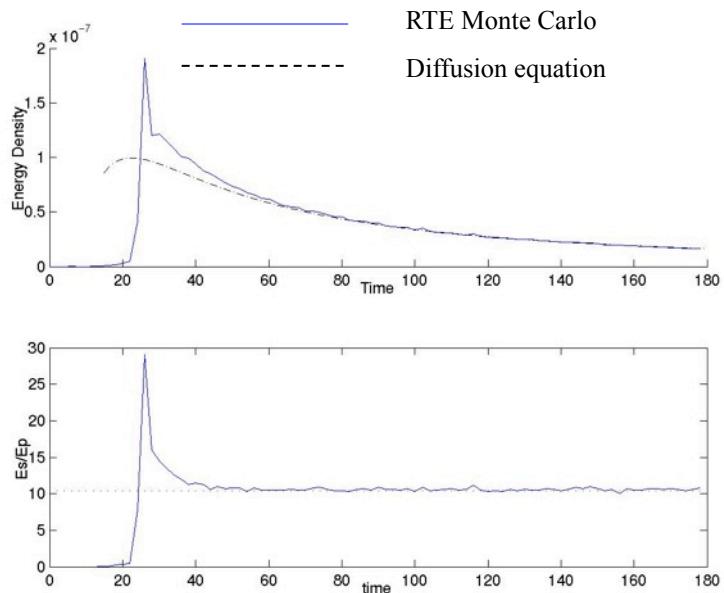
**Diffusion, phase and coherent back scattering**

# Searching for a marker of the regime of scattering...

Equipartition principle for a completely randomized (diffuse) wave-field: in average, all the modes of propagation are excited to equal energy.

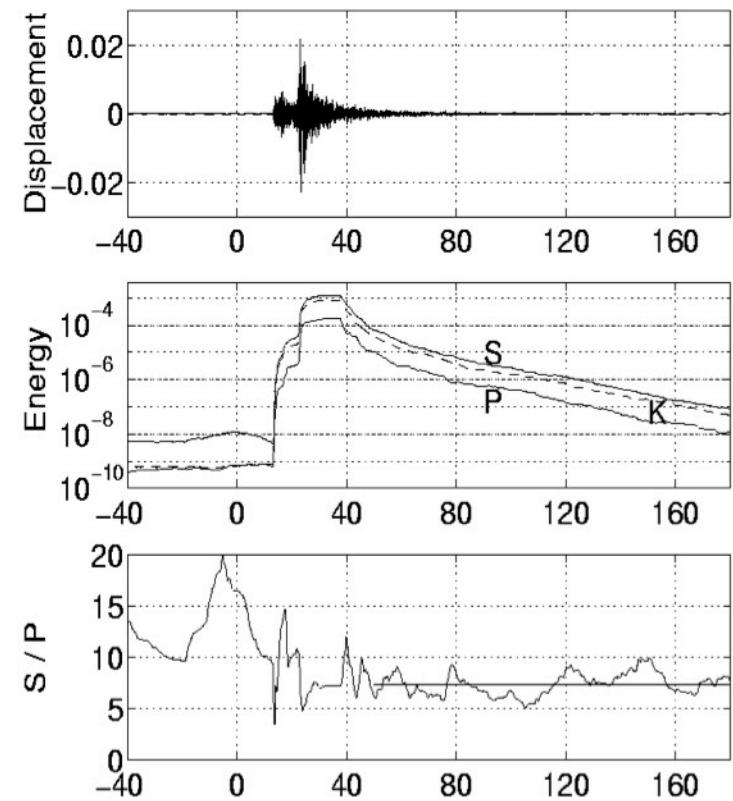
Implication for elastic waves (Weaver, 1982, Ryzhik et al., 1996): P to S energy ratio stabilizes at a value independant of the details of scattering!

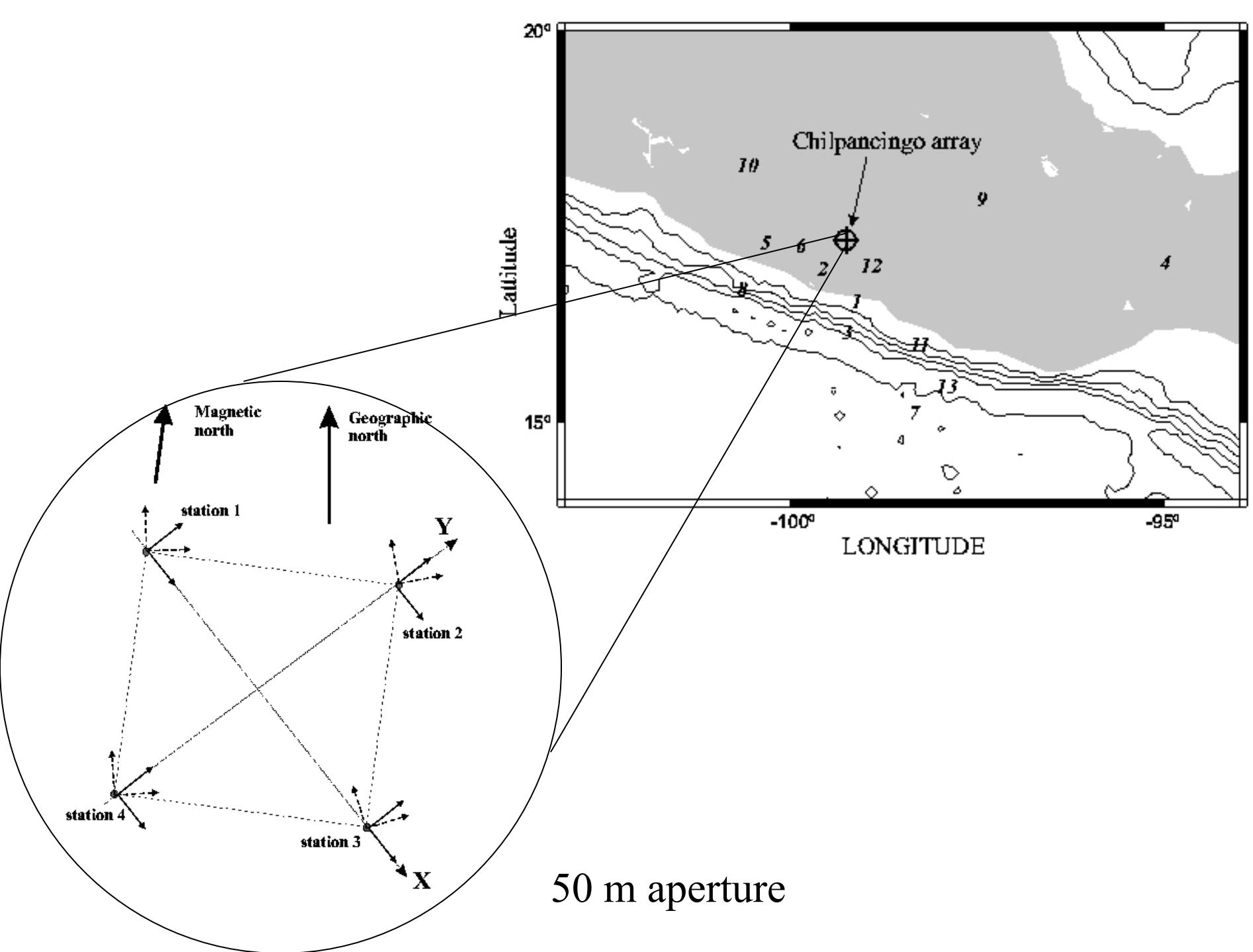
## Numerical simulation



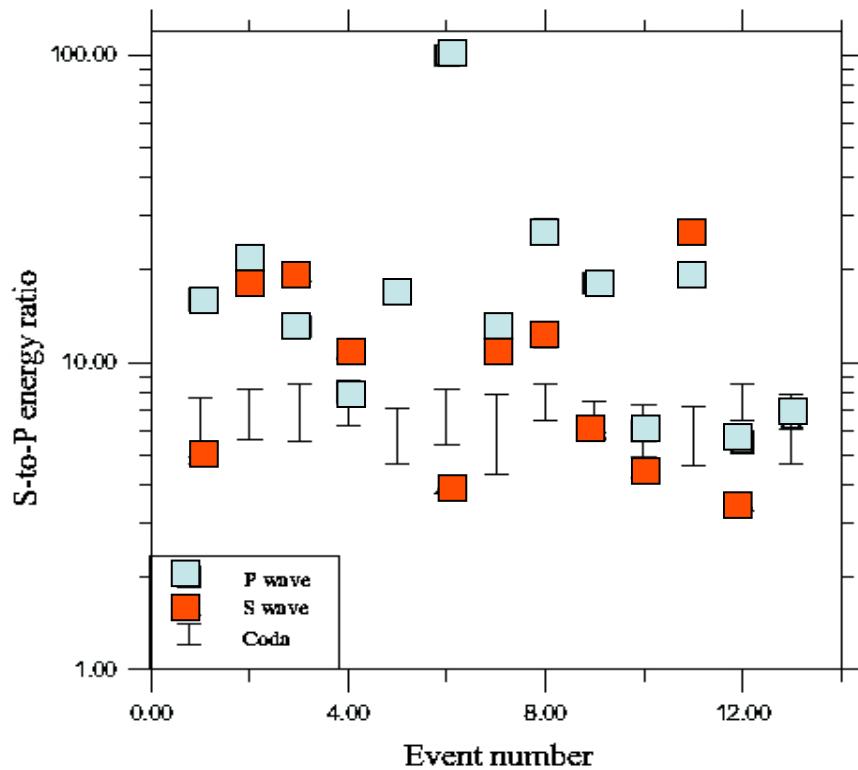
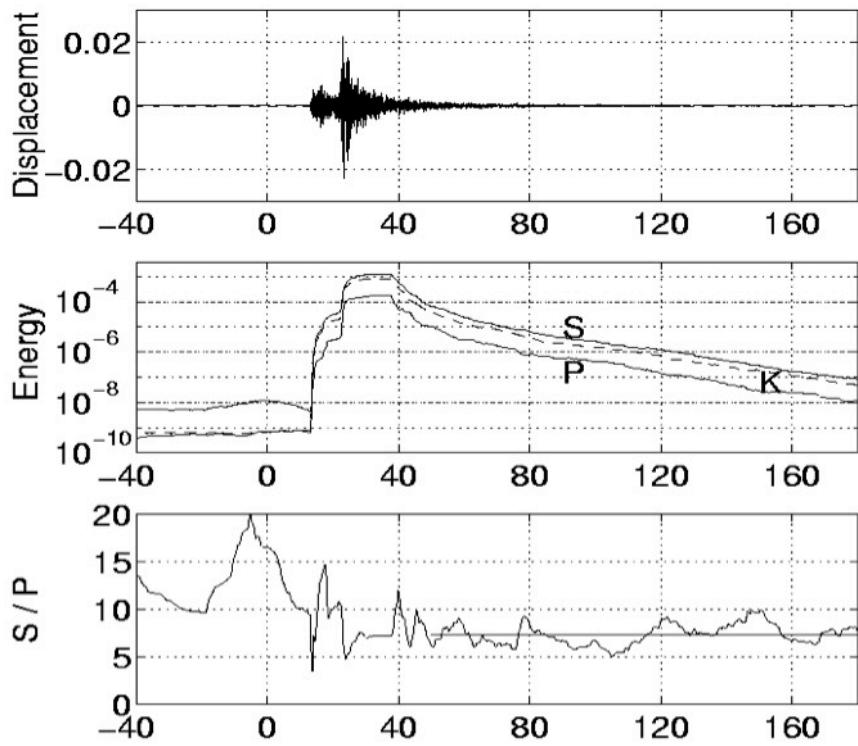
## Observations

Event 11

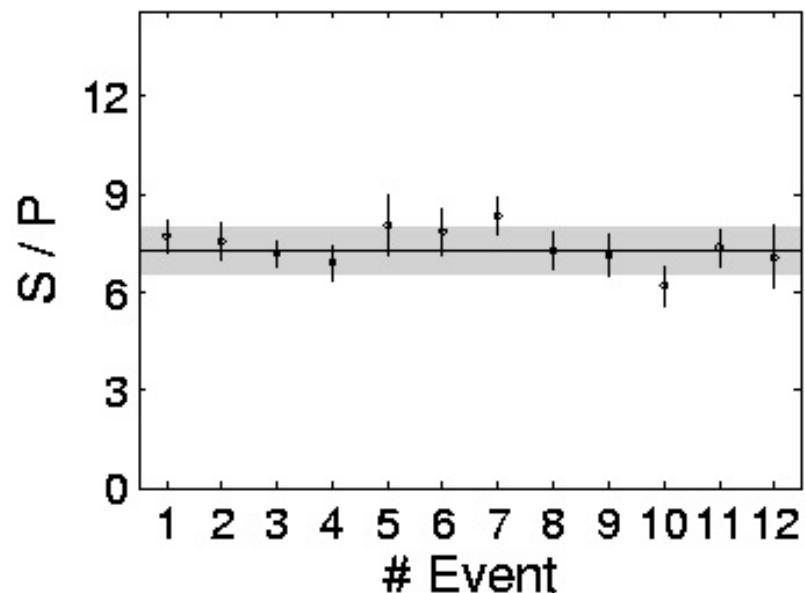




Event 11



Values of stabilization  
for a series of  
events



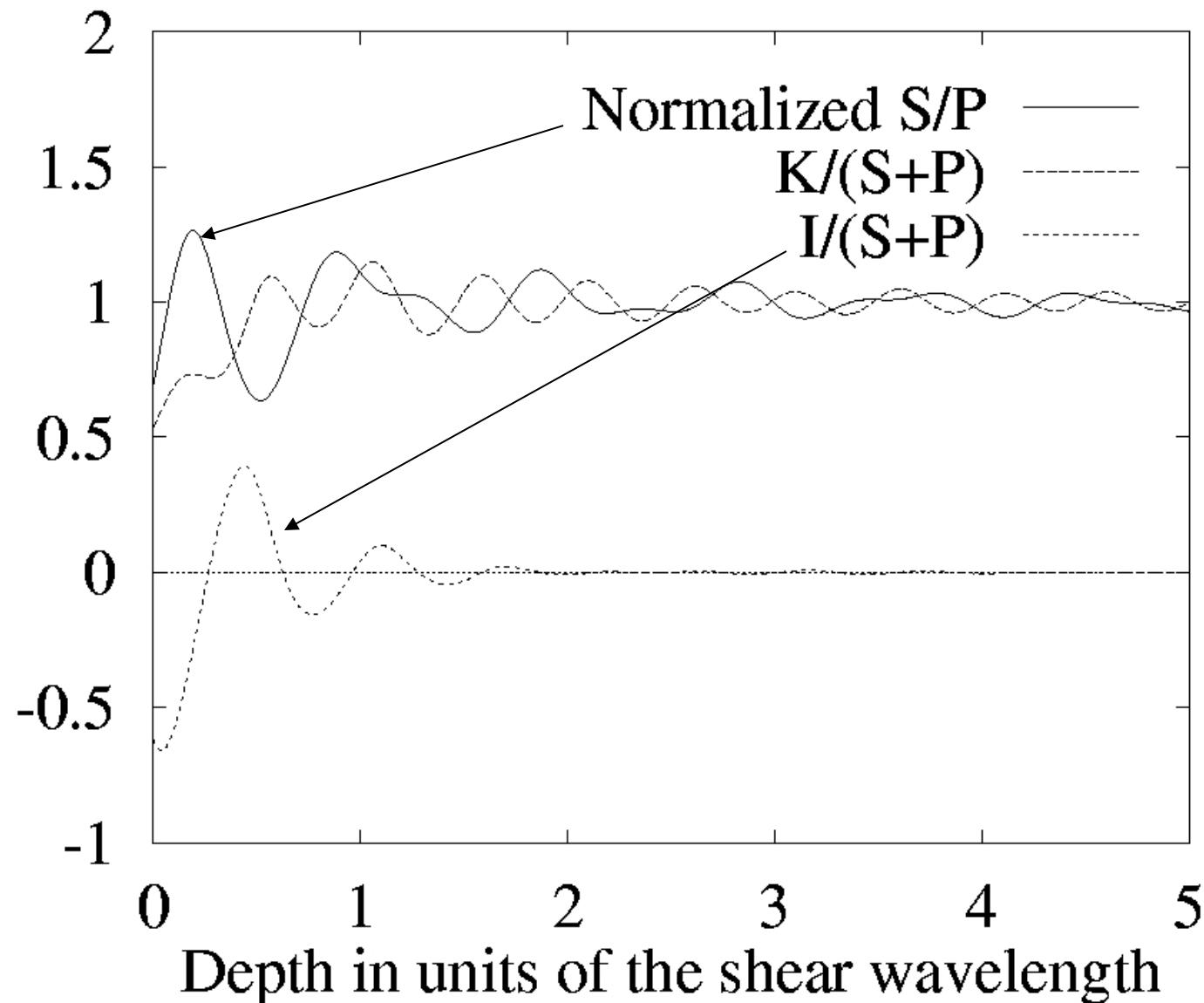
(Shapiro, Campillo, Margerin, Singh, Kostoglodov, Pacheco, BSSA, 2000.)

	OBSERVATIONS	THEORY FULL SPACE	THEORY HALF SPACE (BODY WAVES)	THEORY HALF-SPACE WITH RAYLEIGH
S/P	7.3	10.39	9.76	7.19

(Hennino et al., PRL, 2001)

# Effect of the free surface - A model including Rayleigh waves:

Normal modes of a thick elastic plate





Model 1

# Effect of a ‘thin’ layer: Love and Rayleigh

$$\alpha/\beta = \sqrt{3}$$

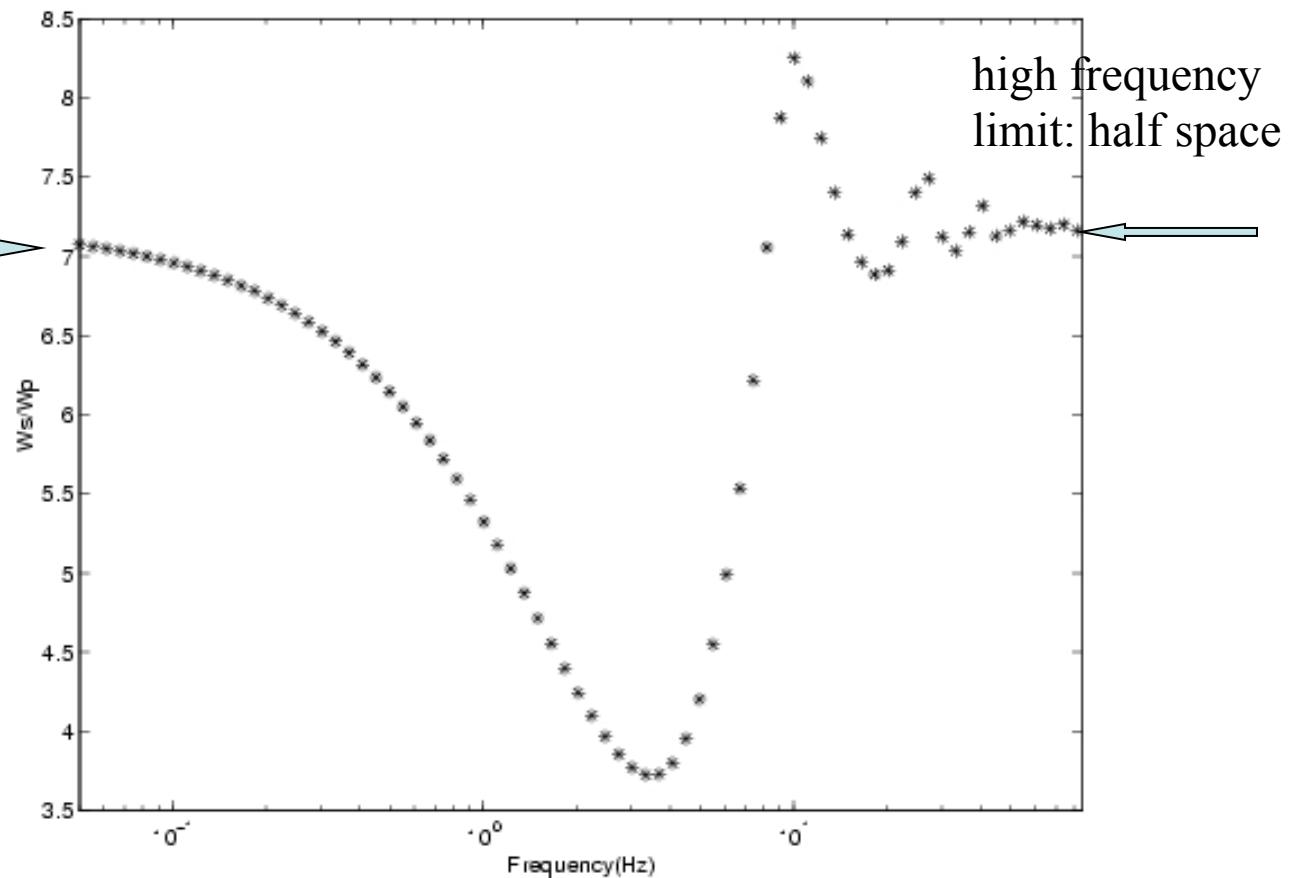
$\alpha = 2.77 \text{ km/s}$   
 $\beta = 1.6 \text{ km/s}$   
 $\rho = 2.7 \text{ km/s}$

65 m

$$\alpha/\beta = \sqrt{3}$$

$\alpha = 5.2 \text{ km/s}$   
 $\beta = 3 \text{ km/s}$   
 $\rho = 2.7 \text{ km/s}$

low frequency  
limit: half space

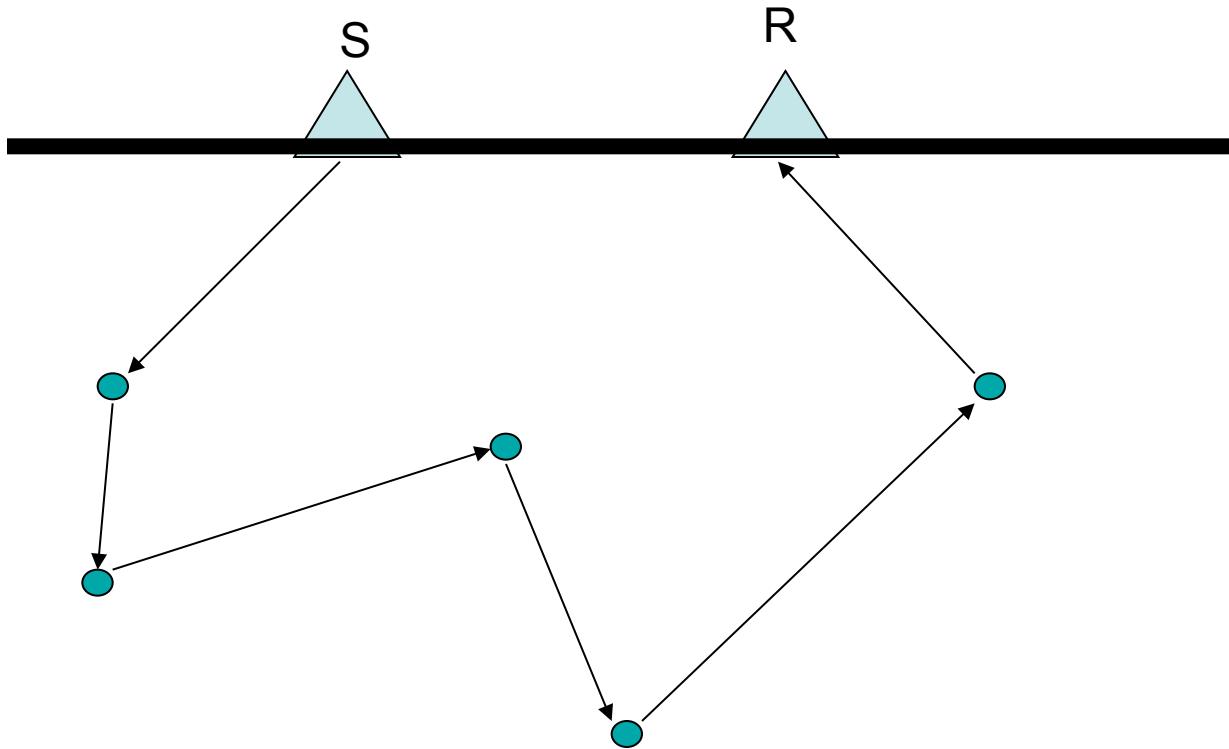
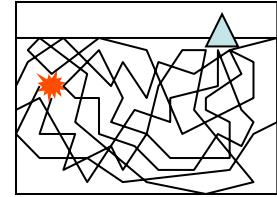


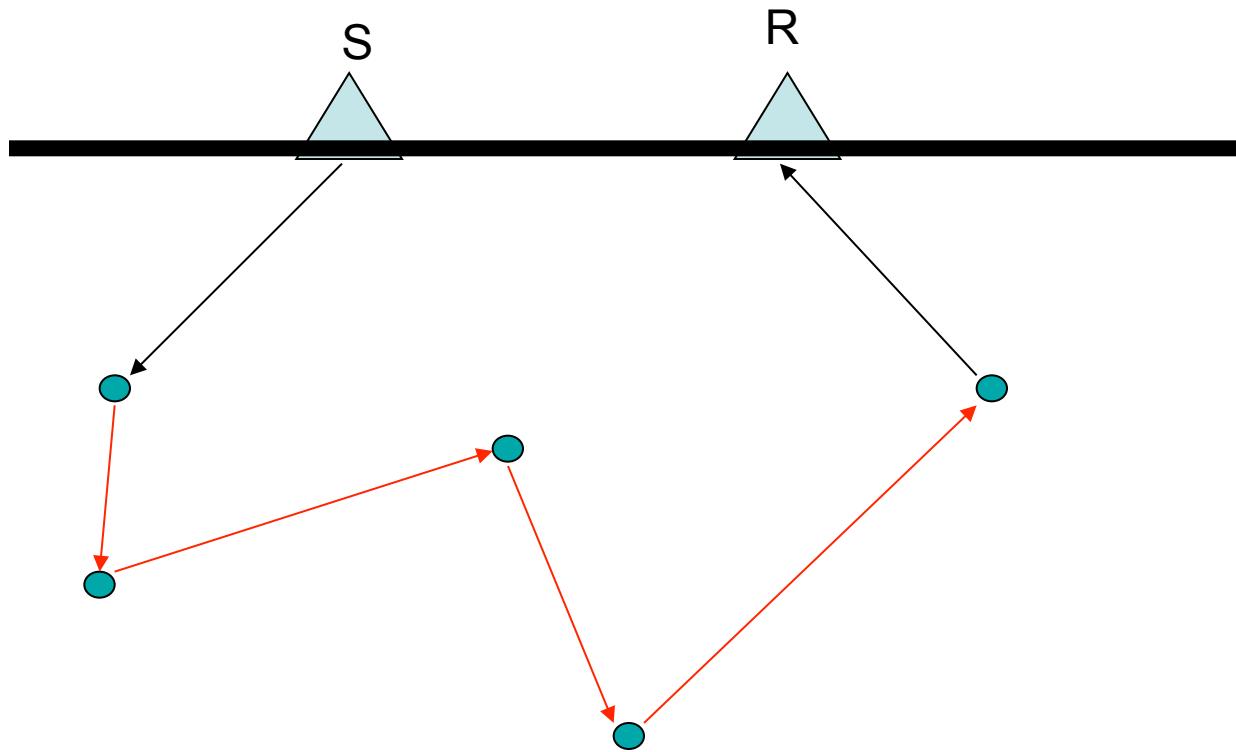
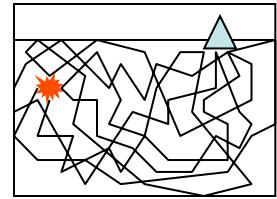
## **Propagation Regimes**

Multiple scattering, diffusion and equipartition

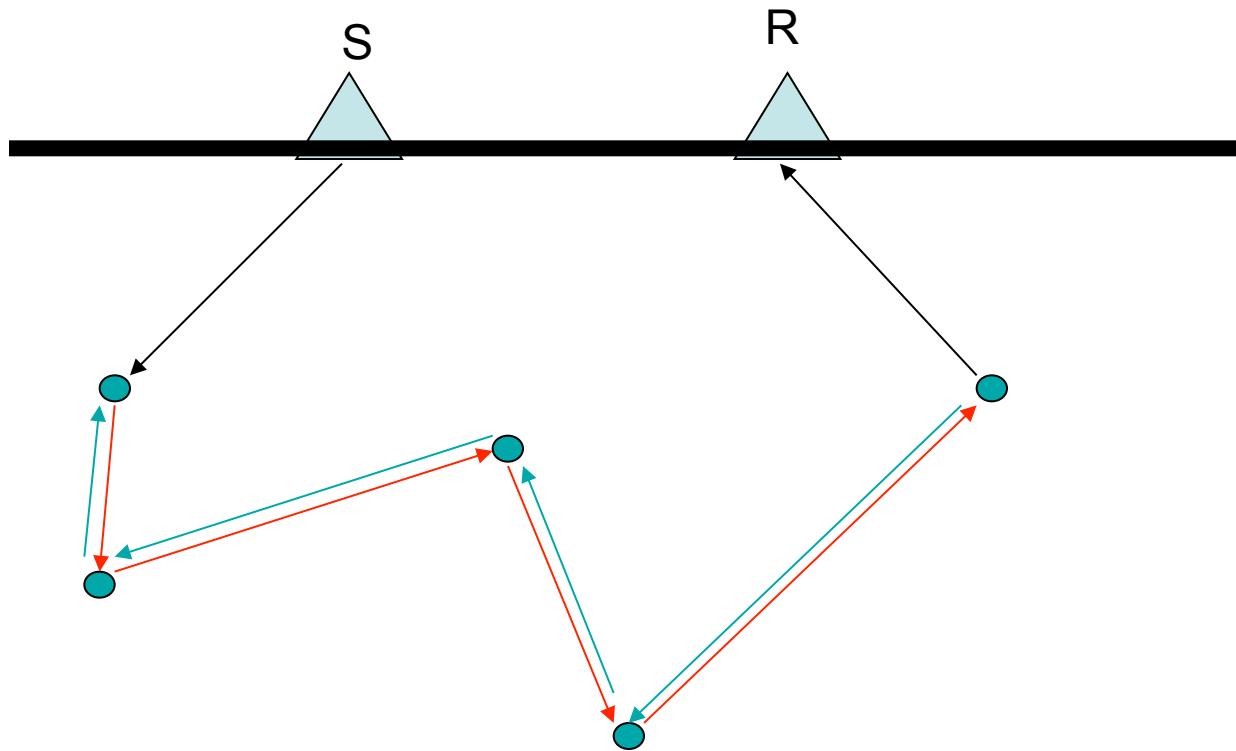
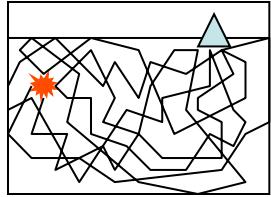
**Diffusion, phase and coherent back scattering**

# Coherent backscattering: multiple scattering

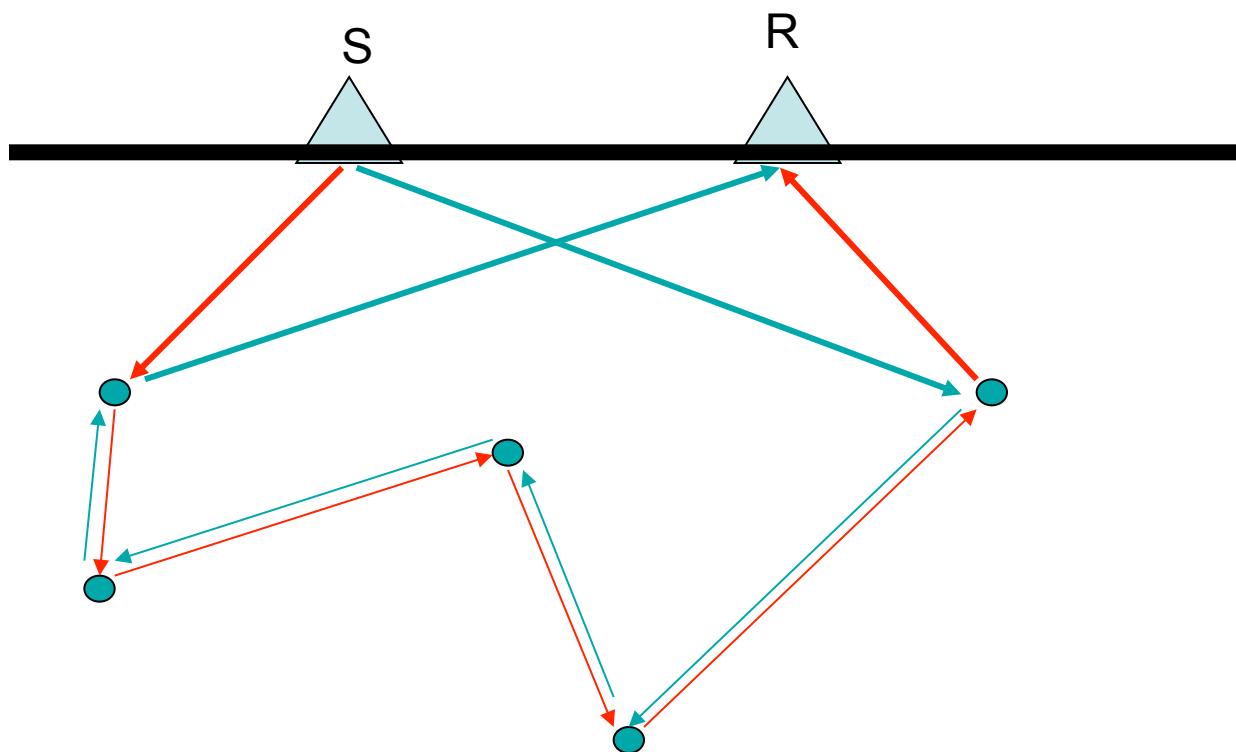
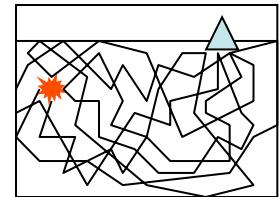




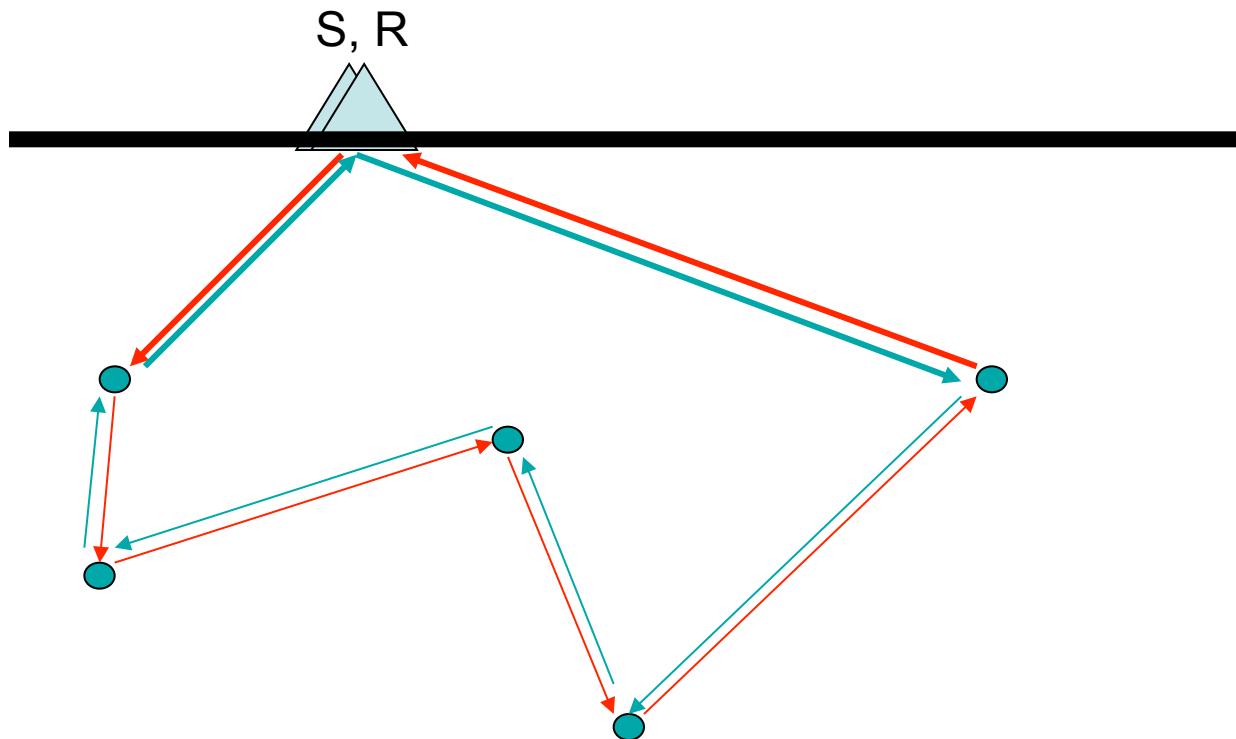
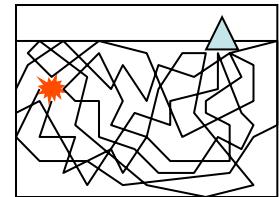
If this path exists..



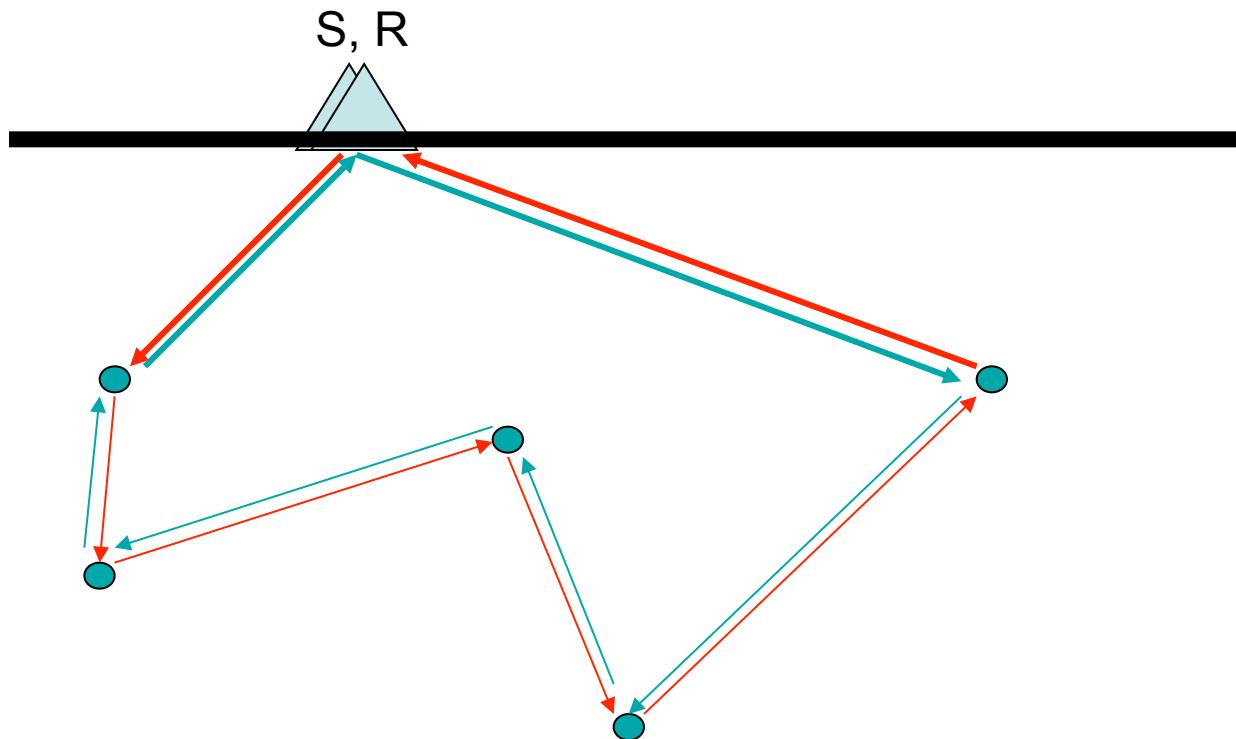
If this path exists, the reciprocal path exists too.



Phase difference: location of the scatterers...

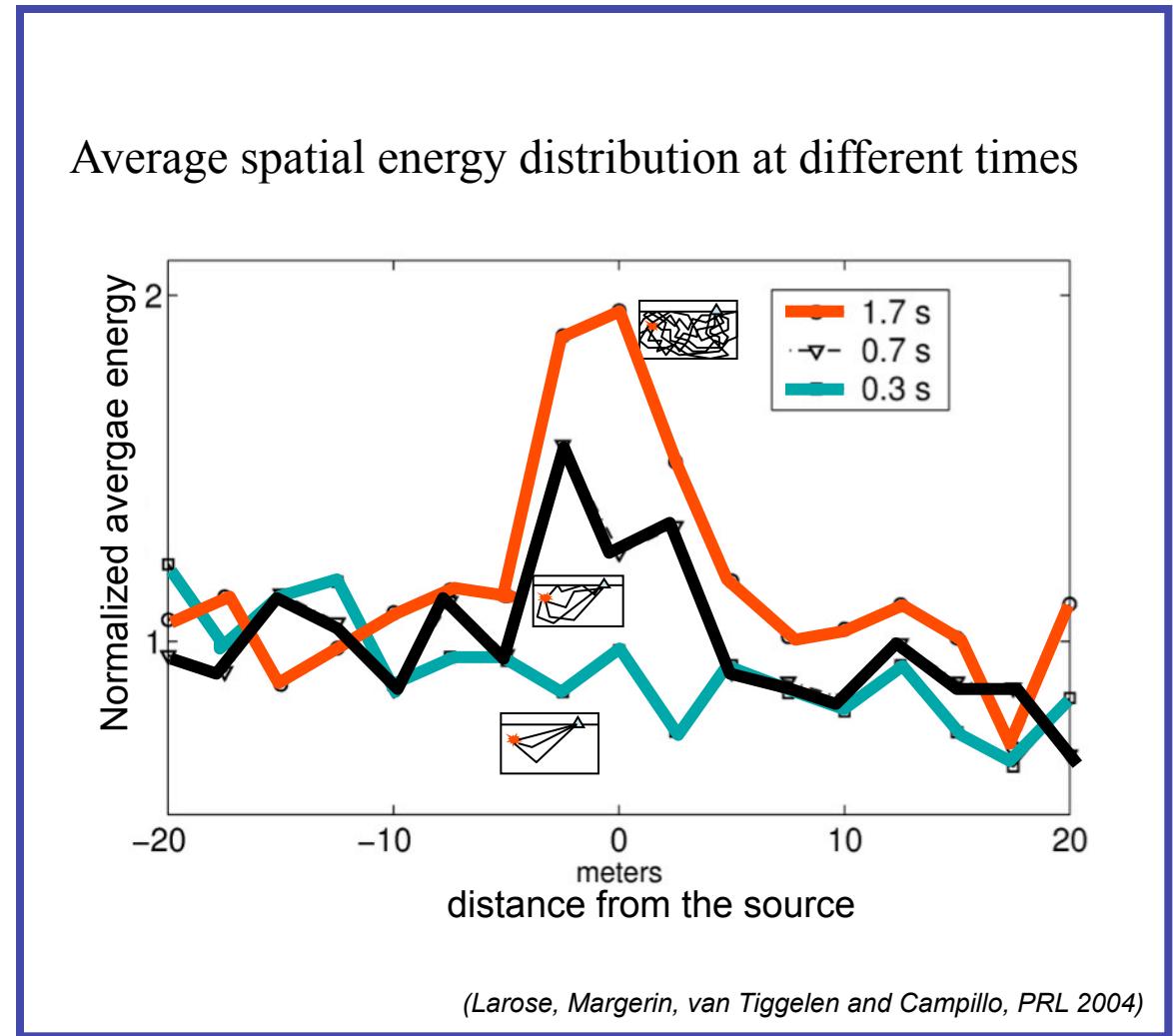
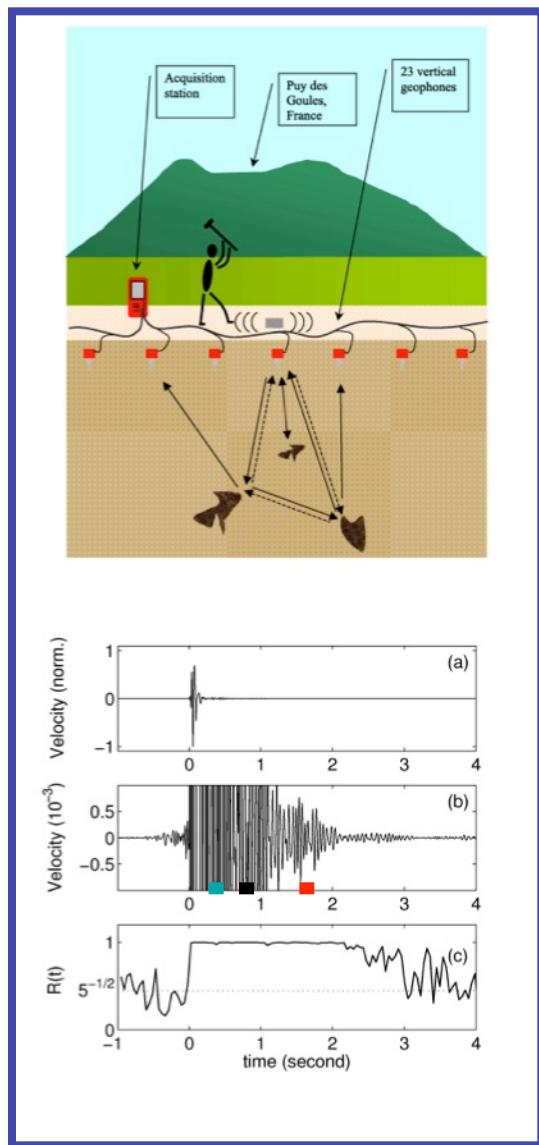


Phase difference: location of the scatterers...  
Except if R and S are at the same place

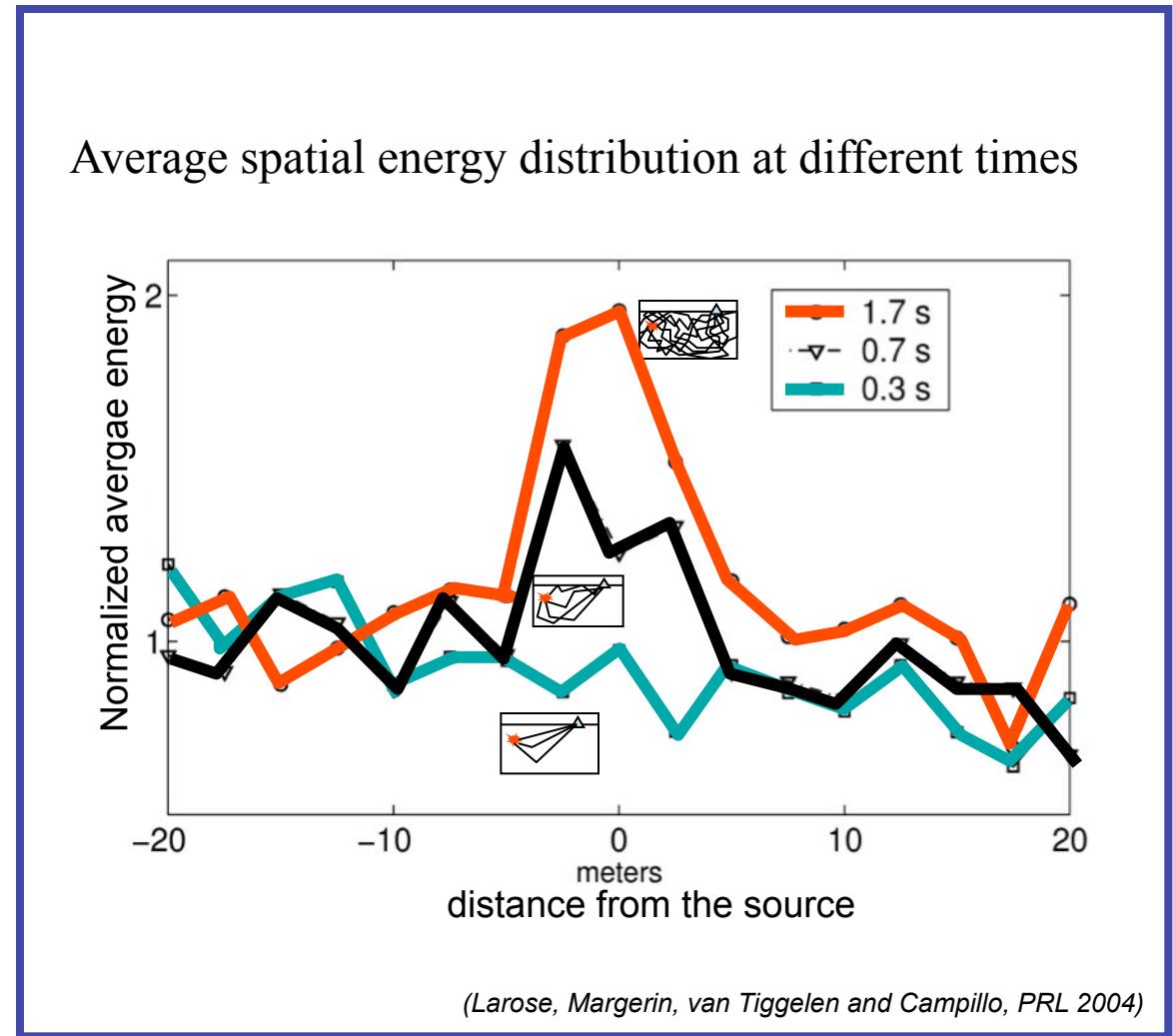
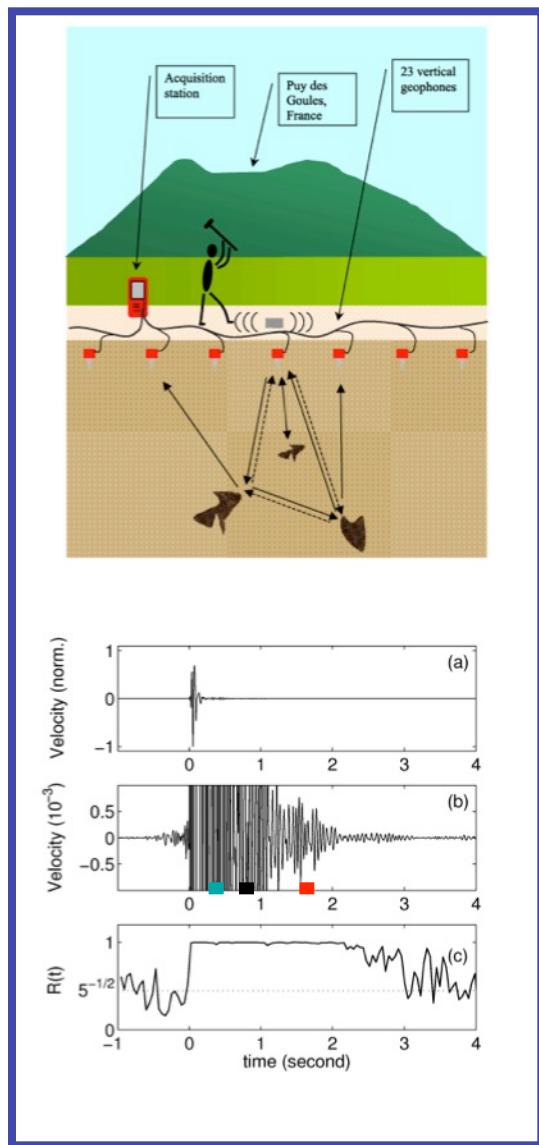


Spot of intensity enhancement: factor 2

# Weak localization: coherent backscattering of seismic waves



# Weak localization: coherent backscattering of seismic waves



|

## **Correlation of diffuse scalar fields and Green function Open medium**

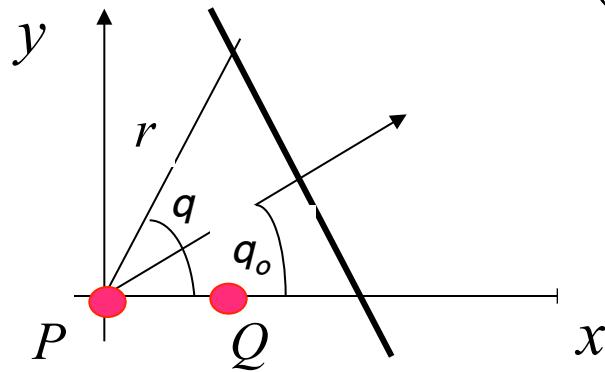
# SPAC

In this method one looks for surface waves velocities in order to find the structure. It is assumed that seismic noise is stationary. Spatial seismic arrays are required to make the azimuthal average.

If moreover, noise is isotropic the same result can be obtained with only two stations by means of the cross correlation stackings for a long time.

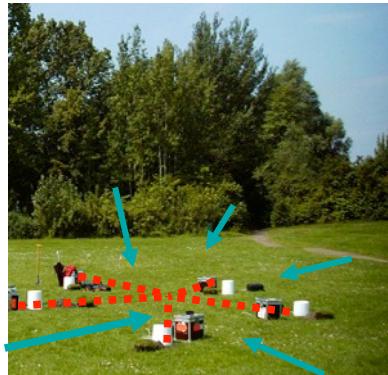
Let's consider the plane wave

$$u(r, \theta, \omega) = F(\omega) \exp(-ikr \cos(\theta - \theta_0))$$



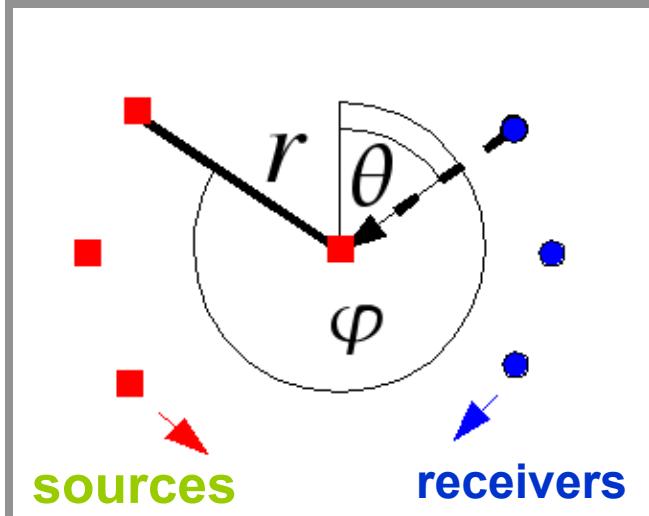
$$\frac{u^P u^Q^*}{|u^P| |u^Q|} = e^{+ikr \cos \theta_0}$$

# Spatial autocorrelation coefficient: Average spatial coefficient (2D circular array)

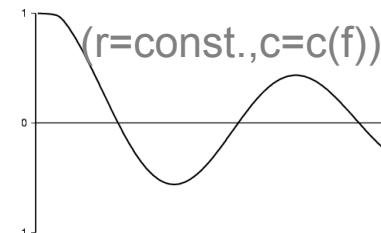


Aki (1957): azimuthal averaged spatial autocorrelation coefficients

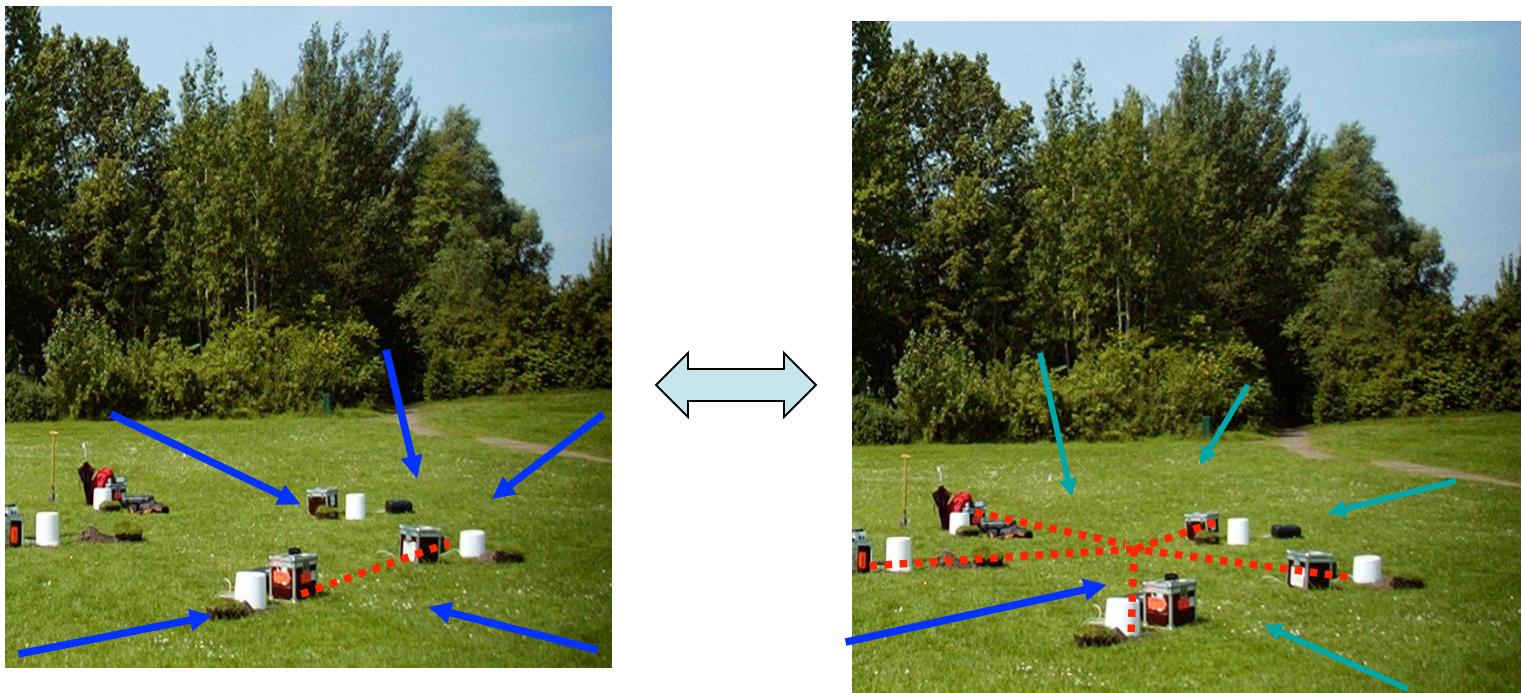
$$\bar{\rho}(\mathbf{r}, \omega_0) = \frac{1}{\pi} \int_0^\pi \rho(\mathbf{r}, \varphi, \omega_0) d(\theta - \varphi)$$



$$\bar{\rho}(r, \omega_0) = J_0\left(\frac{\omega_0 r}{c(\omega_0)}\right)$$



# Spatial autocorrelation coefficient: Average spatial coefficient

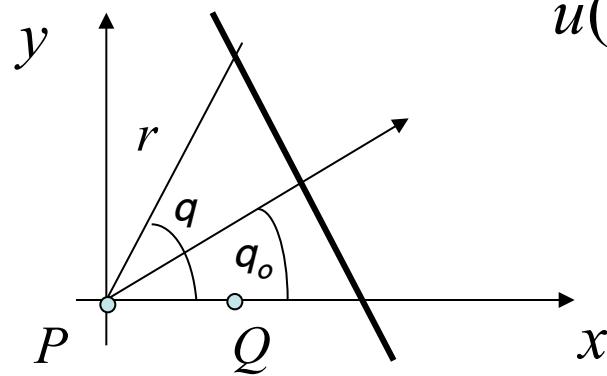


Evaluating  $k$  at different frequencies makes it possible to obtain the dispersion curve  $C(\omega)$ .

The method relies on the hypothesis of the stationnarity of the noise and requires specific array design to perform the azimuthal average.

Another approach consists of using only two points and to rely on long term average to produce the azimuthal average.

Let us consider a plane wave in 2D:



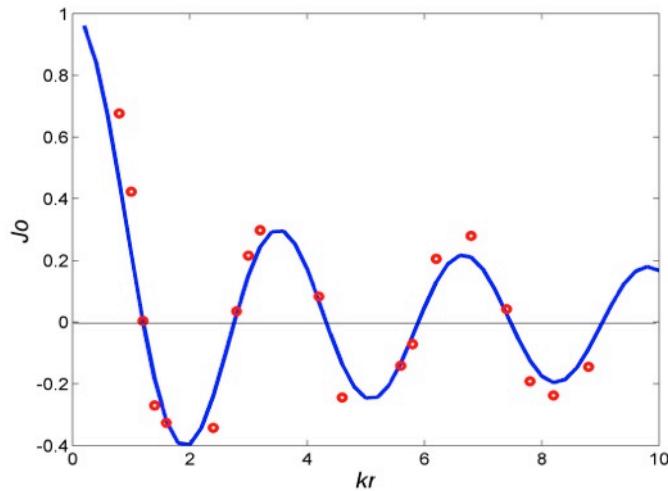
$$u(r, \theta, \omega) = F(\omega) \exp(-i k r \cos(\theta - \theta_0))$$

$$\frac{u^P u^Q^*}{|u^P| |u^Q|} = e^{+ikr \cos \theta_0}$$

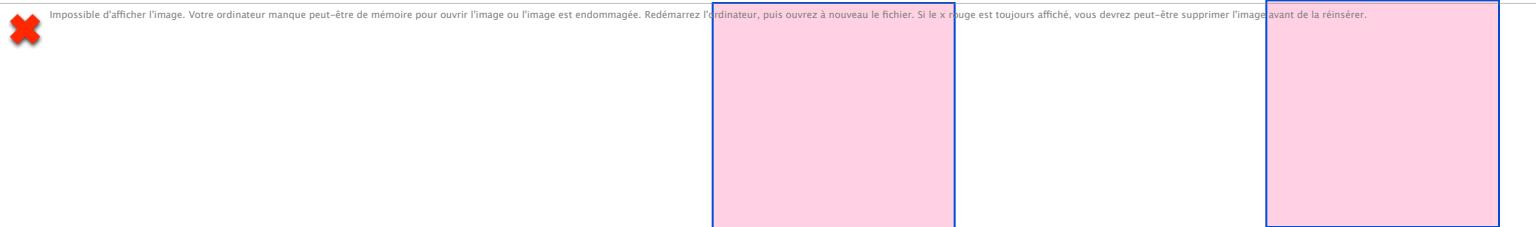
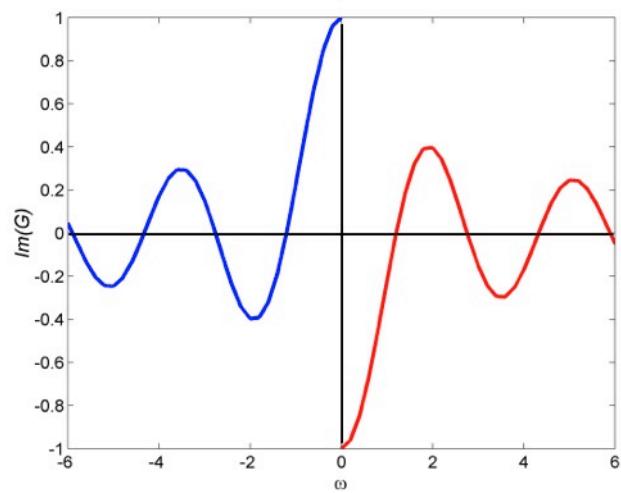
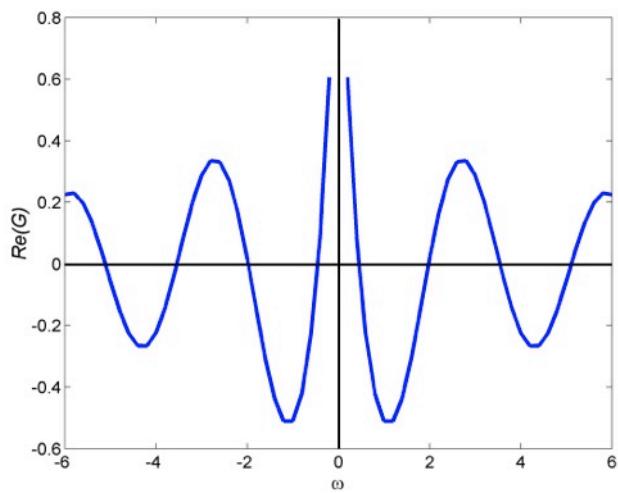
$$\langle \rho(r, \omega) \rangle = \left\langle \frac{u^P u^Q^*}{|u^P| |u^Q|} \right\rangle = \left\langle e^{ikr \cos \theta_0} \right\rangle = \frac{1}{2\pi} \int_0^{2\pi} e^{ikr \cos \theta_0} d\theta_0 = J_0(kr)$$

azimuthal average  
of the spatial  
cross-correlation

$$\frac{1}{2\pi} \int_0^{2\pi} \left( \sum_{m=0}^{\infty} \varepsilon_m i^m J_m(kr) \cos m\theta_0 \right) d\theta_0$$



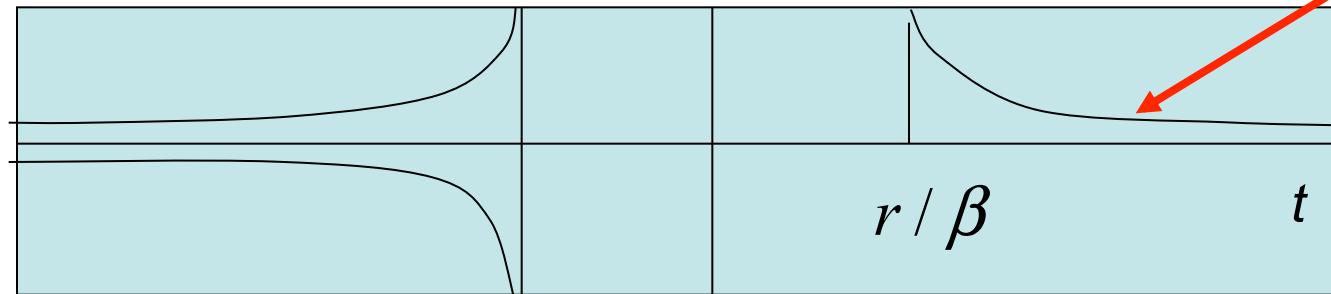
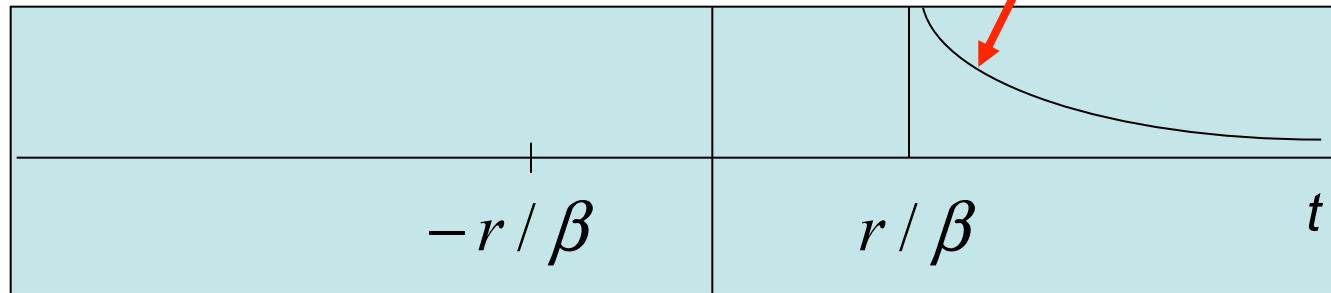
$$G = \frac{1}{4i\mu} H_0^{(2)}\left(\frac{\omega r}{\beta}\right) = \frac{1}{4\mu} \left\{ -Y_0\left(\frac{\omega r}{\beta}\right) - i J_0\left(\frac{\omega r}{\beta}\right) \right\}$$



# Causality

$$G = \frac{1}{4i\mu} H_0^{(2)}\left(\frac{\omega r}{\beta}\right)$$

$$G = \frac{1}{2\pi\mu} \frac{H\left(t - \frac{r/\beta}{\sqrt{t^2 - r^2/\beta^2}}\right)}{\sqrt{t^2 - r^2/\beta^2}}$$

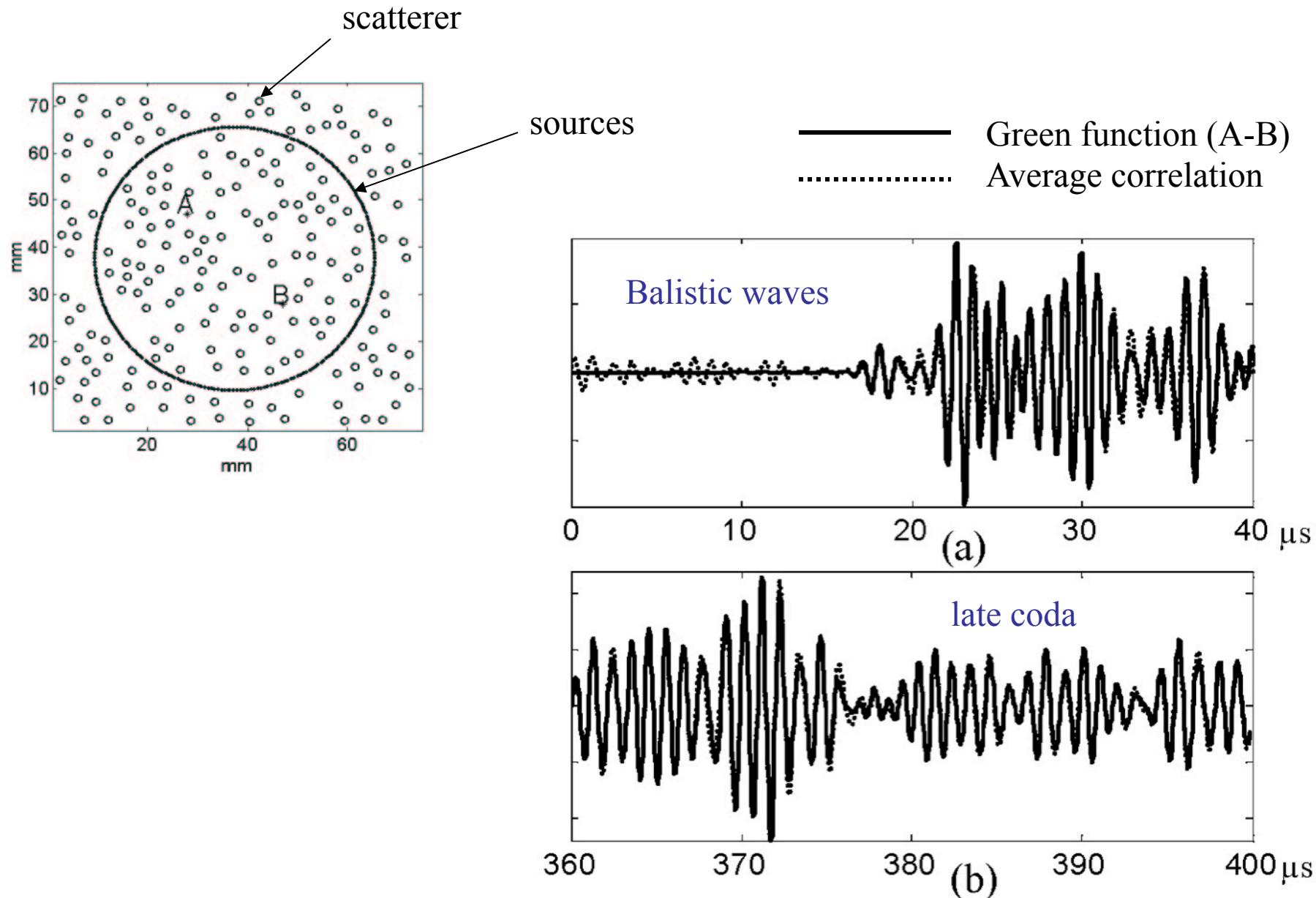


$$G_{22}(r,\omega)\!=\!\frac{1}{4\mu}\!\left\{-Y_0\!\left(\frac{\omega r}{c}\right)\!-\!iJ_0\!\left(\frac{\omega r}{c}\right)\!\right\}$$

$$J_0\!\left(\frac{\omega r}{c(\omega)}\right)\!=\!-4\mu\operatorname{Im}\!\left(G_{22}(r,\omega)\right) \qquad\qquad r=|P,Q|$$

$$\operatorname{Im}\!\left(G_{22}^{PQ}\right)\!=\!\frac{-1}{4\mu}\!\left\langle\frac{u_2(P)u_2^*(Q)}{|u_2(P)\|u_2(Q)|}\right\rangle$$

# A numerical experiment with an open medium (absorbing boundaries):



## **Physical interpretations**

**Time reversal**

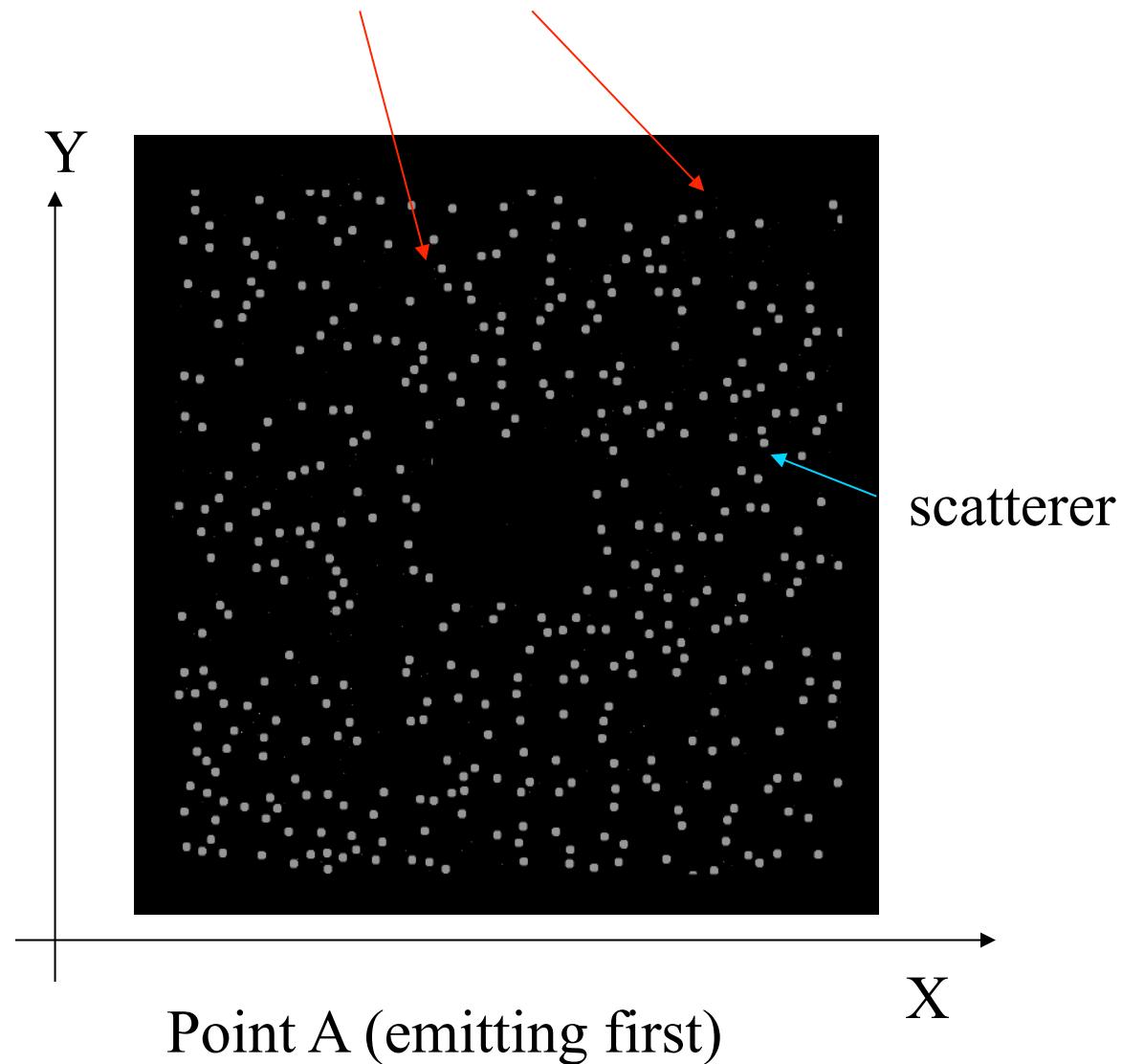
Rays and stationary phase

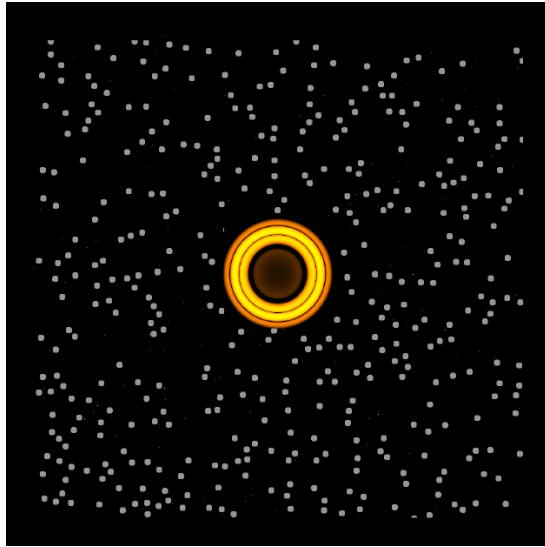
# Correlation vs Time Reversal

- C source
- A and B receivers
- Correlation :  
 $S_{CA}(t) \times S_{CB}(t)$
- A source
- C receiver
- C emits time-reversed field
- B receiver
- Convolution :  
 $S_{CA}(t) \otimes S_{CB}(-t)$

# Numerical 2D FD simulation

200 « sources » C (randomly placed)

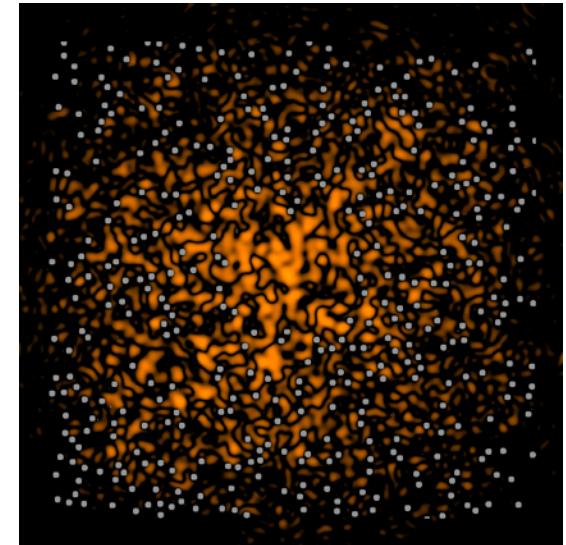


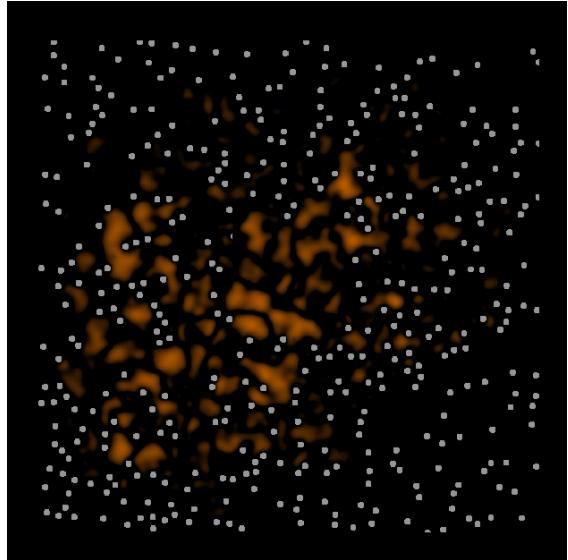


A pulse is emitted in A  
and recorded at point randomly  
distributed

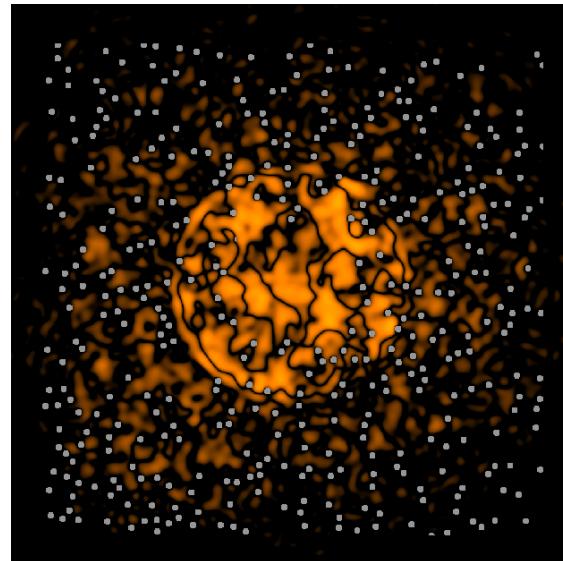


time



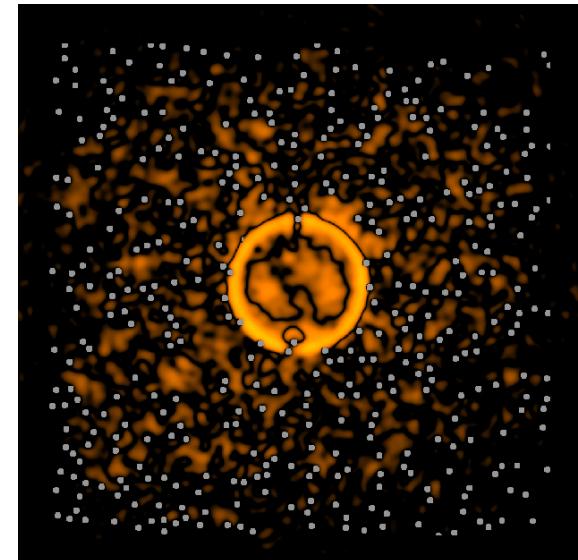


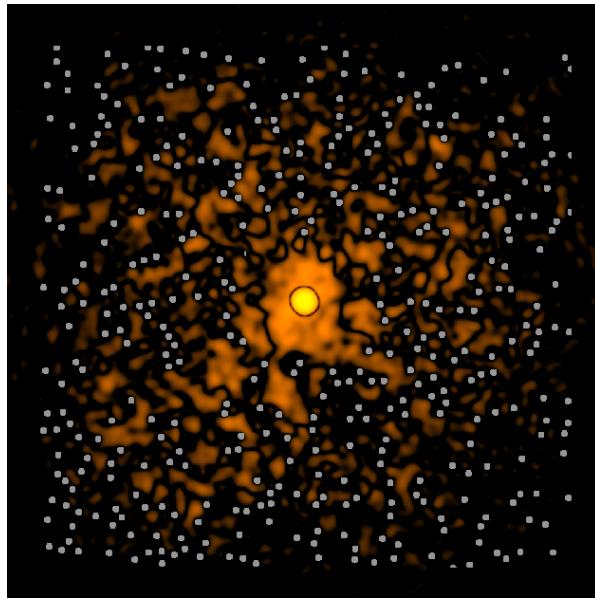
Re-emission from the points ‘C’  
of the time-reversed signals  
(map of cross-correlations)



Constructive  
interferences of time-  
reversed field

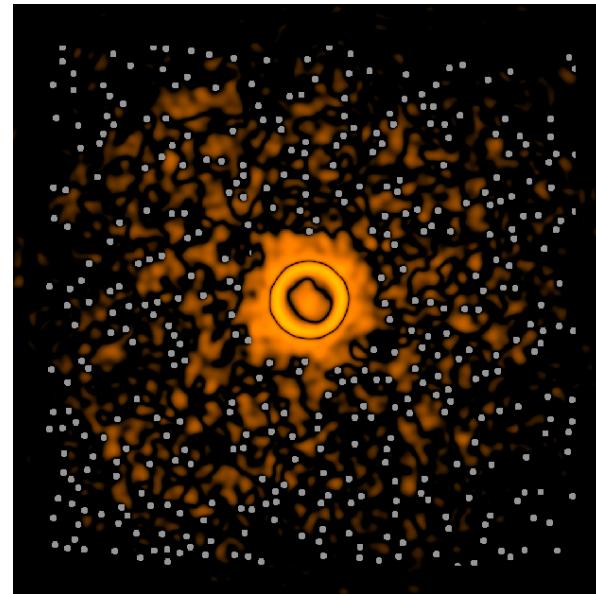
Converging field  
 $: G(-t)$





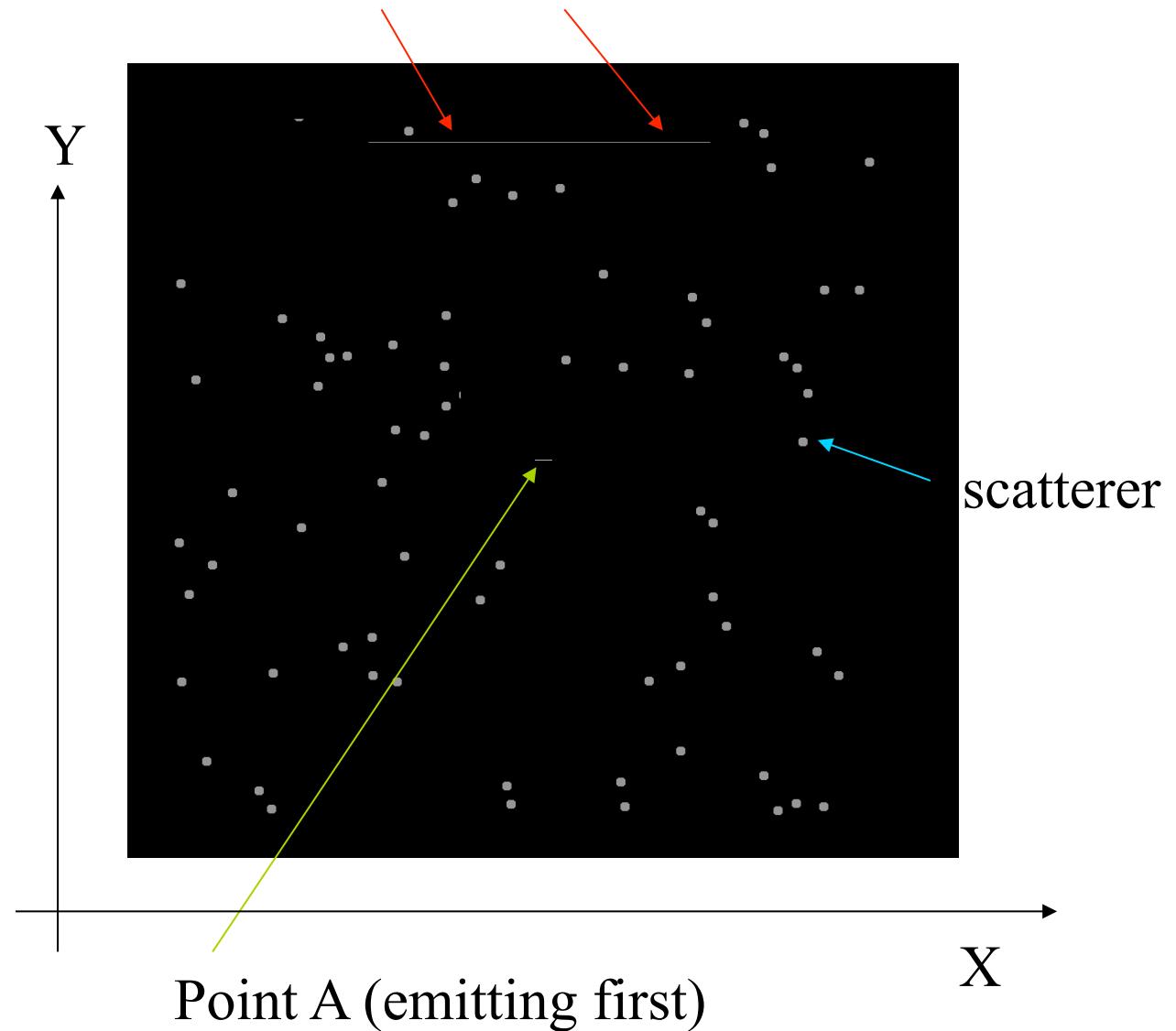
Nearly perfect refocalisation

Re-emission from A :  
 $G(t)$

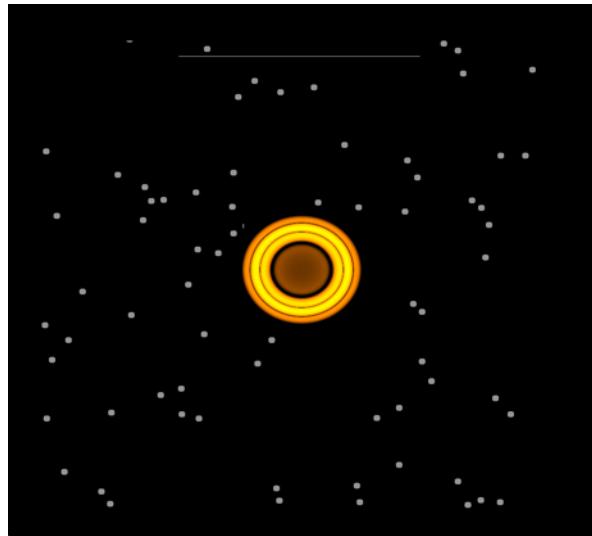


# A more realistic configuration of sources

40 « sources » C (lined-up along a fault...)



# Time reversal experiment

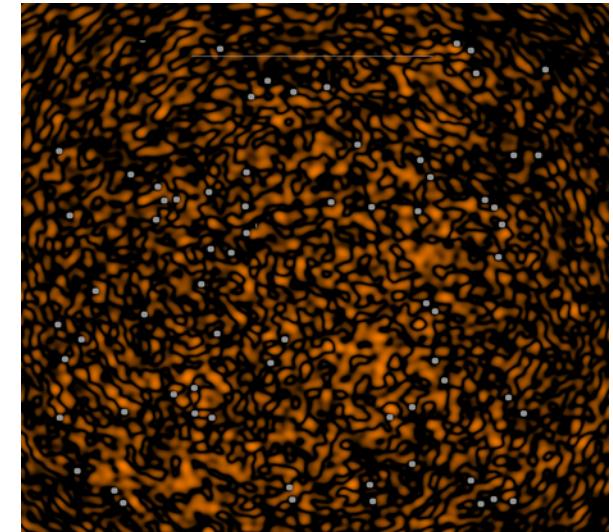


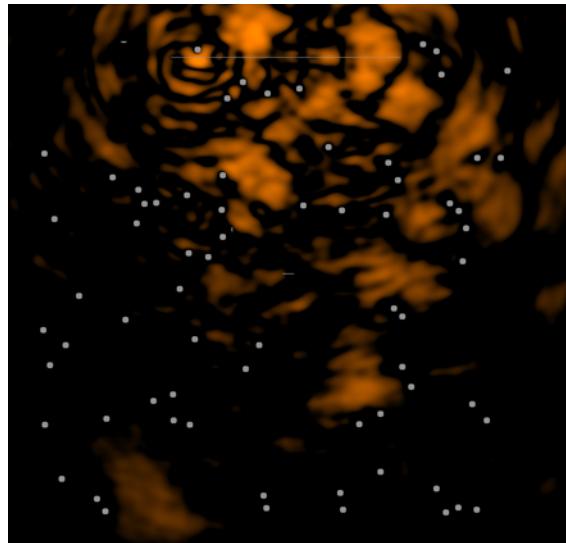
A send a pulse



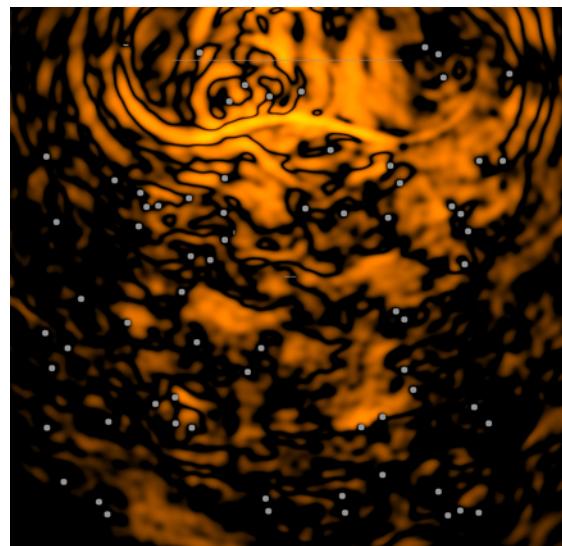
Scattering effects

Diffuse field is  
also recorded



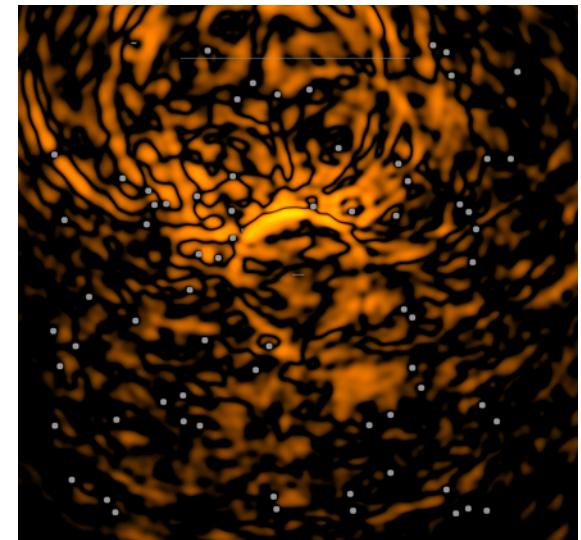


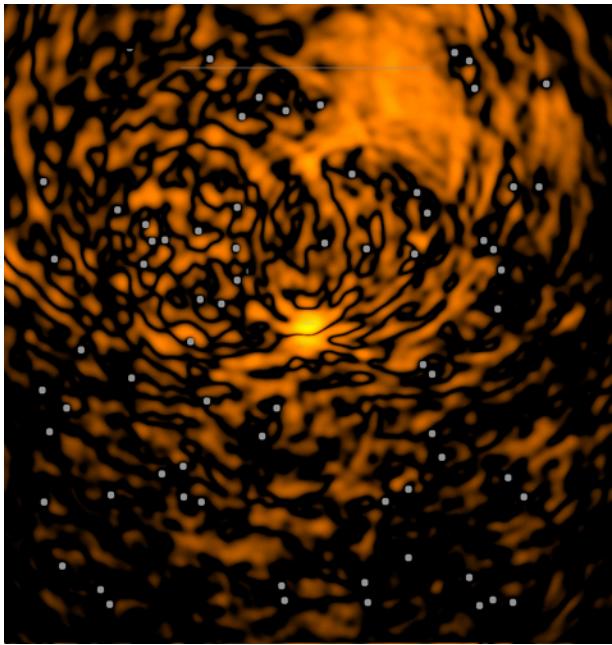
Re-emission



Converging field :  $G(-t)$

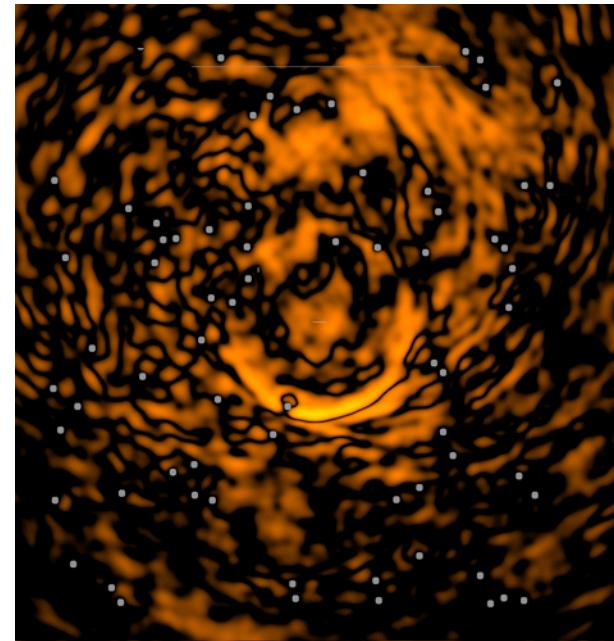
time





Partial focalisation

Diverging field :  $h_{AB}(t)$



The symmetry of the Green function is lost!