



Master Class Senior ISTERre



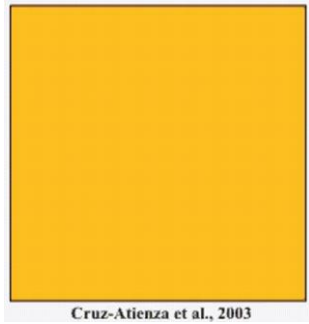
Interaction onde-structure: les enjeux de l'imagerie crustale *Une vision partielle de Jean Virieux*

Direct contributions from
Romain Brossier (ISTERRE)
Ludovic Métivier (LJK/ISTERRE)
Stéphane Operto (GEOAZUR)
Alessandra Ribodetti (GEOAZUR)
Many other contributions from SEISCOPE's team (over 15 years)

Seiscope consortium



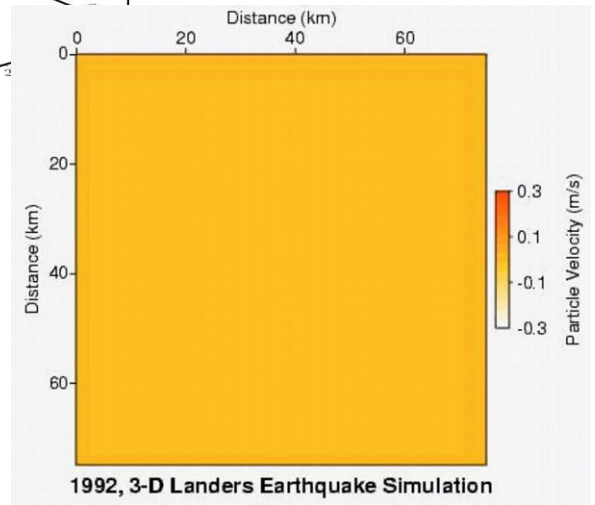
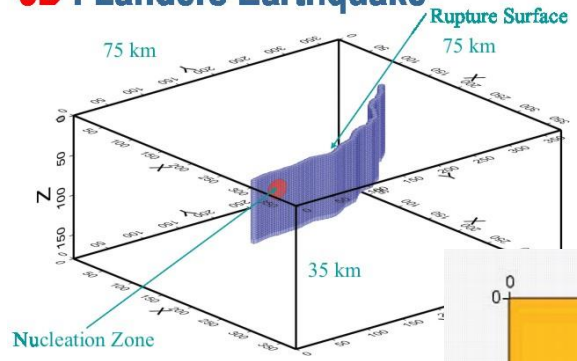
Waves: attractive & vanishing object



Complex wave propagation

Cruz-Atienza et al., 2003

3D : Landers Earthquake



1992, 3-D Landers Earthquake Simulation

My personal scientific interest:

- Waves (optical, acoustic, elastic): unique communication when away
unique probing of hidden objects
very few other left alternatives...

➤ Why Earth Sciences:

& challenging physics: elastic propagation with two speeds
& outdoor activities: key argument for me



work



riding

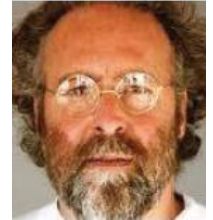
Scientific topics...

PhD topic: earthquake dynamic rupture ...
theoretical investigation using numerical tools
(specifically FD wave propagation modeling)



No data!

aside output: theoretical FWI
with the consortium « GTG »



Post-PhD: travel-time tomography with many people (still on)



Data! Volcanic area (Mt Vesuvius/ Campi Flegrei
nearby Naples) & Rift zone (Gulf of Corinth)



Post-PhD: RAY+BORN imaging with many people (consortium « GRS »)



Data from Oil&Gas enterprises (Elf-Aquitaine – Oseberg field)

Long time after: real applications of FWI (consortium SEISCOPE, still on)



New generation ...



Few Students at GeoAzur ...

Alessandra (1994 – 1998)



Hafedh (2006-2009)



Romain (2006-2009)



Victor (2003-2006)



Vincent P. (2008-2012)



Bernhard (2000-2003)



Celine (2004-2007)



Vincent E. (2007-2011)



Guanghai (2008-2012)



Raffaella (1996-2003)



Francoise (1991-1994)



Pascal (après sa reconversion...)



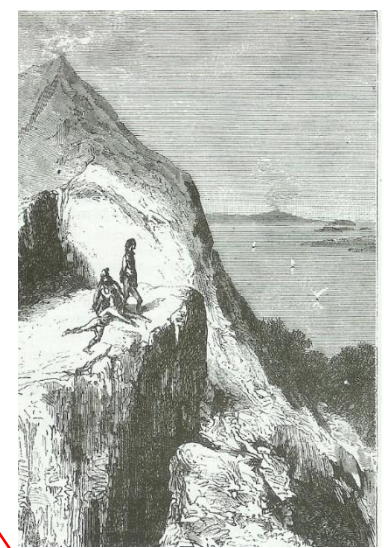
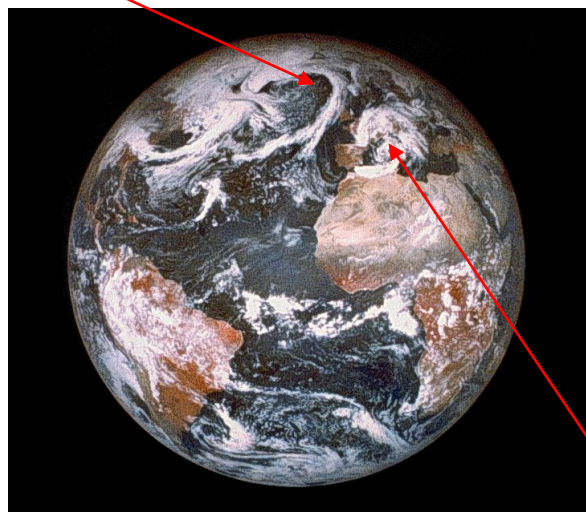
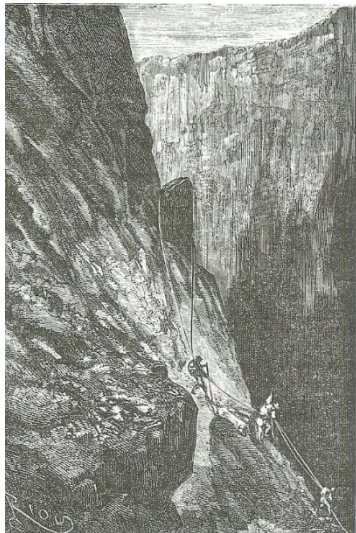
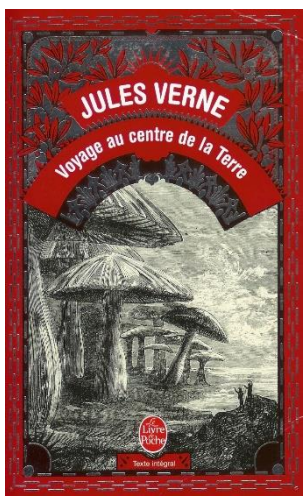
And many others ...

The Earth system: its mysterious interior?

From Snaeffels

Journey inside the Earth

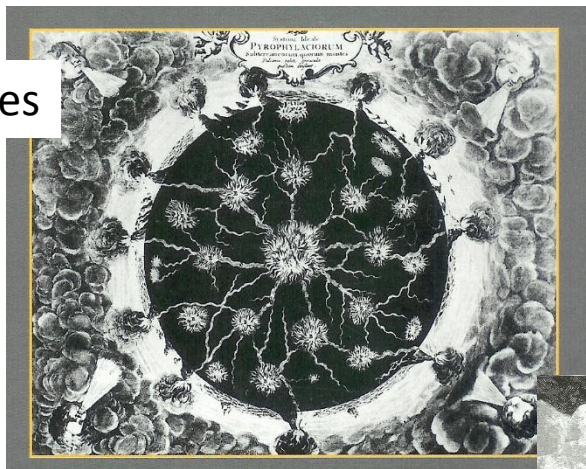
To Stromboli



WHAT DID THEY FIND?



cavities



(Kircher, 1665)

minerals



Over-pressurized water

Phases and waves



lakes

Great discovery at the end of the XIX century
Modern seismology was born!



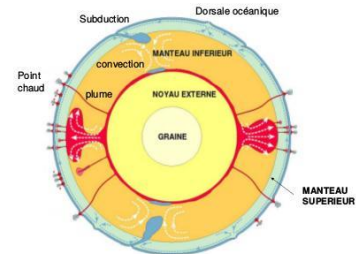
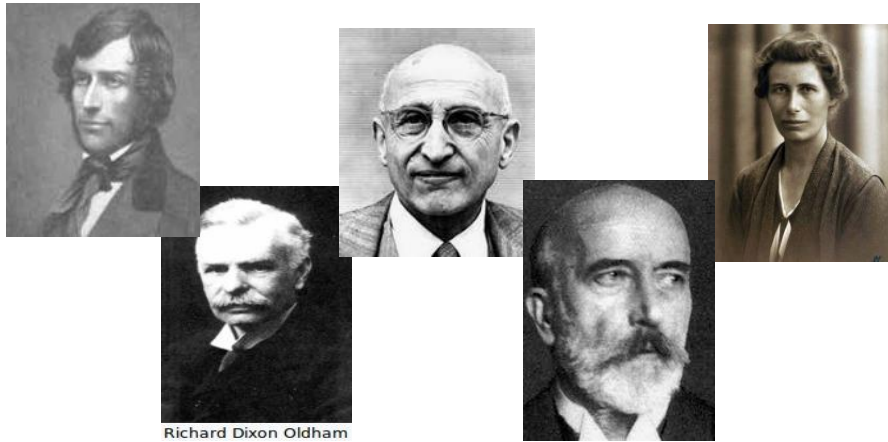
In 1889, a ground vibration in Europe was linked to a Japanese earthquake: waves can travel inside the entire Earth.

Oscillation of a pendulum at Potsdam (Germany)

Telegraph revolution!
Correlation between two far-away objects!

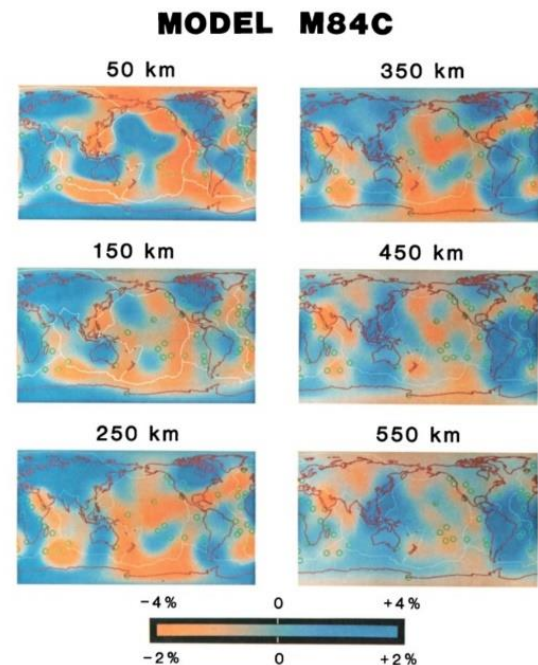
Seismology (including seismics)

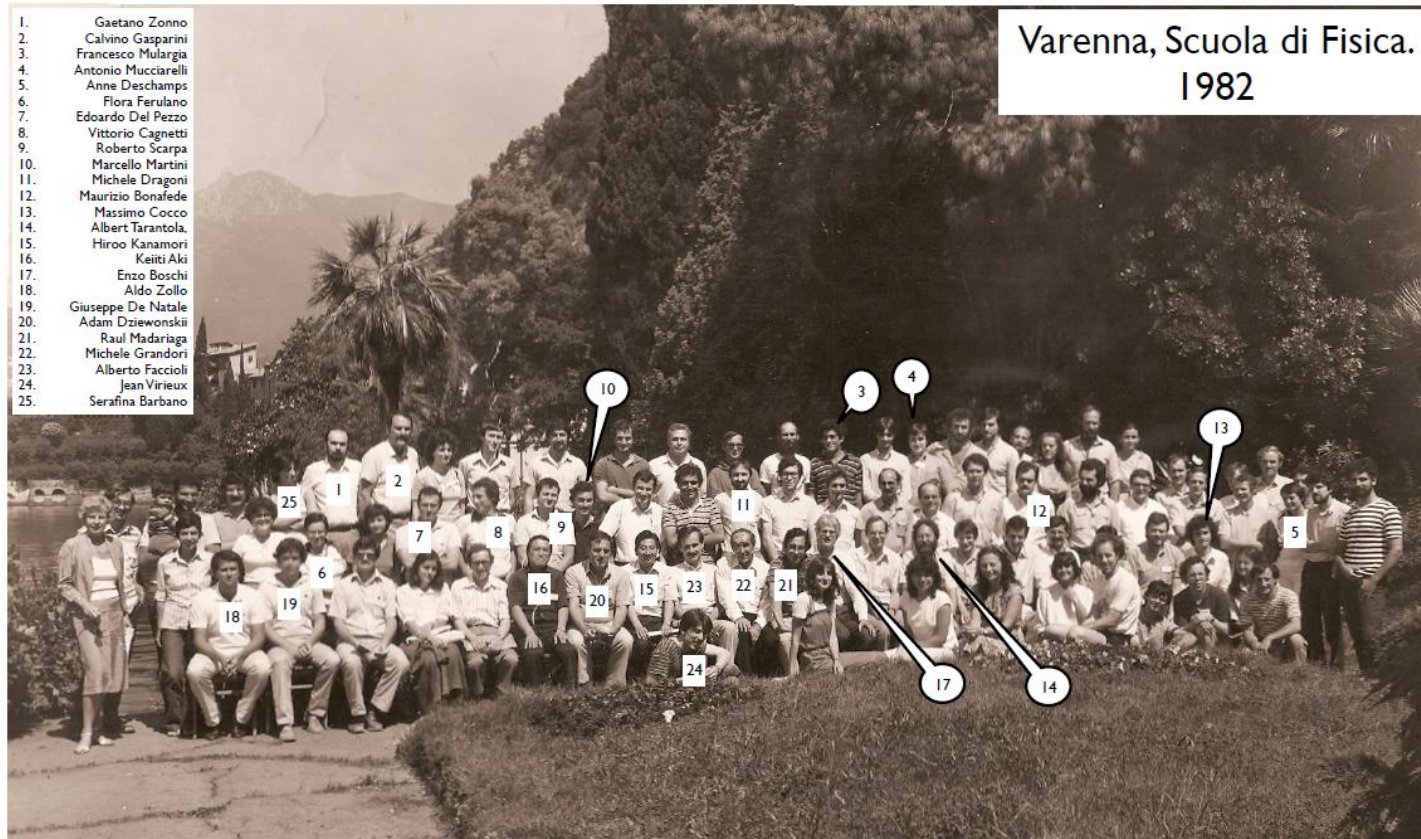
Before 80's, seismic imaging was based on **travel time measurements (direct and reflection signals)**: radial model such as the PREM model (Dziewonski, 1981)



First paper using **amplitude**: M84C models of the Earth interior (Woodhouse & Dziewonski, 1984) competing with models based on **phase/time** data (Dziewonski, 1983, 1984; Dziewonski & Anderson, 1984).

Phase information plays a crucial role



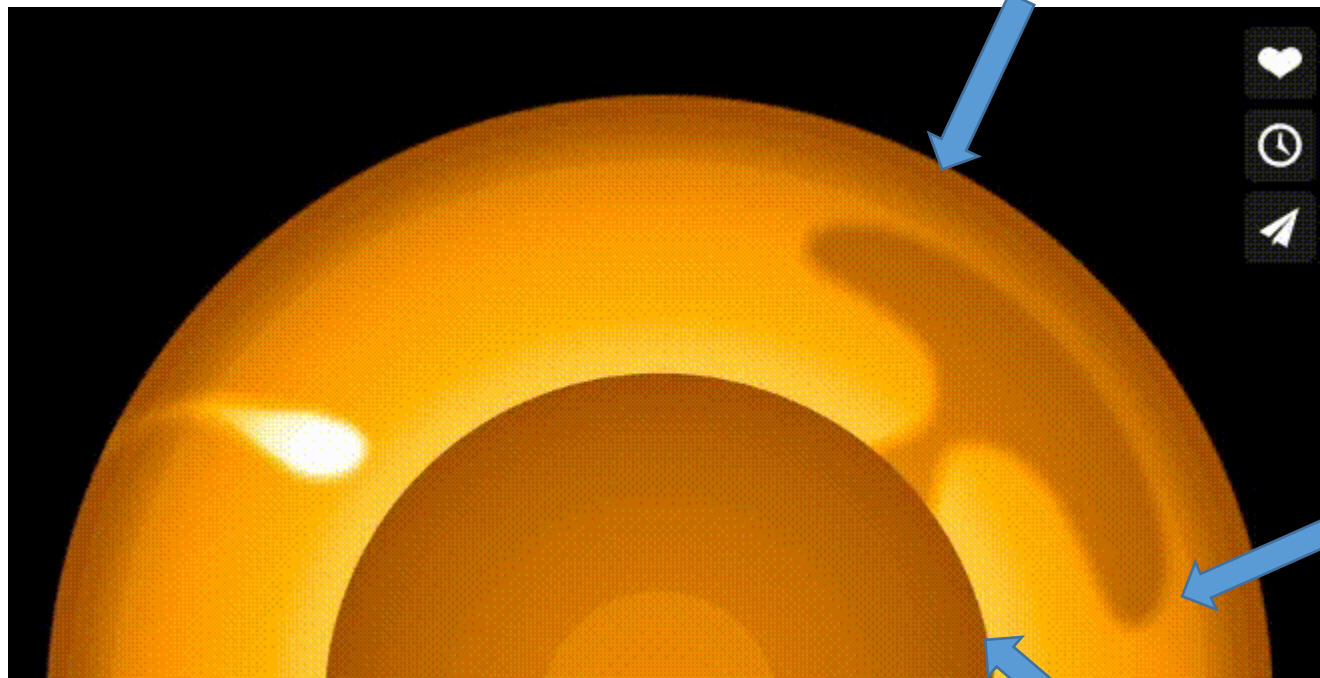


Message of John Woodhouse

Focus your attention on **seismograms** (amplitude analysis)!
They carry a lot of information (normal modes summation)

Focusing targets !!!

Crust: complex nearby object



Mantle: still a rather simple object!

Movie from Stahler & Tarje

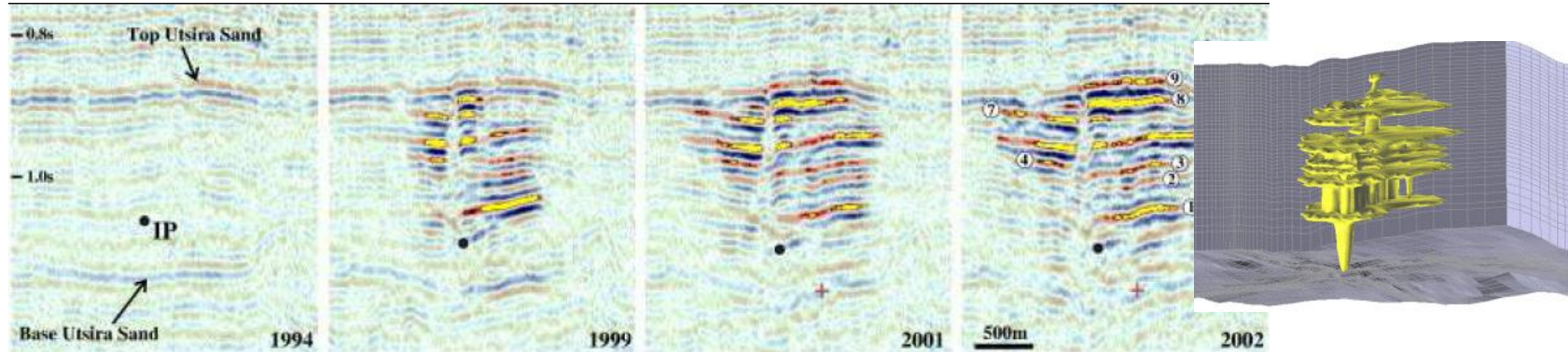
CMB: a far-away object

should be much more complex than available images ...

On the need to image, characterize and monitor the subsurface

CO2 storage: monitoring the Sleipner reservoir ...

Reservoir model



Geological structures for giant storage:

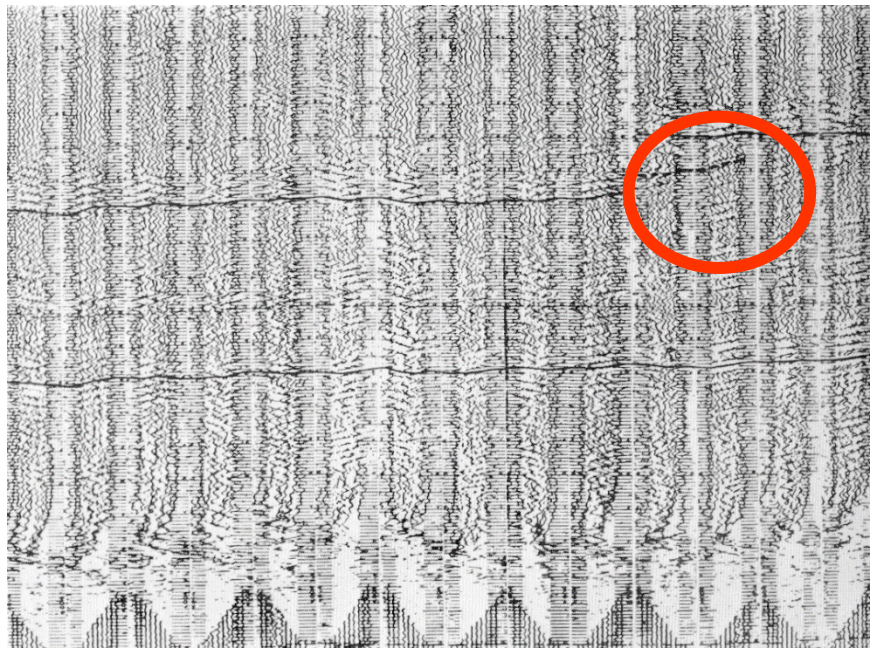
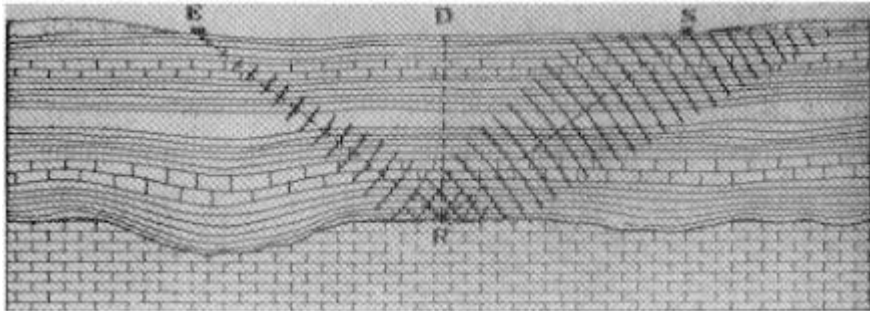
Ultimate storage: CO2 sequestration, chemical deposit, nuclear deposit and so

Temporary storage: heat (geothermal targets), H2, CH4, HP gas over year cycling

Seismic imaging: not only a crucial need for Oil & Gas primary resources, but of critical importance for energy demand.

Young generation should be trained on that topic!

Interpretation cartoon (1921)



Fault detection on paper plots

Fig. 1.8. Refraction recording, 1936. (From Pett, 1976.)

(a) Normal (no salt) record, (b) record showing salt

Source: (a) (b) (c) recording track it was about this time (Courtesy GSI)



Head wave:
Polarity as a
salt proxy



Three sensors and an acquisition truck!

Hard life for field seismologists



sheepfold

but in lovely places

Smooking the drum!



In Greece summer time!

My student time

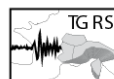
TGRS: STEF station
in my family village



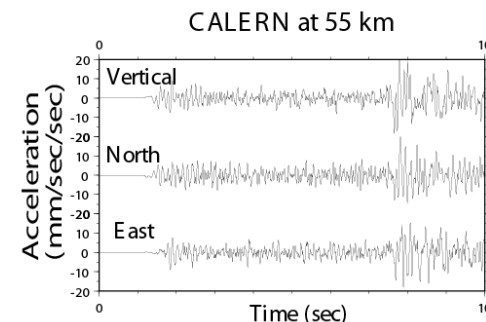
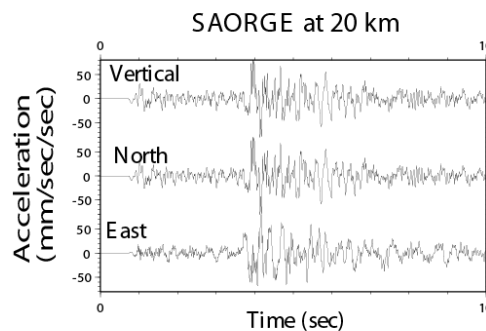
My research time



Radio clock (<1995)
to
GPS clock (> 1995)



True amplitude



INSU/CNRS

MESR/UNSA

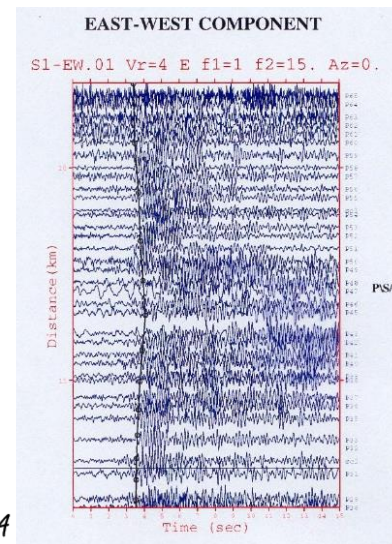
Phases and waves

Hunting for amplitude variation?
Seismic array on Vesuvius summit!



Learning field challenges
Job done but clock synchro. KO!

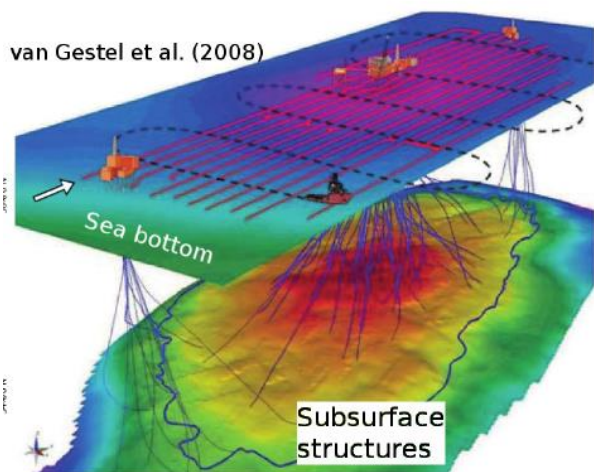
LALA phase!



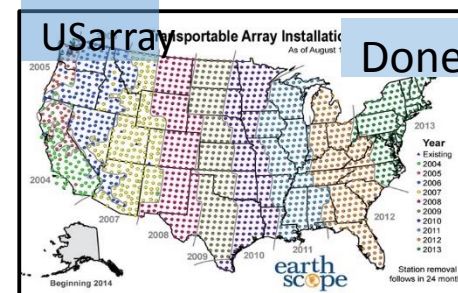
A lot of fun, but
take-away lesson:
professional
experts needed
->
Look at industrial
datasets for crust
imaging

High-quality data with unprecedented density of stations at various scales

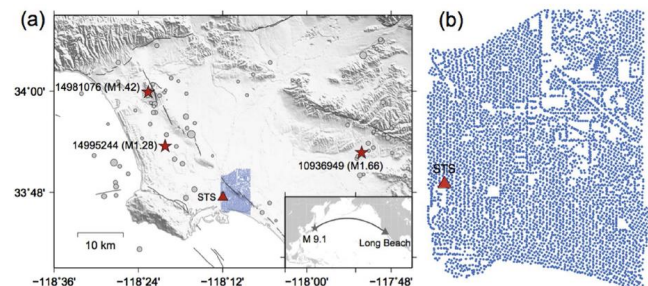
Valhall experiment:
50000 shots; 2300 receivers



New data?
New interpretation?
New methods?



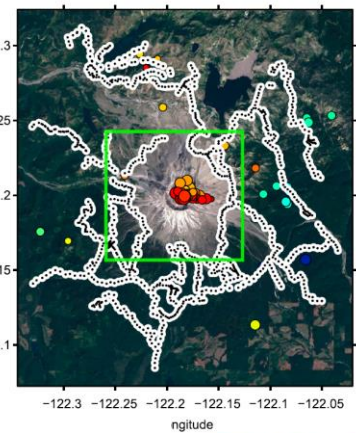
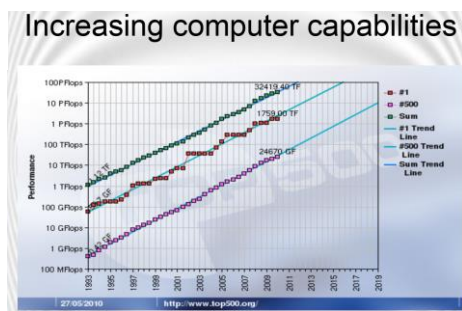
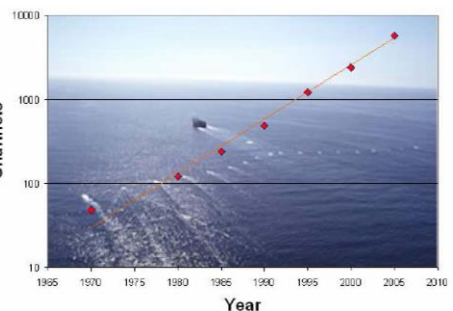
Alparray
Chinesarray



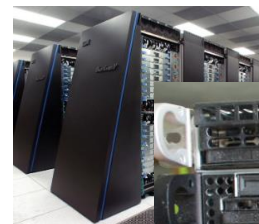
Beach nodal array (Slater & Hollis, 2012; Li et al, 2018)

Mt St Helens (Hansen & Schmandt, 2015)

Moore's Law:
marine seismic channels



❖ Dense seismic data



Sophia Antipolis



❖ High computer capabilities

How lucky are you?



Better knowledge, better management of underground structures and resources (toward technical sustainability ...)

Before (my time) few data and limited computer resources (especially in France):
supercomputer '80s ~ Smartphone '10s



Real medium

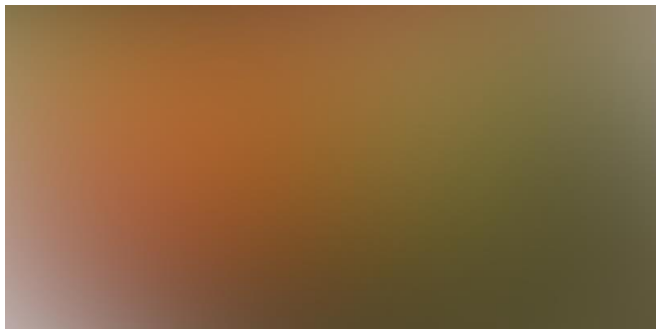


Translucent Earth

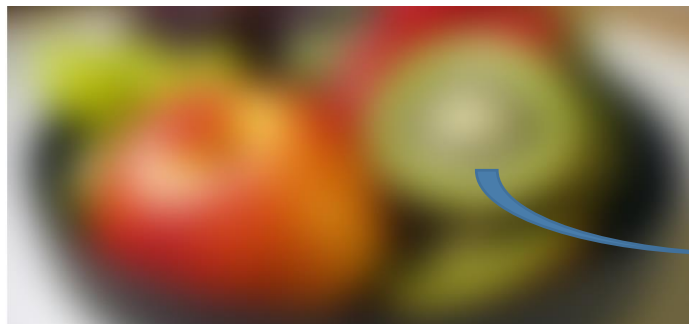
Unaccessible target?
Intrinsic remote sensing limitation
Seismic finite frequency & attenuation

Seismic imaging

Travel time information?



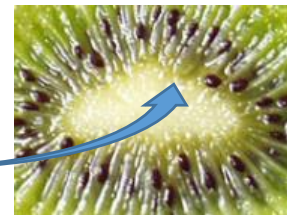
Waveform information?



Additional information?



Interpretation?



Beyond seismic resolution:
dictionary learning

How far should/could we go with seismic waves?

Industrial & Academic communities interact more nowadays !!!

Pending question:

Are we extracting all the information
contained in these massive high-cost datasets?

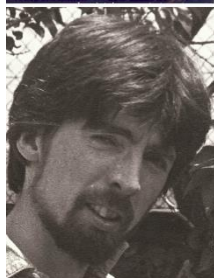
Efficiently, precisely and timely!

Open body (# closed body such as the global Earth: exit normal modes)

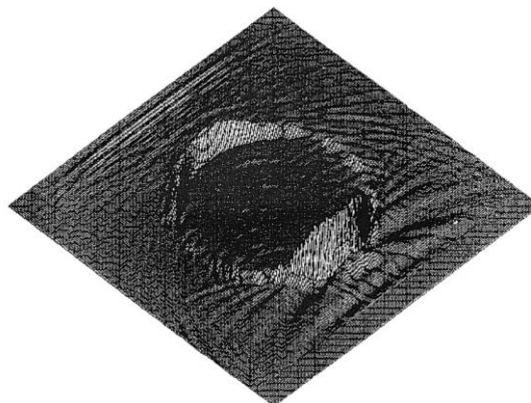
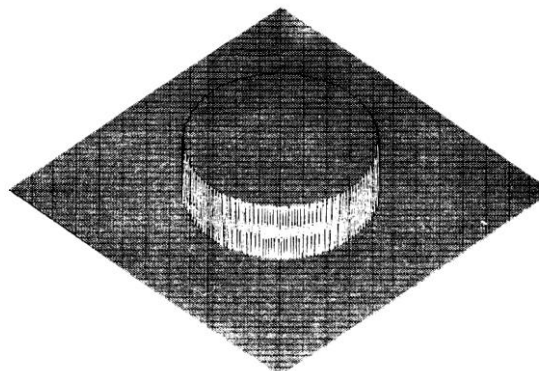
No more guitar!



Odile Gautier



Synthetic model



Reconstructed model using gradient with 5 iterations

From French Mathematical Community

Optimal control of parameters

investigated by Lions (1968),
Chavent (1974), Bamberger et al (1977),
Lailly (1984), and finally Tarantola (1984)

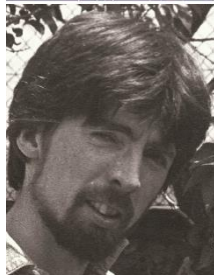
More than ten years for a slow
percolation from mathematics to geology

...

CRAY-1 resources:

Only acoustic case; not enough
memory; a lot of I/O (my own virtual
memory tool ...); few iterations using
gradient method ...

Open body (# close body such as the global Earth: exit normal modes)



Synthetic model



Reconstructed model using gradient with 5 iterations

FULL WAVEFORM INVERSION!
A dream at that time (1985)

No more
Solution of
Computation
parameters
(1968), Chavent (1974),
Enger et al (1977), Lailly (1984),
and finally Tarantola (1984)

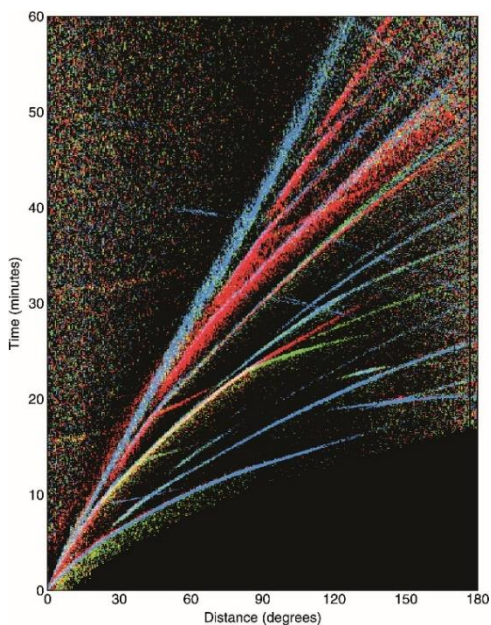
Ten years for a slow percolation from mathematics to geology ...

CRAY-1 resources:
Only acoustic case; not enough memory; a lot of I/O ...; few iterations using gradient method ...

Travel-time tomography: phase information of wave propagation

Robust observable based on clock precision (GPS) and not on instrument calibration

Half-century recording
Earth phase events



For mantle tomography
detection/association phase events OK



For crust tomography
overlapping phase events ??? ←



passive tomography (local earthquake tomography)
active tomography (upper crust tomography)

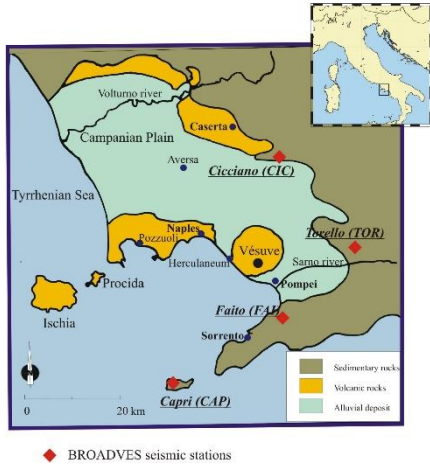
Warning

You may hear that I am using in a rather equivalent way the word « phase » and the word « traveltime ».

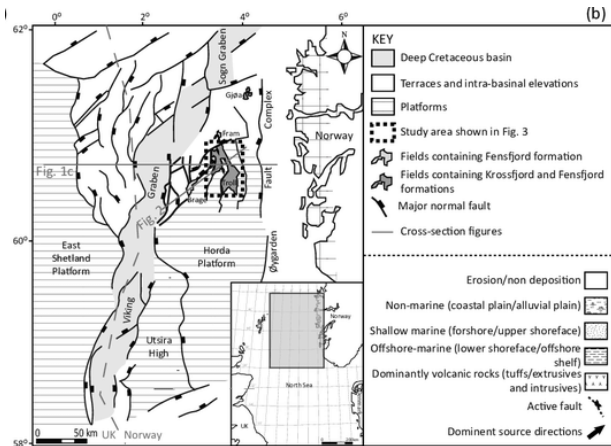
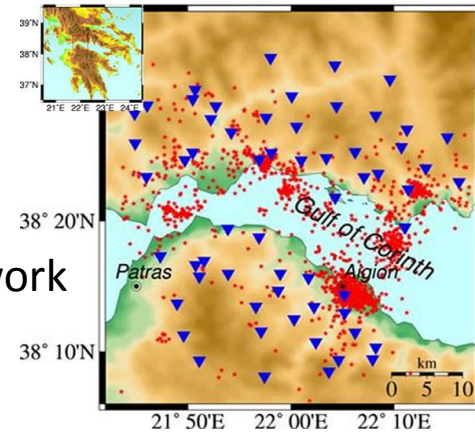
They describe different ways of propagation for dispersive waves. For ballistic waves, which are almost non-dispersive, they are equivalent.

Back to basics for three targets!

- Naples volcanic structures: Mt Vesuvius & Campi Flegrei: local earthquake tomography (LET) and active seismics. More than 20 years of investigation!
Hunting amplitude variation (Oh LALA !!!)

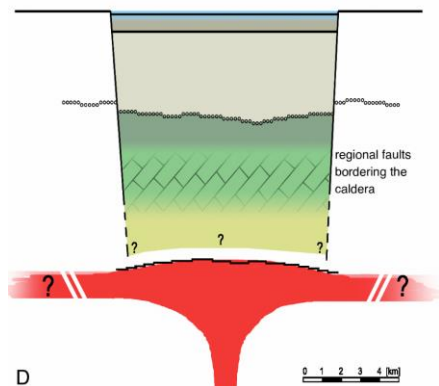


- Corinth Rift:
LET 1991 (H. L-C), 2001 (3F-Corinth F.C. & I.M.) -> CRL (P. B.)
Two seismic temporary experiments before permanent network

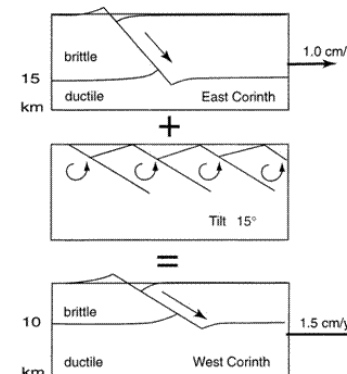
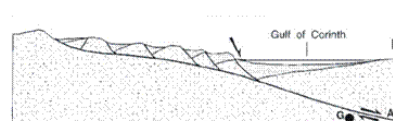
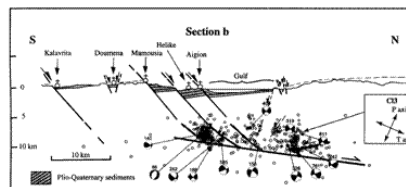


- Viking Graben (Oseberg field of Norway) : active seismics (Ray-Born Imaging) 1992 (R. Madariaga consortium GBS supported by Elf-Aquitaine and CNRS)

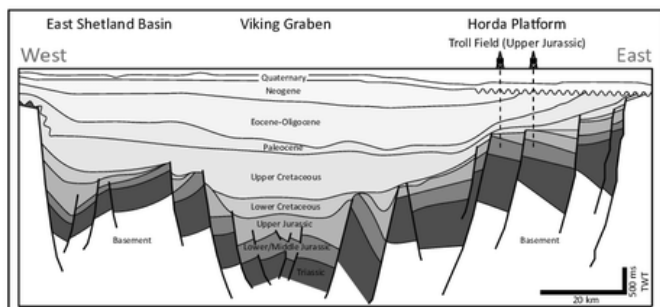
- Mt Vesuvius & Campi Flegrei: are these volcanic systems connected at depth?



- Corinth Rift: extension interpretation?



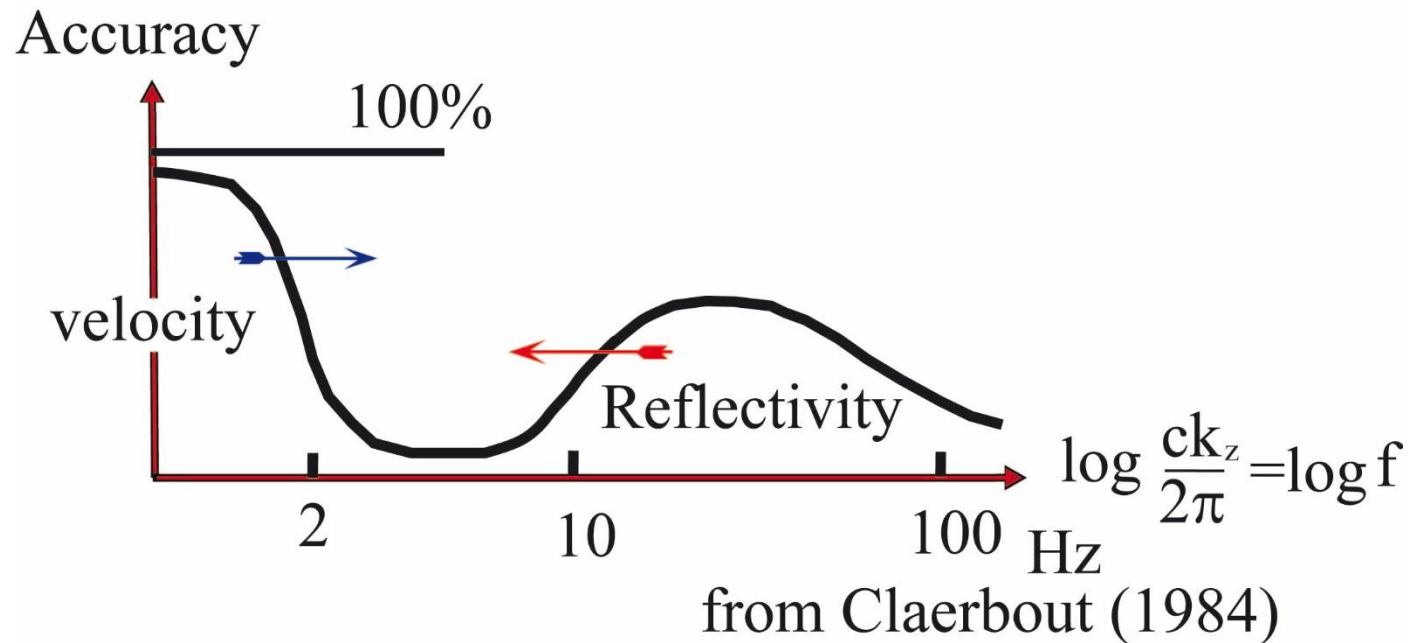
- Viking graben: fluid migration?



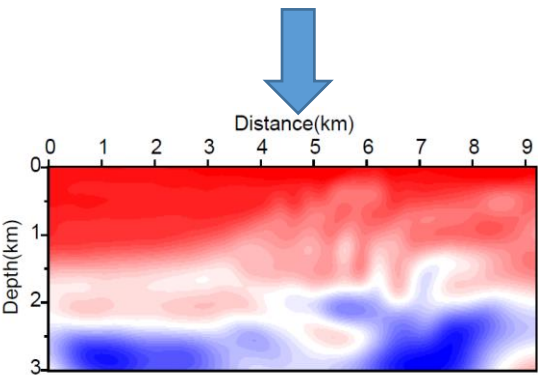
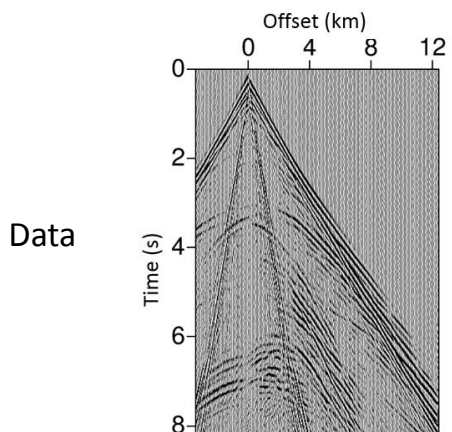
Seismic data and interpretation aside other geological, geochemical and geophysical data ... (no black/white answer)

Two scales show up in the model reconstruction

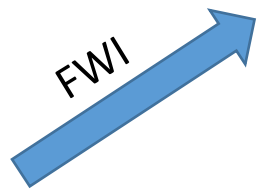
Recorded seismic traces bring different information from medium properties



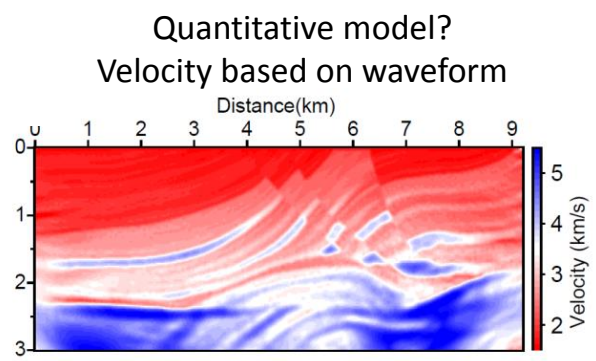
Phase (traveltime for non-dispersive wave) drives imaging



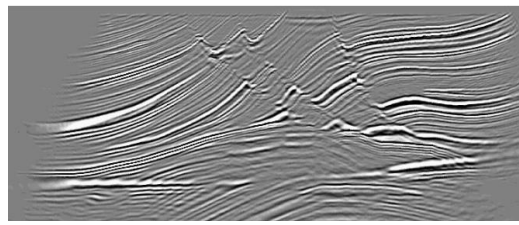
Macro-model based on velocity analysis



RTM
or
PDM



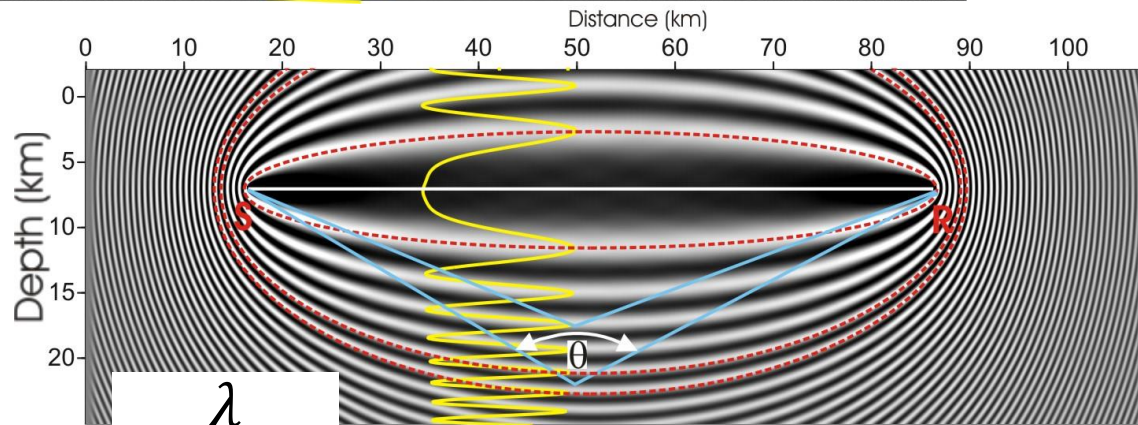
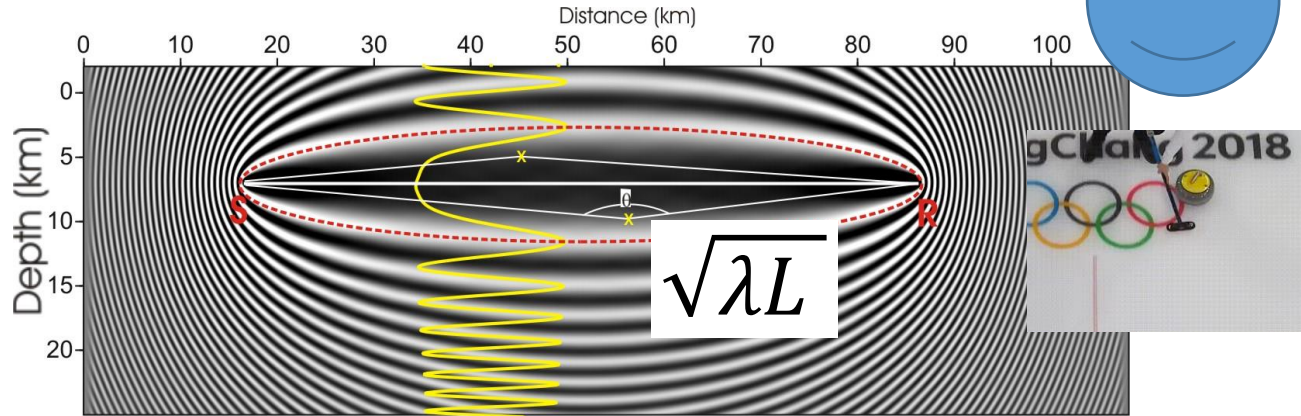
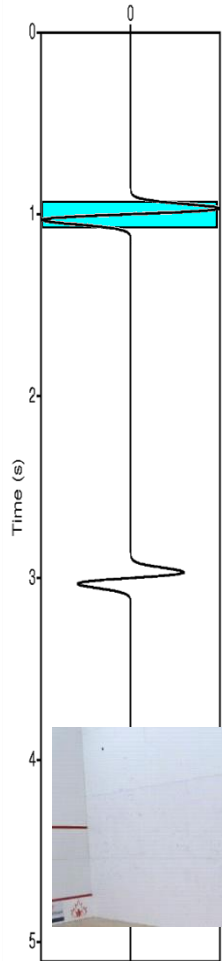
LSRTM



Qualitative model?
Reflectivity based on migration

Wavepath: sensitivity kernel

From transmission information (velocity analysis) ...

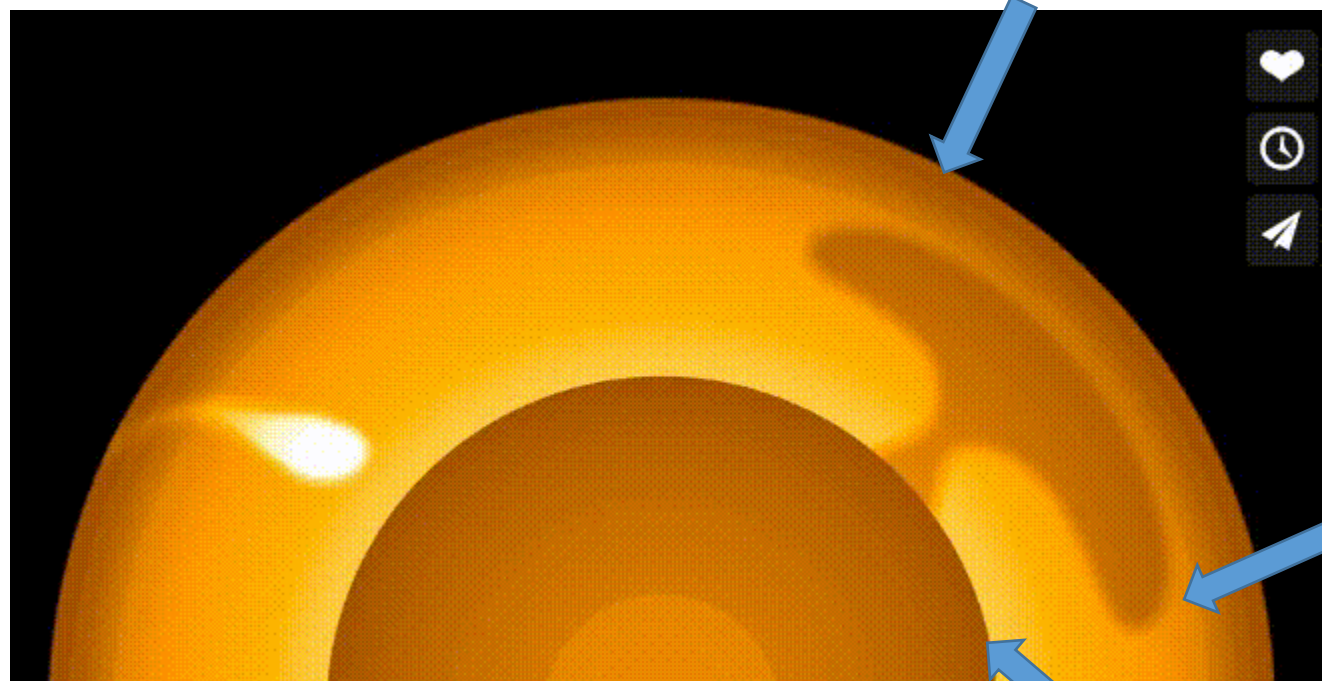


Inverse of K_z

... to reflection information (migration, fwi)

Crust and Transition zone (CMB) are two expected complex targets

Crust: combined transmission and reflection regimes and all intermediate regimes



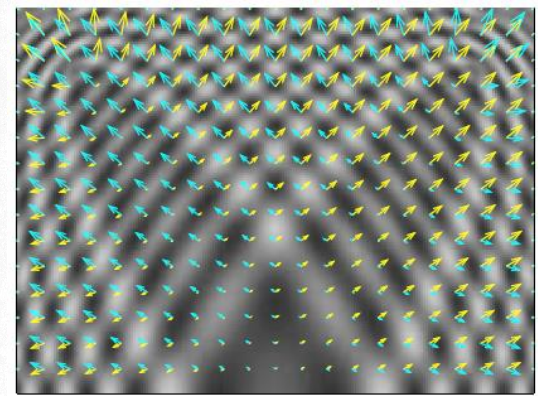
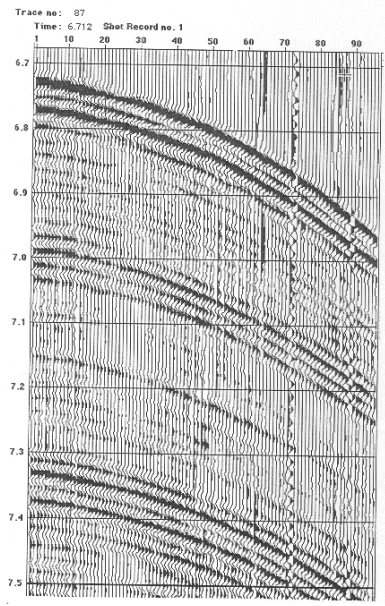
Mantle: only transmission regime for now on!

Movie from Stahler & Tarje

CMB: mainly probed by waves through transmission regime while expecting more complex interaction!

Macro-model reconstruction

Kinematic correction: NMO velocity



NMO velocity: estimation based on the identification of hyperbolae lines of reflections



Wave-matter interaction:
Transmission regime based on reflection events

Collecte des traces en points miroirs

Common source profile CSP

Common midpoint profile CMP

Offset X

PM

temps double

$\Delta T = T_x - T_0$

indicatrice

$$T_x^2 = T_0^2 + \frac{X^2}{V_{nmo}^2}$$

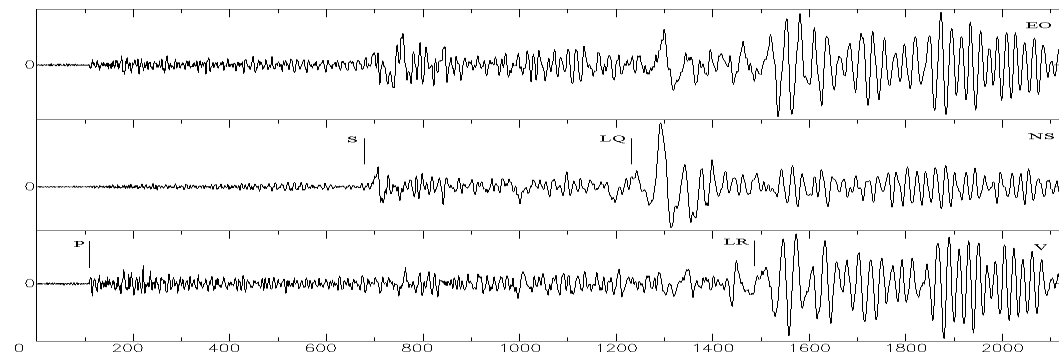
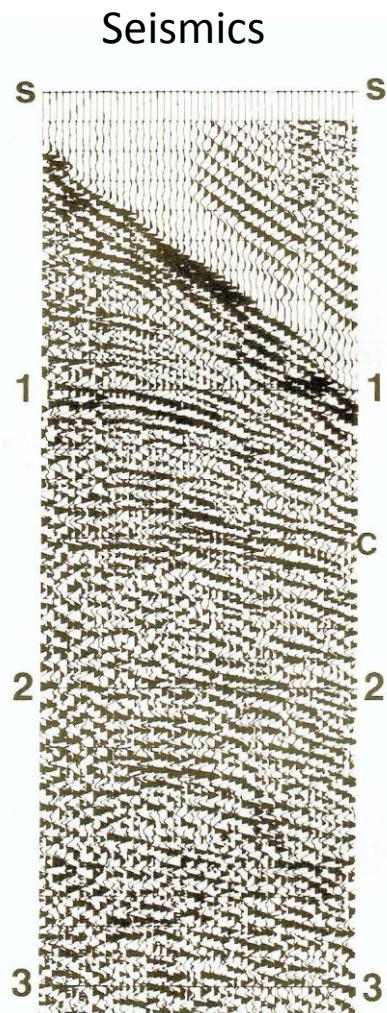
In the Earth

Three time scales (for ballistic waves)

Three spatial scales ...

Two options:

- a. **extracting specific observables (time/phase)**
- b. ***using records as they are (calibrated amplitude!)***

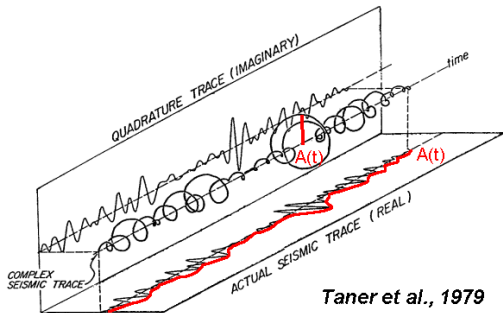


Seismology

Extracting specific observables from records

Kinematic observables (as for oceanographic and infrasonic tomographies)

- times or (instantaneous) phases (used in seismic time tomography)
- angles (back-azimuth and slowness angle: source, receiver, diffractor)
- time curvature (example of velocity analysis)



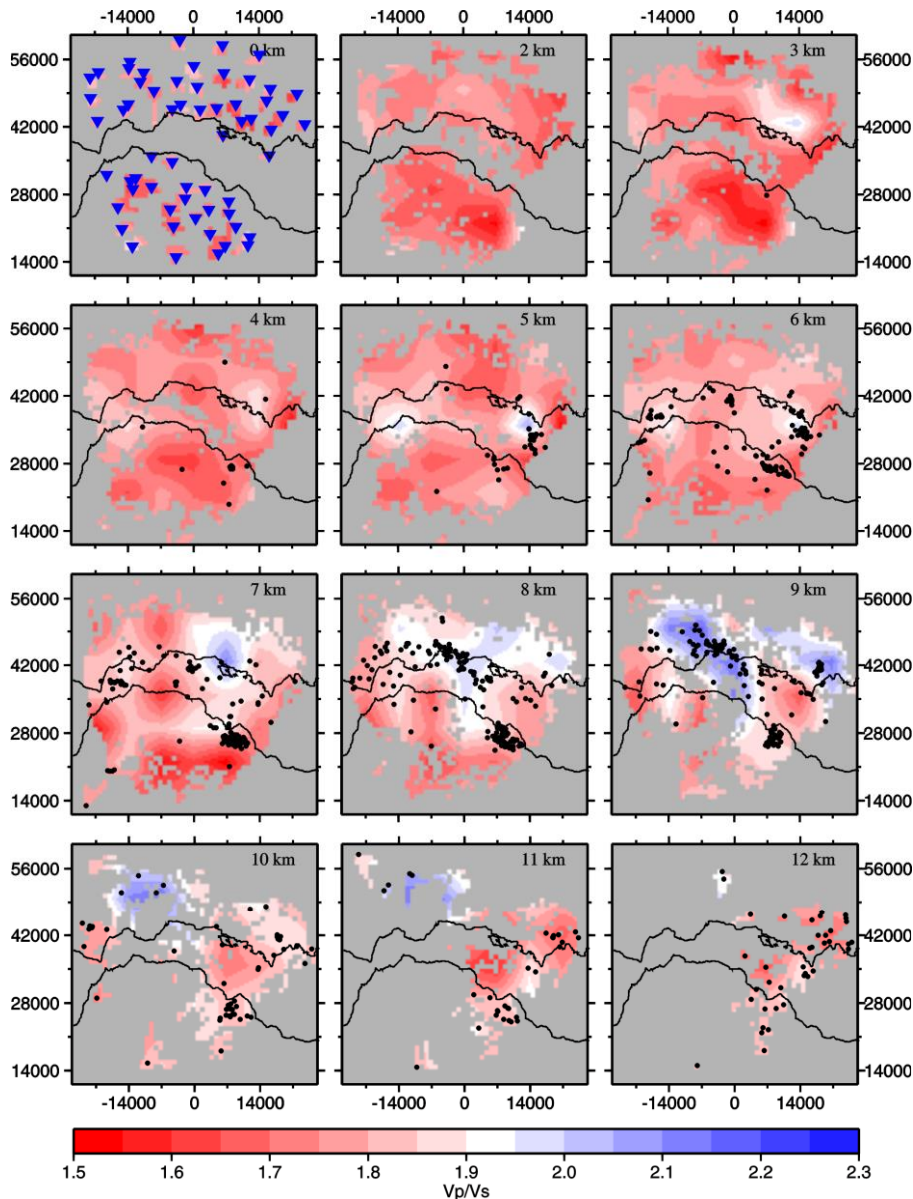
Correct physical description of time/phase data
Tools for predicting accurately these observations

Dynamic observables (option A+)

- (Instantaneous) envelopes
- amplitude decrease with offset (AVO)
- amplitude decrease with frequency (attenuation – broadband records)



Correct physical description of amplitude variation
Tools for predicting accurately these observations

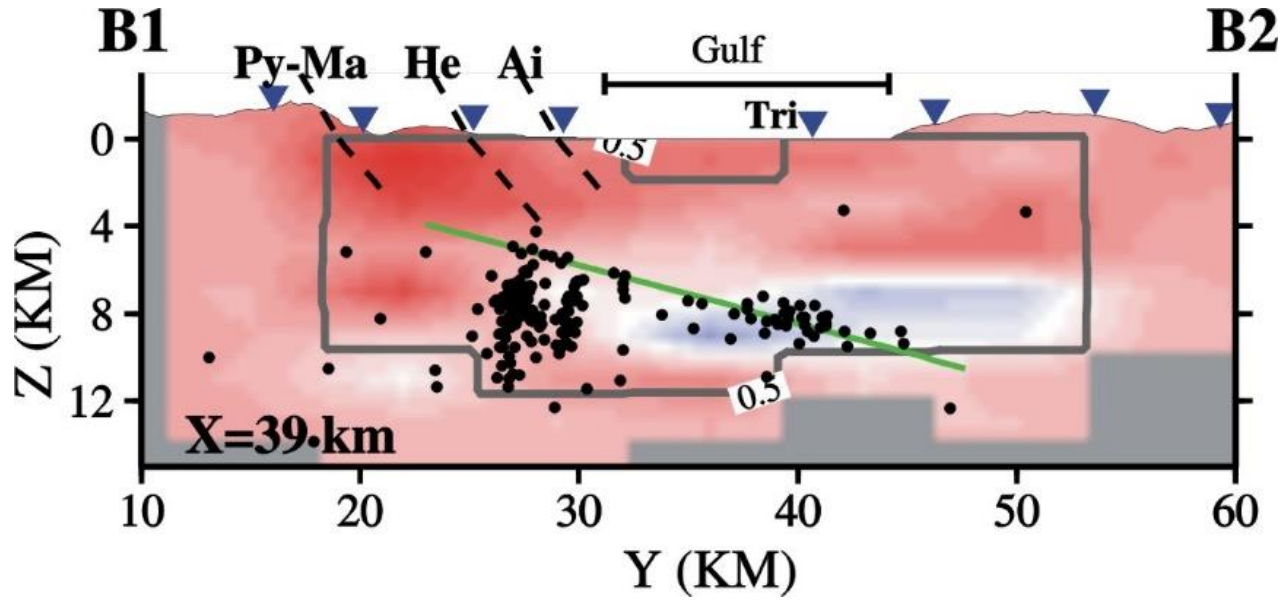


PhD of Diana Latorre (2004)

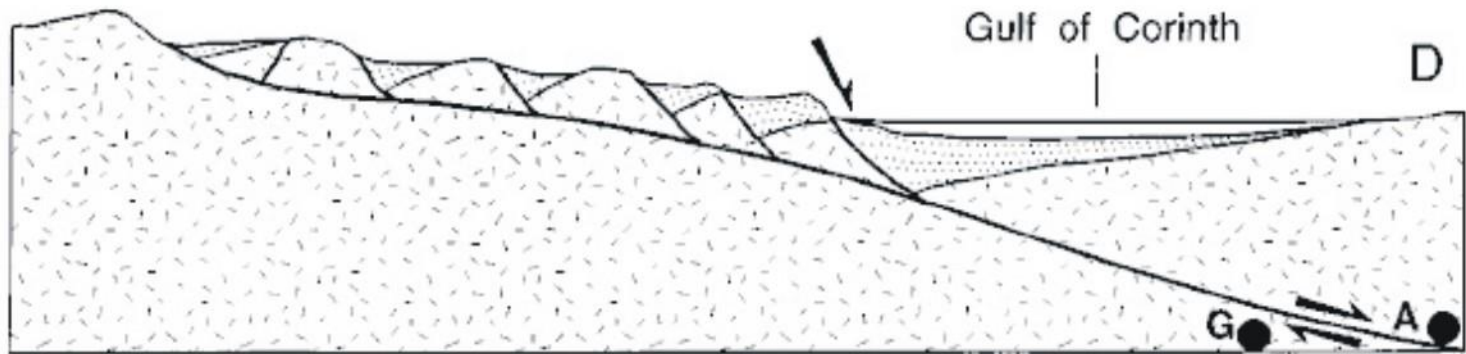
Vp/Vs image deduced from the Vp and Vs tomographic models

- Vp/Vs value may discriminates lithology.
- Variations depend on porosity and fluid saturation.

Fluid combined with soft material may help sliding a thick block over another one. The tilting mechanism is not mandatory (the extension factor β of oceanic margin)



?



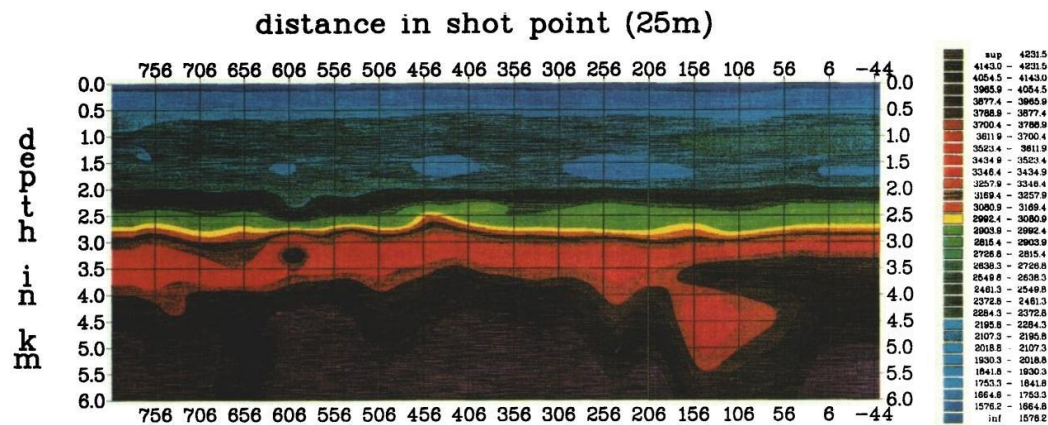


FIG. 8. Background velocity model for the Oseberg Field profile selected for asymptotic inversion. Velocities were obtained by standard velocity analysis provided to us together with the data set.

Macro-velocity model built from velocity analysis ...

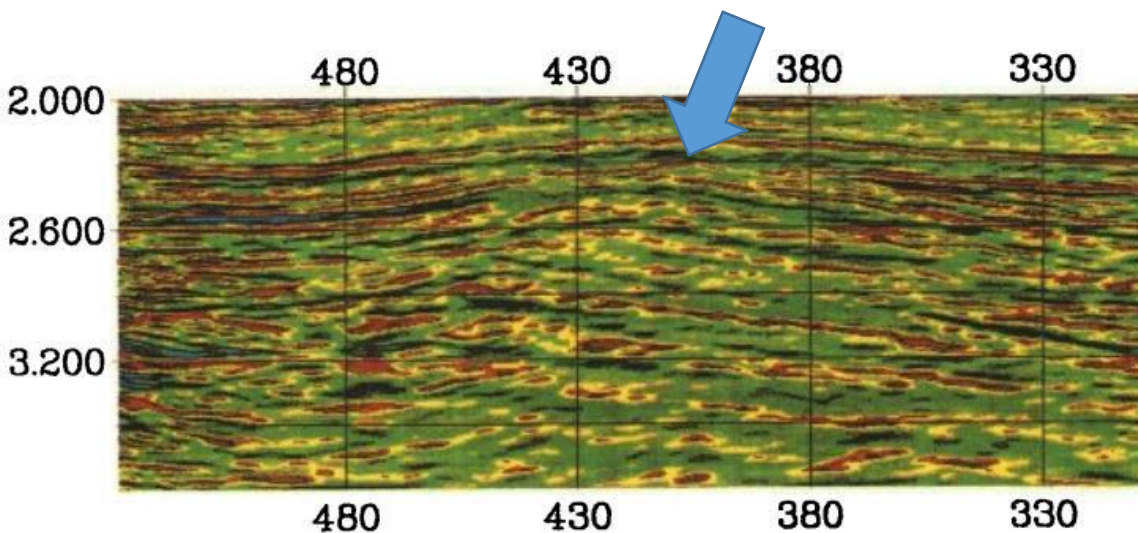
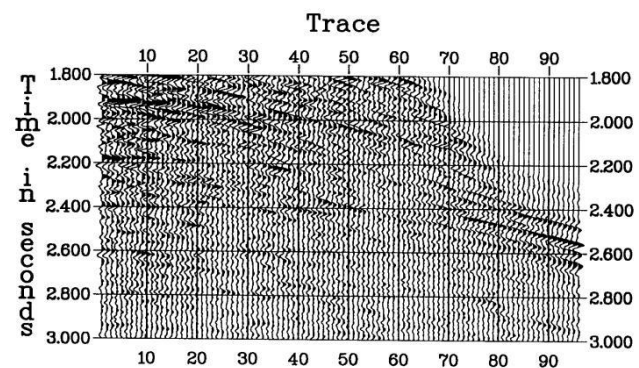
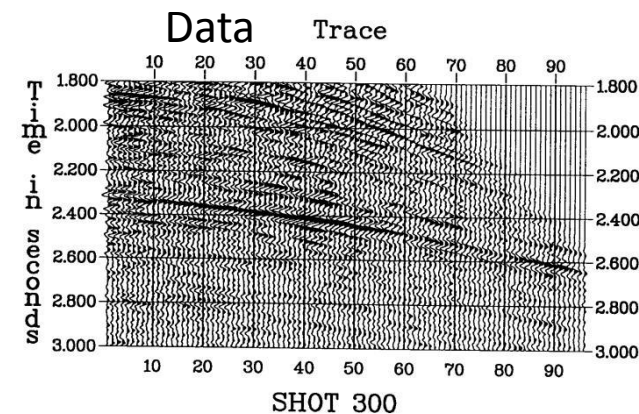


FIG. 10. Velocity perturbations inverted in the smaller target zone of the Oseberg profile; in different panels.

Lambaré et al (1992)



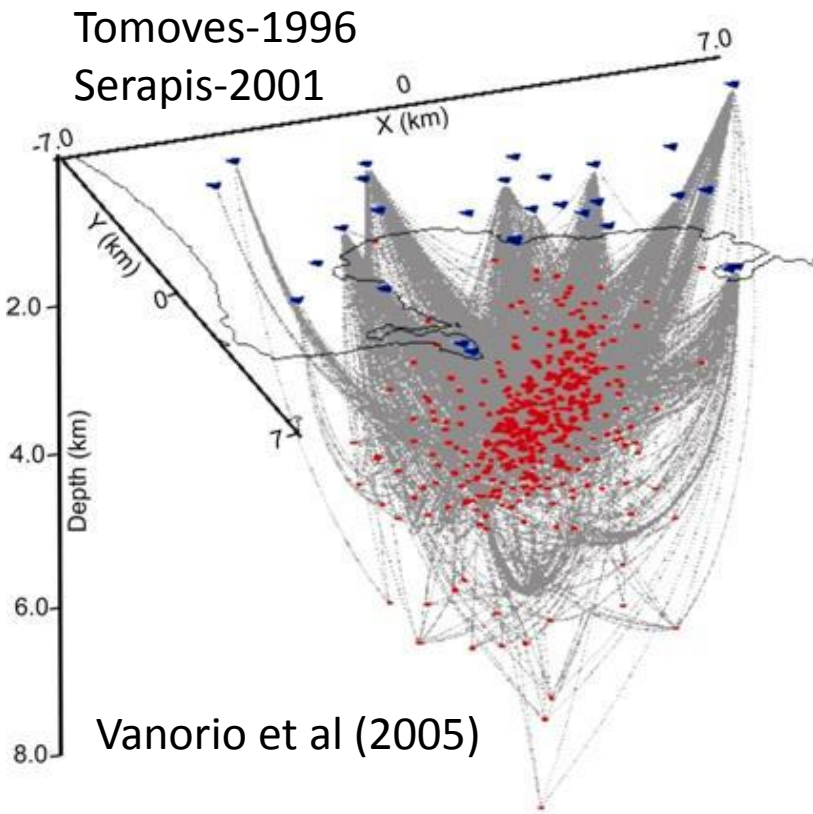
RESIDUALS SHOT 300 3d ITERATION

Energy has been removed ...

Puzzle of Napolitean volcanoes



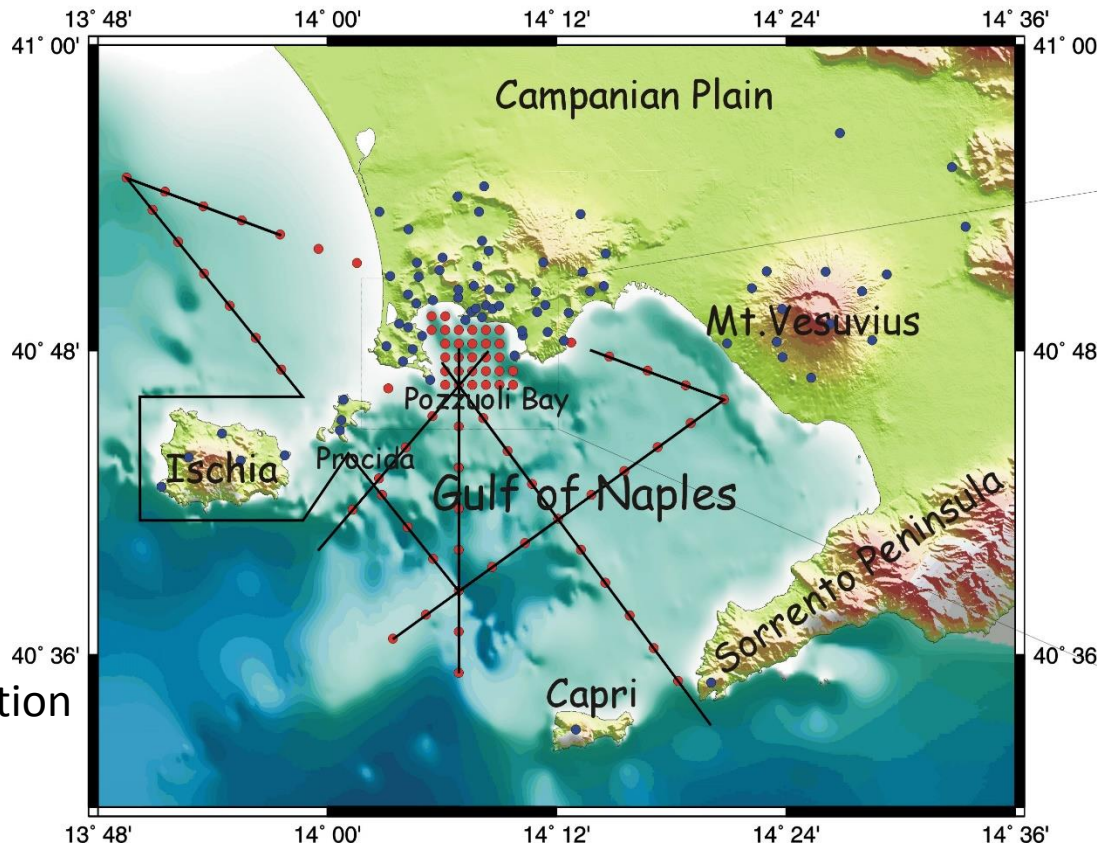
A long-term adventure
by Aldo Zollo ... and many people



The 1982-1984 bradyseismic crisis

Passive and active seismic experiments

Three attempts for amplitude interpretation
on this target: we partially succeed!



Travelttime tomography: still strongly based on 90's methodology (as far as I can see) ...
small computer requirement ...

Questions: Grand-daddy research

why images are so smooth!

how to use differential information (DD strategy)!

debates regarding theoretical features (understanding)!



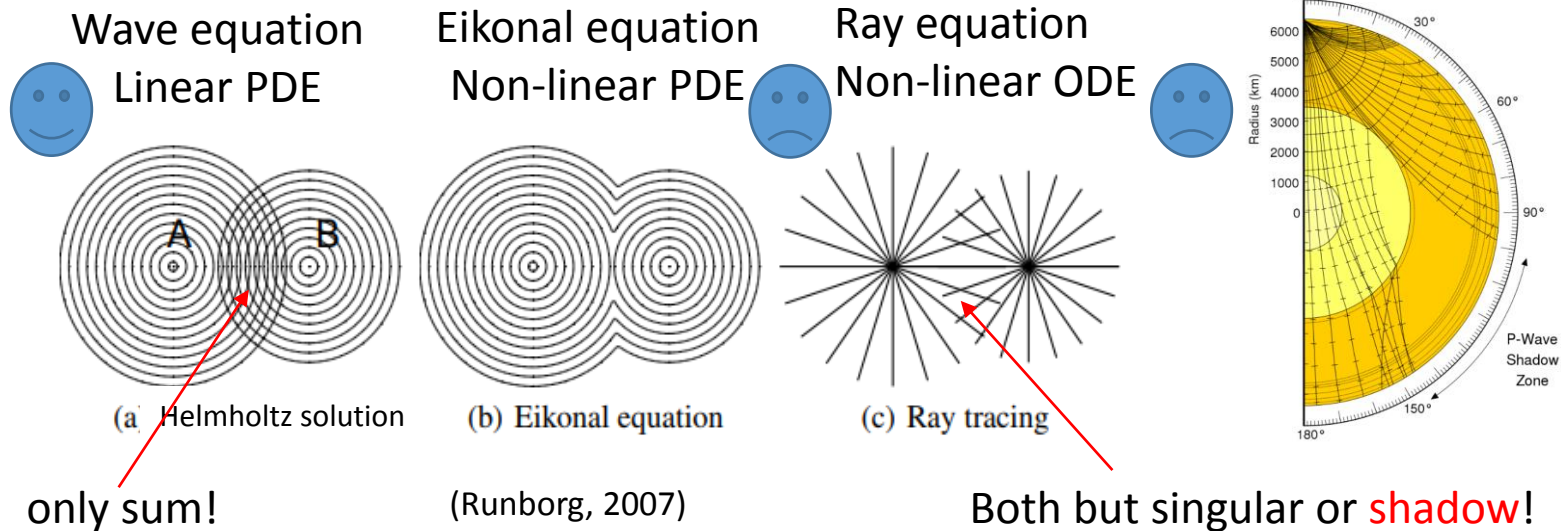
Full waveform inversion: intense search activity

many groups are involved ...

HPC implications ...

Fashionable research





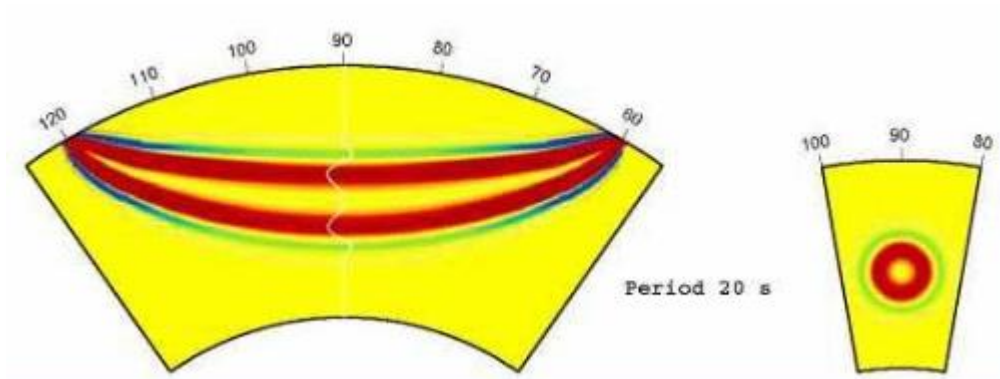
❑ **Traveltime tomography cannot consider beating waves!**

❑ **Full waveform inversion does!** as well as wave-equation traveltime tomography

No free lunch: not only computer resources but possible leakage between model parameters (speed versus attenuation, compression versus shear interpretation and so on ... anisotropy, acoustic approx.)

(Almomin & Biondi, 2012; Li & Chauris, 2017; Sun et al, 2017; Djebbi & Alkhalifah, 2020 ...)

<https://www.geoazur.fr/GLOBALSEIS/nolet/BDdiscussion.html>



(Dahlen et al, 2000; Dahlen & Nolet, 2005; Nolet, 2008)

Phase/time tomography based on Eikonal does not provide the so-called BD shape promoted by Dahlen & Nolet ... !!!

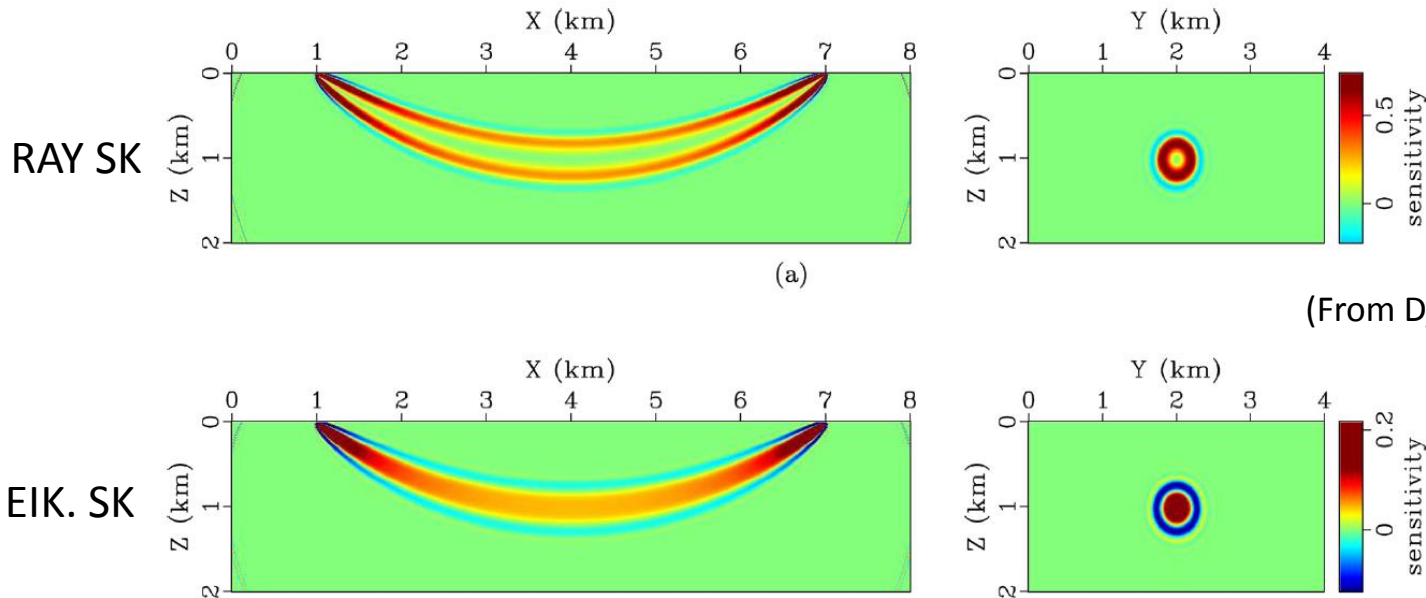
Such shape is related to **ray concept** (and not to the **time/phase concept** for which the zero-sensitivity along the ray does not exist)

Many authors have tried to understand better this theoretical contradiction ...

When considering phase information

DIFFRACTION EFFECT NOT INCLUDED IN RAY APPROACH

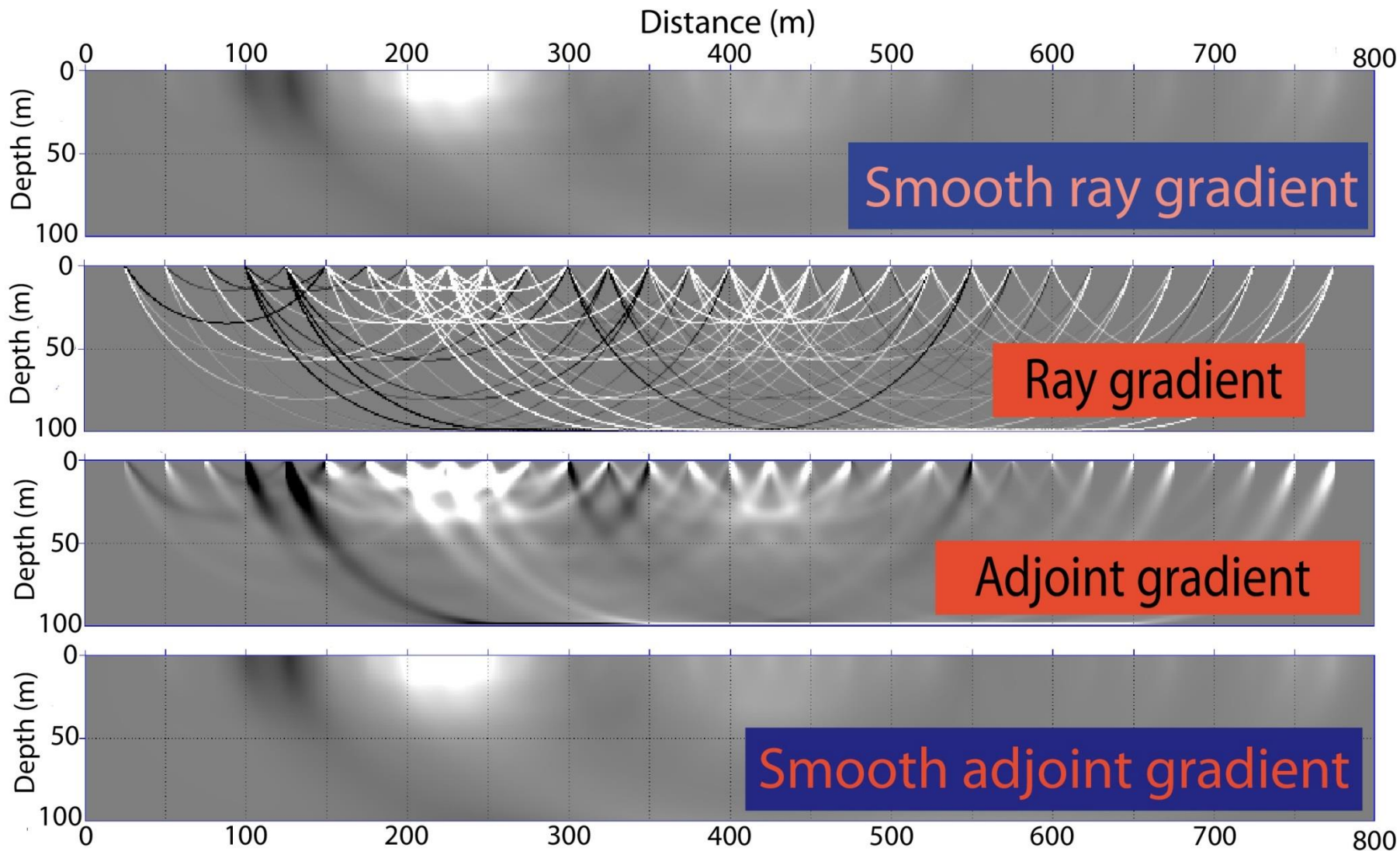
DIFFRACTION EFFECT CAN BE INCLUDED IN THE SO-CALLED EIKONAL APPROACH



Outcome of this debate: is it really important for applications?

My negative option (2014) has moved to a more positive one (2019) ...

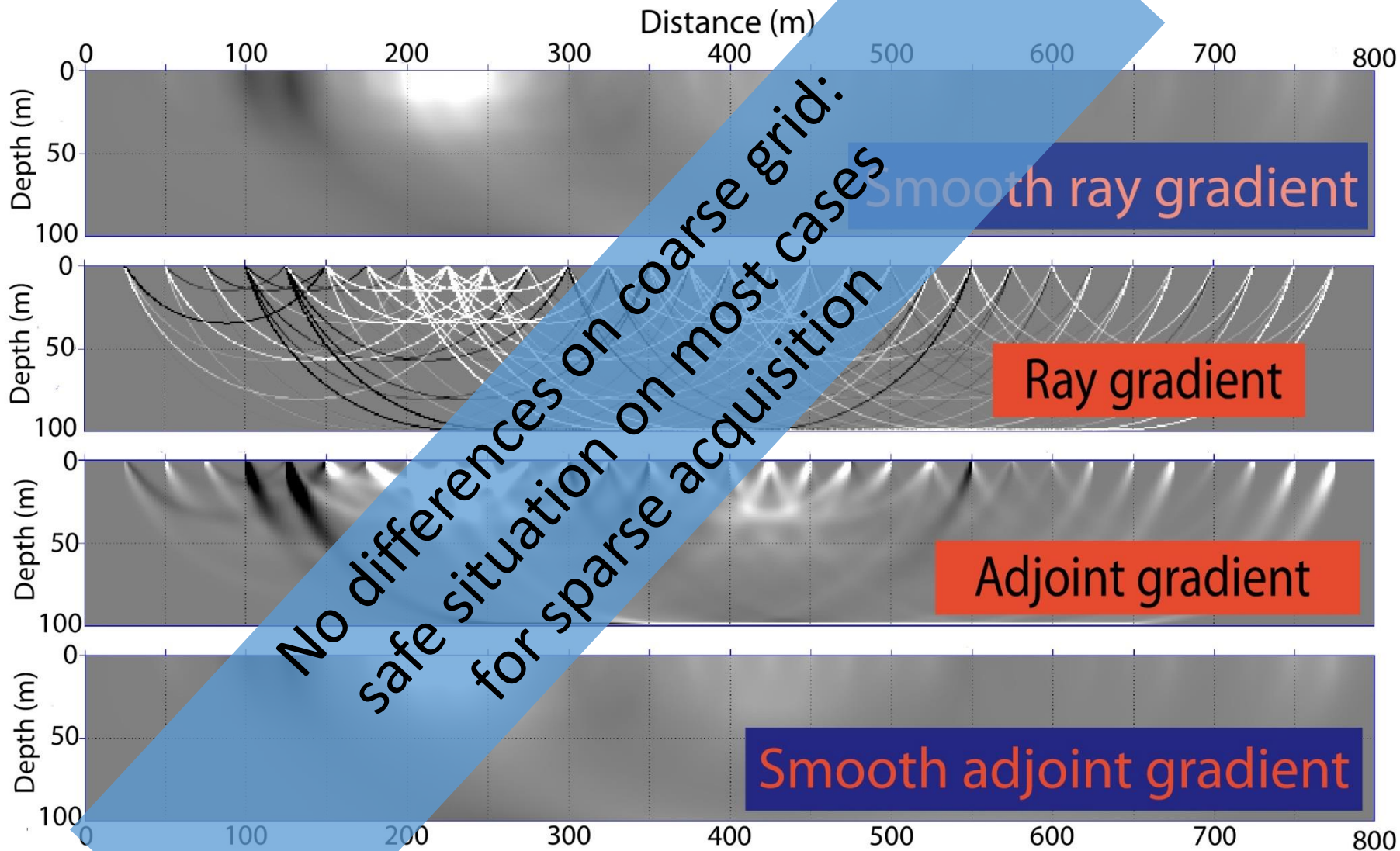
Projection of gradients on discrete model



Eikonal=Adjoint

(Bretaudiou et al, 2014)

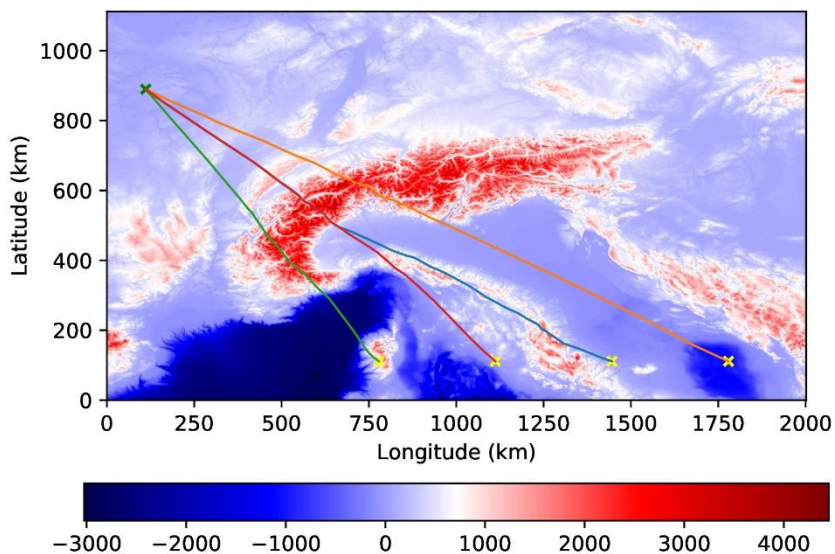
Projection of gradients on discrete model



(Bretaudeau et al, 2014)

Dense array: is it still true?

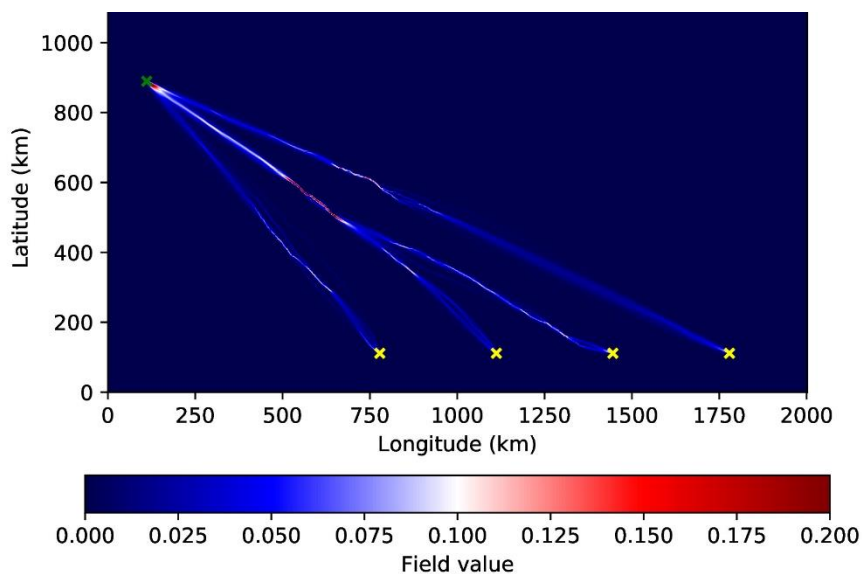
Ray SKs



Toy example:
topography taken as velocity perturbation
(mimicking HF pattern)

Rays: same value (ray length) of the sensitivity
(velocity impacts only the ray path itself)

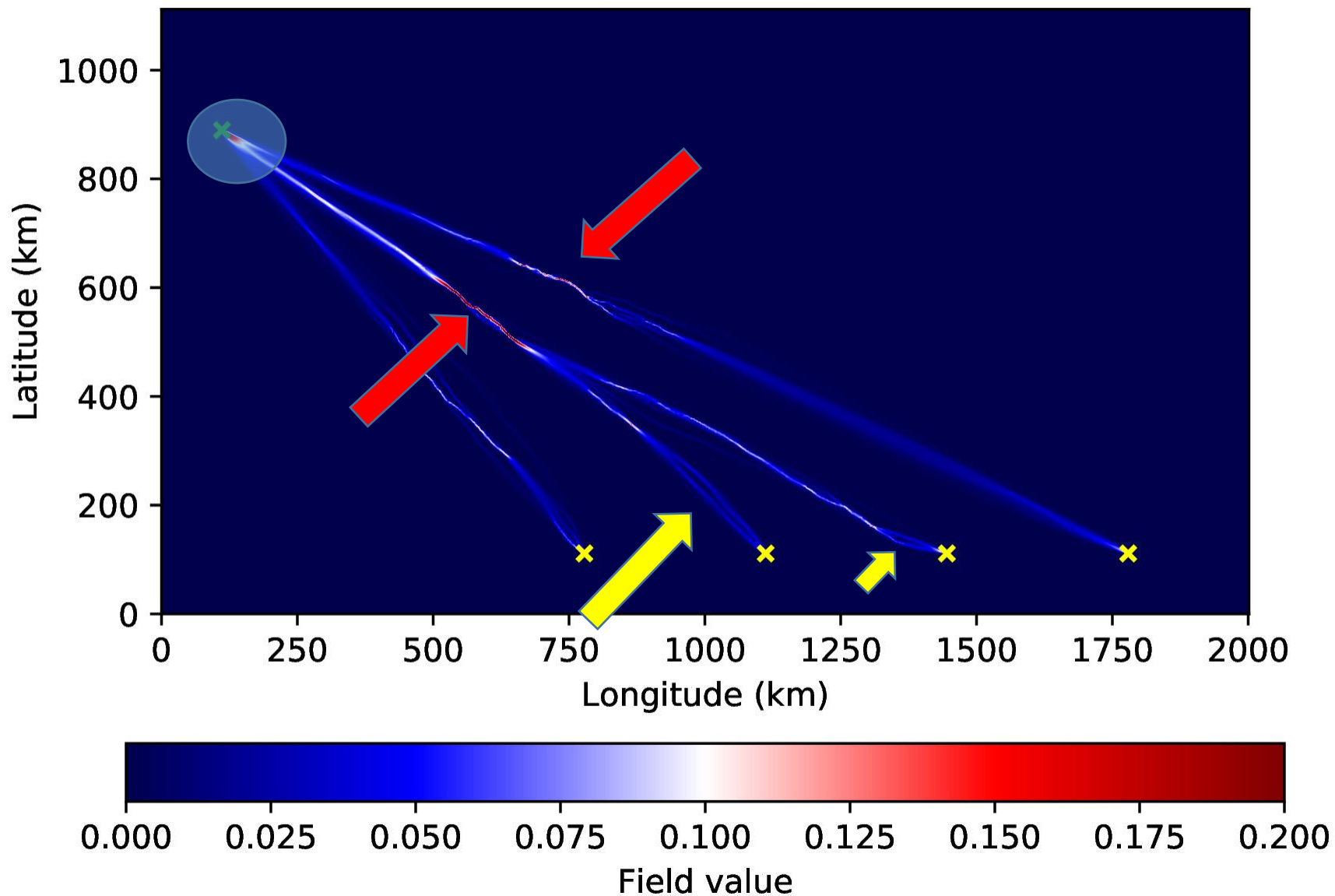
Adjoint= Σ Eikonal SKs



Eikonal SK exhibits more
complex pattern:
variable values along the
trajectory !!!

SKs: different values with possible different
paths before reaching the station

SKs

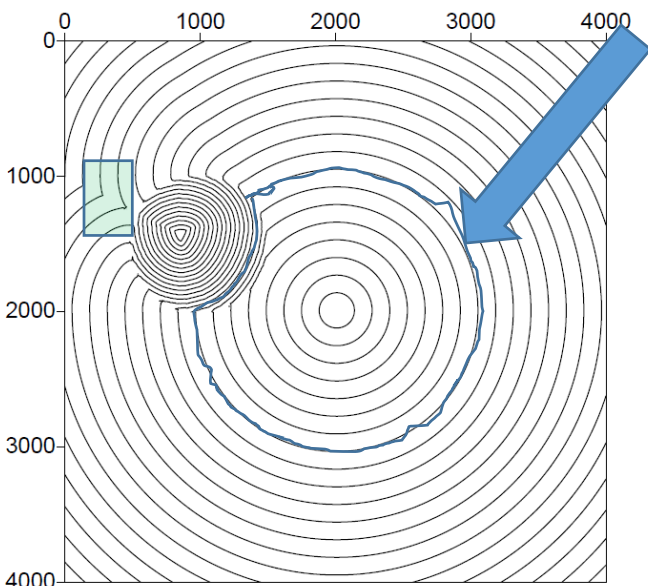
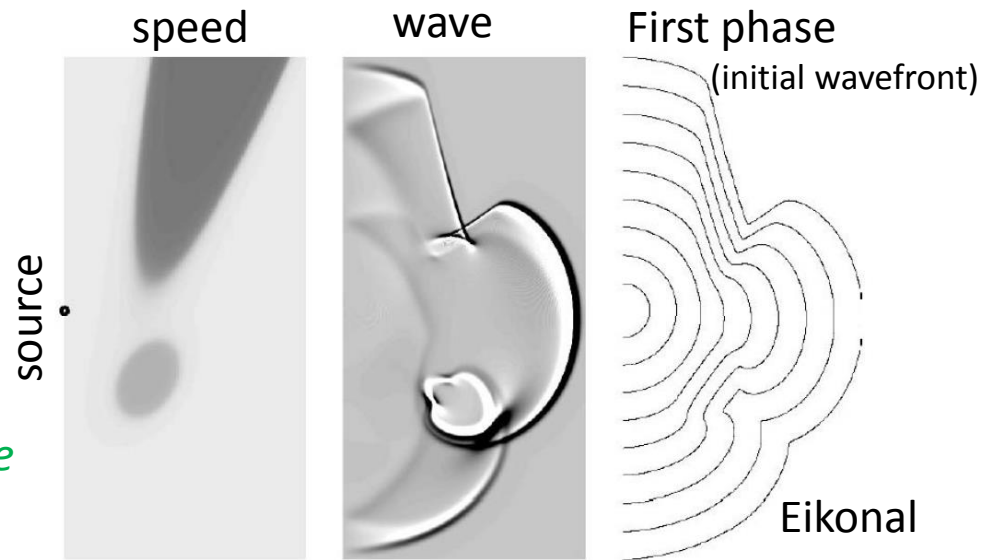


Eikonal: a more general concept?

A non-familiar interpretation of phases
(often associated to HF approximation)

Wave disturbance (field discontinuity)
valid for any media
single value and **always an answer**
observable: **continuous wavefronts**
but discontinuous derivative

(Runborg, 2007)



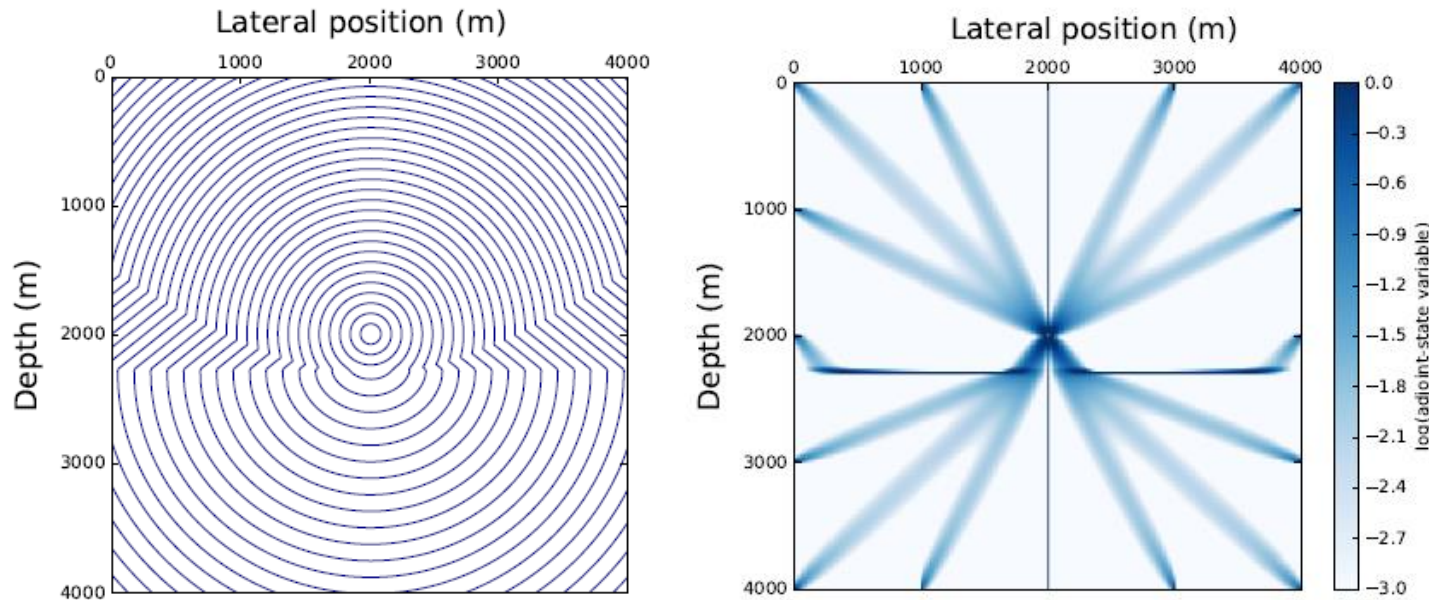
Line of same dancers
Wavefront: particles
moving a synchronized
way. They are in phase.



(Le Bouteiller, 2018)

Diffraction effects included ($u \propto \omega^{1/2}$)

(Le Bouteiller, 2019)



RD:
SK field

Reverse differentiation gives Eikonal sensitivity kernels:
where to insert velocity anomalies to match time data (sum of kernels over receivers) (Taillandier, 2009; Lelièvre et al, 2011; Tavakoli F. et al, 2017, 2019; Sambolian et al, 2019, 2021; Tong, 2021)

Sensitivity kernel defines zones of velocity perturbation affecting the time/phase at the receiver (regardless of the frequency content of waves!)

Forward differentiation (FD): slowness perturbation \rightarrow time perturbation

For a given point x , $\delta u(x) \rightarrow \delta T(\cdot)$ everywhere

Reverse differentiation (RD): time perturbation \rightarrow slowness perturbation

For a given point x , $\delta T(x) \rightarrow \delta u(\cdot)$ everywhere

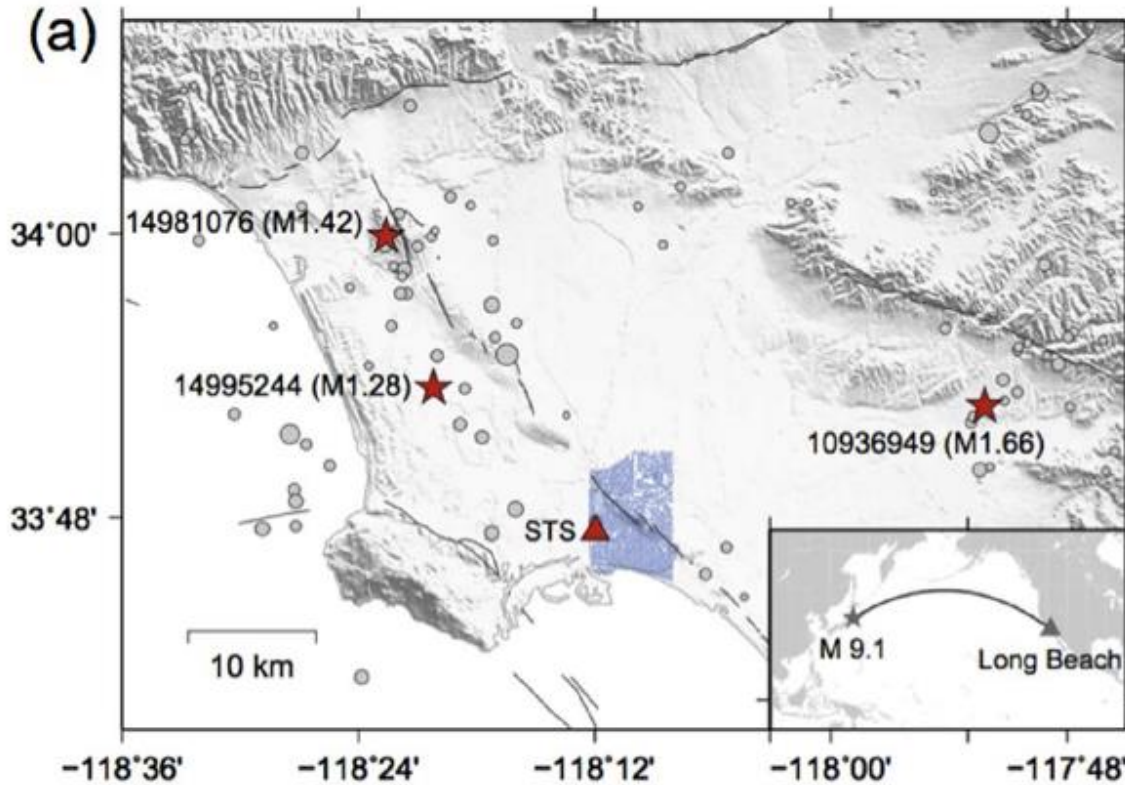


i.e. time residual at a given station

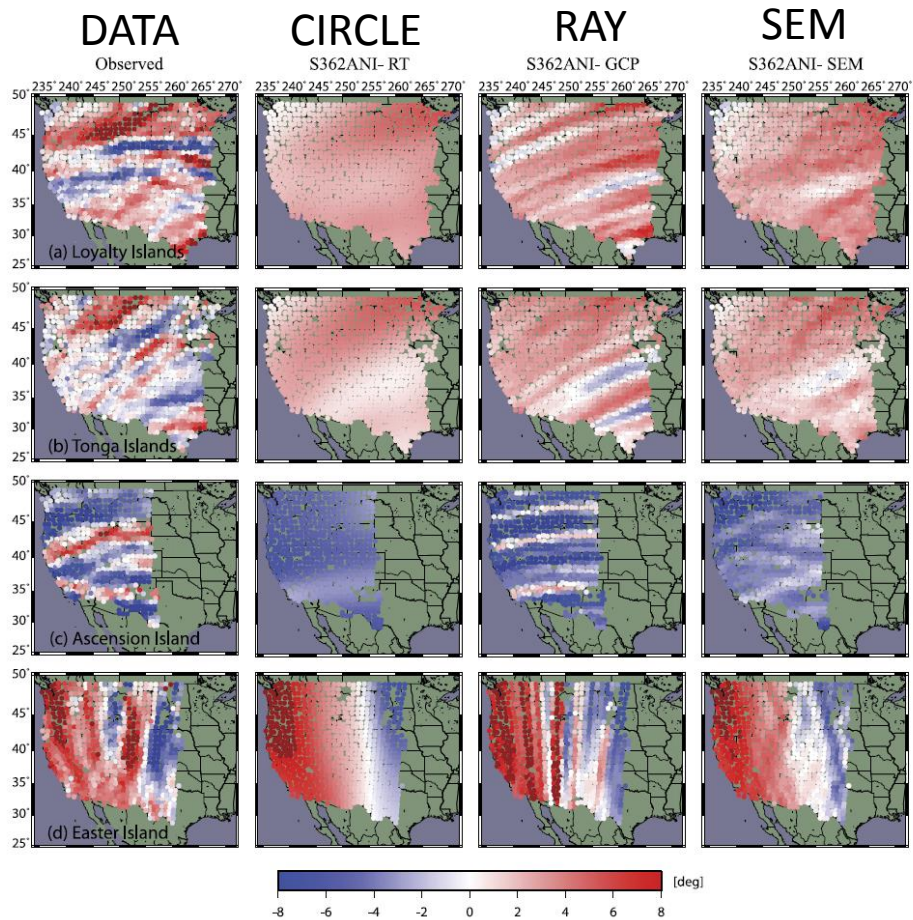
Ray linear relation

$$T(x, u + \varepsilon \delta u) = T(x, u) + \varepsilon \int_{\text{source}}^{\text{receiver}} \delta u \, dl$$

Locally, the ray sensitivity is the length of the ray inside the cell of the grid
only a geometrical value: arclength



How to take benefit of such density?



For 4 earthquakes, measured arrival angles across the USArray network & predicted angles for a S362ANI model using

- (1) great-circle approximation,
- (2) ray tracing and
- (3) SEM modeling (interference issue)

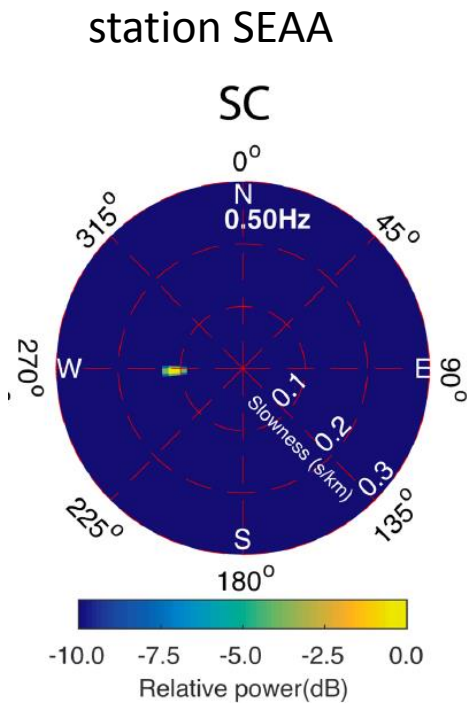
Observables have more information than those predicted by built Earth models

Figure 7 from Foster, Ekström & Hjörleifsdottir (EPSL, 2014)

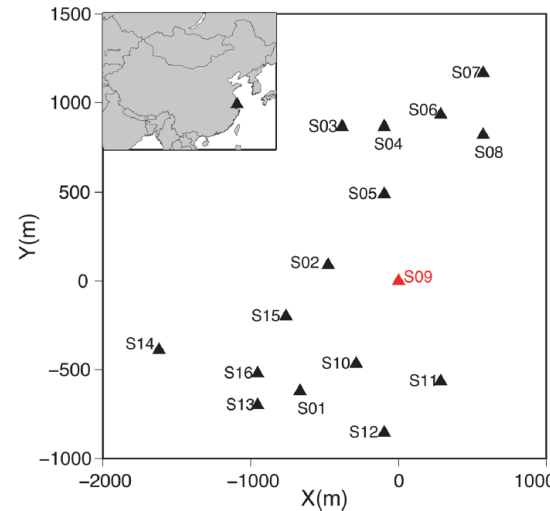
Improving tomographic approaches for honoring both traveltimes (or phases) and azimuthal (and/or slowness) measurements?

(Hu & Menke, 1992)
(Yanoskaya, 1996)

for Wenchuan 2008 earthquake
SEAA: sub-array of 16 sensors



ChineseArray



Azimuthal angle and apparent slowness

(from figure 16 of Hu, Zhang & Yu; GJI 2019)

Slowness=ray parameter=slope

Increasing density of stations offer us new observables such as slowness and back-azimuth observables aside traveltimes observables (nonlinear time series analysis)

Using dynamic observables (option A+)

Or

Using records as they are (option B)

- Deep understanding of the wave propagation physics (PDE?)
- High-quality data: reinforced acquisition protocol
- Accurate prediction of true wavefields (beyond waveforms)



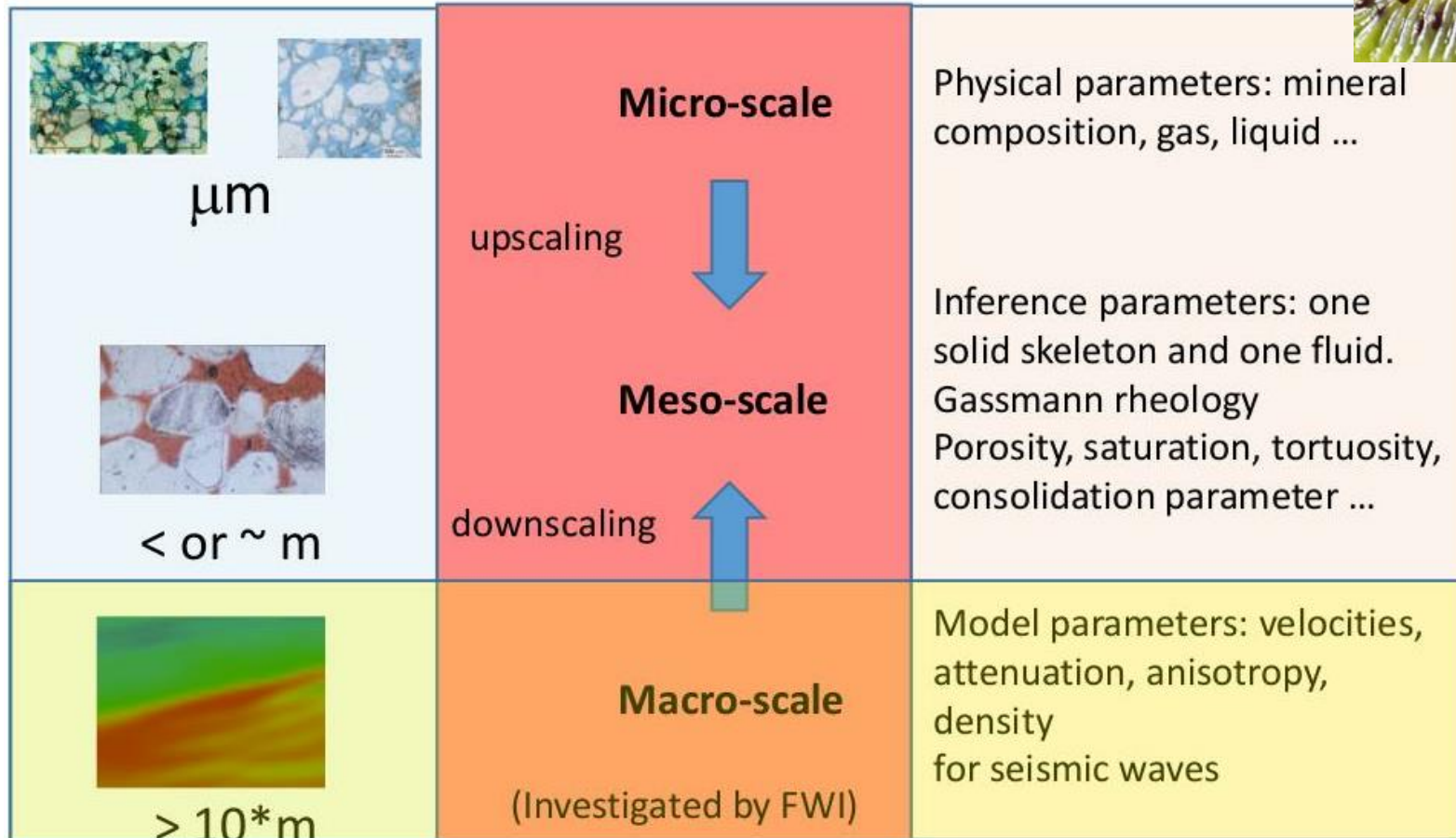
Another story, next time

- Migration: an efficient linear process with always an answer you may like it or not (interpreters are key experts). *(very robust to amplitude errors)*
- FWI: a costly non-linear process with often no answers! This makes students rather frustrated... *(amplitude influence)*
(easier life for interpreters 😊)
- FWI: when it works (needing not only smart students but high-quality data), potential quite interesting interpretation with physical quantities!



Final goal of deterministic seismic imaging
(not yet crust/reservoir characterization)

Model # Medium



From Dupuy's PhD (2011)

- ❖ Crust is a complex body where waves interact with matter with # illuminations
- ❖ Propagation in the crust may go beyond simple partial differential equation modeling.
- ❖ Kinematic observables: robust measurements
(improved quality, thanks to multiscale arrays!)
- ❖ Zeroth-order robust physics (accurate modeling of rather stable wave propagation quantities)

- ❖ Room for progress in kinematic tomography (**times**, **angles**, *time curvature*)
- ❖ Dense arrays: **new imaging** approaches (time reversal concept of Claerbout)!
- ❖ Mitigating the curse of computer resources opening the road to UQ

For kinematic tomography, we should proceed with the same effort than the one devoted to FWI since the beginning of this century

Could also help FWI to the right path of imaging (especially for multiple parameters): another story 😊

- ❖ Massive data
- ❖ Computer resources

How lucky are you?



How responsible you are!

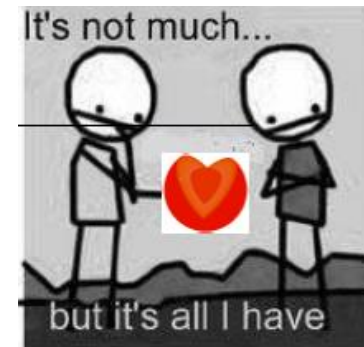
Smart (and well-trained) people for mastering these data with sophisticated and efficient approaches

For better knowledge of the shallow crust of the Earth (needed for our modern society)

Subsurface should not be our garbage can ... but the unique place for resilient storage!

Many thanks for this masterclass initiative giving me the opportunity to draw potential research topics on kinematic attributes tomography without disregarding dynamic attributes inversion ...

Many thanks to sponsors of the SEISCOPE consortium supporting FWI research,



*Thank you very much
For your attention*

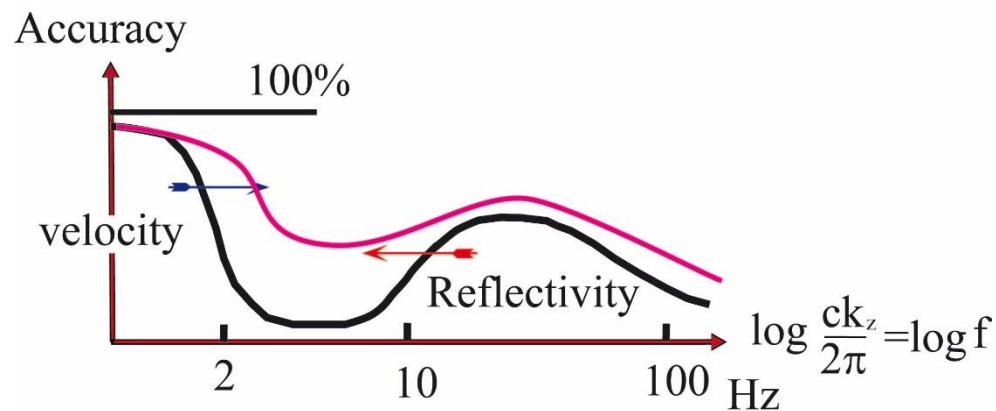
Questions?

Broadband sources Complete illumination of targets

The dream ...

Merging the velocity-model building and the migration into one integrated task when considering full waveforms

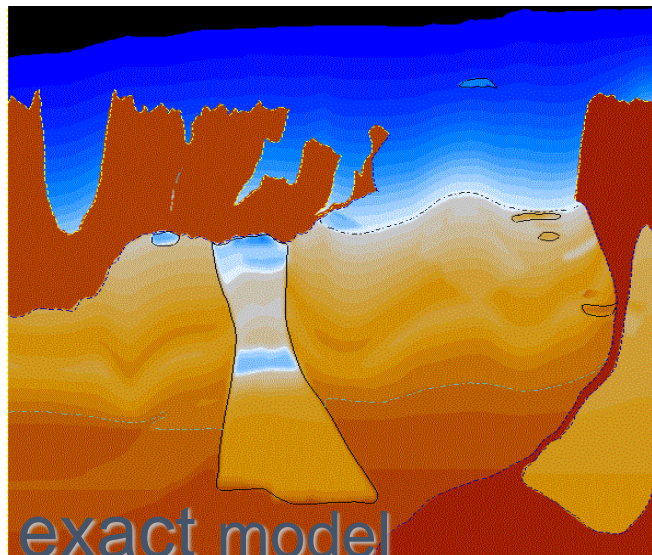
Continuous sampling of the wavenumber spectrum and filling the intermediate wavelength gap.



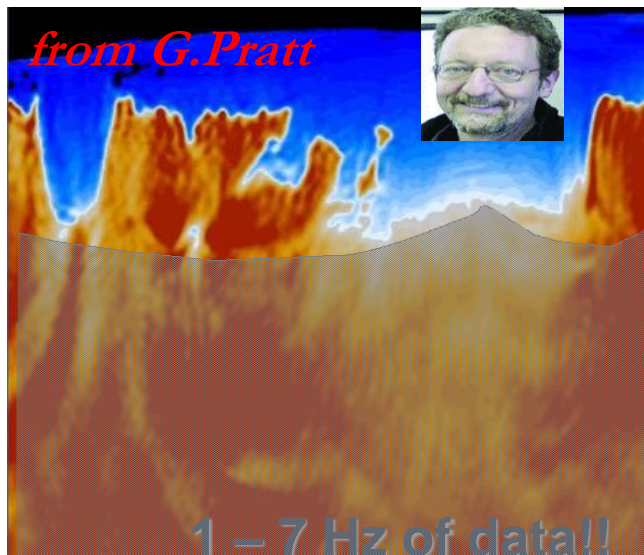
Tarantola (1984, 1987)

Amplitude: essential contribution

Low frequency information



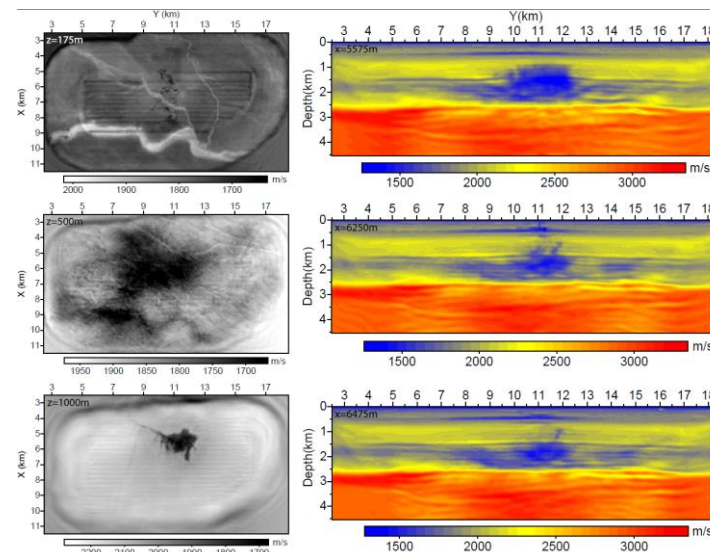
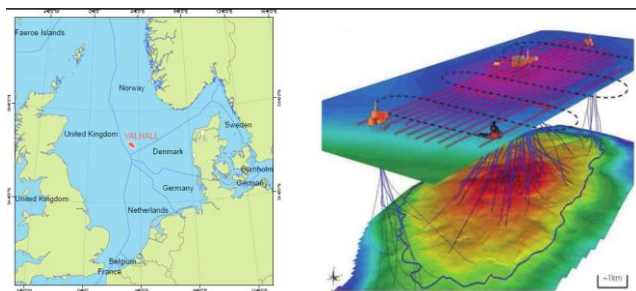
2D synthetic example



Pratt & Brenders (2004)

Benefit from waves, even when phases interfere

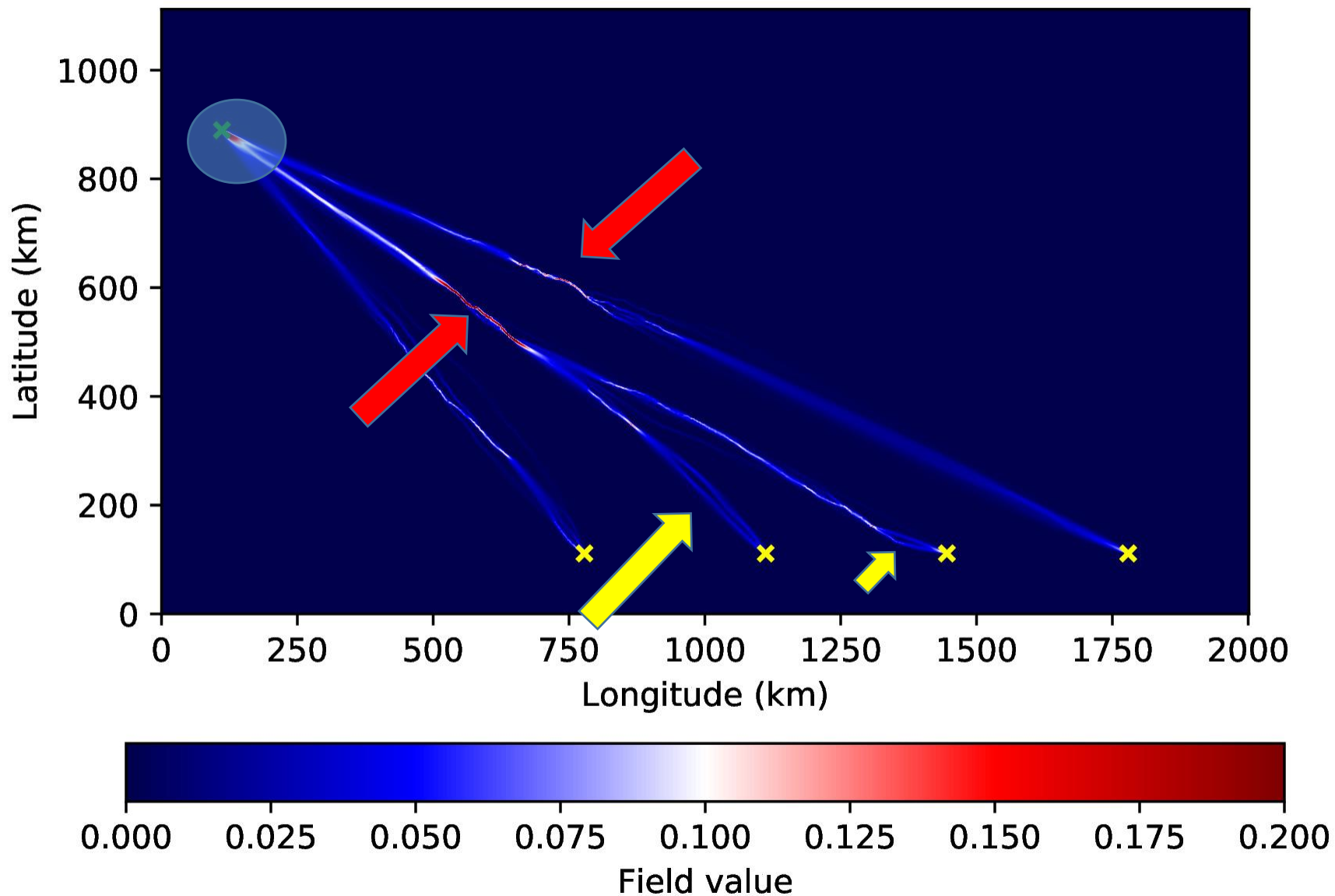
3D real example:
Valhall oil field



Operto et al. (2013, 2015)

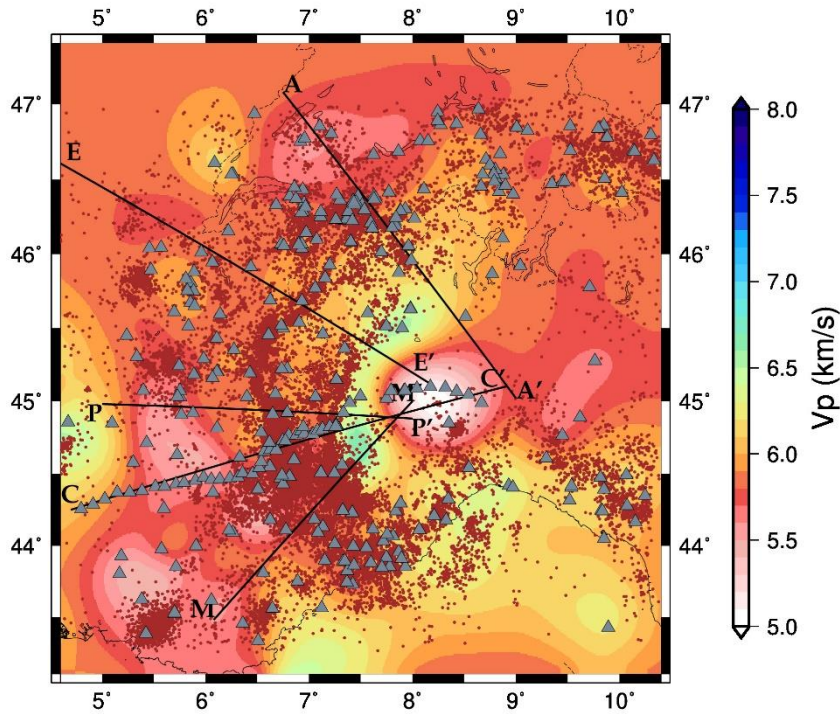
(Warner et al., 2008; Sirgue et al., 2010 ...)

SKs



Traveltime tomography: resolution!

Section 10 km



Alps Zone: Traveltime tomography

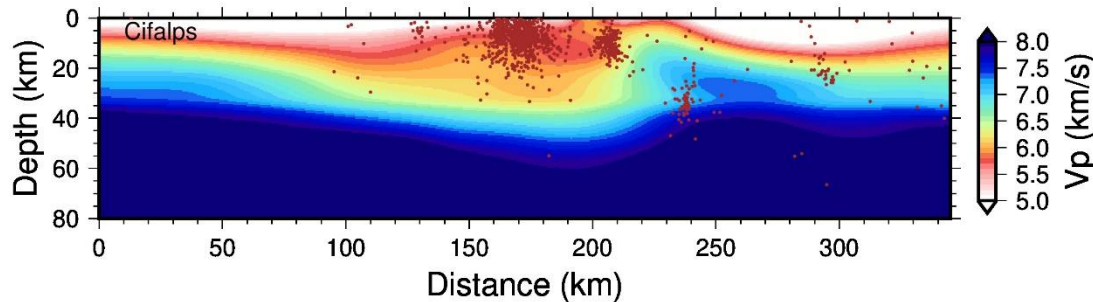
Seismic networks become denser and denser!

Picking strategy: manual, automatic?

Other kinematic observables:
slowness and azimuth angles?

or more!

CIFALPS profile (CC')



I am running now this 90's tool
over 1 500 000 picks coming
from 30 years of data on a PC
(big thanks to many people)
for these manual picks)

Travel time tomography: we need

- Eikonal solver (time; non-linear PDE)
- RD solver (time SK; linear PDE) (often approximated by rays!)

« Slope » tomography (based on time and time gradient): we need

- Eikonal solver (time; non-linear PDE)
- RD solver (time SK; linear PDE)
- *FD solver (time gradient; linear PDE)*
- *RD solver (time gradient SK; linear PDE)*

90's tentatives got
poor impacts...

sometimes known as polarization tomography
using backazimuth measurements (arrays!)

Automatic strategy was mandatory for industrial datasets (80's!)

Massive Dense volumetric time & slope picking strategy
for reflection traveltimes tomography (beyond velocity analysis)

Automatic strategy is fastly percolated into academic seismological community with increasing dense networks

From **manual strategy** (SEISAN software?) to automatic one (but often single-station with xC components) picking strategy (EarthWorm software?) for academic datasets (00's)

Local arrays with few sensors (at least one 3C and many 1C):

9C = 3C + 6 1C seems to become a standard in seismic microzoning

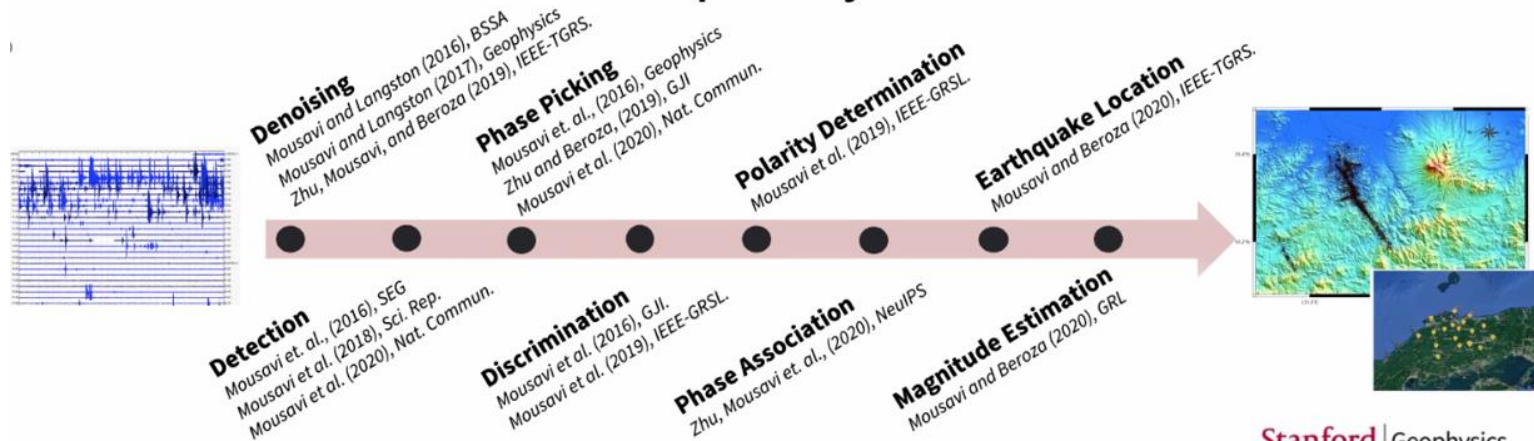
Very exciting novel methods based on unsupervised or supervised machine-learning approaches

Summary

Deep Learning + Deep Thinking

can help us to process seismic data ...

more effectively
more efficiently
more precisely



From a WS at LANL (2021)

S.M. Mousavi & G. Beroza (ArXiv:1912.01144v1, 2019)

Pr. A. Tarantola

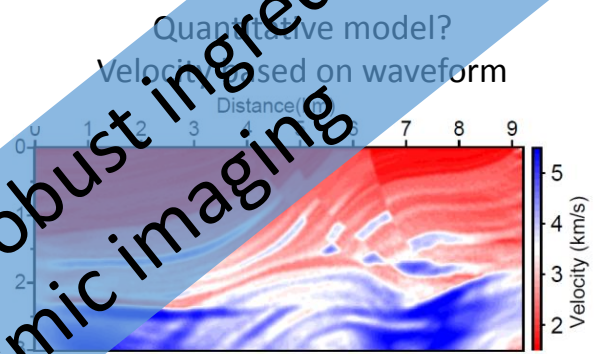
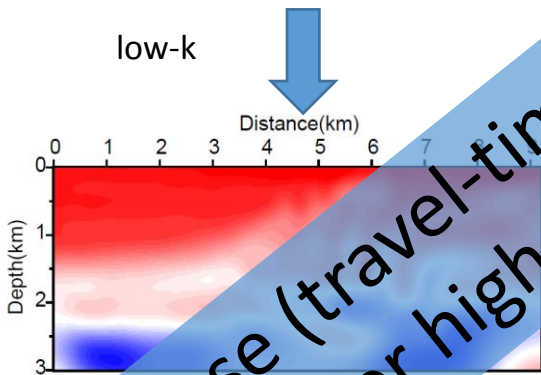
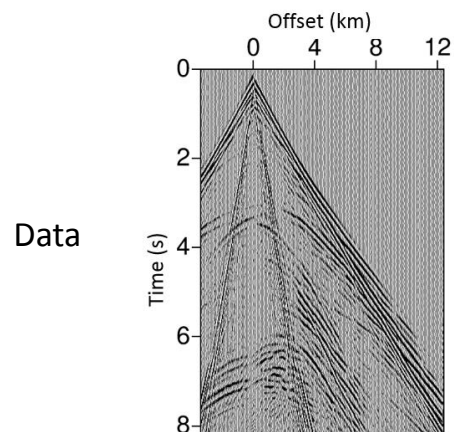


1st paper on FWI: Geophysics, 1984

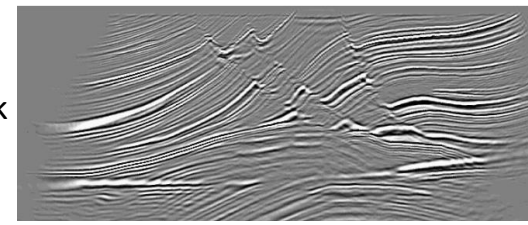
Full Waveform Inversion FWI

No prior identification ->
No prior scale separation!

Phase (traveltime for non-dispersive wave) drives imaging



LSRTM



Phase (travel-time) is still the robust ingredient for high-resolution seismic imaging

Macro model based on velocity analysis



RTM
or
PDM

Qualitative model?
Reflectivity based on migration

ACRONYMS FOR FWI VARIATIONS: AWI, WRI, TFWI, EWI, JFWI, MBTT, AEWI ...

Using dynamic observables Or Using records as they are

Acquisition design (source & receiver positioning & orientation?)

Time synchronization (GPS now! Clock precision < 1 ms \rightarrow 0.1ms)



Amplitude calibration (instrument!; installation coupling!)

relative values more than absolute values (still not differential)

(wavefield \rightarrow waveform) relaxes installation & experimental procedure !

Source knowledge (source signature; radiation pattern)

quite challenging for natural sources (earthquakes: focal mechanism)

still difficult for active sources (shots, blasts: source signal $s(t)$)

(wavefield \rightarrow waveform) mitigates these issues

Deep understanding of the physics of wave propagation

Tools of accurate prediction of waveforms (wavefields!)