

Seismic Anisotropy

Monitoring of Seismogenic and Volcanic zones

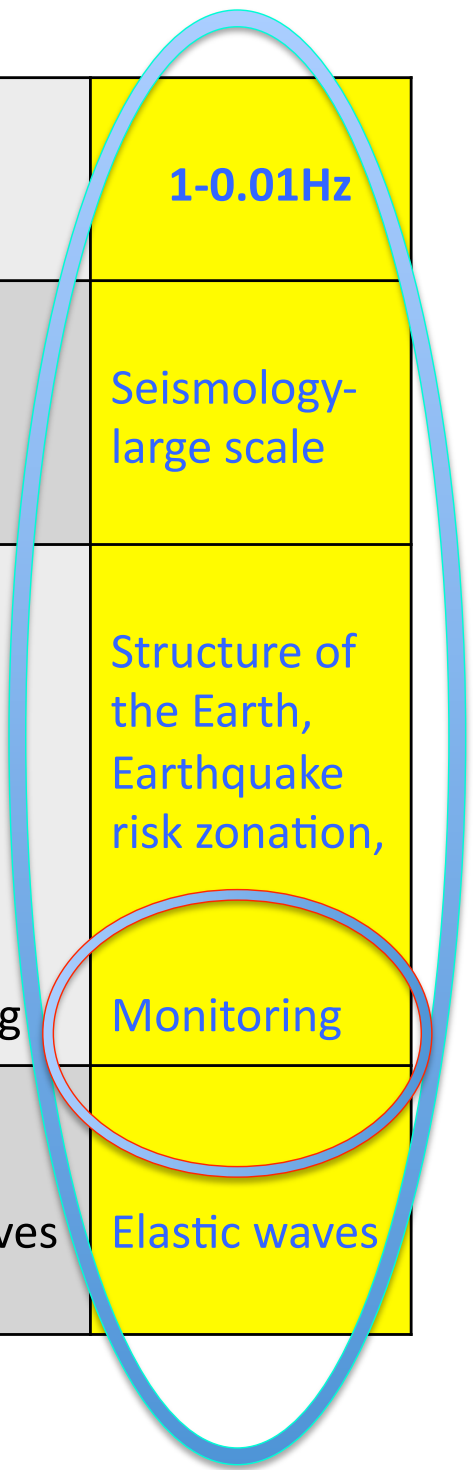
Jean-Paul Montagner

Maria Saade, Philippe Roux, Florent Brenguier,
Edouard Kaminski, Stéphanie Durand, Yanick Ricard,
Kohtaro Araragi, Yosuke Aoki, Paul Cupillard

...

IPG-Paris, ISTerre-Grenoble, U. Nancy, ENS-Lyon, ERI-Tokyo

Frequency	Mhz-kHz	10000-100Hz	100-10Hz	10Hz-1Hz	1-0.01Hz
Domain	Laboratory acoustics	Underwater acoustics	Shallow seismic imaging	Seismic imaging	Seismology-large scale
Applications	NDT Monitoring	Tomography Source detection	Structure of shallow layers Geotechnical applications, land slides Monitoring	Natural resources Natural hazards Monitoring	Structure of the Earth, Earthquake risk zonation, Monitoring
Wave type	Acoustic/elastic waves	Acoustic waves	Elastic waves	Elastic waves	Elastic waves





OUTLINE

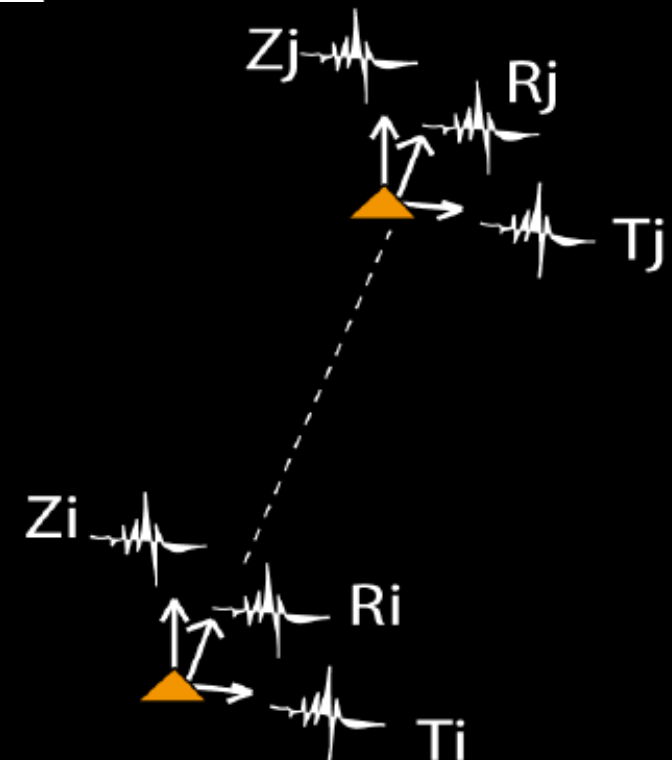
- **Data driven field: Seismic Data**
- **Cross-correlation tensor: Seismic Anisotropy?**
- **Seismic Anisotropy: many processes**

- **Scientific Issues:**
 - 3D- Anisotropic Structure of the Earth
Upper mantle – LAB
 - **Seismic monitoring:**
Temporal changes of anisotropy in
 - seismogenic zones (Northern Honshu – Japan)
 - volcanic zones (Mt Fuji)

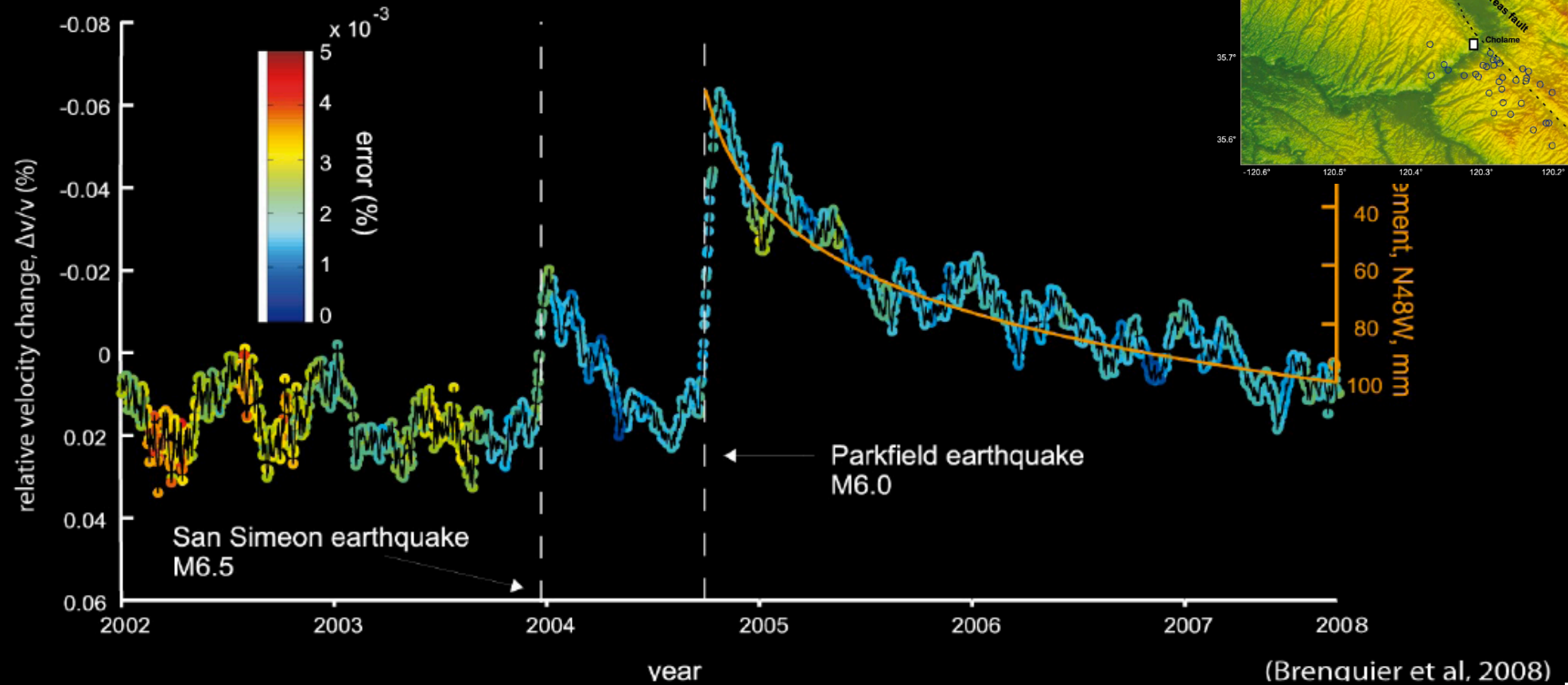
Cross-correlation tensor C_{ij}
 for 2 stations i, j and 3 components k, l
 Seismic signals $S_{ik}(t), S_{jl}(t)$

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT

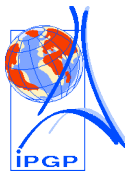


Parkfield High Resolution Seismic Network (HRSN)

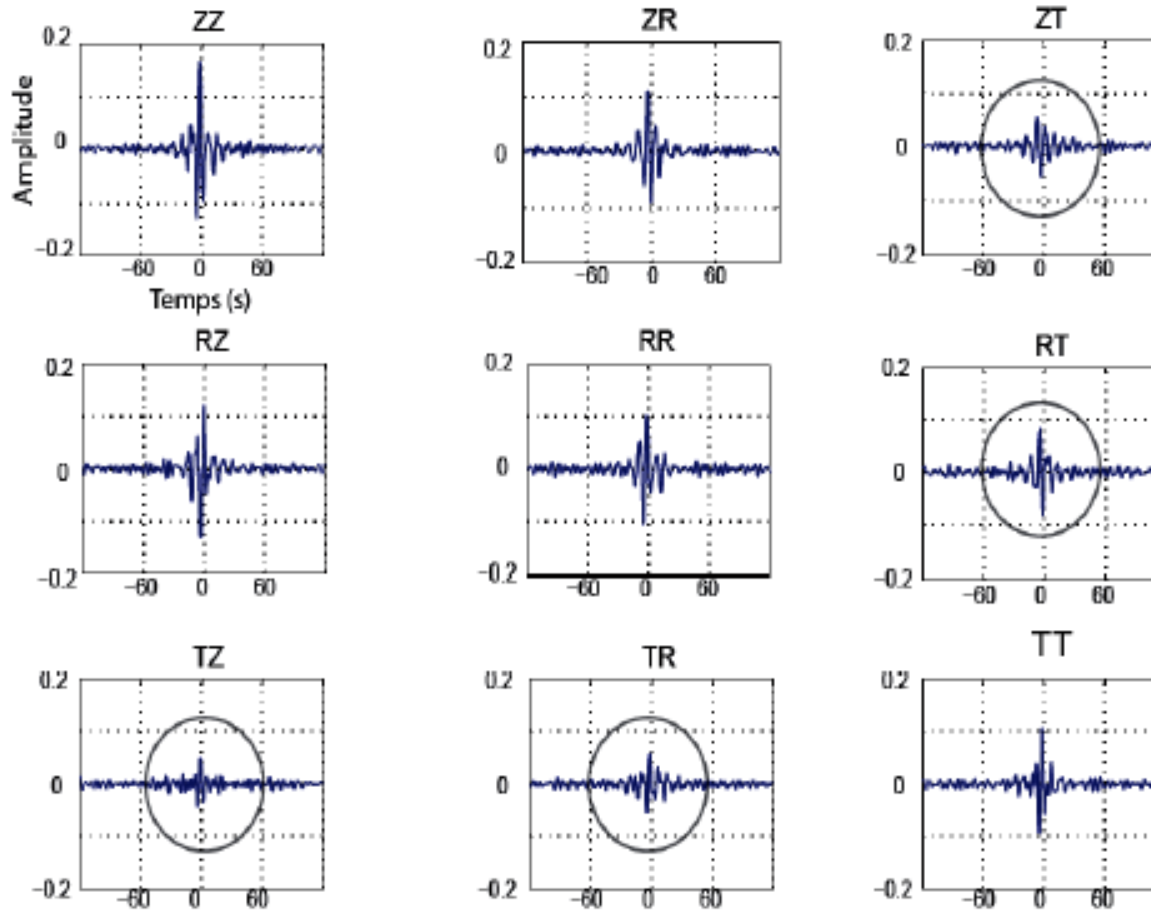


ZZ: Co-seismic and post-seismic relative velocity change
(Brenquier et al., 2008)

Example of cross-correlation tensor



Parkfield HRSN – Stack (30days)

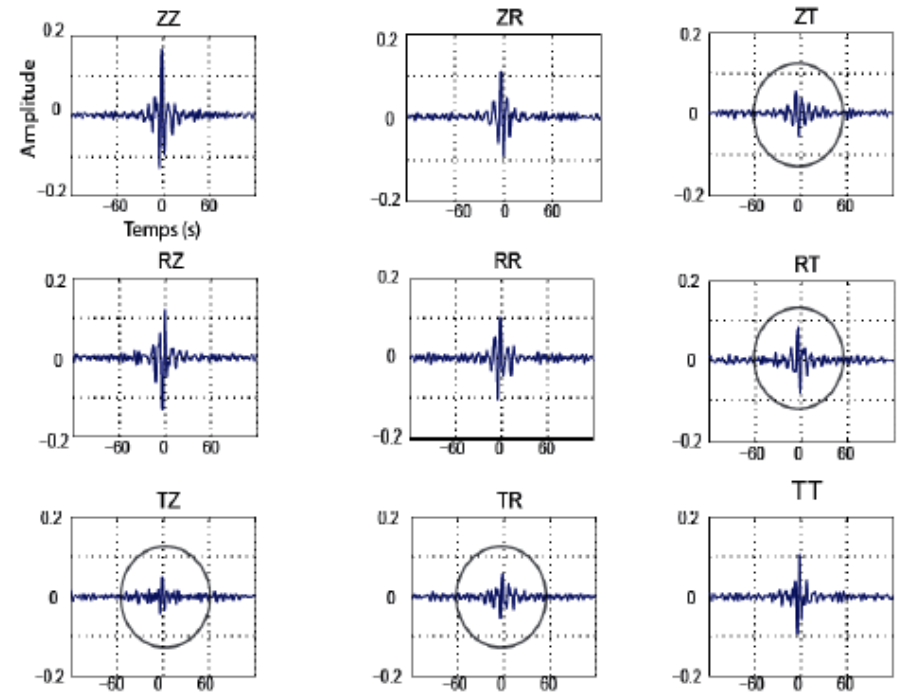


TZ, TR, ZT, RT \neq 0

How can we explain the off-diagonal terms of the cross-correlation tensor?

TZ, TR, ZT, RT \neq 0

- Non uniform distribution of seismic noise sources?
- Lateral heterogeneities of Velocities?
- Seismic anisotropy?**

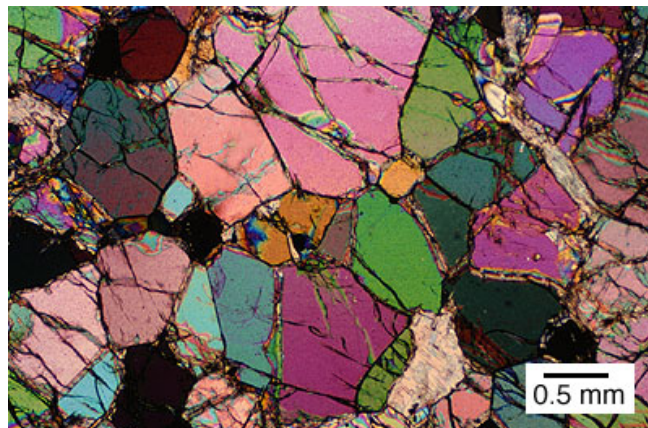


To explain seismic data:
heterogeneities isotropic / anisotropic, anelasticity

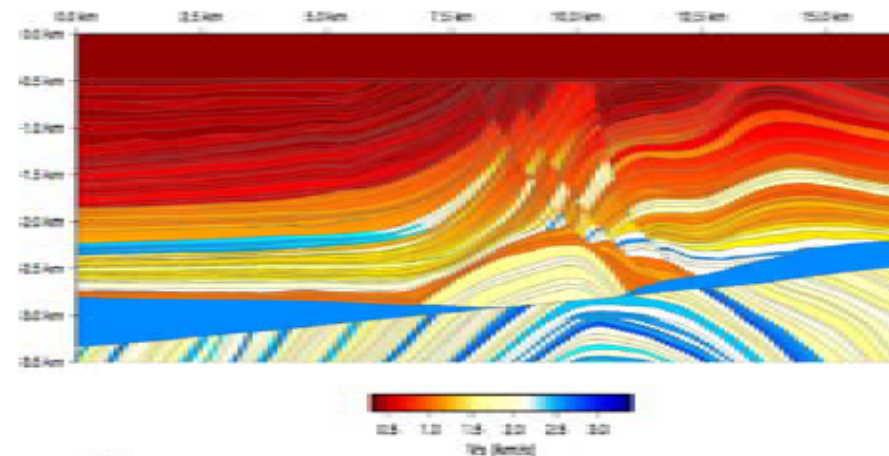
Confusion: Heterogeneity \neq Anisotropy

-Homogeneous, anisotropic
(Olivine aggregates)

-Heterogeneous, isotropic

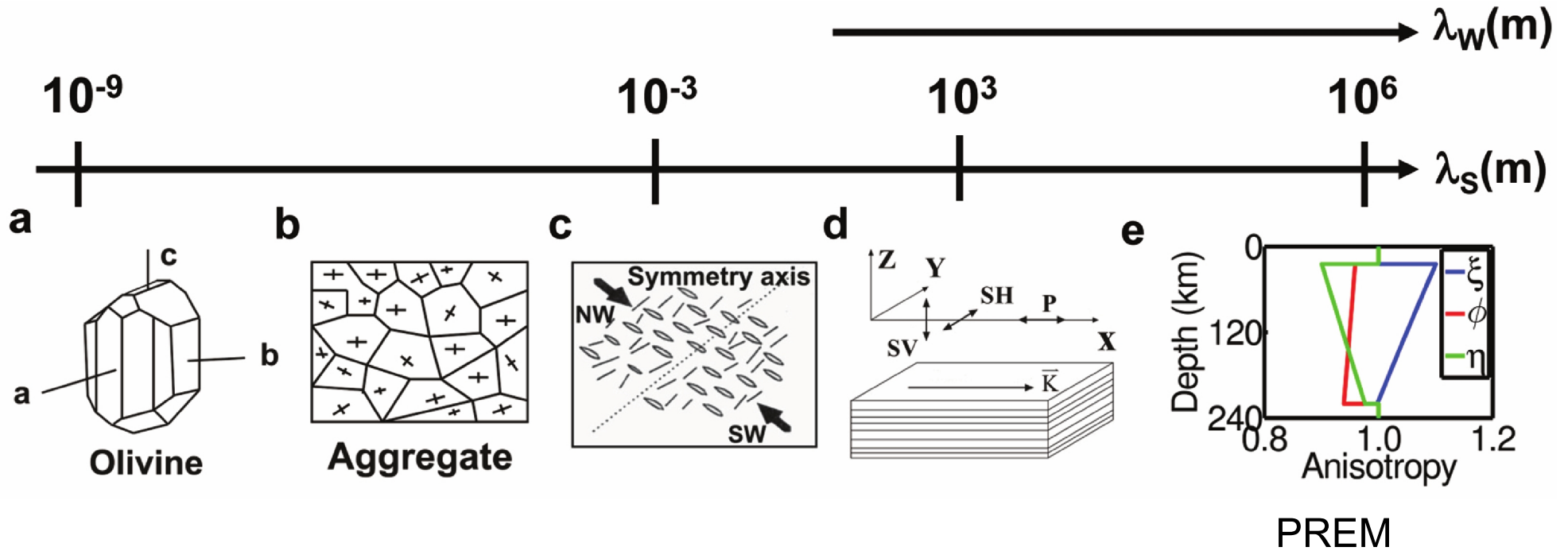


Marmousi



Solid Earth: heterogeneous + anisotropic + anelastic

Seismic Anisotropy at all scales



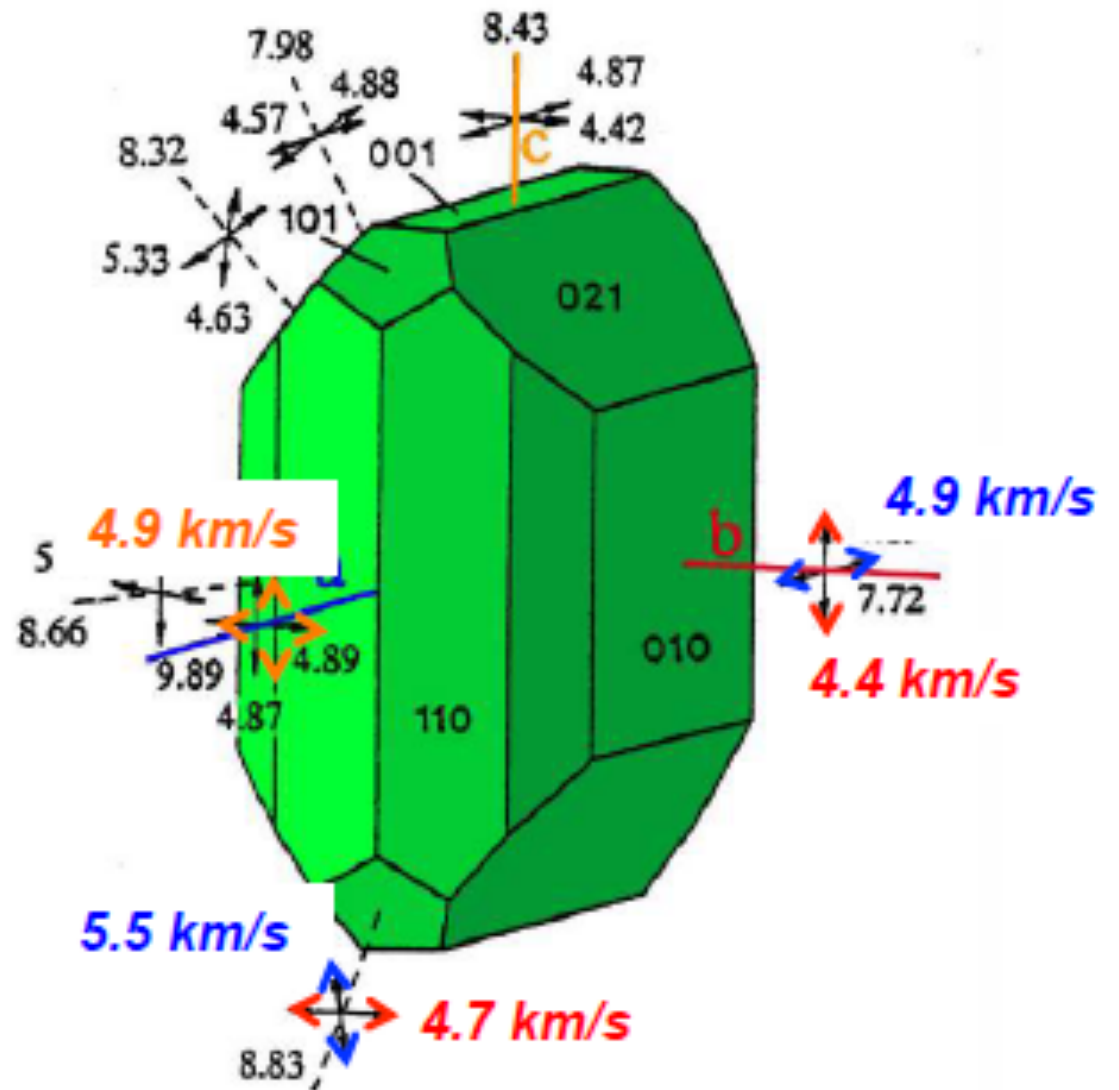
PREM: radial anisotropy: up to 10%

λ_W seismic wavelength

λ_S spatial scale

(Wang et al., 2013)

Olivine : 20% anisotropy

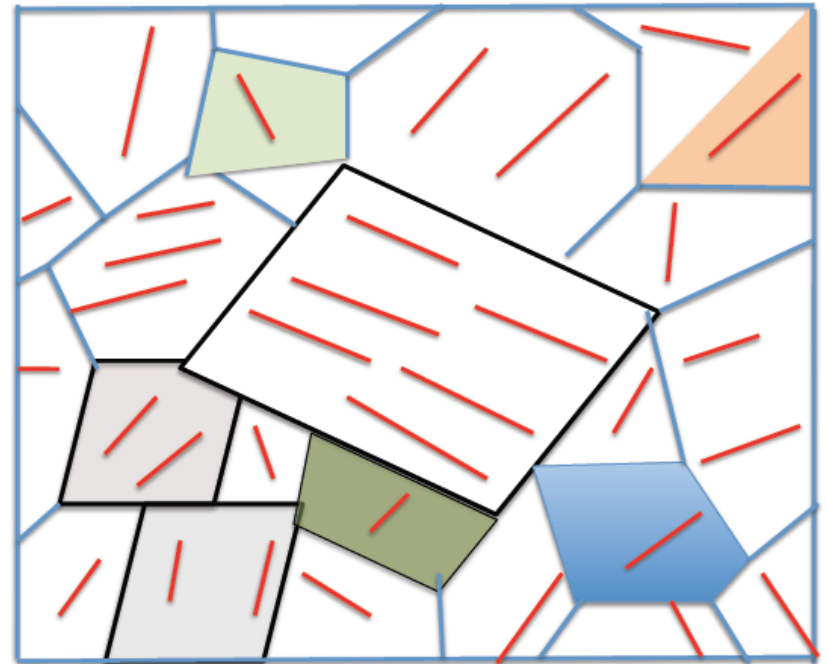
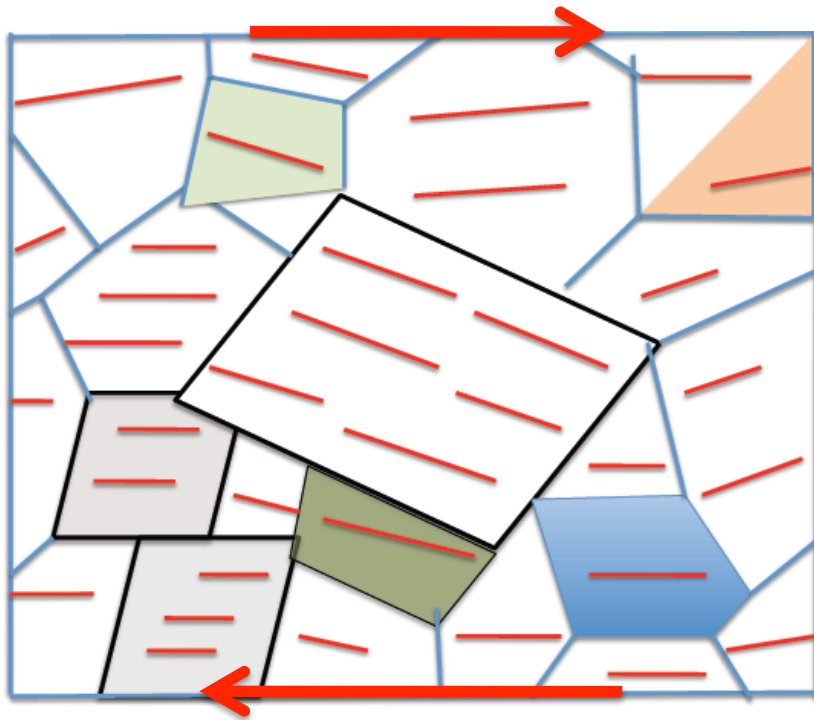


Coherent Strain field (convective process)
(L.P.O.: lattice preferred orientation)

Incoherent Strain field
(random orientation)

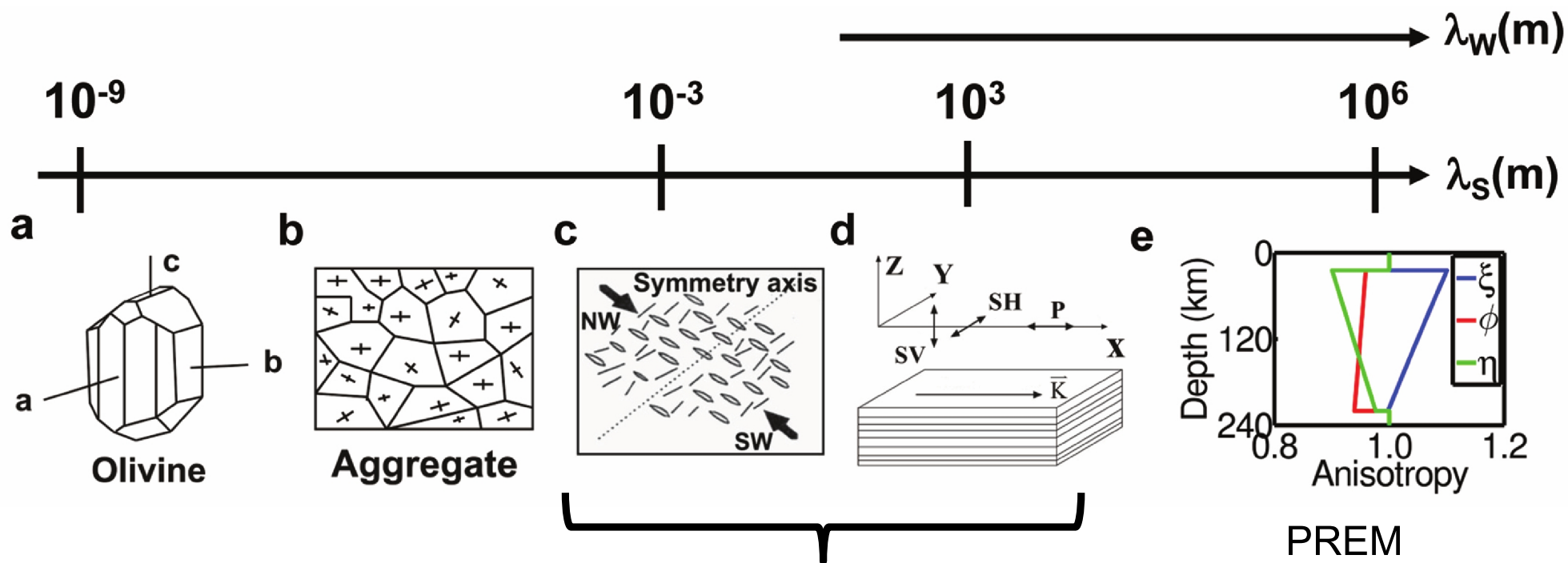
Large-scale Seismic anisotropy $\neq 0$

Large-scale Seismic anisotropy ≈ 0



Inner Organization

Seismic Anisotropy at all scales



PREM: radial anisotropy: up to 10%

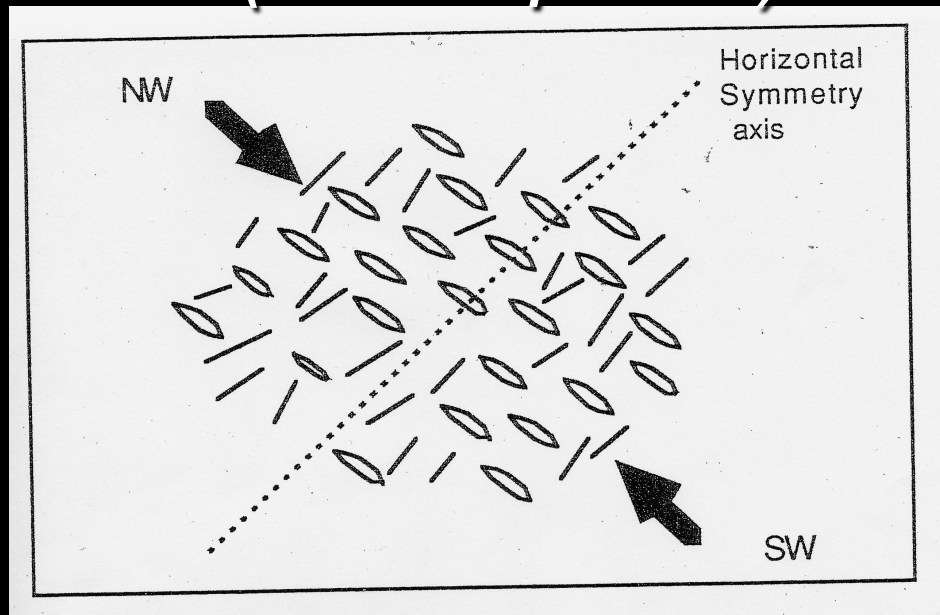
λ_W seismic wavelength

λ_S spatial scale

(Wang et al., 2013)

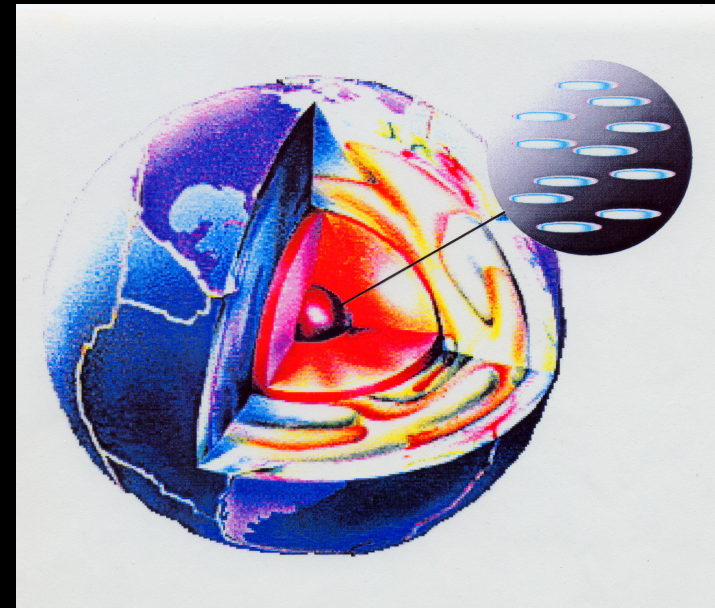
S.P.O.: Shape Preferred Orientation **Cracks, fluid inclusions, ... (Stress field)**

Crust (+lithosphere)



(Babuska and Cara, 1991)

Inner core



(Singh et al., 2001)

STRESSMETER

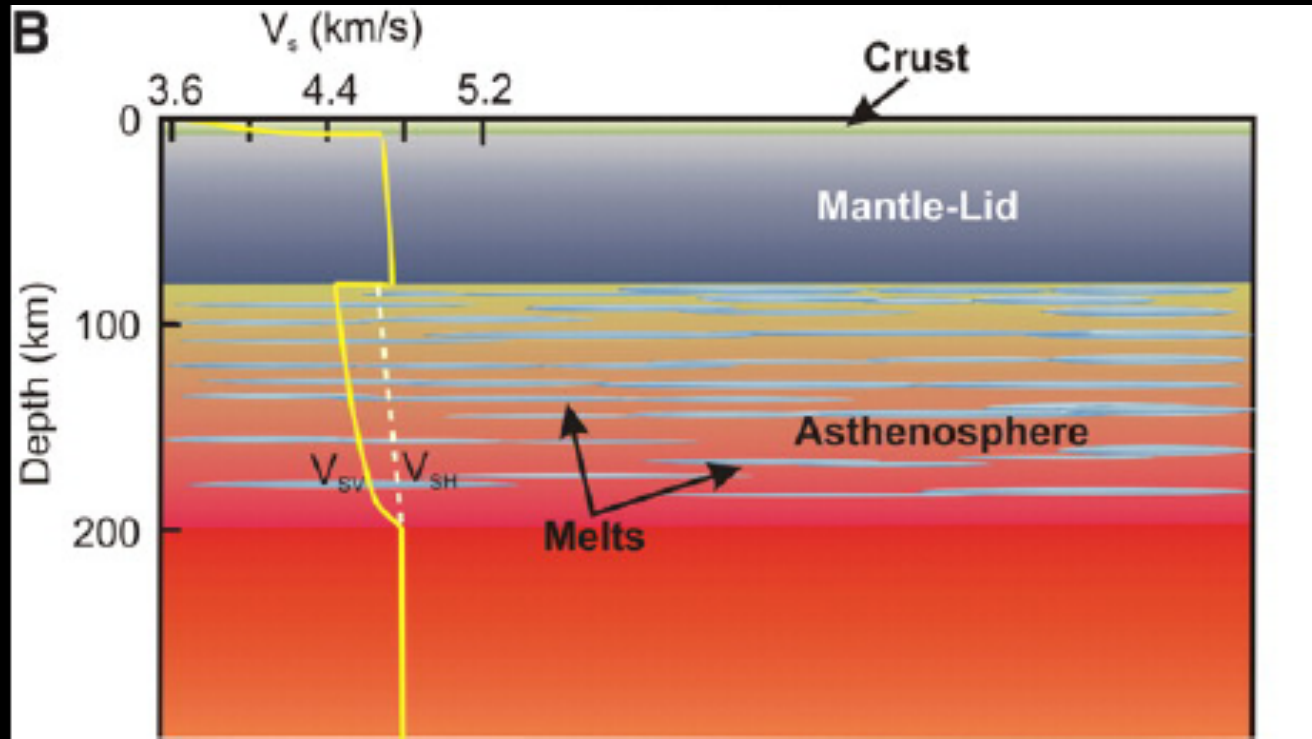
Temporal variations of anisotropy?

Monitoring of cracked, fractured zones

(seismogenic zones: Durand et al. 2011; Saade et al., 2014, 2017)

FINE LAYERING: Stratification Anisotropy

Mille-feuilles model (partial melting)



Effective medium

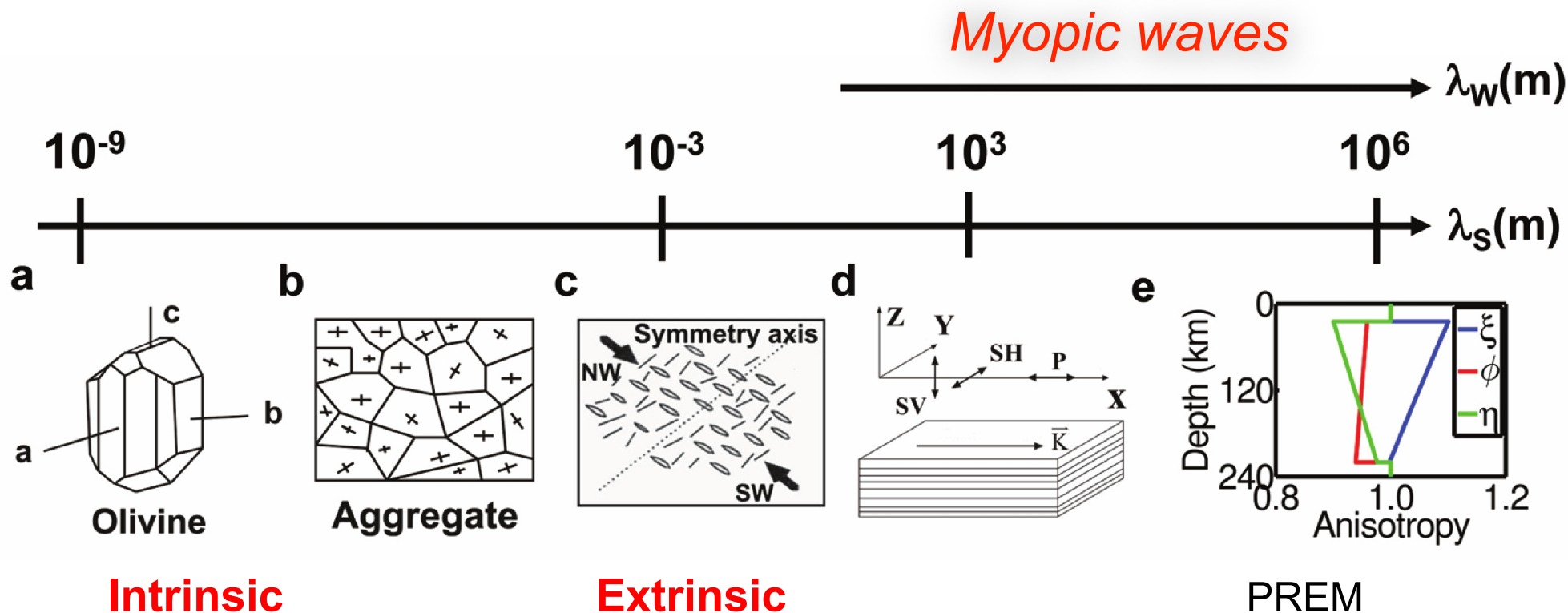
Radial anisotropy (Kawakatsu et al. 2009)

V.T.I. Vertical Transversely Isotropic medium:

5 parameters

$$(A = \rho V_{PH}^2, C = \rho V_{PV}^2, F, L = V_{SV}^2, N = V_{SH}^2)$$

Seismic Anisotropy at all scales



Observed (apparent) anisotropy
Intrinsic versus Extrinsic anisotropy

$$\alpha = p\alpha^{\text{int}} + (1-p)\alpha^{\text{ext}}$$

NOT A SECOND ORDER EFFECT, different interpretations



Wave propagation in anisotropic medium

- Isotropic medium

$$\rho \frac{\partial^2 u_i}{\partial t^2} = \lambda \frac{\partial \Theta}{\partial x_i} + \mu \frac{\partial}{\partial x_j} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

Where u is seismic displacement,
 λ and μ Lamé parameters

$$\Theta = \nabla \cdot \mathbf{u}$$

- Anisotropic medium

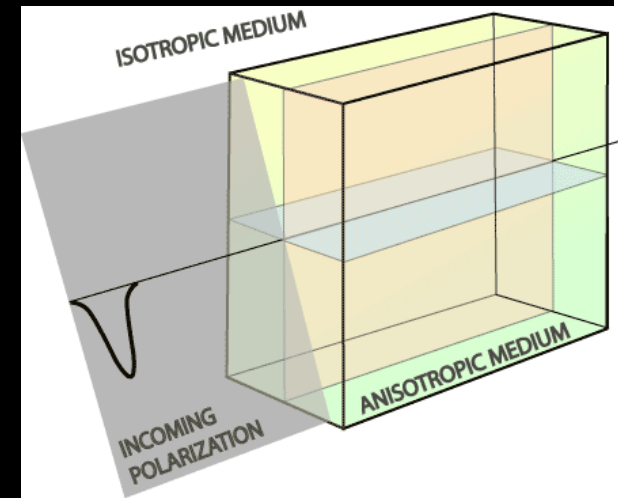
$$\rho \frac{\partial^2 u_i}{\partial t^2} = \frac{1}{2} \gamma_{ijkl} (u_{k,lj} + u_{l,kj})$$

γ_{ijkl} 4th-order elastic tensor



Different kinds of anisotropy effects on seismic waves

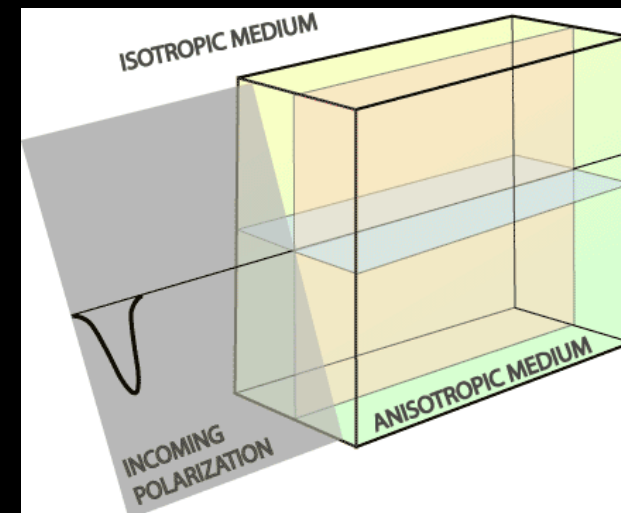
- Body waves: Shear wave splitting (birefringence)



Courtesy of Ed. Garnero

Different kinds of anisotropy effects on seismic waves

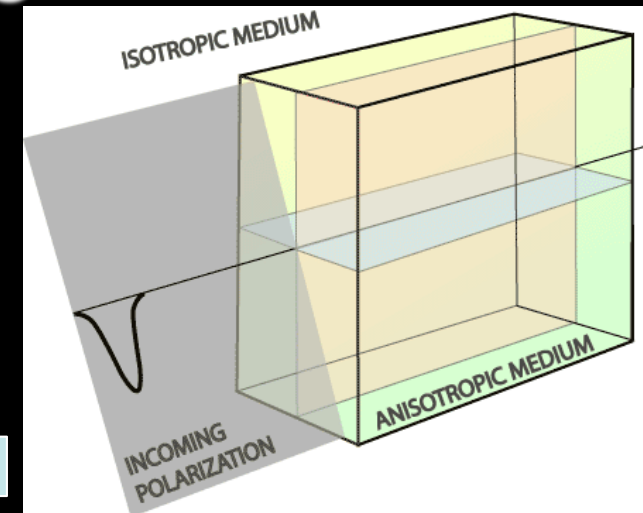
- Body waves: Shear wave splitting (birefringence)
- Surface waves (Rayleigh and Love):
 - Rayleigh-Love discrepancy
 - Azimuthal variations of phase (or group) velocities, radial anisotropy
 - Quasi-Rayleigh, Quasi-Love polarization anomalies



Courtesy of Ed. Garnero

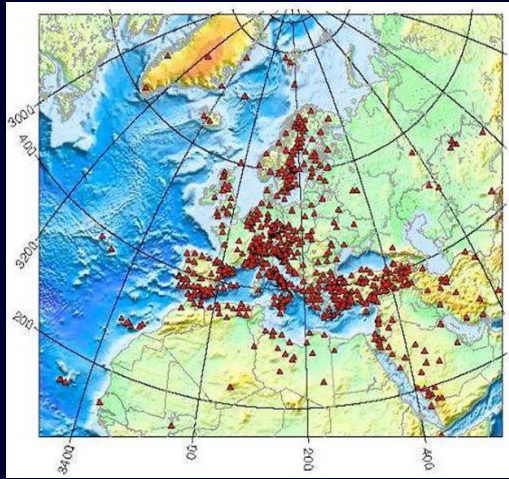
Different kinds of anisotropy effects on seismic waves

- Body waves: Shear wave splitting (birefringence)
- Surface waves (Rayleigh and Love):
 - Rayleigh-Love discrepancy
 - Azimuthal variations of phase (or group) velocities, radial anisotropy
 - Quasi-Rayleigh, Quasi-Love polarization anomalies

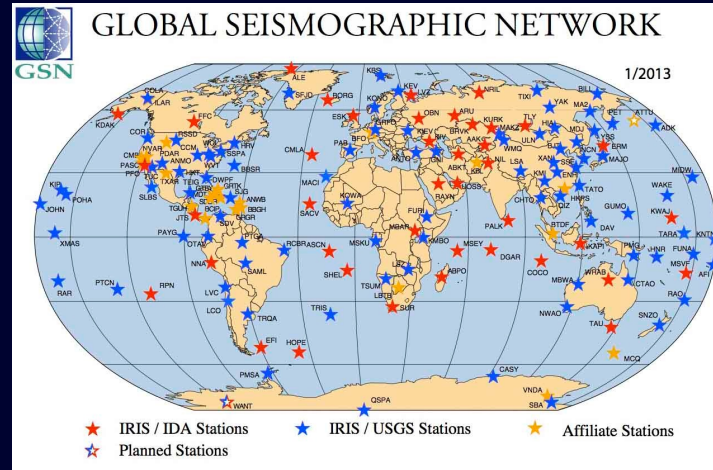


Courtesy of Ed. Garnero

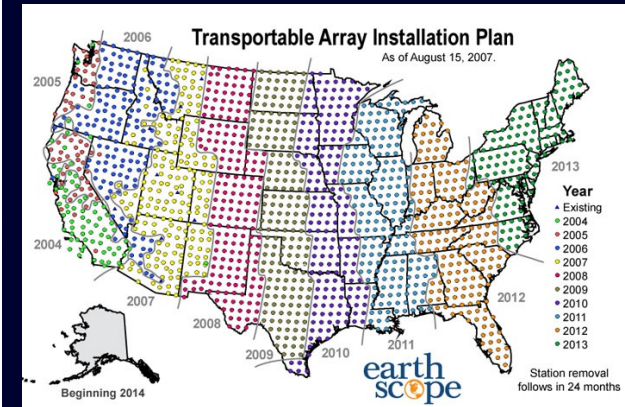
Data in Global & Regional Seismology



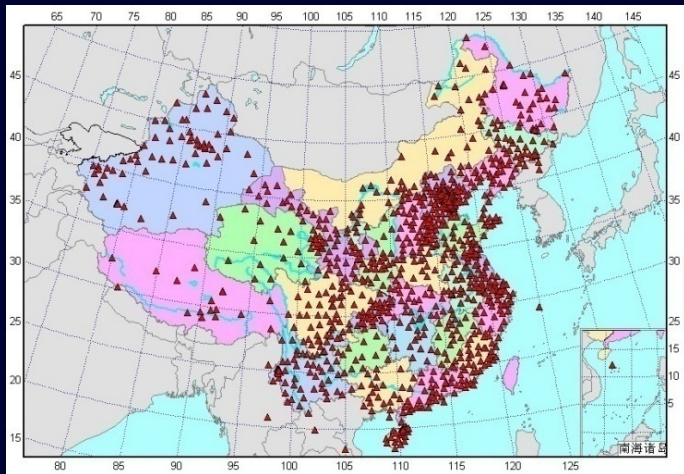
[www.geo.uib.no]



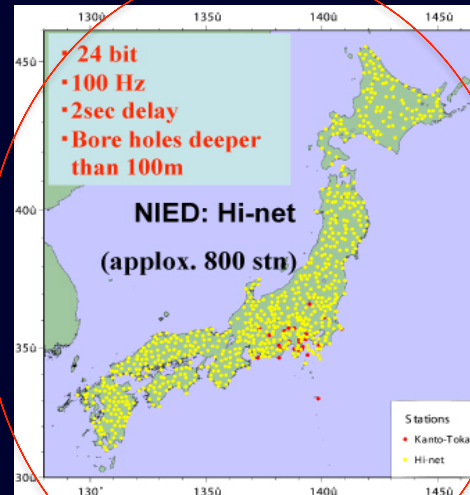
[www.iris.edu]



[web.mst.edu]



[data.earthquake.cn]



[drh.edm.bosai.go.jp]



Barruol et al., 2013



OUTLINE

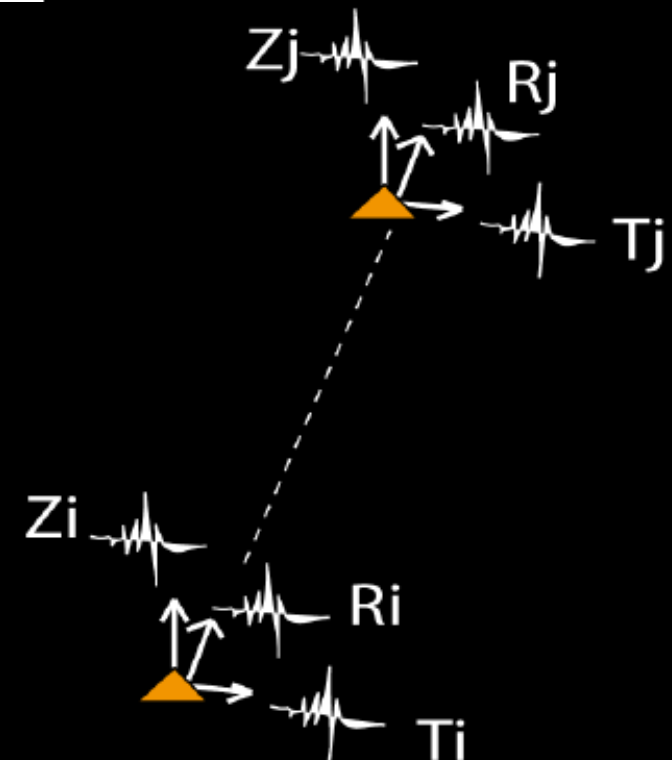
- Data driven field: Seismic Data
- Cross-correlation tensor: Seismic Anisotropy?
- Seismic Anisotropy: many processes

- **Scientific Issues:**
 - Seismic monitoring:**
 - Temporal changes of polarization anisotropy in:**
 - **seismogenic** (Parkfield, Cal., USA; Iwate-Miyagi, Japan)
 - Volcanic zones (Mount Fuji, Japan)

Cross-correlation tensor C_{ij}
 for 2 stations i, j and 3 components k, l
 Seismic signals $S_{ik}(t), S_{jl}(t)$

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT



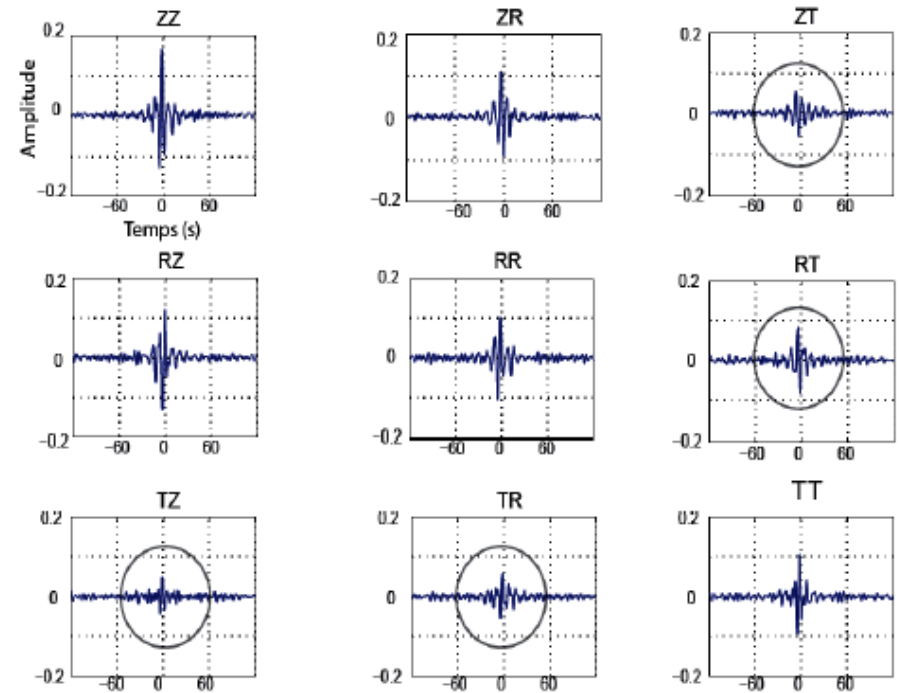
How can we explain the off-diagonal terms of the cross-correlation tensor?

$$TZ, TR, ZT, RT \neq 0$$

-Non uniform distribution of seismic noise sources?

-Lateral heterogeneities of Velocities?

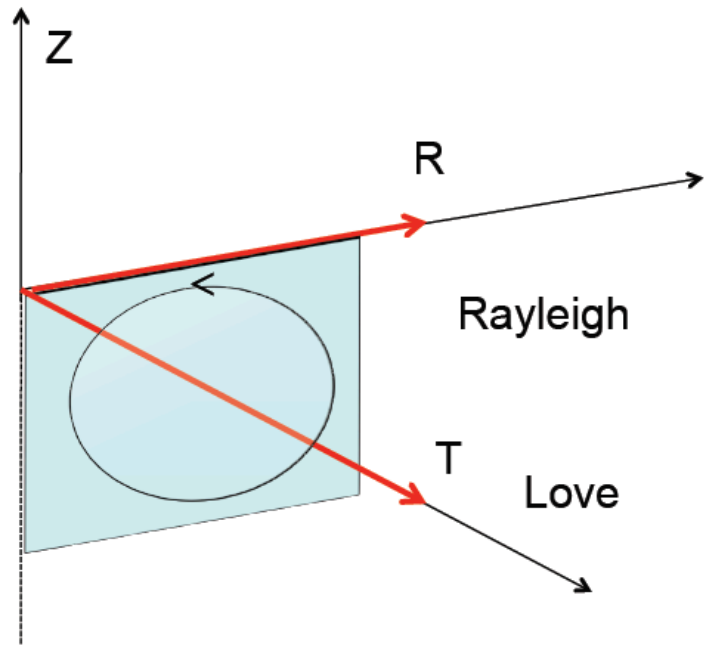
-Seismic anisotropy?



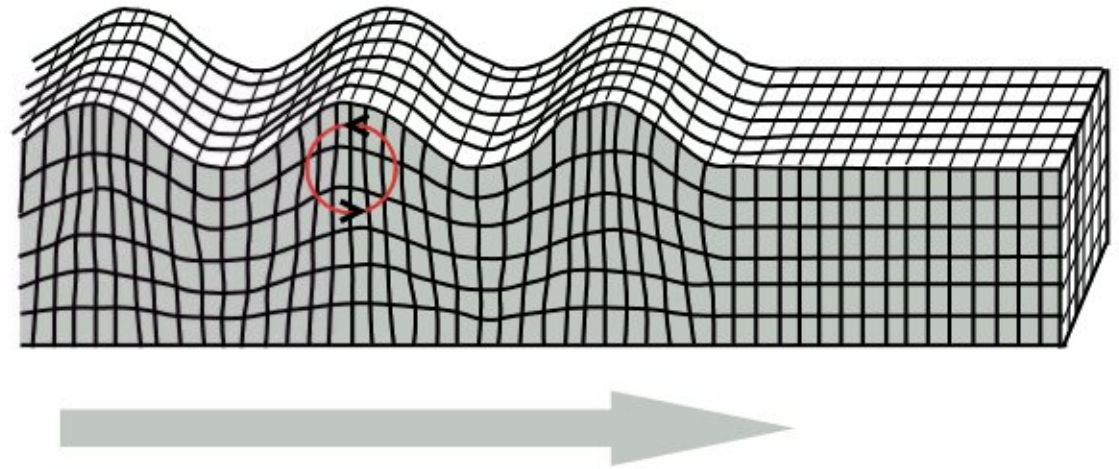
Effect of anisotropy on amplitude

Polarization of surface waves

ISOTROPIC MEDIUM

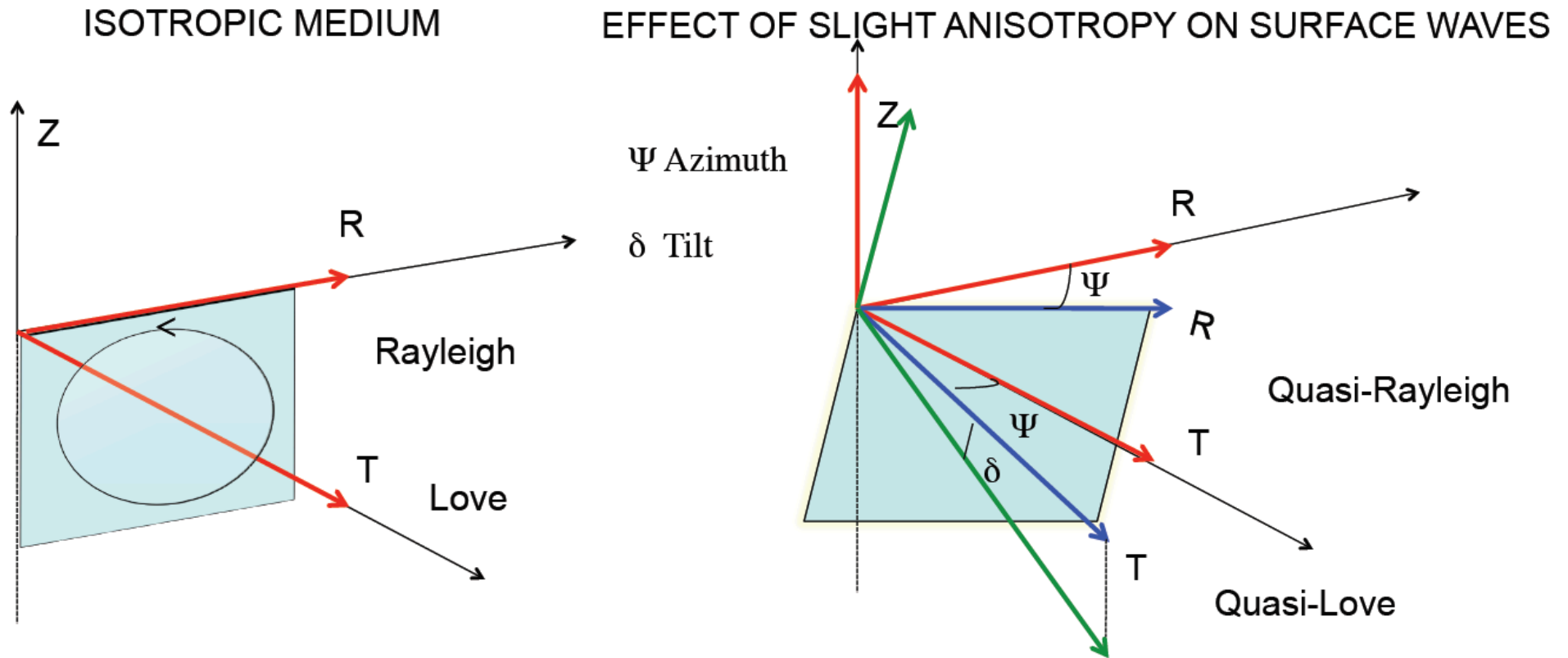


Rayleigh Wave



Effect of anisotropy on amplitude

Polarization of surface waves



New Observables Ψ, δ

THEORY

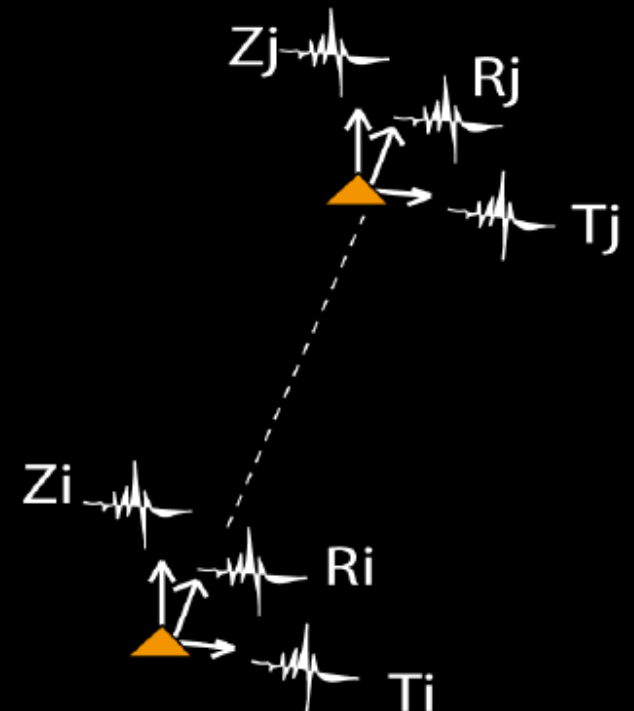
Cross-correlation for 2 stations i, j and 3 components k, l

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}}$$

Random sources:
Related to Green's tensor i,j (Medium response)

Cross-Correlation Tensor

ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT



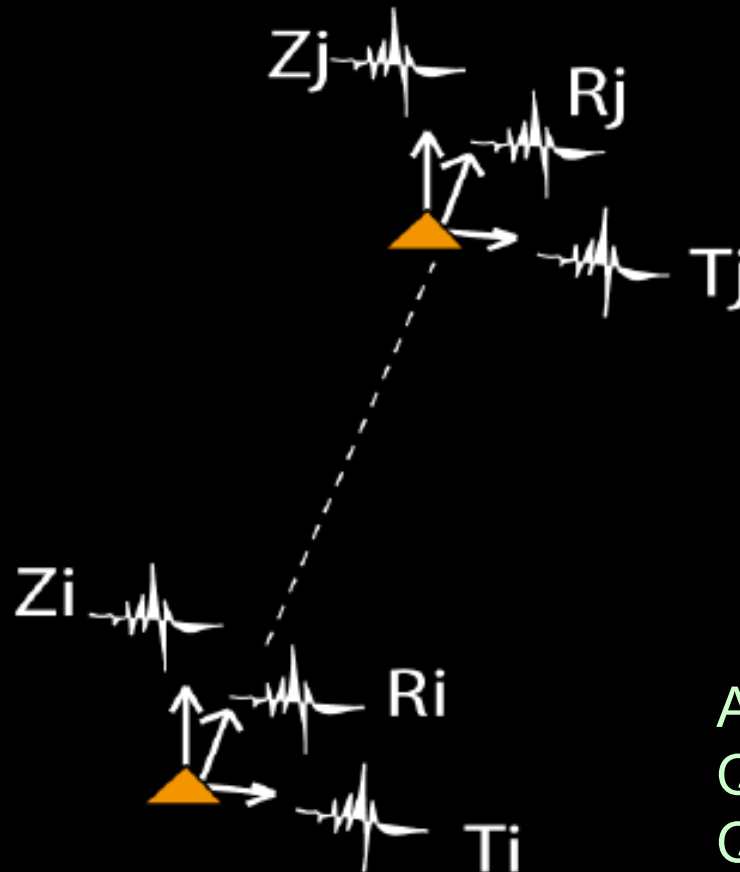
THEORY

Cross-correlation
for 2 stations i, j
and 3 components k, l

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

ZZ	ZR	~0
RZ	RR	~0
~0	~0	TT

ISOTROPIC MEDIUM
Rayleigh wave
Love wave



ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT

ANISOTROPIC MEDIUM
Quasi-Rayleigh wave
Quasi-Love wave

THEORY

Cross-correlation
for 2 stations i, j
and 3 components k, l

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

Azimuthal
Anisotropy
of Rayleigh
waves

Effect of anisotropy on C_{ij}

ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT

Joint Inversion
of Rayleigh
and Love waves:
Radial anisotropy

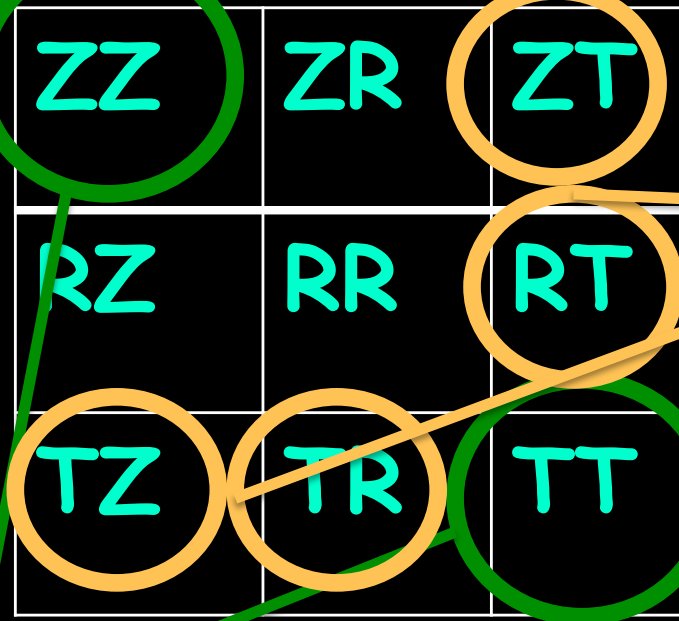
THEORY

Cross-correlation
for 2 stations i, j
and 3 components k, l

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

Azimuthal
Anisotropy
of Rayleigh
waves

Effect of anisotropy on C_{ij}



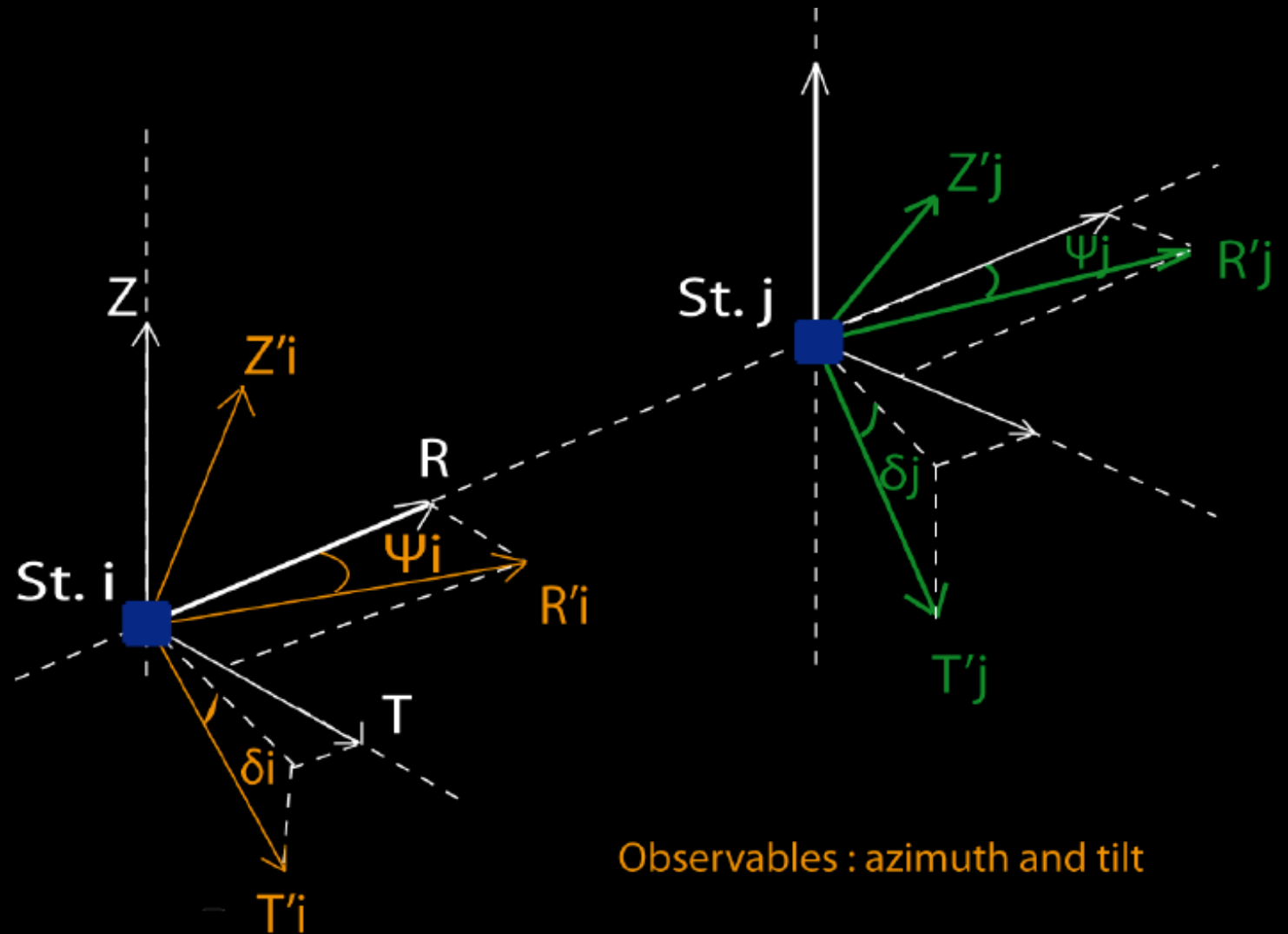
Quasi-Rayleigh
Quasi-Love
waves:
polarization
anomalies

Joint Inversion
of Rayleigh
and Love waves:
Radial anisotropy

ORA: Optimal Rotation Algorithm (Roux, GJI, 2010)

Minimization of the RT, TR, ZT and TZ components

ZZ	ZR	ZT
RZ	RR	RT
TZ	TR	TT



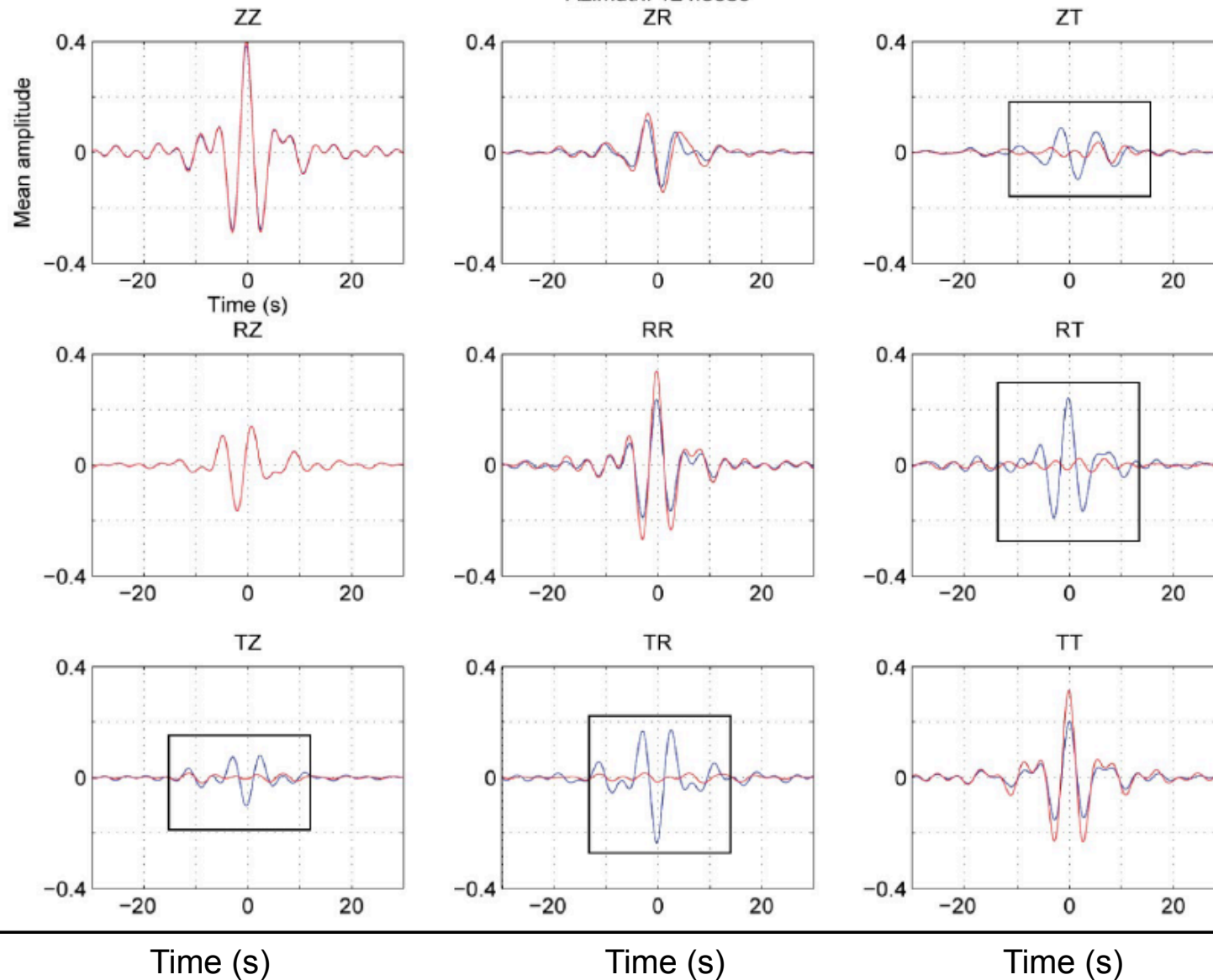
Observables : azimuth and tilt

Ψ δ

Efficiency of Optimal Rotation Algorithm (ORA)

GREEN'S TENSOR

Station pair 1-11
Azimuth: 124.5883



— Before ORA
— After ORA

=> Quasi-Rayleigh

New observables:
 Ψ and δ

Temporal Changes
of Ψ and δ ?

Temporal changes of Cross-correlations (polarization angle Ψ)

2 effects:

- Non-random distribution of seismic sources



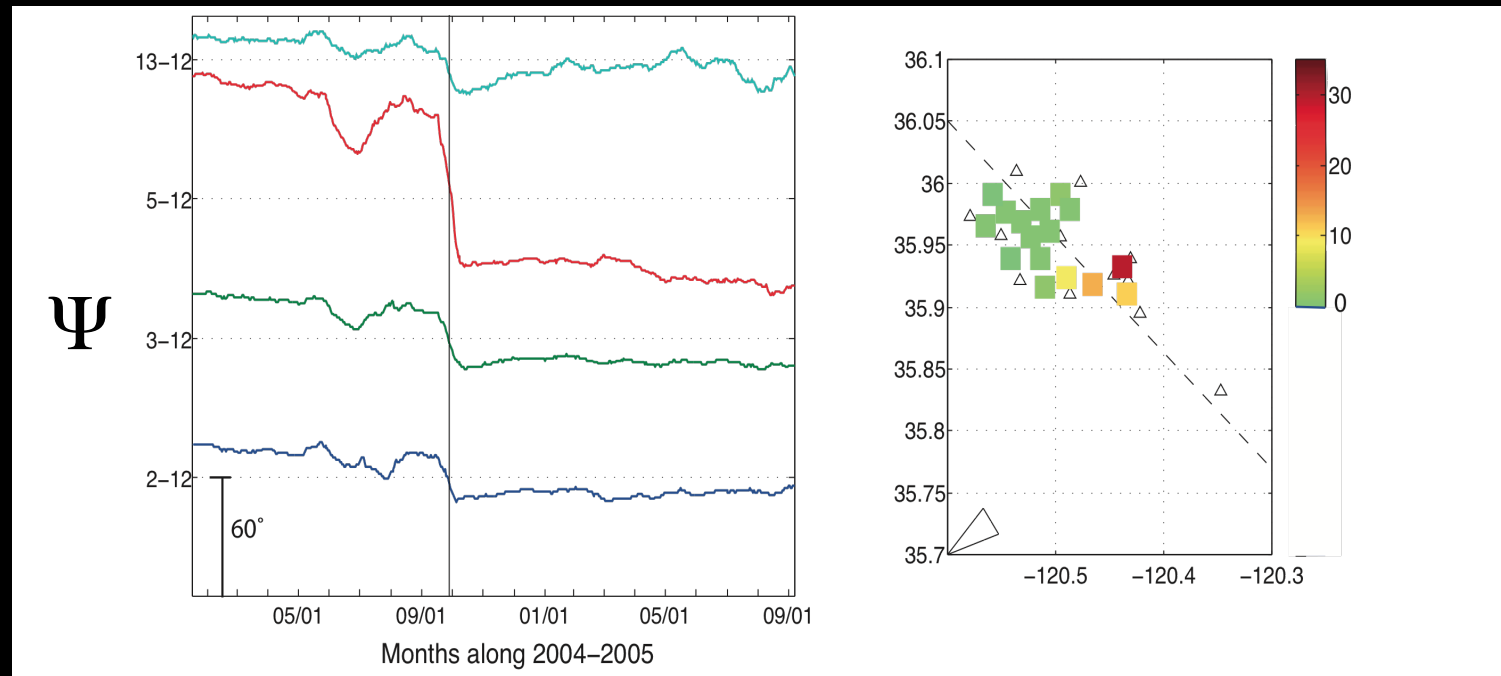
seasonal variations
(beamforming analysis)

- ANISOTROPY changes

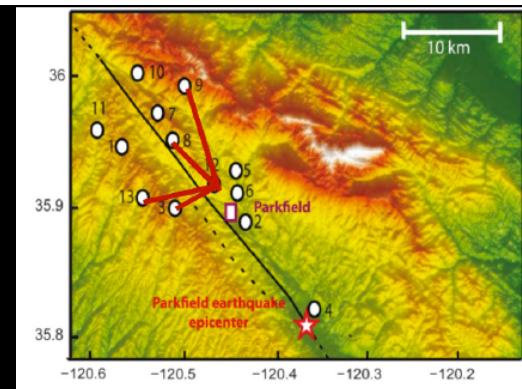


Stress field temporal variations

Time variations of Ψ angle after noise removal



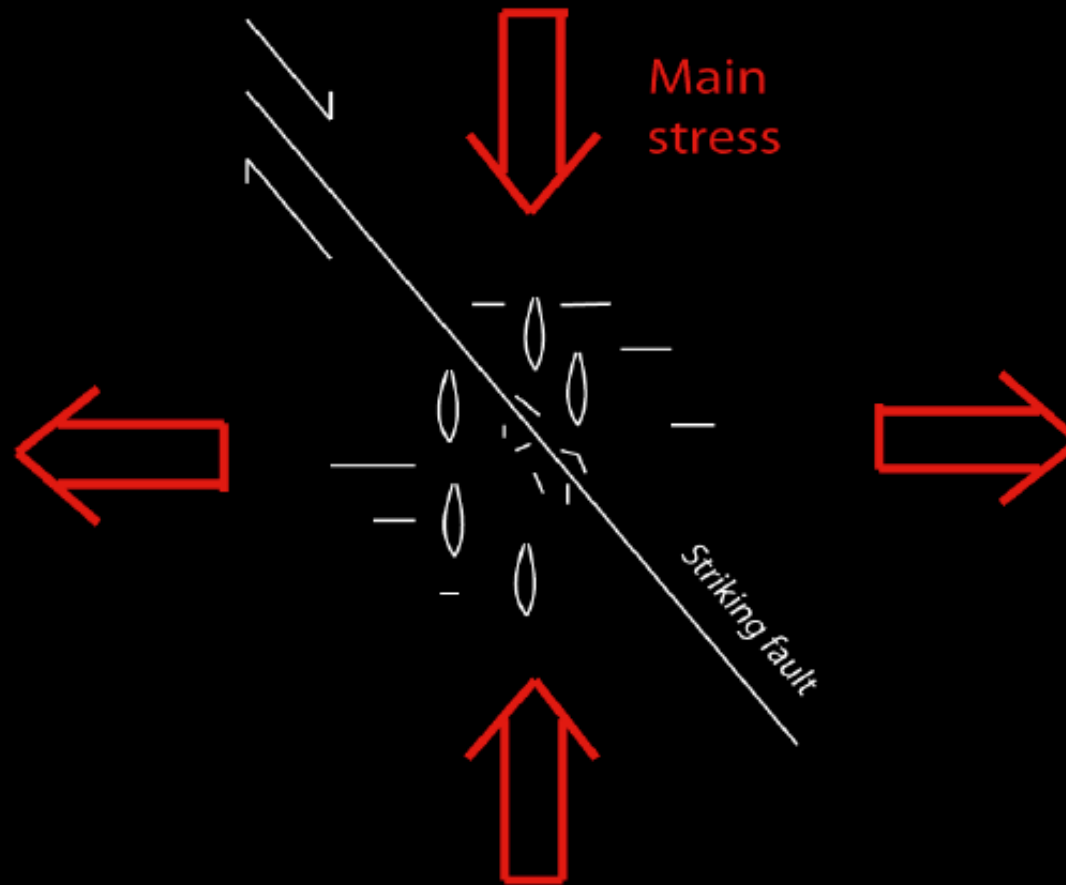
Significant co-seismic jumps for station pairs containing station 12



Tentative (reasonable) interpretation:
stress rotation => apparent rotation of the crack distribution

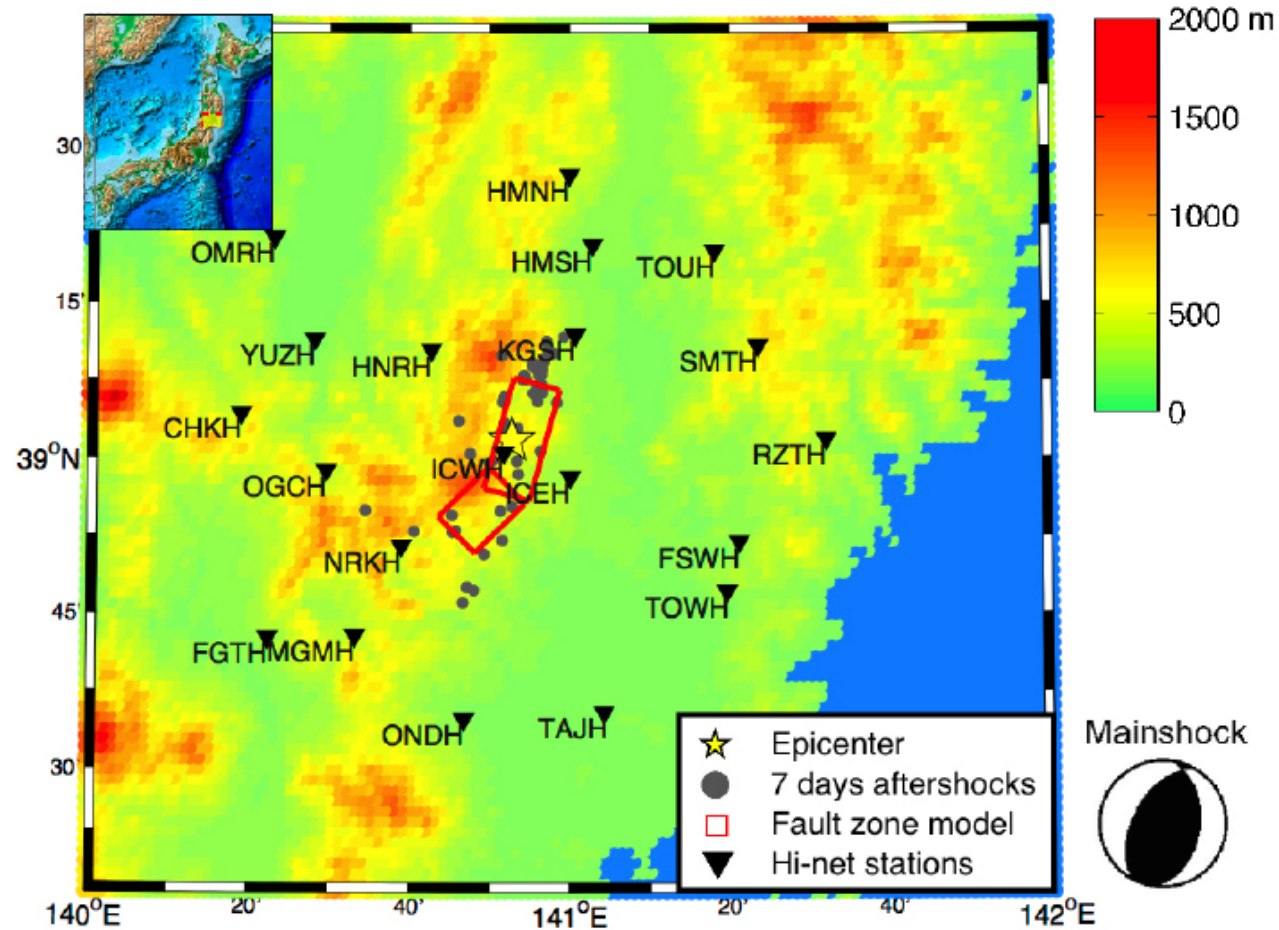
Seismic Anisotropy: Cracks, fluid inclusions

stress field rotations in the crust
⇒ temporal variations of velocity
and anisotropy during seismic cycle?



Other Tectonic context

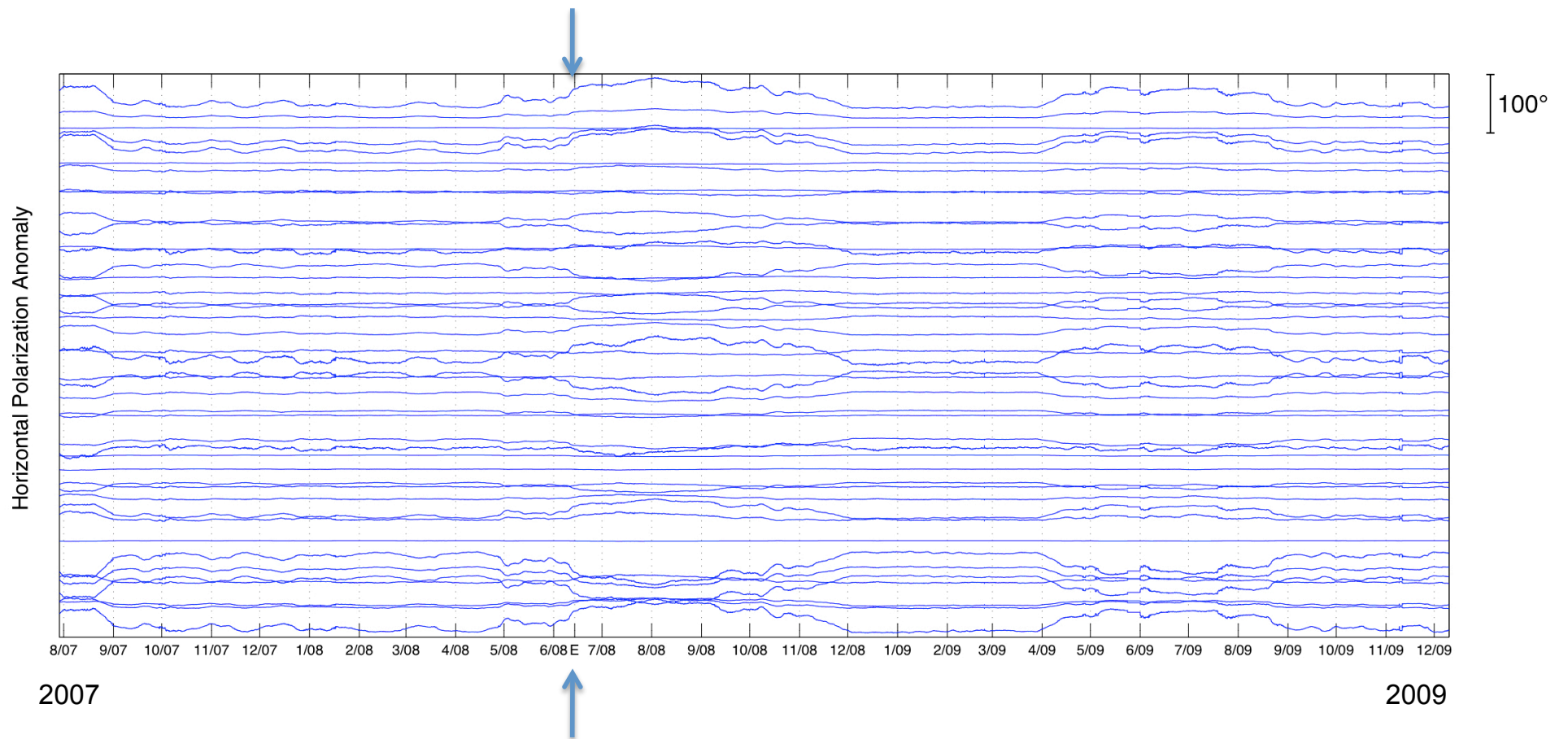
Iwate – Miyagi earthquake (14/06/08)



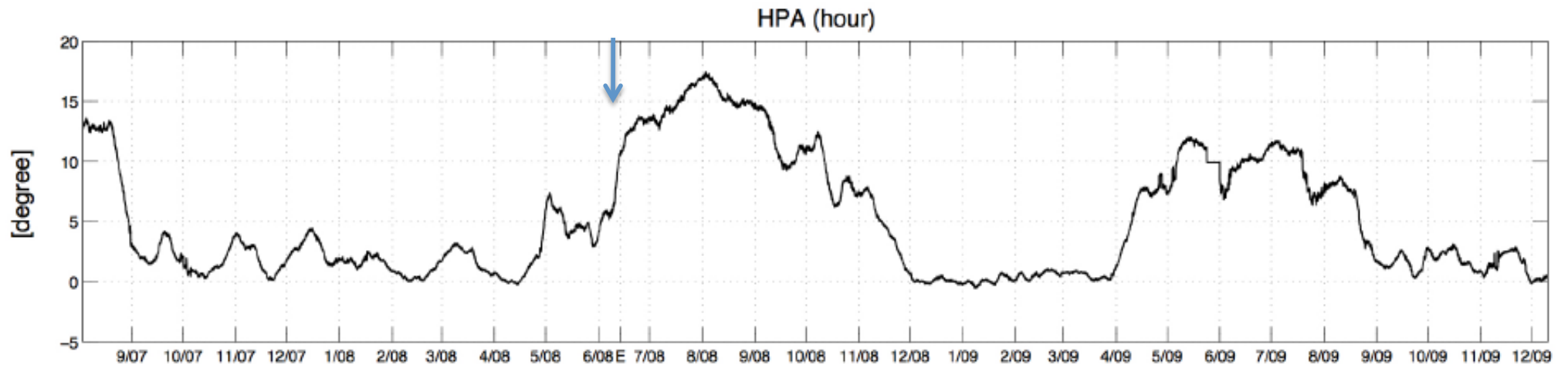
Data: Jul. 2007- Dec. 2009

NIED

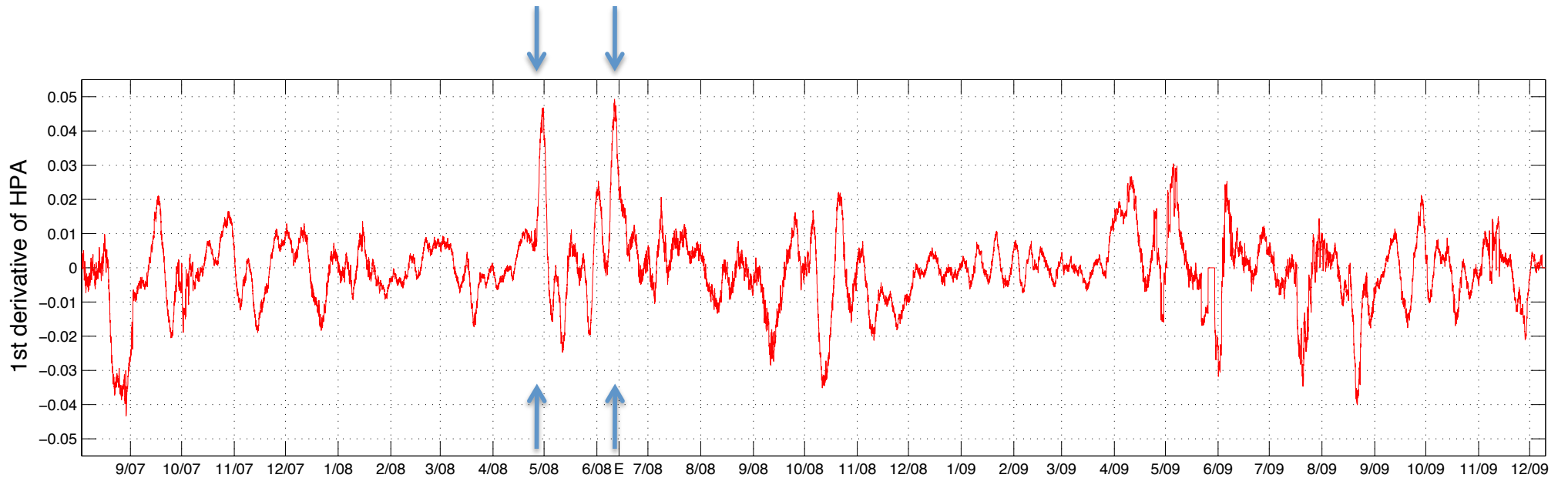
Horizontal Polarization Anomaly Ψ



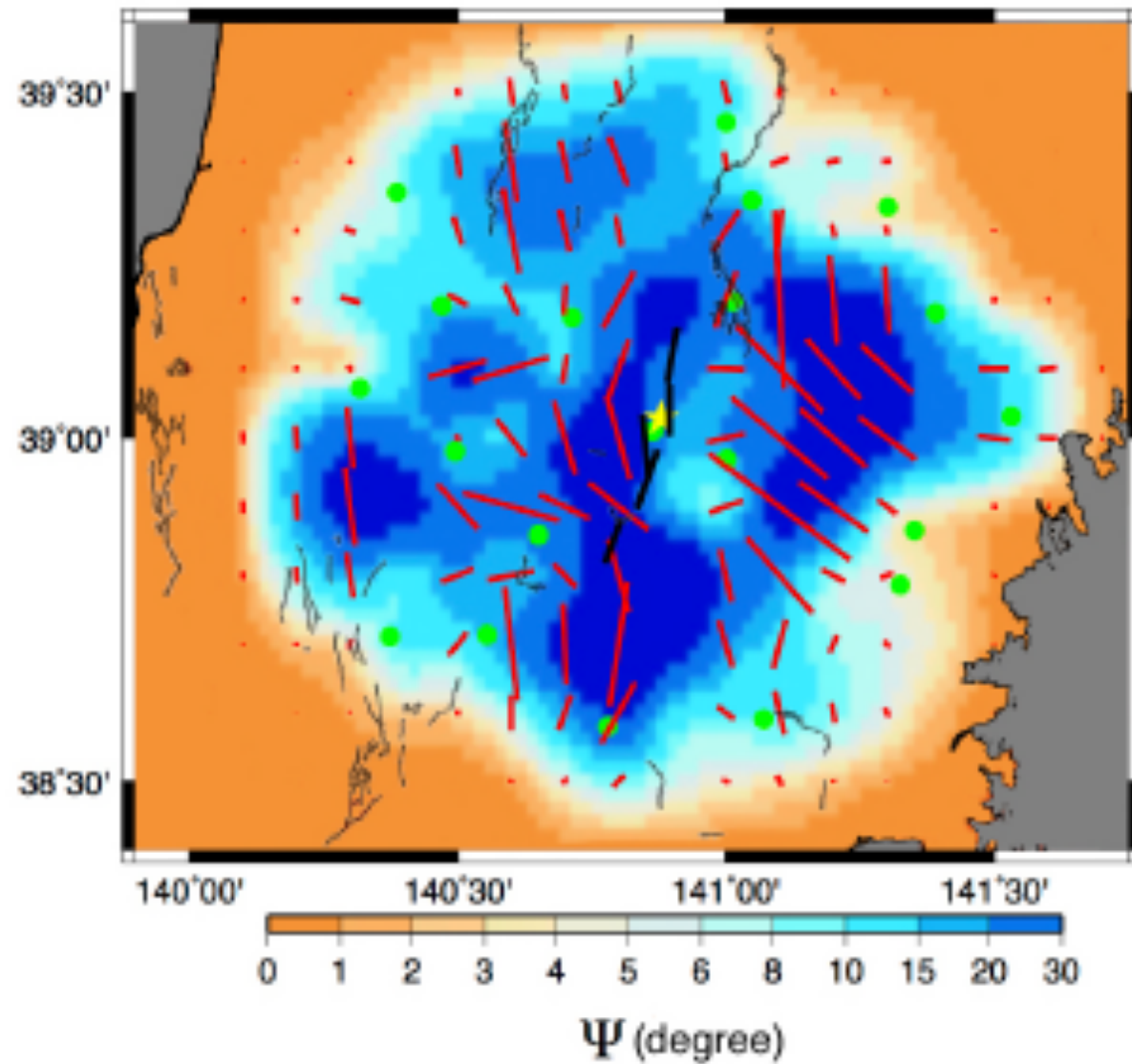
Average of the Horizontal Polarization Anomaly Ψ (HPA)



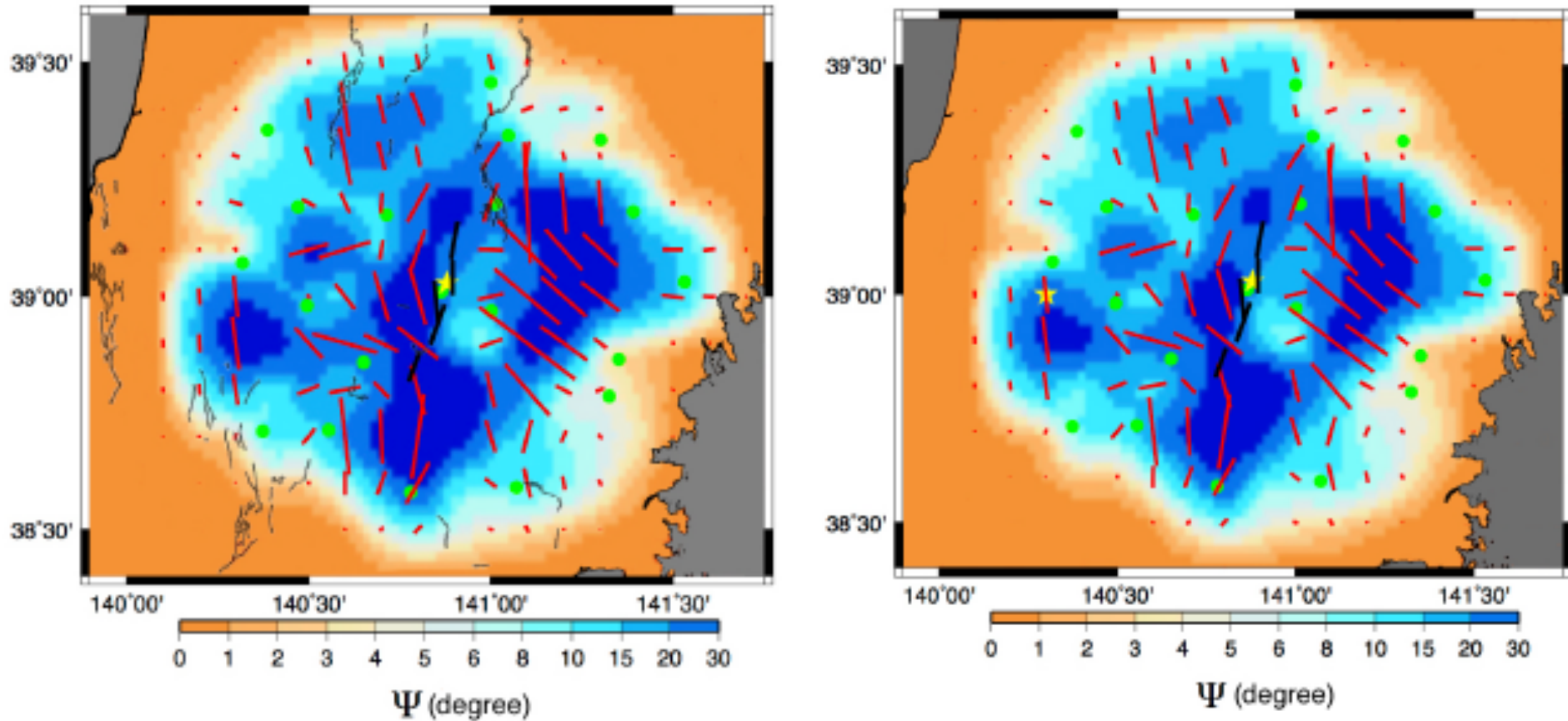
Time derivative of HPA



Horizontal Polarization change Ψ at the time of the EQ June 13, 2008



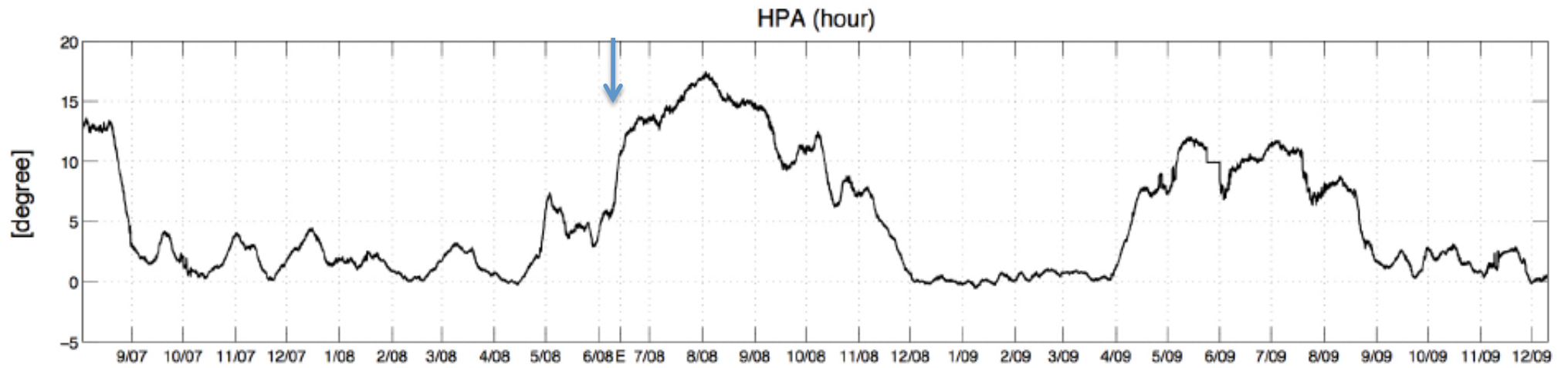
Horizontal Polarization change Ψ at the time of the EQ and at the end of April



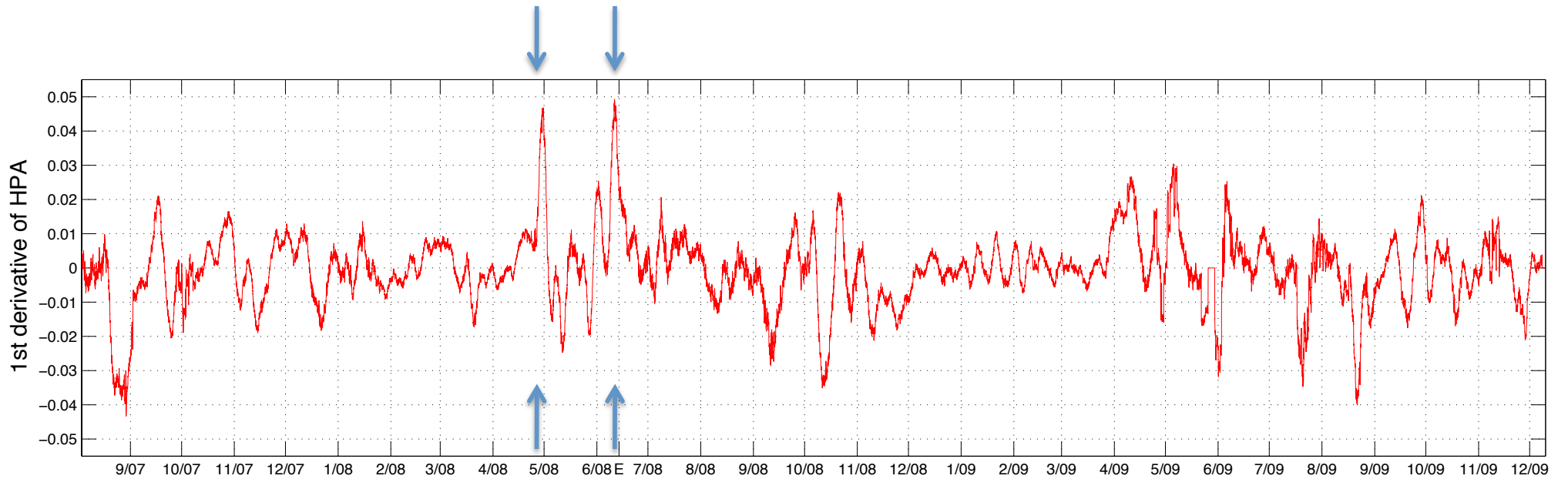
Many similarities for variations of the polarization anomaly for both events
(deep earthquake on April 17, 2008; Mw= 5.7)

Z! correlation does not mean causality

Average of the Horizontal Polarization Anomaly Ψ (HPA)



Time derivative of HPA





OUTLINE

- Seismic Data
- Seismic Anisotropy: many processes

- **Scientific Issues:**

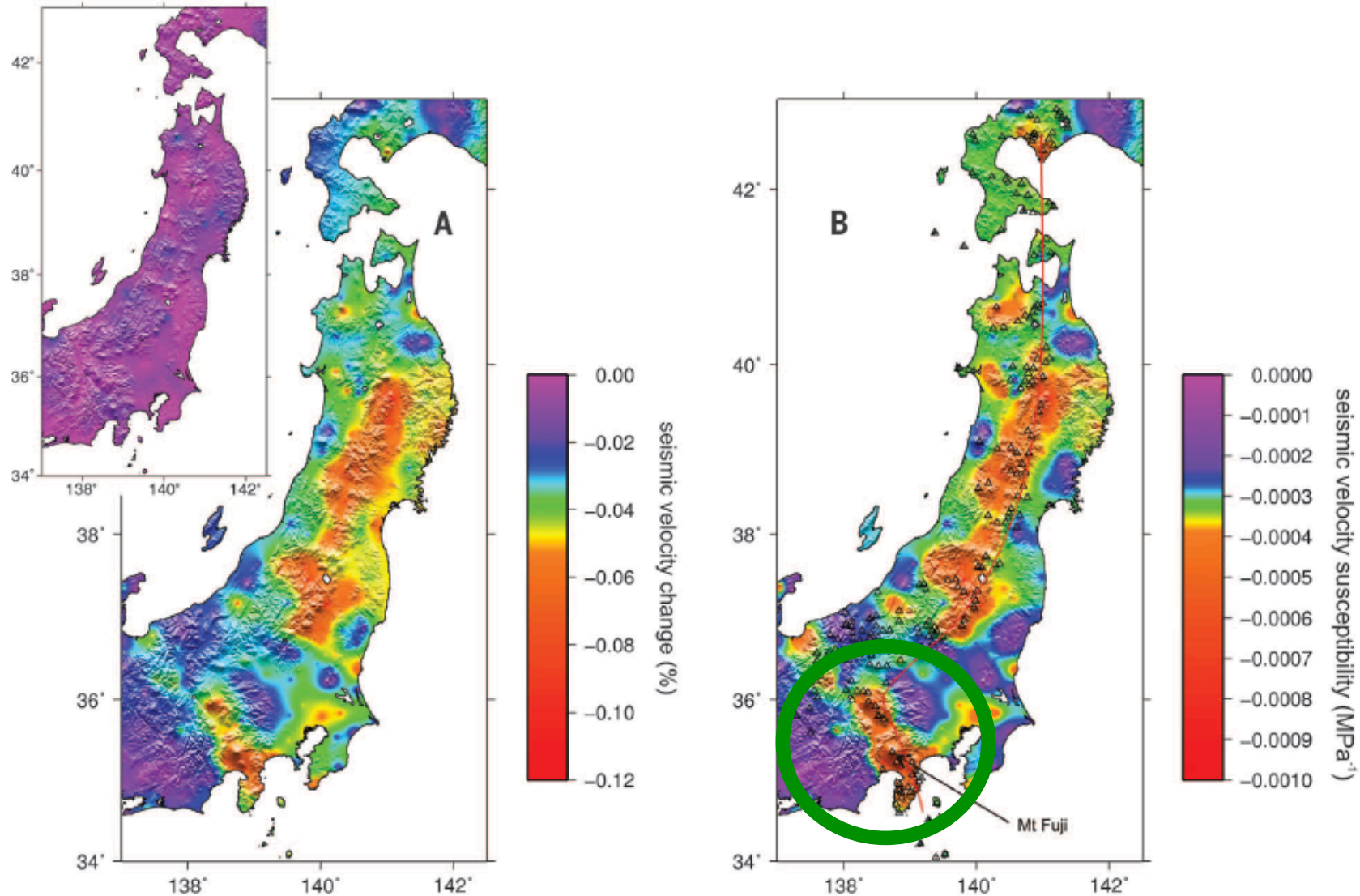
- Seismic monitoring:**

- Temporal changes of polarization anisotropy in:**

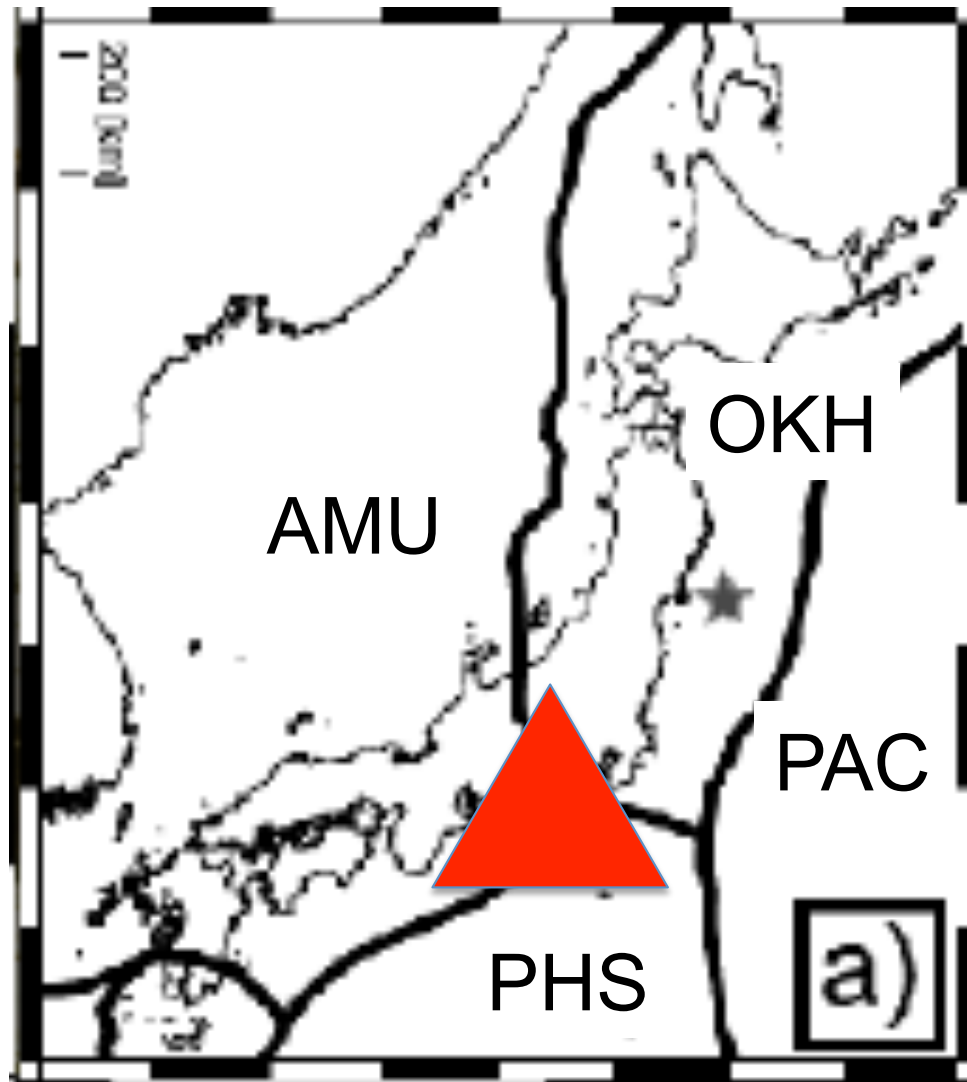
- seismogenic (Parkfield, Cal., USA; Iwate-Miyagi, Japan)

- **Volcanic zones (Mount Fuji, Japan)**

Volcanic zones: after the Tohoku-oki earthquake (11 march 2011)

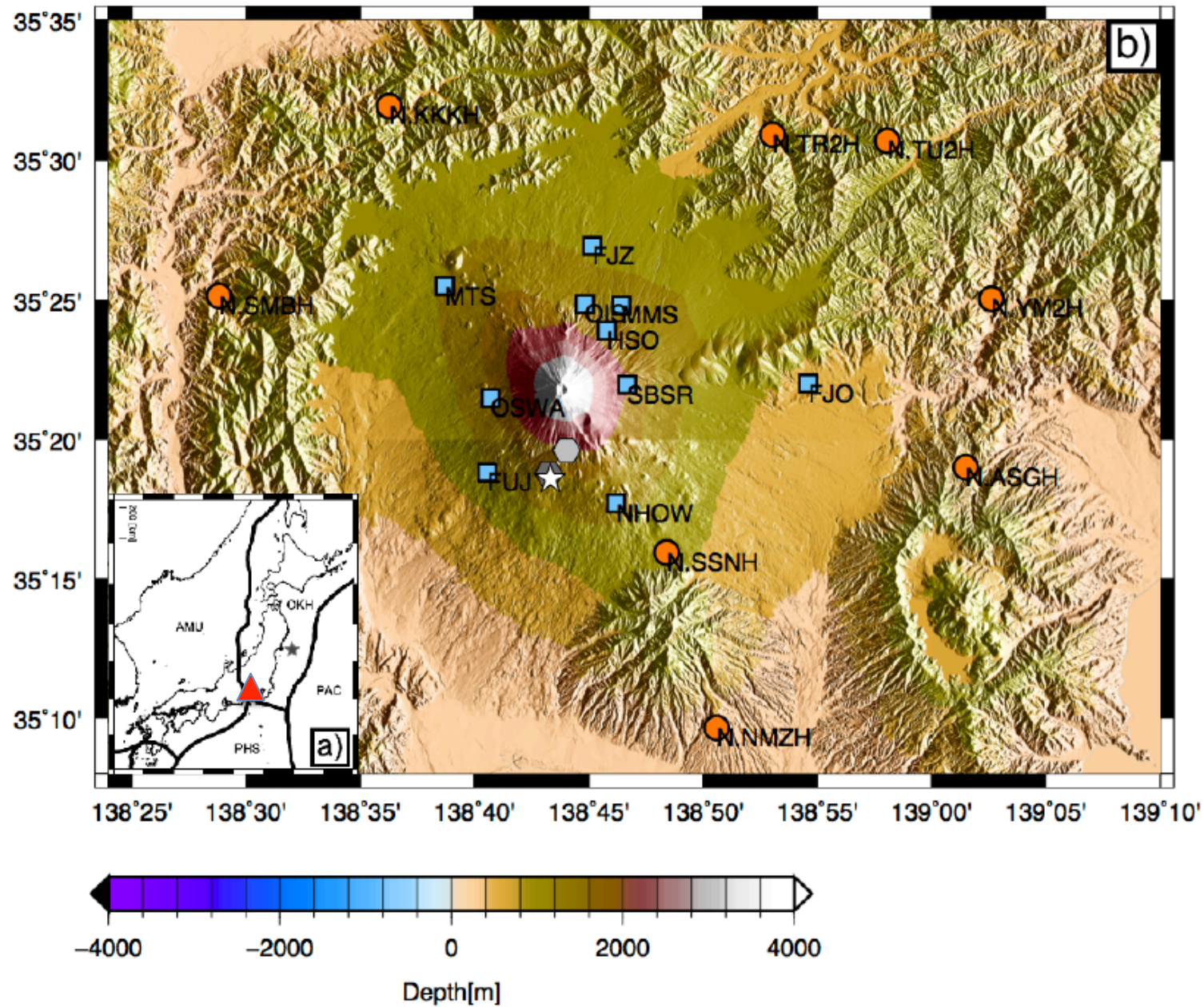


Monitoring of volcanic zones: Mount Fuji (+Hakone)



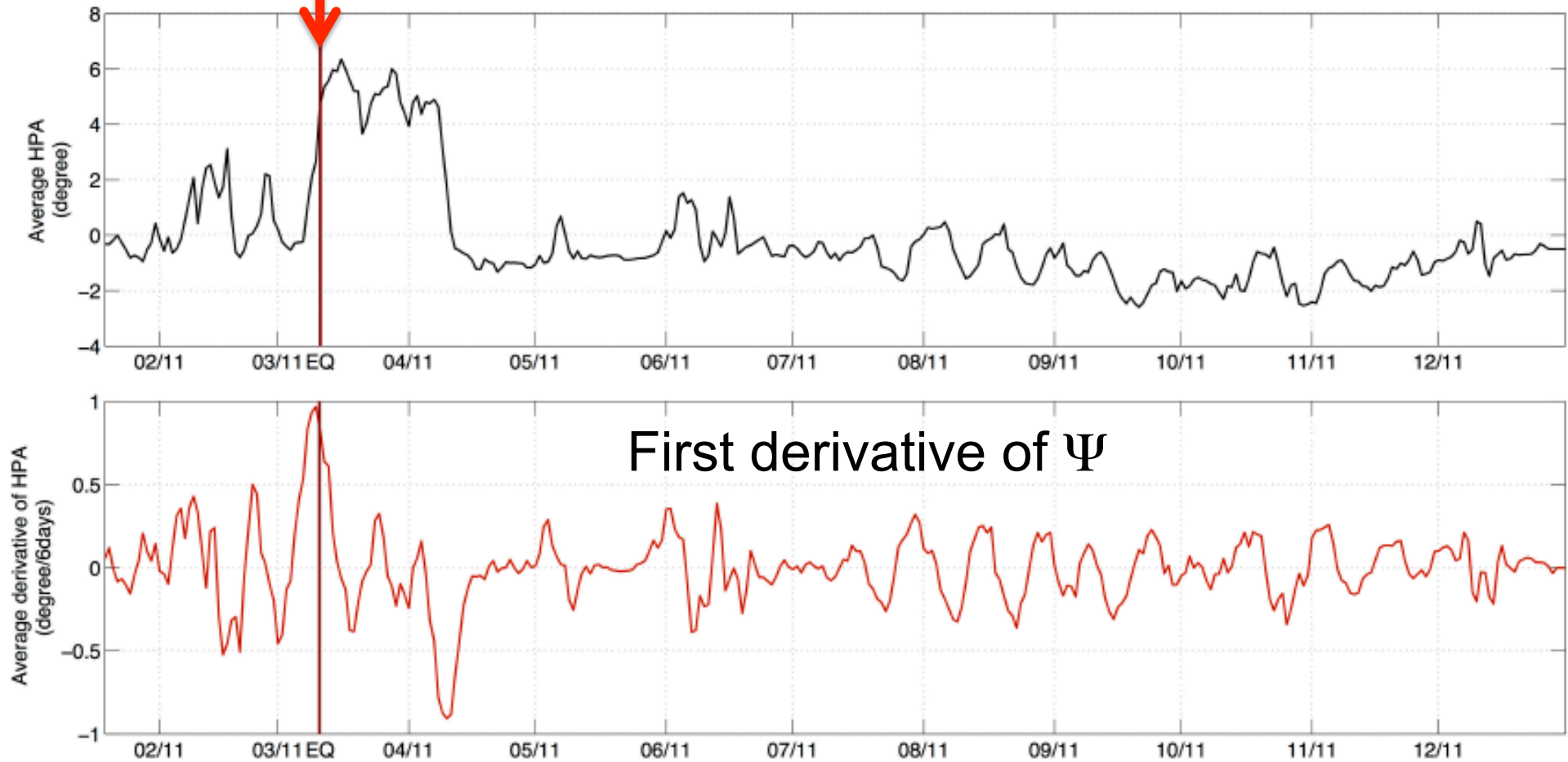
Hiroshige, 1858

Monitoring of volcanic zones: Mount Fuji



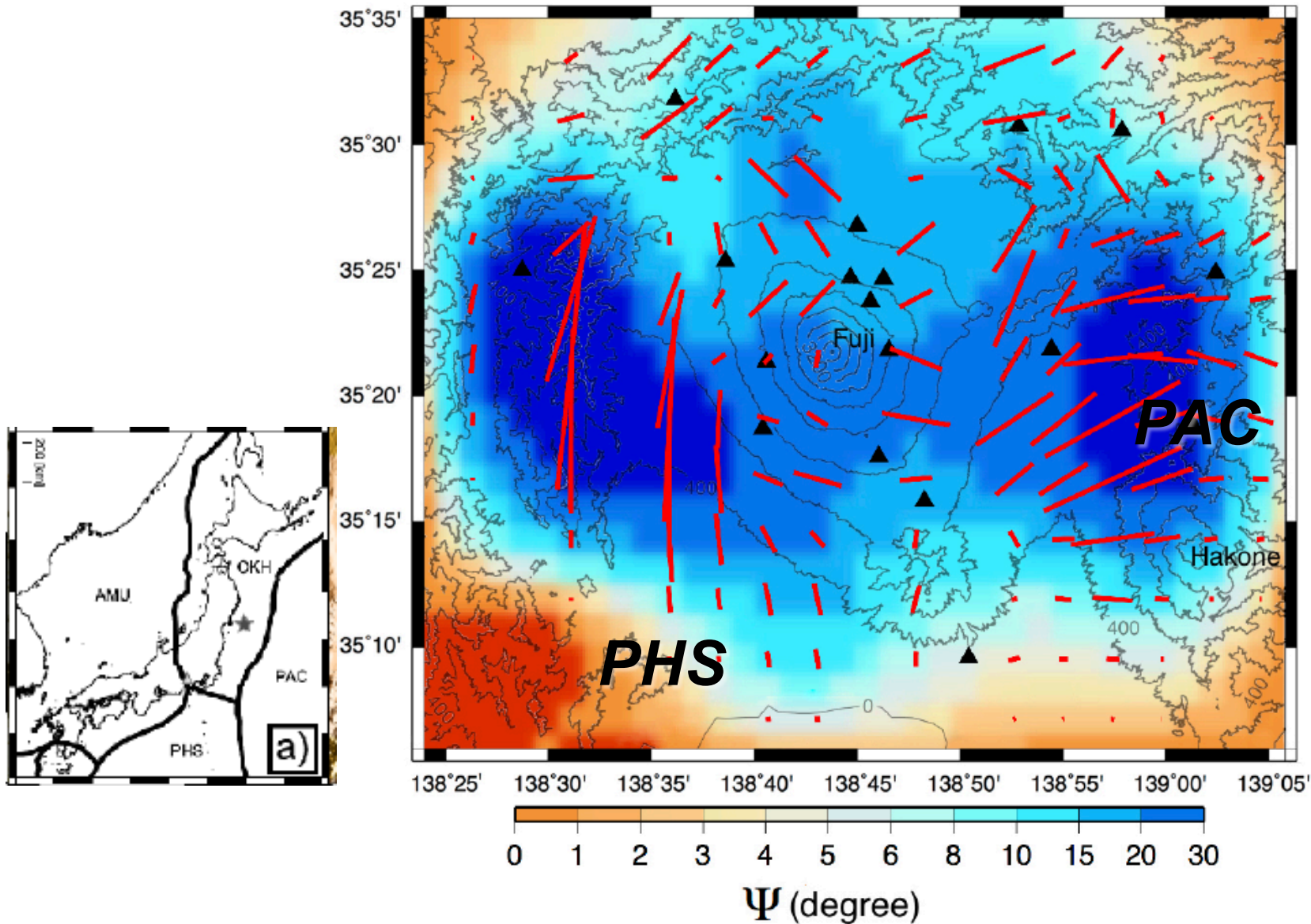
Monitoring of volcanic zones

Average of the horizontal polarization anomaly Ψ



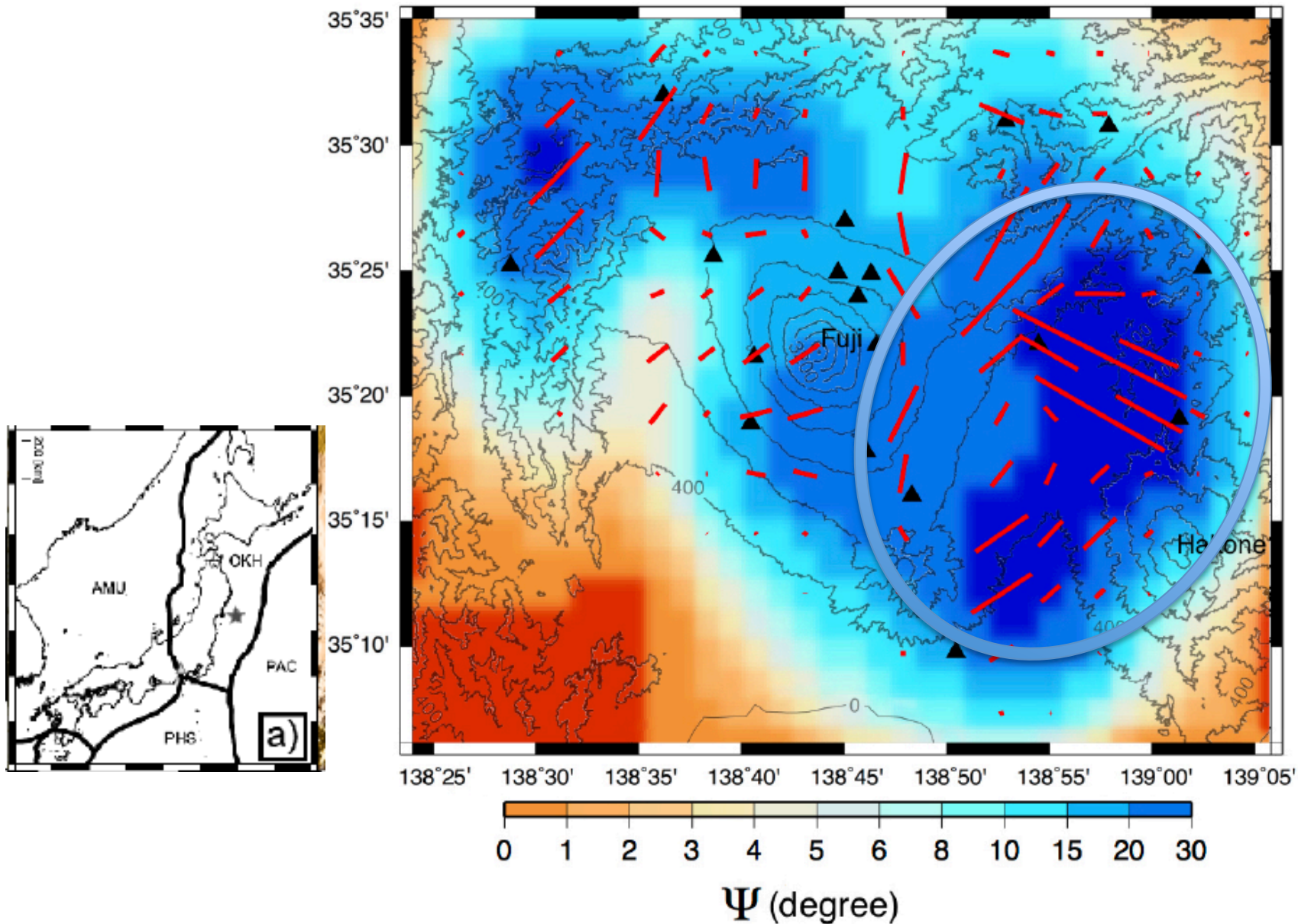
Monitoring of volcanic zones

Average polarization change + Orientation of anisotropy



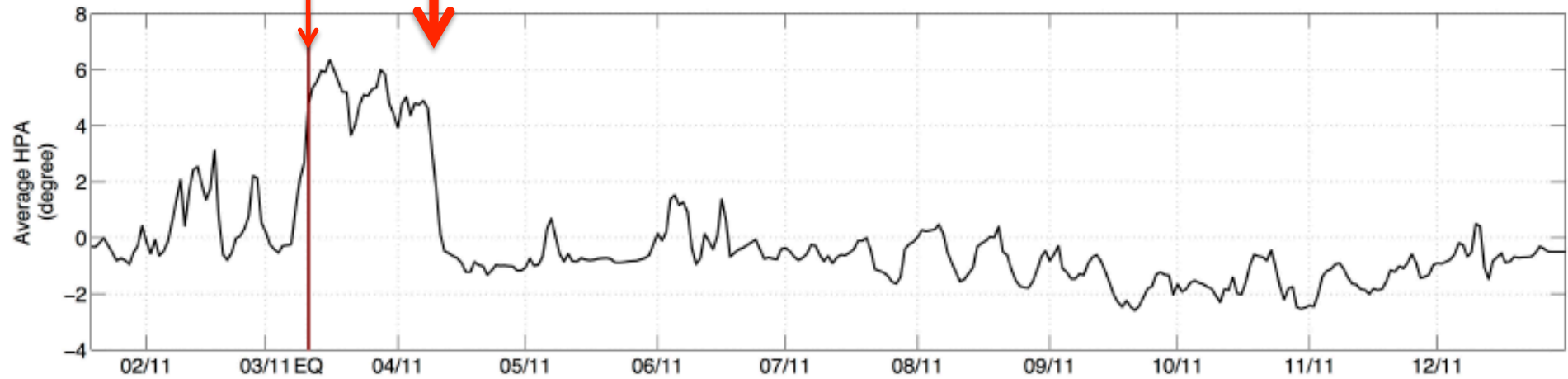
Monitoring of volcanic zones

Polarization and anisotropy change at the time of the EQ

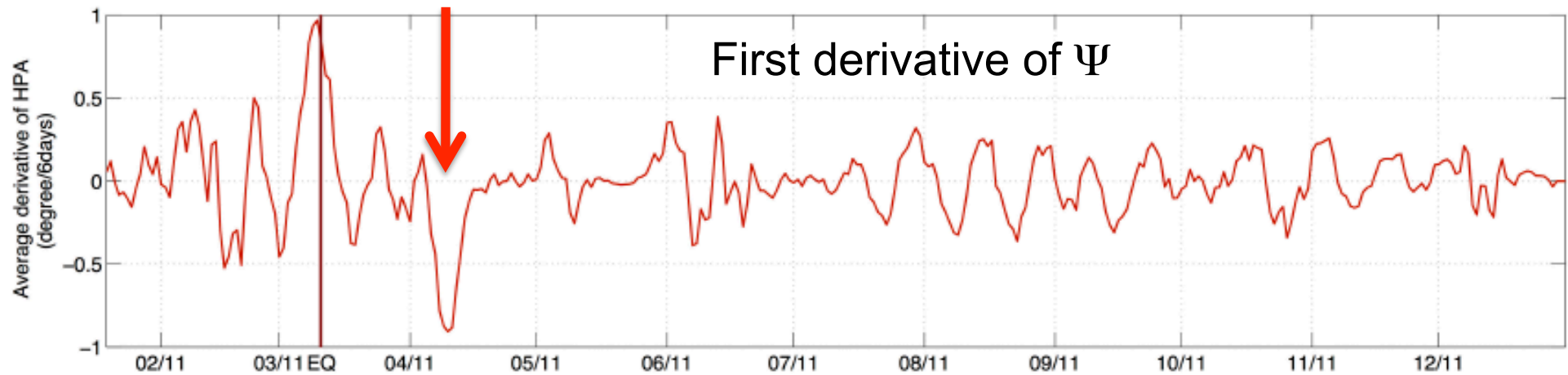


Monitoring of volcanic zones

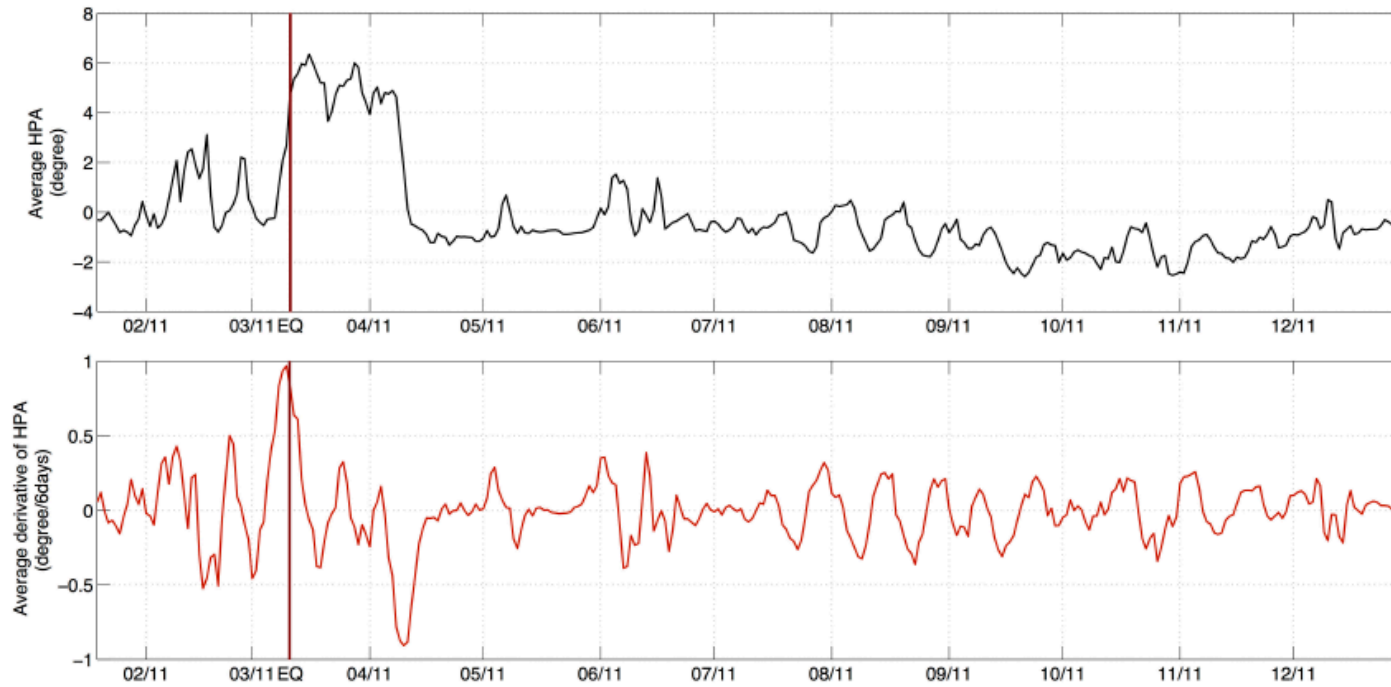
Average of the horizontal polarization anomaly Ψ



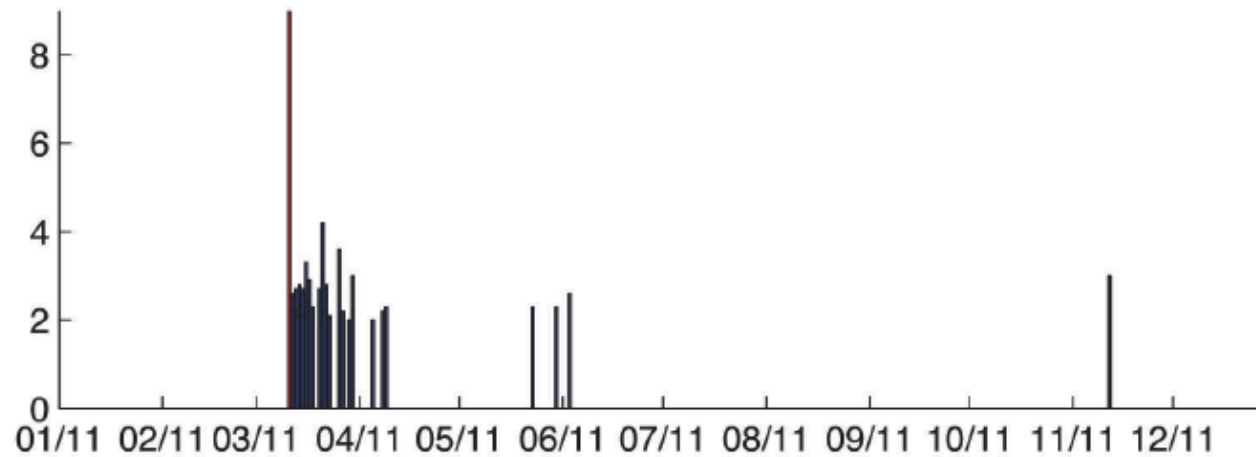
First derivative of Ψ



Monitoring of volcanic zones

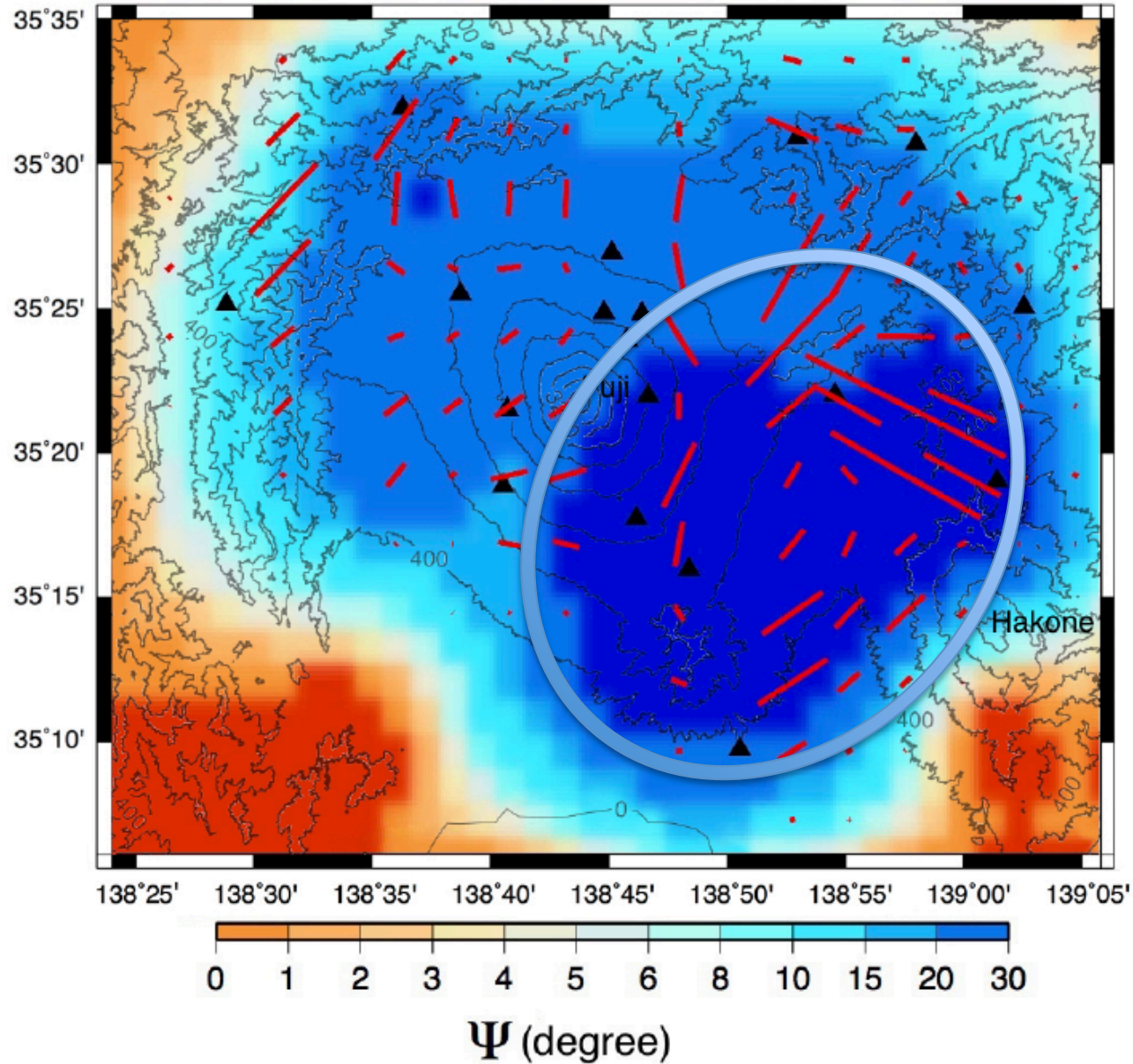


Seismicity in Hakone Area – $M_w > 2$ & depth < 10 km

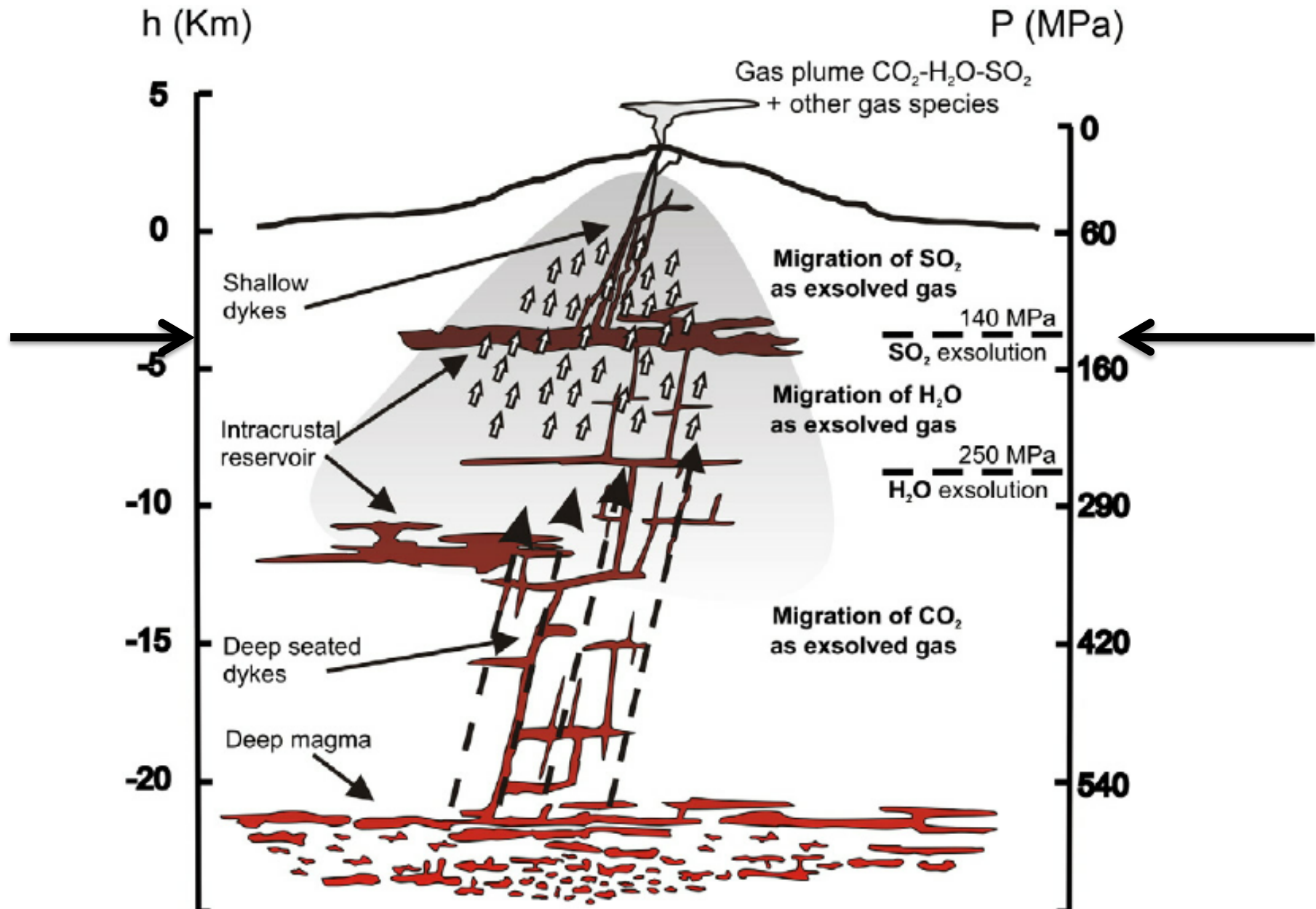


Monitoring of volcanic zones

Polarization drop and anisotropy change in April



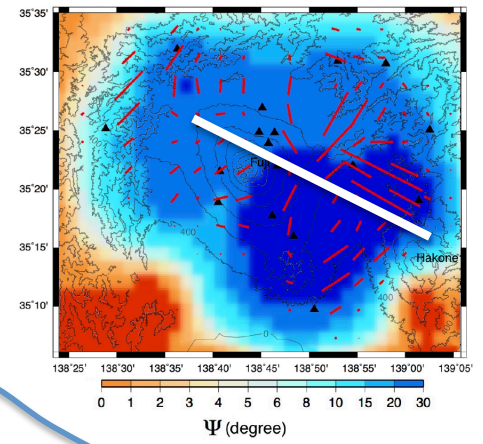
Plumbing System



Tentative interpretation

Fuji

Hakone



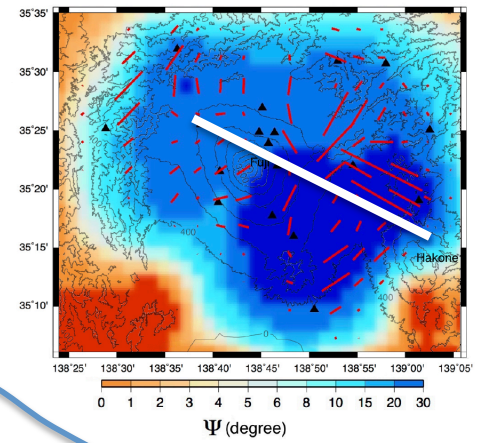
~3km



Tentative interpretation

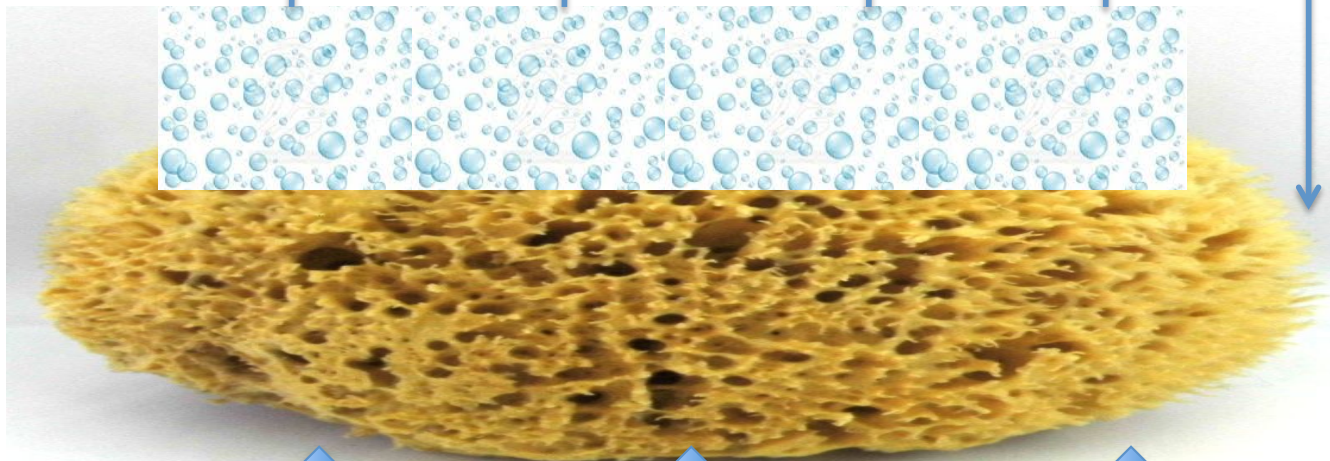
Fuji

Hakone



Porous wave

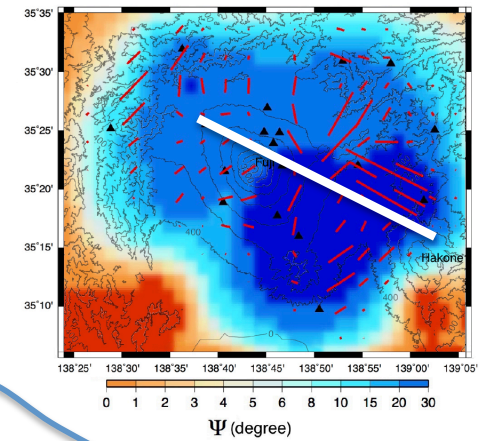
~3km



Tentative interpretation

Fuji

Hakone



$$v = 10^{-3} \text{m/s}$$

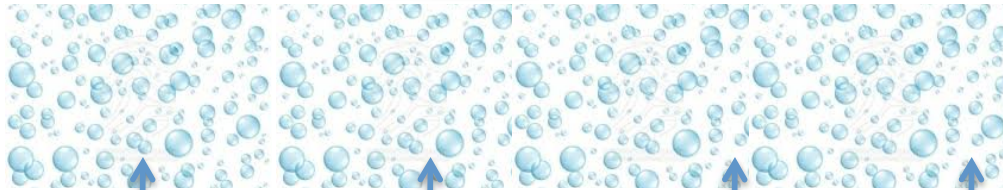
$$v = \frac{k \Delta \rho g}{\mu \Phi}$$

μ viscosity $\approx 10^{-3} \text{ Pa}\cdot\text{s}$

k permeability

Φ porosity

a inclusion dimension $\approx 0.1 \text{ m}$

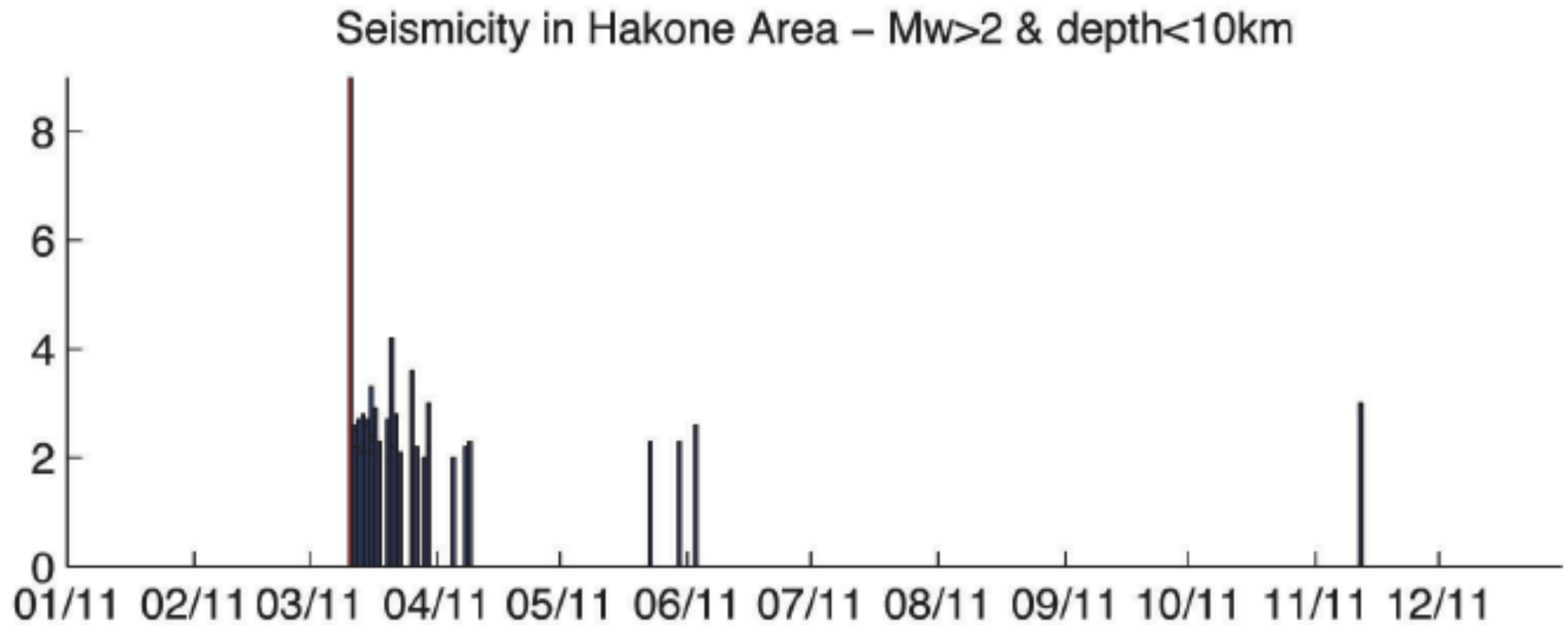


Porous wave



$\approx 3 \text{ km}$

Monitoring of volcanic zones



Conclusions

Seismic anisotropy is necessary to explain seismic data
It provides new information on many geophysical processes

•New Method for continuous monitoring of the stress field in:

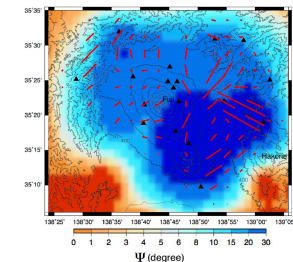
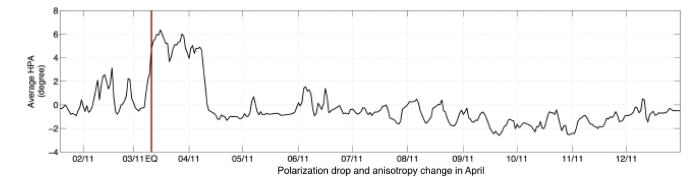
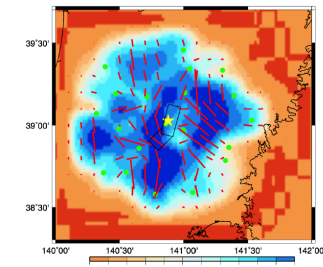
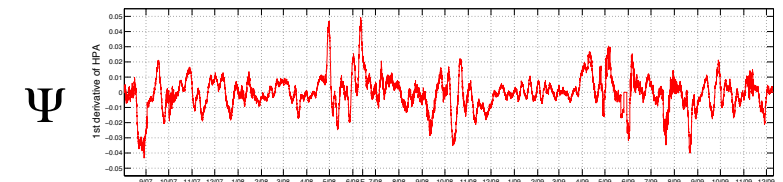
-Seismogenic zones

Significant temporal change of Ψ observed in parts of the cracked zone

- Volcanic zones: Porous wave

-Interpretation in terms of anisotropy variations and stress changes

- Other applications in oil/gas reservoirs
(Mordret et al., 2013; Tomar et al., 2016)



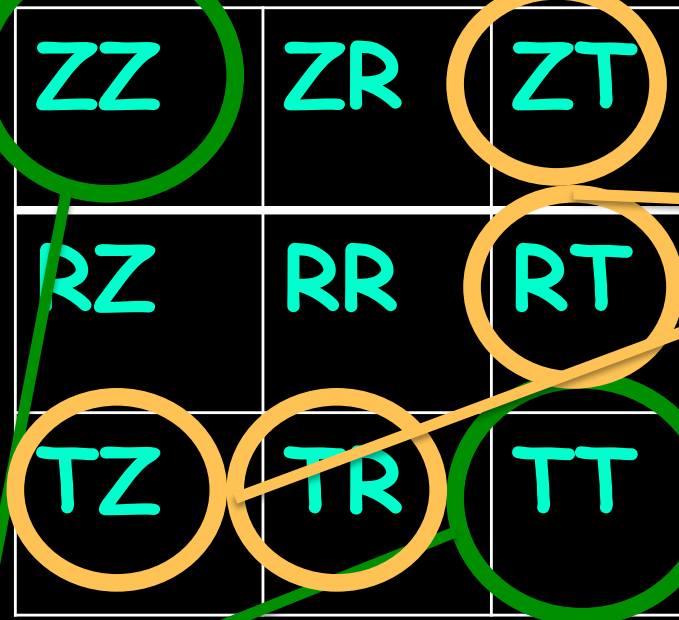
THEORY

Cross-correlation
for 2 stations i, j
and 3 components k, l

$$[C_{ij}(t)]_{kl} = \frac{\int_0^T S_{ik}(\tau) S_{jl}(t + \tau) d\tau}{\sqrt{\int_0^T S_{ik}^2(\tau) d\tau \int_0^T S_{jl}^2(\tau) d\tau}},$$

Azimuthal
Anisotropy
of Rayleigh
waves

Effect of anisotropy on C_{ij}

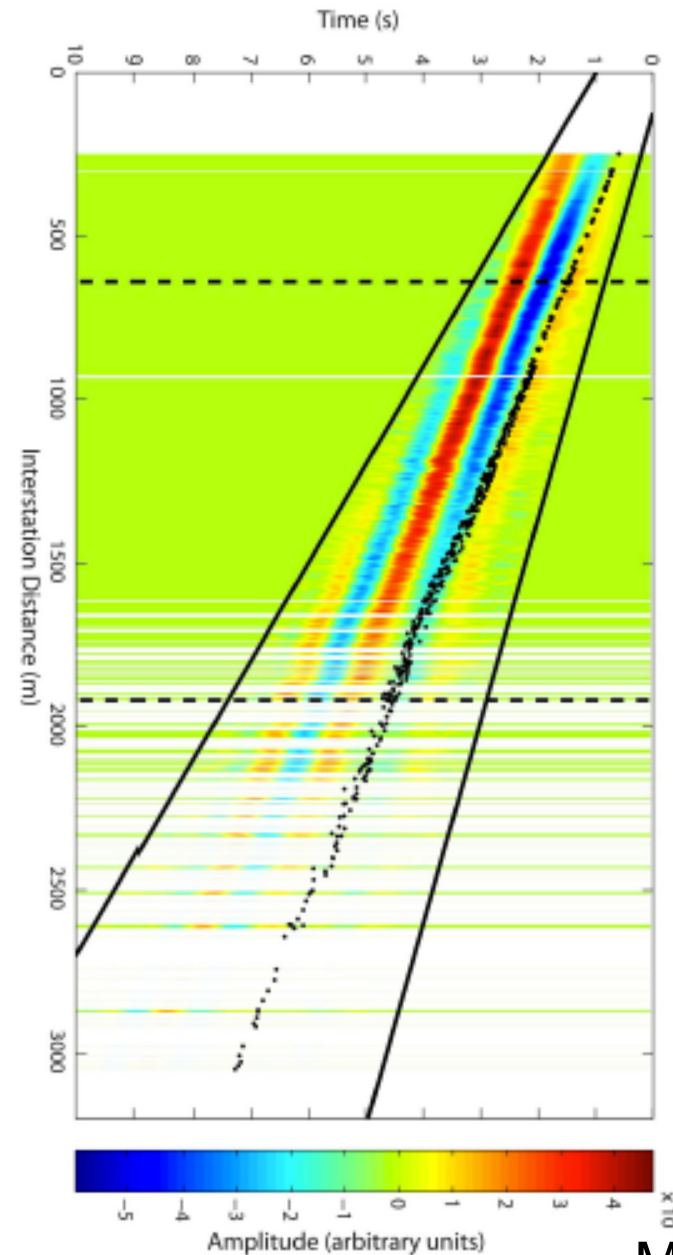
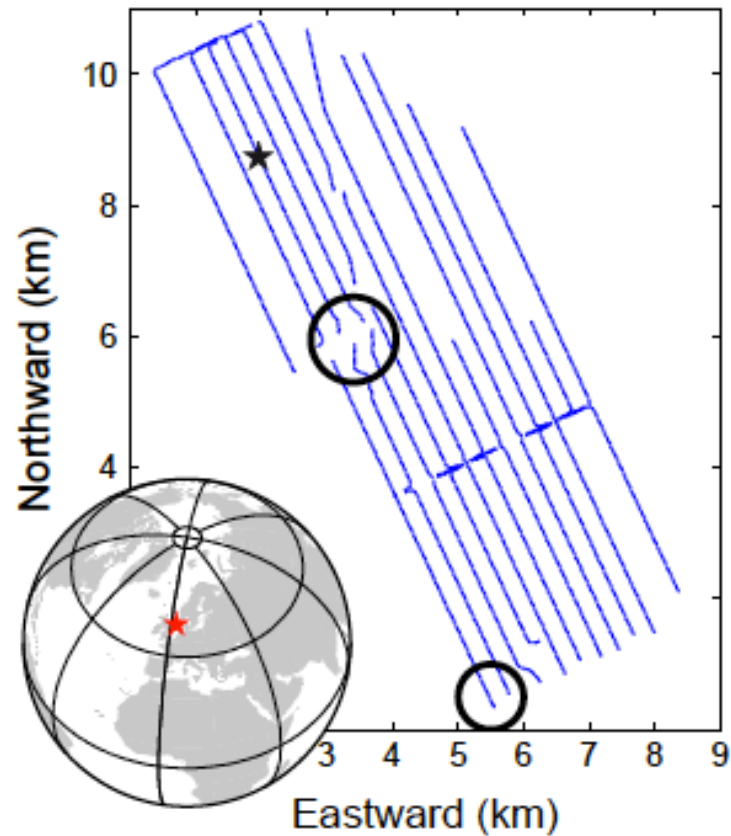


Quasi-Rayleigh
Quasi-Love
waves
polarization
anomalies

Joint Inversion
of Rayleigh
and Love waves:
Radial anisotropy

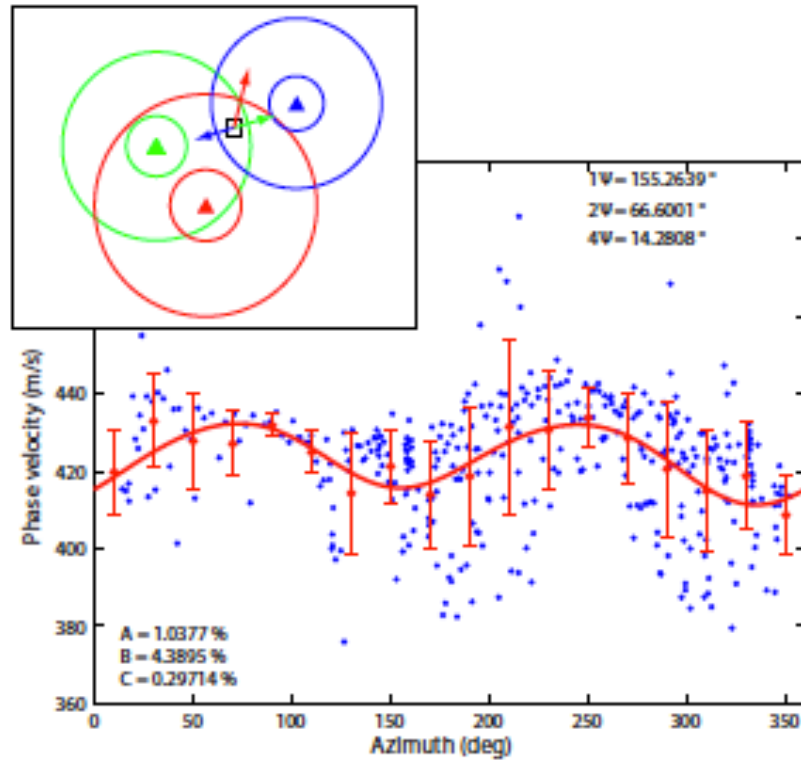
Rayleigh wave (seismic noise): Azimuthal variation on ZZ-component

Valhall LoFS network



Rayleigh wave: Azimuthal variation on ZZ-component (1-2-4- Ψ terms)

At T=0.8s



Valhall LoFS: 2- Ψ term

