

Slip rate and slip magnitudes of past earthquakes along the Bogd left-lateral strike-slip fault (Mongolia)

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SUMMARY

We carried out morphotectonic studies along the left-lateral strike-slip Bogd Fault, the principal structure involved in the Gobi-Altay earthquake of 1957 December 4 (published magnitudes range from 7.8 to 8.3). The Bogd Fault is 260 km long and can be subdivided into five main geometric segments, based on variation in strike direction. West to East these segments are, respectively: the West Ih Bogd (WIB), The North Ih Bogd (NIB), the West Ih Bogd (WIB), the West Baga Bogd (WBB) and the East Baga Bogd (EBB) segments. Morphological analysis of offset streams, ridges and alluvial fans—particularly well preserved in the arid environment of the Gobi region—allows evaluation of late Quaternary slip rates along the different faults segments. In this paper, we measure slip rates over the past 200 ka at four sites distributed across the three western segments of the Bogd Fault. Our results show that the left-lateral slip rate is $\sim 1 \text{ mm yr}^{-1}$ along the WIB and EIB segments and $\sim 0.5 \text{ mm yr}^{-1}$ along the NIB segment. These variations are consistent with the restraining bend geometry of the Bogd Fault.

Our study also provides additional estimates of the horizontal offset associated with the 1957 earthquake along the western part of the Bogd rupture, complementing previously published studies. We show that the mean horizontal offset associated with the 1957 earthquake decreases progressively from 5.2 m in the west to 2.0 m in the east, reflecting the progressive change of kinematic style from pure left-lateral strike-slip faulting to left-lateral-reverse faulting. Along the three western segments, we measure cumulative displacements that are multiples of the 1957 coseismic offset, which may be consistent with a characteristic slip. Moreover, using these data, we re-estimate the moment magnitude of the Gobi-Altay earthquake at M_w 7.78–7.95.

Combining our slip rate estimates and the slip distribution per event we also determined a mean recurrence interval of ~ 2500 – 5200 yr for past earthquakes along the different segments of the western Bogd Fault. This suggests that the three western segments of the Bogd Fault and the Gurvan Bulag thrust fault (a reverse fault bounding the southern side of the Ih Bogd range that ruptured during the 1957 earthquake) have similar average recurrence times, and therefore may have ruptured together in previous earthquakes as they did in 1957. These results suggest that the western part of the Bogd Fault system, including the Gurvan Bulag thrust fault, usually behaves in a ‘characteristic earthquake’ mode.

Keywords: Seismic cycle; Geomorphology; Earthquake dynamics; Continental neotectonics; Continental tectonics: strike-slip and transform; Asia.

1 INTRODUCTION AND TECTONIC SETTING

Western Mongolia is undergoing NNE–SSW compression related to the India–Asia collision (Molnar *et al.* 1977). The Altay mountain range and its southeastern extension, the Gobi–Altay, represent the northernmost active compressional belt in central Asia (Tapponnier & Molnar 1979; Cunningham 1998, 2005; Vassallo *et al.* 2007a). Fission track analyses carried out within the Gurvan Bogd mountain range, the easternmost part of the Gobi–Altay, show that Cenozoic transpressional tectonics started 5 ± 3 Ma (Vassallo *et al.* 2007a), uplifting an ancient neoplain that had formed during Jurassic time (Jolivet *et al.* 2007).

During the last century, between 1905 and 1957, four $M8$ earthquakes occurred in western Mongolia and the immediately adjacent area of China (Molnar & Deng 1984; Baljinnyam *et al.* 1993; Schlupp 1996) making this region one of the most tectonically active intracontinental domains in the world (Fig. 1A). These earthquakes were produced by large strike-slip faults reactivating pre-Cenozoic structures (see Florensov & Solonenko 1965; Baljinnyam *et al.* 1993; Jolivet *et al.* 2007).

Several morphotectonic studies show that the Bogd and Bulnay left-lateral strike-slip faults have late-Quaternary slip rates of ~ 1 mm yr⁻¹ (Ritz *et al.* 1995, 2006; Rizza 2010) as do strike-slip faults in Altay (Walker *et al.* 2006; Vassallo 2006; Nissen *et al.* 2009a), consistent with geodetic measurements (Calais *et al.* 2003; Vergnolle *et al.* 2003). These low slip rates imply long recurrence intervals for large earthquakes, consistent with palaeoseismological investigations (Schwartz *et al.* 1996; Prentice *et al.* 2002; Schwartz *et al.* 2007; Ritz *et al.* 2009; Rizza 2010). Considering these long recurrence intervals, the 20th century Mongolian seismicity has been described as a seismic cluster involving mechanical coupling between large strike-slip faults via viscoelastic post-seismic stress transfer in the lower crust (Chery *et al.* 2001; Pollitz *et al.* 2003; Vergnolle 2003).

Among the four historical earthquakes, the 1957 Gobi–Altay event is the best known because it is the most recent event, was exceptionally mapped shortly after the event (Florensov & Solonenko 1965) and is associated with a complex pattern of surface rupture. The earthquake occurred along the fault zone that bounds the northern Gobi–Altay range and the seismotectonics features were first described in Florensov & Solonenko (1965). The main 1957 rupture is along a 260-km-long left-lateral strike-slip Bogd Fault with an average trend of N100°E. About 100 km of reverse faulting along the Ih Bogd and Baga Bogd massifs also occurred simultaneously with the Bogd Fault in 1957 (Fig. 1B).

Kurushin *et al.* (1997) resurveyed and updated the offsets along the entire 1957 surface rupture providing a detailed map of the slip distribution throughout the entire region of surface breaks, including 260 km of rupture along the Bogd Fault and 100 km of ruptures along other faults. Kurushin *et al.* (1997) estimated an average horizontal displacement of 3–4 m, with a maximum offset of 7 m. The overall 1957 fault surface ruptures are consistent in terms of geometry, kinematics and slip distribution with the accumulated topography (Bayasgalan *et al.* 1999a; Ritz *et al.* 2006).

Morphological analysis of offset streams, ridges and alluvial fans—particularly well preserved in the arid environment of the Gobi region—allows evaluation of late-Quaternary slip rates along the different faults segments (e.g. Ritz *et al.* 2006). Vertical slip rates have been estimated and are similar (0.1 – 0.2 mm yr⁻¹) on several of the reverse fault segments, at Bitut (Vassallo *et al.* 2007b), Dalan Turuu (Hanks *et al.* 1997) and Gurvan Bulag (Carretier *et al.* 2002;

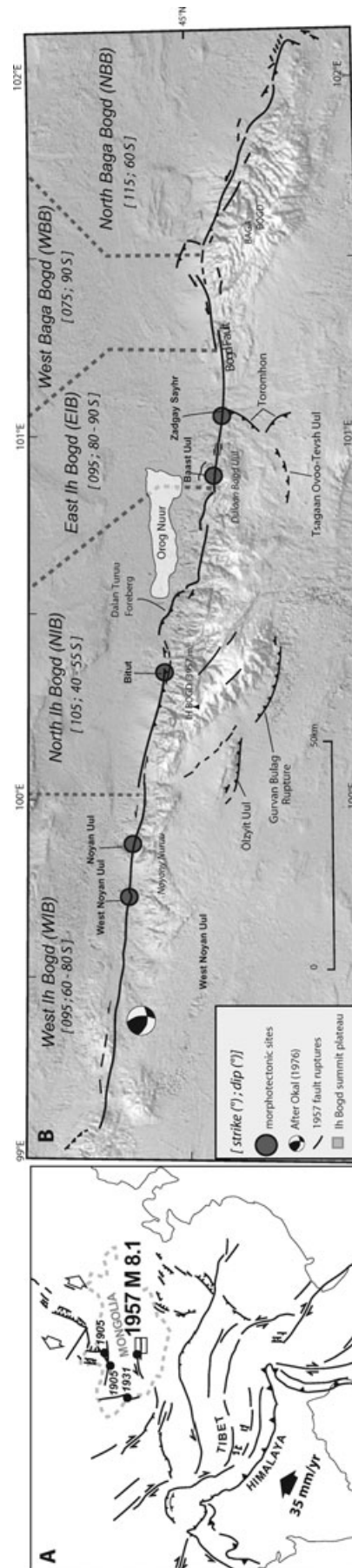


Figure 1. (A) Principle active faults within the India–Eurasia collision. Black dots depict the epicentre of earthquakes with $M > 8$ in Mongolia during the last century. The box represents the study area presented in this paper. (B) Map for the 1957 Gobi–Altay surface rupture (modified from Baljinnyam *et al.* 1993). We subdivided the fault into five segments based on the fault orientation (from Kurushin *et al.* 1997). Gray dots represent morphotectonic sites.

Ritz *et al.* 2003; Vassallo *et al.* 2005). In addition, the horizontal slip rate has been estimated on the Bogd Fault at Noyan Uul, where an alluvial fan displaced by ~ 110 m was dated at ~ 80 ka yielding a maximum left-lateral slip rate of ~ 1.2 mm yr $^{-1}$ (Ritz *et al.* 1995).

In this paper, we provide new estimates on the horizontal slip rate along the Bogd Fault at four sites: Noyan Uul, West Noyan Uul, Bitut and Baast Uul (Fig. 1B). We refine previous morphotectonics interpretations at Noyan Uul (i.e. Ritz *et al.* 1995, 2006) and present new field-based measurements of lateral offset features within the three other sites. This work also refines published *in situ* ^{10}Be data in Noyan Uul and Bitut (i.e. Ritz *et al.* 1995, 2006; Vassallo *et al.* 2007a) together with new *in situ* ^{10}Be and luminescence age constraints. We also study the slip distribution along the fault at Noyan Uul, Bitut and Baast Uul and Zadgay Sayhr, a site located 11 km east of Baast Uul. Comparing 1957 offsets with cumulative displacements allows us to estimate the distribution of slip associated with past earthquakes. Combining slip rate estimates and the slip distribution per event allows us to estimate the mean recurrence intervals of past earthquakes.

2 SLIP RATE ANALYSIS ALONG THE BOGD FAULT

The 260-km-long Bogd Fault can be subdivided into five geometric segments based on the variation in strike direction (Fig. 1B and Table 1). West to east these segments are, respectively: the West Ih Bogd segment (WIB), the North Ih Bogd segment (NIB) which is characterized by a significant reverse component of slip, the East Ih Bogd segment (EIB), the West Baga Bogd segment (WBB) and the North Baga Bogd segment (NBB). The epicentre of the 1957 earthquake is located along the West Ih Bogd segment (Florensov & Solonenko 1965; Okal 1976).

In this study, we focus our morphotectonic analyses within the WIB, NIB and EIB segments. To determine horizontal slip rates along these three segments, we analyse cumulative displacements of alluvial surfaces (~ 20 to ~ 300 m) that we dated using *in situ* produced ^{10}Be (e.g. Bierman 1994; Ritz *et al.* 1995; Braucher *et al.* 1998; Siame *et al.* 2000) and Optically Stimulated Luminescence (OSL, e.g. Le Dortz *et al.* 2009; Nissen *et al.* 2009b; Rizza *et al.* 2011). We also use published ^{10}Be data from two sites along the WIB and NIB segments to refine our previous slip rate estimates at Noyan Uul (Ritz *et al.* 1995; Vassallo *et al.* 2005; Ritz *et al.* 2006) and to estimate the horizontal slip rate at Bitut (Vassallo *et al.* 2007b).

Because previously published data were obtained at the Tandatron [the French Accelerator Mass Spectrometer (AMS, Raisbeck & Yiou 1987)], they have been recalculated to be compared with the newly acquired data acquired from ASTER, the new French AMS (Arnold *et al.* 2010). This means the use of the same standard value, the NIST Standard Reference materials with an assigned $^{10}\text{Be}/^9\text{Be}$ ratio of $2.79 \cdot 10^{-11}$, and a ^{10}Be half-life of $1.387 \cdot 10^6$ yr $^{-1}$ (Chmeleff

et al. 2010; Korschinek *et al.* 2010). Beryllium production rates have been calculated using the Stone (2000) scaling polynomial with a high-latitude sea level (SHL) spallation production rate of 4.49 at $\text{g}^{-1} \text{yr}^{-1}$. All calculations are performed using attenuation lengths of 160, 1500 and 5300 g cm^{-2} with associated relative contributions to the total production rate of 97.85 per cent, 1.50 per cent and 0.65 per cent for neutrons, slow muons and fast muons, respectively (Braucher *et al.* 2003).

To estimate surface-exposure ages, we use both ^{10}Be concentrations at the surface and its distribution with depth (See Appendix A). Exposure ages are better defined using depth profiles, allowing the determination of denudation rate and inheritance (e.g. Repka *et al.* 1997). Exposure ages obtained from surficial samples were used only if ages given by depth profiles were poorly constrained. When sample concentrations appear to be at steady state, an integration of effective exposure time was determined (Lal 1991). The results for the ^{10}Be analysis are presented in Appendix A.

The slip rate ranges from minimum to maximum values corresponding to $v_{\min} = (\Delta_x - \sigma_{\Delta_x}) / (\Delta_t + \sigma_{\Delta_t})$ and $v_{\max} = (\Delta_x + \sigma_{\Delta_x}) / (\Delta_t - \sigma_{\Delta_t})$ and, where Δ_x is the measured offset of a geomorphic feature across the fault and Δ_t is the age estimate for the feature, with the corresponding errors σ_{Δ_x} and σ_{Δ_t} , respectively. When the slip rate is estimated using offset terrace risers, the ages of the upper and lower terraces provide maximum and minimum ages for the risers, respectively (e.g. Hanks & Thatcher 2006; Cowgill 2007; Gold *et al.* 2009).

2.1 The West Ih Bogd segment

2.1.1 West Noyan Uul site

The West Noyan Uul site is situated along the eastern part of the WIB segment within an area of lower relief (Fig. 1B). The rupture pattern of the Bogd Fault becomes more complex at this location, with several obliquely trending segments (Fig. 2A). As described by Kurushin *et al.* (1997), there are two branches of the fault at this location. The main fault is expressed as a west-trending, north-facing scarp, which cuts alluvial surfaces. Farther south, the second branch forms the northern boundary of the massif, and trends in a southwesterly direction for about 3 km, and then curves towards the northwest for about 1.5 km before rejoining the main fault. This fault network forms a releasing bend (east side of Fig. 2A), which is consistent with the numerous normal fault scarps described by Kurushin *et al.* (1997) in this area. The region farther west, where the two faults strands rejoin may be described as a restraining bend, consistent with the thrust fault scarp in this area.

Estimating offsets. The West Noyan Uul site is located along the main EW trending fault where a system of alluvial terraces is left-laterally displaced with a small reverse component. Figs 2(B) and (C) show the aerial photograph of the site and the corresponding digital elevation model (DEM) obtained from our kinematic Global Positioning Systems (GPS) survey. On the west bank of the stream, we identified a series of four displaced terraces (T3, T2, T1 and T0). T3 terrace has a flat morphology with some boulders standing above the surface. T2 and T1 terraces present similar morphology with relatively flat surfaces, above which stand large boulders encased in the alluvial sediments that underlie the surfaces. T0, the youngest terrace, is characterized by a well-preserved bar-and-swale morphology. South of the fault, we have subdivided the T0 surface into two subunits: T0-1 and T0-2. The T0-2 surface is more eroded and shows evidence of incision by T0-1.

Table 1. Characteristics of the five main segments of the Bogd Fault. The dips reported along the fault are from the Kurushin *et al.* (1997) estimates.

Segment	Longitude position	Length (km)	Strike direction	Dip ($^{\circ}$)
WIB	99.09 $^{\circ}$ E–100.03 $^{\circ}$ E	74	N095 $^{\circ}$ E	60 $^{\circ}$ –80 $^{\circ}$
NIB	100.03 $^{\circ}$ E–100.86 $^{\circ}$ E	67	N105 $^{\circ}$ E	40 $^{\circ}$ –55 $^{\circ}$
EIB	100.86 $^{\circ}$ E–101.16 $^{\circ}$ E	23	N095 $^{\circ}$ E	80 $^{\circ}$ –90 $^{\circ}$
WBB	101.16 $^{\circ}$ E–101.53 $^{\circ}$ E	28	N075 $^{\circ}$ E	80 $^{\circ}$ –90 $^{\circ}$
NBB	101.53 $^{\circ}$ E–102.07 $^{\circ}$ E	50	N115 $^{\circ}$ E	60 $^{\circ}$

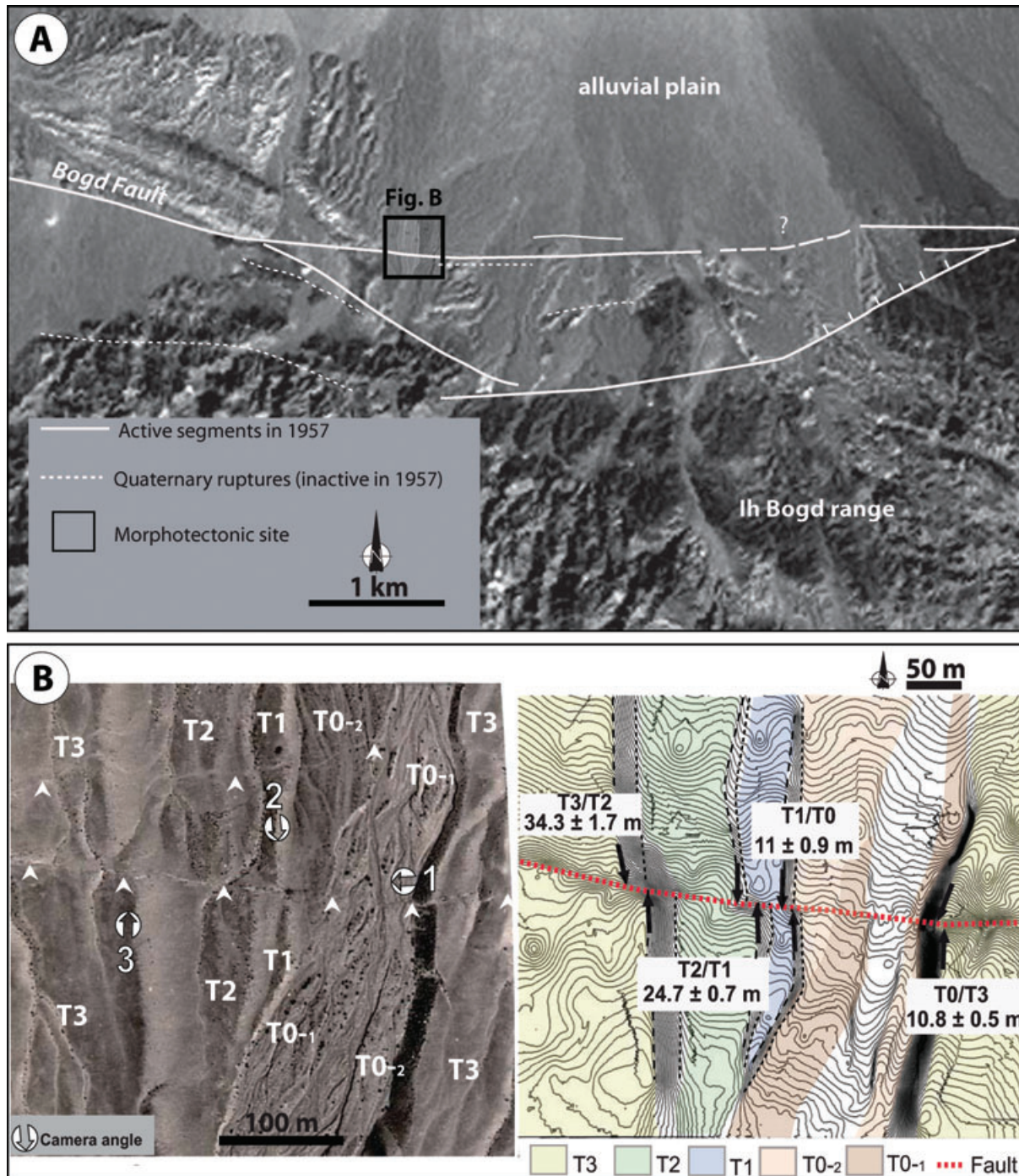


Figure 2. Morphological analyses at West Noyan Uul. (A) Hillshaded SRTM image (<http://srtm.csi.cgiar.org>) of the West Noyan Uul site, with 1957 surface ruptures mapped in white. The black box represents the study area. (B) Aerial photography (1958) showing the terraces T0, T1, T2 and T3 displaced along the Bogd Fault. Note that we subdivided the T0 terrace into two units T0-1 and T0-2. The locations of field photos in Fig. 3 are shown here: (1) Fig. 3(A), (2) Fig. 3(B) and (3) Fig. 3(C). (C) Digital elevation model and the morphotectonic interpretations, based on aerial photographs and reconnaissance field mapping, with the piercing point used to measure the offset risers (black arrows).

On the west bank of the stream channel there is a series of three risers between the four terrace surfaces. Using the DEM that we generated from our survey, we estimate cumulative displacements of 34.3 ± 1.7 m, 24.7 ± 0.7 m and 11 ± 0.9 m for T3/T2, T2/T1 and T1/T0 risers, respectively (see Fig. 2C). The east bank has experienced more severe erosion, and the only preserved terrace riser is between T3 (the highest and oldest terrace) and T0-1 (the youngest surface terrace). T2 and T1 terraces have been com-

pletely removed. T3/T0-1 riser offset is 10.8 ± 0.5 m, similar to the T1/T0-2 riser offset measured on the west bank. This shows that the offset of the T3/T0 riser accumulated after that the lower terrace was abandoned.

Dating markers. To date the offset terrace risers, we measured ^{10}Be concentrations surface samples collected in the four alluvial terraces and one soil pit dug into the T2 terrace. Fig. 4(A) shows the results of ^{10}Be surface concentrations for the different

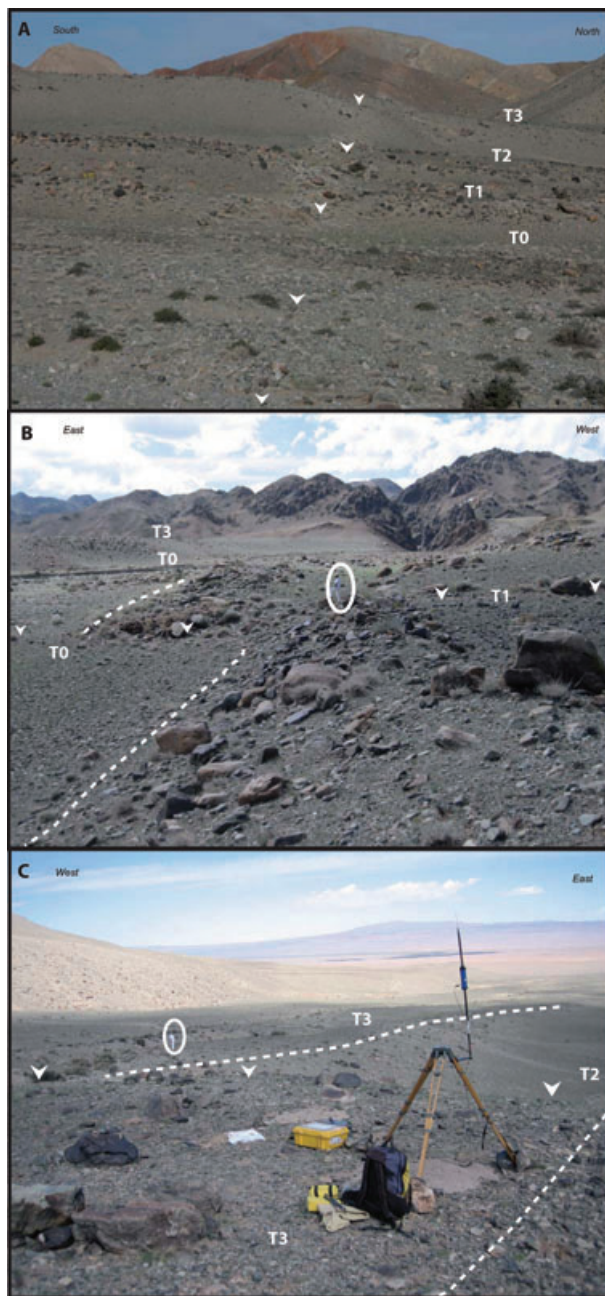


Figure 3. Photographs depicting field relationships at the West Noyan Uul site. (A) Field photo looking west from T0 terrace in the right bank showing the stepped terraces displaced along the 1957 rupture surface. (B) Field photo (looking south) from T1 terrace showing the T0/T1 riser along the 1957 rupture, with piercing lines (white dashed lines) corresponding to the base of riser T0/T1. Note that a person making scale is surrounded by the white circle. (C) Field photo (looking north) from T3 terrace showing the T3/T2 riser along the 1957 rupture, with piercing lines corresponding to the top of riser T3/T2. Note that a person making scale is surrounded by the white circle.

terraces. For each surface, we collected pieces of rock samples from the tops of large quartz-rich boulders (diameter > 1 m) and one amalgamated sample made of ~40 pebbles of quartz with centimetre size of ~1–2 cm (i.e. Repka *et al.* 1997). The T3 terrace has the highest ^{10}Be concentrations, which range from $1.75 \pm 0.07 \text{ Mat g}^{-1}$ to $1.58 \pm 0.03 \text{ Mat g}^{-1}$ (considering sample M07–01 as an outlier). The ^{10}Be concentrations in T2 terrace are between $1.61 \pm$

0.02 Mat g^{-1} and $1.02 \pm 0.02 \text{ Mat g}^{-1}$, considering sample M07–11 as an outlier. We found ^{10}Be concentrations between $0.96 \pm 0.03 \text{ Mat g}^{-1}$ and $0.49 \pm 0.04 \text{ Mat g}^{-1}$ for T1, and ^{10}Be concentrations between $0.65 \pm 0.01 \text{ Mat g}^{-1}$ and $0.25 \pm 0.01 \text{ Mat g}^{-1}$, for T0 terraces (T0–1 and T0–2 surfaces and considering sample M07–26 as an outlier). T0, T1 and T2 terraces present high dispersions in the ^{10}Be surface concentrations that could be due to variable inheritances. ^{10}Be surface concentrations on T3 show less dispersion and are close to T2 concentrations, suggesting that T3 is approaching steady-state equilibrium.

Fig. 4(B) shows the depth distribution of ^{10}Be concentrations within the T2 terrace. Two different stratigraphical units are observed in the soil pit: a 1-m-thick debris flow unit made of quartz-rich cobbles (10–50 cm diameter) in a sandy-silty matrix and overlays a gravelly unit made of centimetre-scale clasts in a sandy-silty matrix. We collected at different depths one single clast (or fragment of clast) along a 130-cm-depth profile. The concentrations do not show an exponential decrease with the depth as expected. Starting at 1.00 Mat g^{-1} at the surface, the ^{10}Be concentration decreases rapidly with depth and reaches $0.6\text{--}0.7 \text{ Mat g}^{-1}$ at the bottom of the first unit at 95 cm depth. The concentration then increases to 0.85 Mat g^{-1} at 130 cm in the second unit. This depth distribution of ^{10}Be concentrations is difficult to interpret, and is probably due to the fact that the concentrations correspond to individual rock samples (pebbles) and not amalgamated samples. The increase at 130 cm depth is consistent with the change observed in the stratigraphy at 100 cm depth, and can be interpreted either as pre-exposure before the deposition of the debris flow unit or as a different inheritance.

Taking into account that the surface samples and the profile present variable amount of inheritance, we use the lowest ^{10}Be concentrations for each surface to minimize the effect of exposure history prior to deposition (e.g. Vassallo *et al.* 2007a). We then determine minimum ages of $14.6 \pm 0.7 \text{ ka}$, $28.7 \pm 2.4 \text{ ka}$ and $59.8 \pm 1.2 \text{ ka}$ for the T0, T1 and T2 terraces, respectively. For the T3 terrace, considering that the surface has reached a steady-state concentration, we calculate a maximum denudation rate of 8.6 m Myr^{-1} . We then determine an effective irradiation time (T_{eff}) of 85.3 ka , that is the minimum time required to reach the steady state for a given denudation rate (Lal 1991).

Estimating slip rates. As mentioned in Section 2, a conservative approach to calculate the slip rates from offset terrace risers is to divide the offsets by the ages of the upper and lower terraces. We estimate maximum horizontal slip rates by dividing the riser offsets by the ages of the lower terraces. Offsets recorded by the T3/T2 ($34.3 \pm 1.7 \text{ m}$), T2/T1 ($24.7 \pm 0.7 \text{ m}$), T1/T0–2 ($11 \pm 0.9 \text{ m}$) and T3/T0–2 ($10.8 \pm 0.5 \text{ m}$) risers yield maximum slip rates of $0.57 \pm 0.04 \text{ mm yr}^{-1}$, $0.87 \pm 0.1 \text{ mm yr}^{-1}$, $0.76 \pm 0.1 \text{ mm yr}^{-1}$ and $0.74 \pm 0.07 \text{ mm yr}^{-1}$, respectively. We then estimate minimum slip rates by dividing the riser offsets by the ages of the upper terraces. Minimum slip rates of $0.39 \pm 0.07 \text{ mm yr}^{-1}$, $0.13 \pm 0.01 \text{ mm yr}^{-1}$, $0.41 \pm 0.02 \text{ mm yr}^{-1}$ and $0.40 \pm 0.02 \text{ mm yr}^{-1}$ are calculated for the T1/T0–2, T3/T0–2, T2/T1 and T3/T2 risers, respectively. Note that the T3 surface is at steady state, and therefore the age is a minimum, and we have no maximum age estimate for this surface.

However, at West Noyan Uul, offset risers among the terrace pattern show that the offsets were recorded after the abandonment of the lower terraces. For instance the T1/T0 riser offset in the west bank is equal to the T3/T0 riser offset in the east bank showing that these offsets were recorded after the abandonment of T0. For these reasons, we give preference to our slip rate estimates using the ages of the lower terraces.

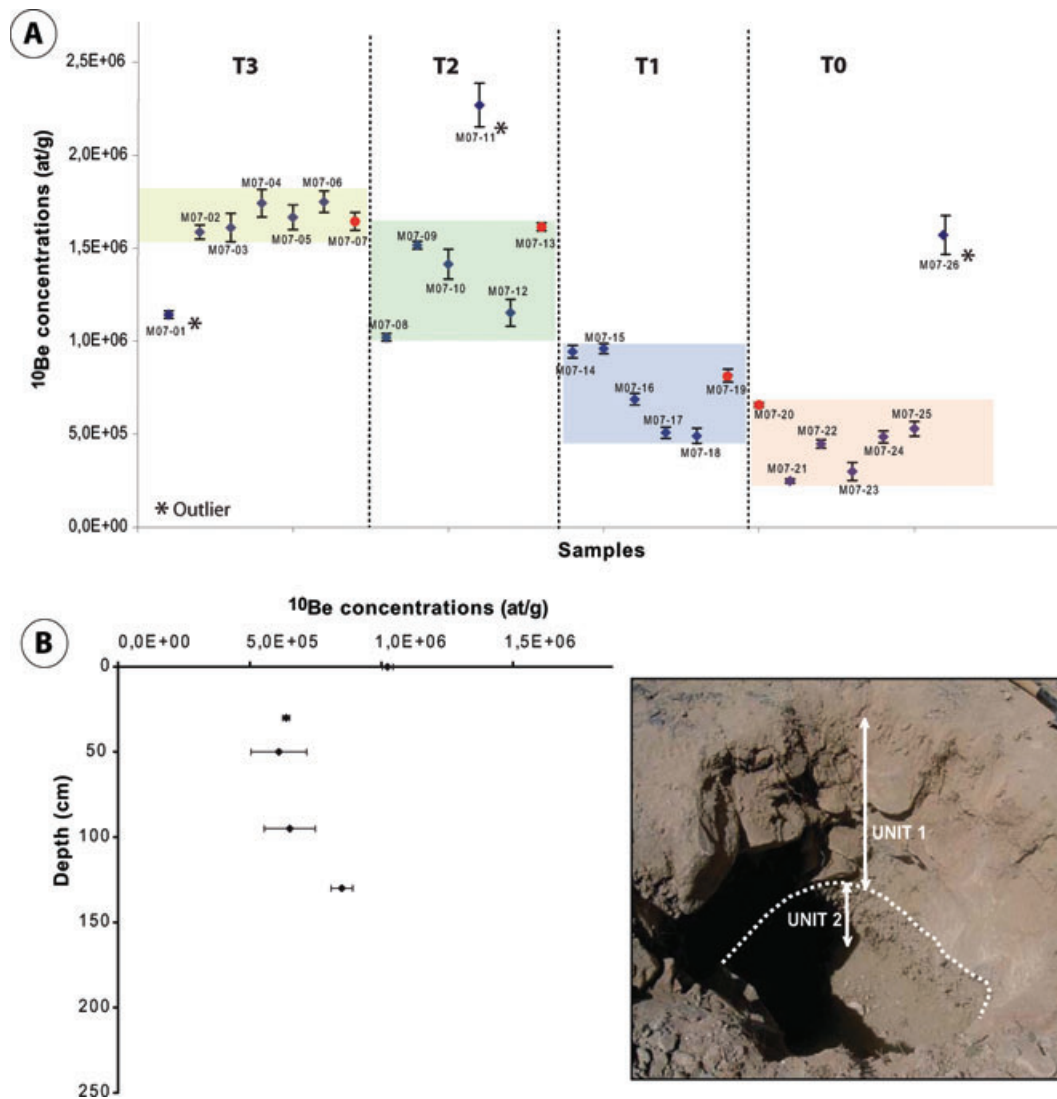


Figure 4. ^{10}Be analyses at West Noyan Uul. (A) ^{10}Be surface concentrations for the T3, T0, T2 and T1 terraces. The blue dots are for the boulders and red dots are for the amalgamated samples collected during the 2007 fieldtrip. (B) Depth distribution of ^{10}Be concentrations within T2 terrace, at West Noyan Uul site. Field photography of the soil pit within the T2 terrace at West Noyan Uul (taken in 2007). Two units have been identified by a stratigraphic contact, marked by the white dashed line.

2.1.2 Noyan Uul site

The Noyan Uul site is situated within the Noyony Nuruu massif, immediately west of the Ih Bogd massif (Fig. 1B). Here, the N095°E trending Bogd Fault separates a mountainous area to the south from the alluvial plain to the north (Fig. 5A). Along the fault, three alluvial fan surfaces (S0, S1 and S2) are displaced. S0 surface, the youngest one, still shows a bar-and-swale morphology whereas the S1 and S2 surfaces are flatter, deeply incised and there are few remaining standing boulders. The incisions are deeper and more widely spaced on S2 than on S1. The overall pattern reflects the classic evolution of alluvial surfaces with a short-wavelength/low-amplitude topographic signal covered by boulder fields on S0, evolving towards long-wavelength/high-amplitude topographic signal on which the number of standing boulders diminishes gradually on S1 and S2 (e.g. Ritz *et al.* 2006). On the south side of the fault, within the canyon incision, we mapped remnant surfaces corresponding to S0, S1 and S2 surfaces (Fig. 5B).

Estimating offsets. Using 1:25 000 scale aerial photographs from a 1958 Russian–Mongolian airborne survey, Ritz *et al.* (1995; 2006) estimated the displacements of S1 and S2 surfaces to be 110 ± 10 m and 220 ± 10 m, respectively. To estimate more accurately the cumulative displacements, we surveyed the area with Real Time Kinematic (RTK) GPS to produce a DEM that we combine with aerial photographs (Fig. 5C). We estimated offsets using the offset risers observed on both sides of the fault. Using both tops and bottoms of S1/S0 and S2/S1 risers, we calculated mean values (see Fig. 2B, note that the riser S2/S1 south of the fault is wider than S1/S0).

The riser between the S0 surface and the present streambed is offset by 19 ± 1.0 m (Fig. 5C). Using the S1/S0 riser and the S2/S1 riser, we measured offsets of 51 ± 6 m and 136 ± 14 m, respectively. To estimate the total offset of S1 and S2 surfaces, we also use their apices (defined by both the radial rill patterns incising the surfaces and the contours lines) with respect to the width of the catchment of

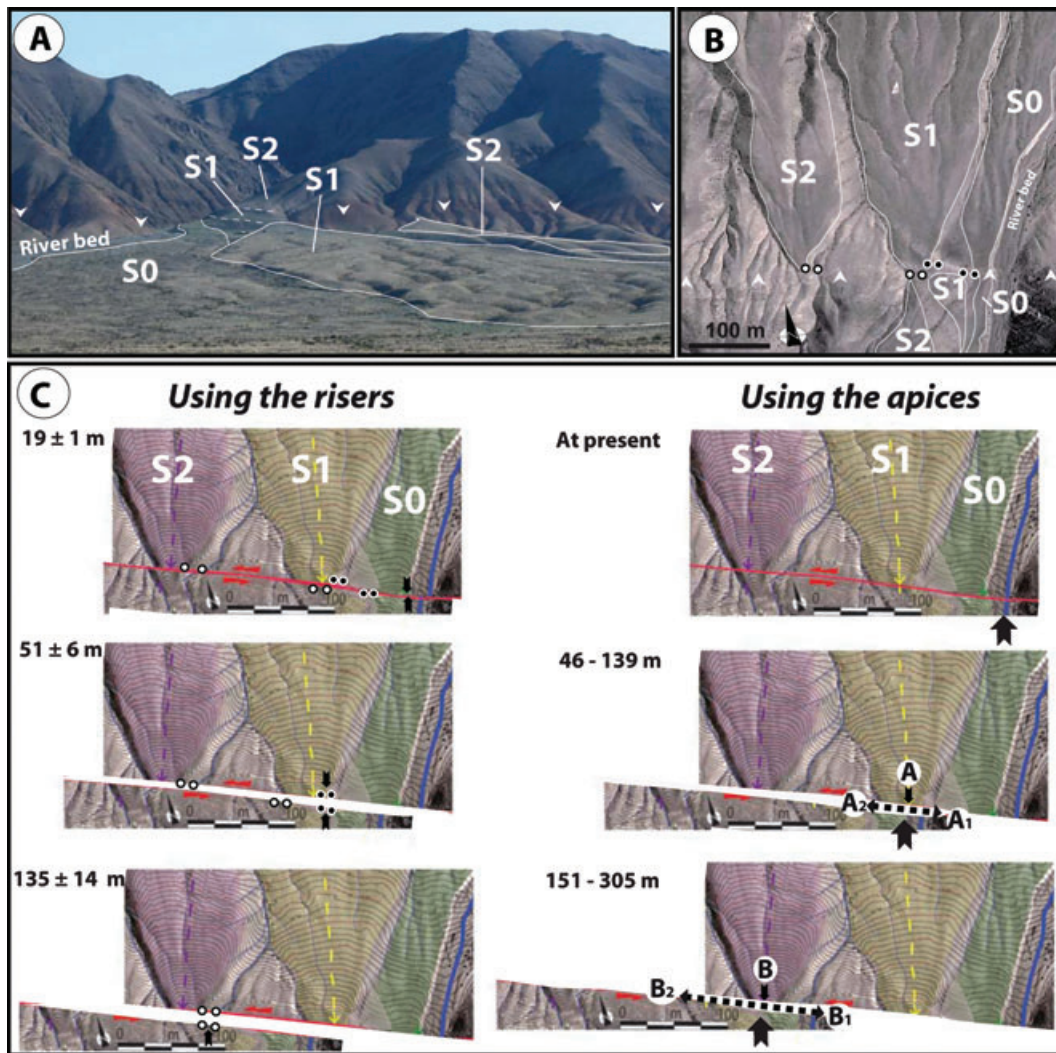


Figure 5. Morphological analyses at Noyan Uul. (A) Photography showing the Noyony Nuruu Massif and the three alluvial fans (S0, S1 and S2) at the Noyan Uul site. White triangles represent the fault line in the landscape. (B) Aerial photography (1958) showing the three alluvial fans displaced along the Bogd Fault. Three alluvial surfaces S0, S1 and S2 are in the north part of the fault while the remnant surface of S1 and S2 surfaces are identified in the western part of the canyon incision, south of the fault. Piercing points for the S1/S0 riser correspond to the white dots filled in black and the piercing points for the S2/S1 riser correspond to the black dots filled in white. (C) Digital elevation models and the morphotectonic reconstruction using the risers or apices of the alluvial fans. Note that the DEM was calculated by interpolation of 2671 data points.

the outlet axis. The maximum value is defined by the incision of the stream within the bedrock in the eastern part of the canyon, and the minimum value by the riser between alluvial surfaces in the western part of the canyon incision. Indeed, the centre of the drainage at the time of S1 deposition may have been located west of the present-day centre of the drainage, and the drainage may have incised along the eastern side of the S1 remnant surface fan. For the S1 surface, displacements range from 139 m to 46 m and correspond to the piercing points A-A₁ and A-A₂, respectively (Fig. 5C). For the S2 surface, displacements range from 305 m to 151 m and correspond to the piercing points B-B₁ and B-B₂, respectively.

Our estimate for the S1 surface displacement using the axis of incision is similar to the previously published estimate of ~ 110 m (Ritz *et al.* 1995, 2006), but with higher uncertainties. This can be explained by the distortion of the aerial photographs that we corrected using the DEM. On the other hand, the previous estimate for S2 cumulative displacement was based on the offsets of two large stream incisions in the alluvial fans relative to streams upslope of

the fault. Ritz *et al.* (1995, Fig. 2) considered these two large stream incisions as contemporaneous to S2 deposition.

Dating markers. To date the fan surfaces, we use *in situ* produced ^{10}Be data from Ritz *et al.* (1995) and Vassallo *et al.* (2005) for surfaces S1 and S2 together with new data from two amalgamated surface samples from S1 and S2 alluvial fans and one depth profile within S2 surface. These two amalgamated samples were made of ~ 40 quartz pebbles with centimetre size of 1–2 cm. Fig. 6(A) shows that the ^{10}Be concentrations in S1 surface range between $1.04 \pm 0.11 \text{ Mat g}^{-1}$ and $1.37 \pm 0.14 \text{ Mat g}^{-1}$. For the S2 surface, the ^{10}Be concentrations range between $1.56 \pm 0.17 \text{ Mat g}^{-1}$ and $1.83 \pm 0.12 \text{ Mat g}^{-1}$, considering sample DVII-1 as an outlier. The ^{10}Be concentrations for the two amalgamated samples are $1.49 \pm 0.04 \text{ Mat g}^{-1}$ and $2.08 \pm 0.06 \text{ Mat g}^{-1}$, on S1 and S2 surfaces, respectively. The offset of surface S2 being more than twice that of surface S1, the S2 concentration should be much larger than the S1 concentration if we assume a constant slip rate. However, the difference in ^{10}Be concentrations between S1 and S2 surfaces is only

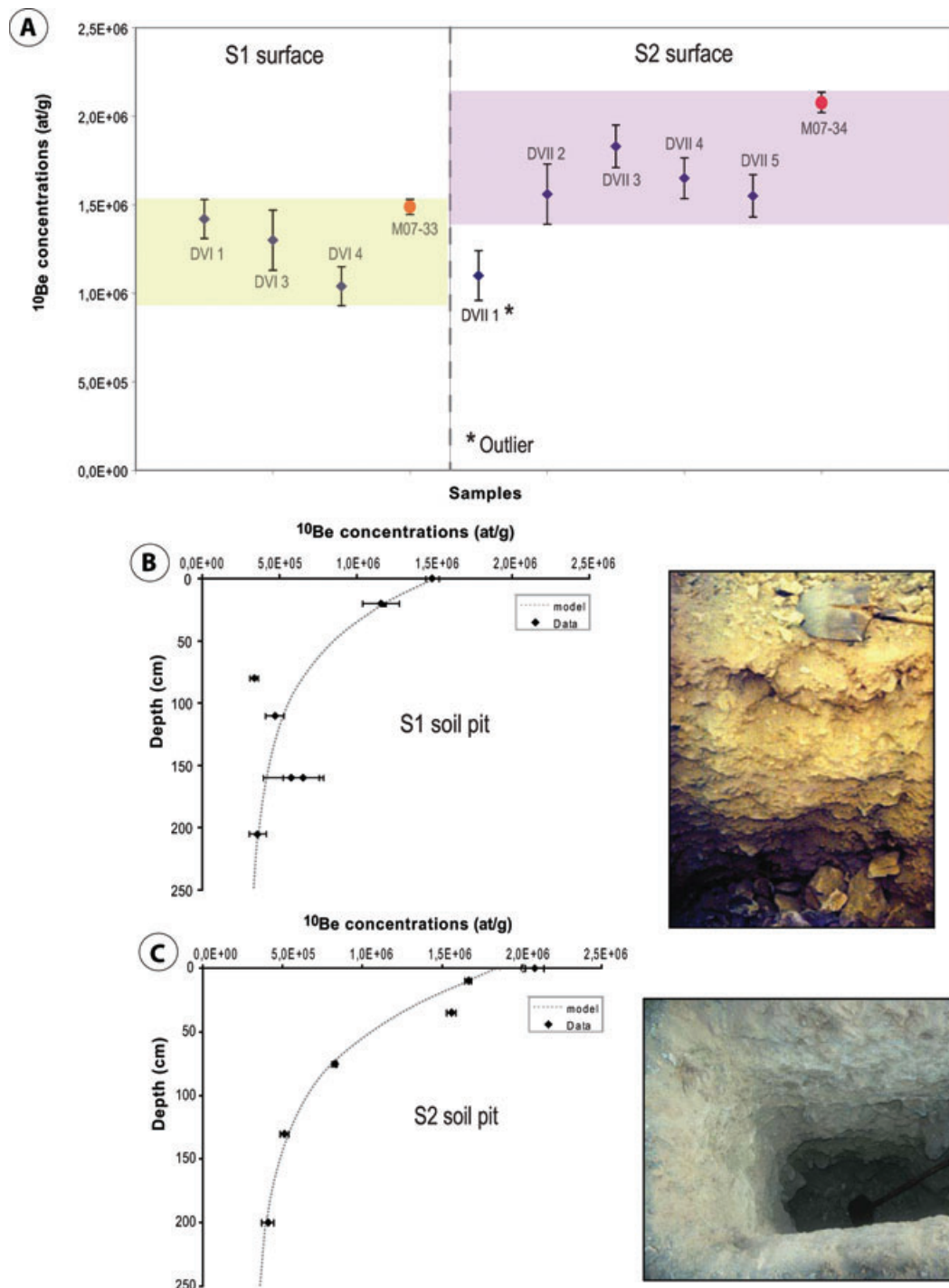


Figure 6. ^{10}Be analyses at Noyan Uul. (A) ^{10}Be surface concentrations for S1 and S2 surfaces. The blues dots are for the samples collected in 1995 (from Ritz *et al.* 1995) and restandardized here. The red dots are for the two amalgamated samples collected during the 2007 fieldtrip. (B) Depth distribution of ^{10}Be concentrations in S1 surface. Black squares are for the data; the dashed lines represent the best fit using a chi-squared inversion. Field photography of the soil pit in the S1 fan surface (taken in 1997). (C) Depth distribution of ^{10}Be concentrations in S2 surface. Black squares are for the data; the dashed lines represent the best fit using a chi-squared inversion. Field photography of the soil pit in the S2 fan surface (taken in 2007).

0.38 Mat g^{-1} suggesting that S2 has probably reached steady-state equilibrium and that S1 may be approaching steady-state concentrations. Assuming surface concentrations at steady state, we calculate maximum denudation rates to be on the order of 8.2 m Myr^{-1} and 6.8 m Myr^{-1} for S1 and S2 surfaces, respectively.

Figs 6(B) and (C) show the depth distribution of ^{10}Be concentrations in samples from surfaces S1 and S2. We collected at different

depths one single clast (or fragment of clast) along an $\sim 2 \text{ m}$ depth profile. Both profiles show a decrease in ^{10}Be concentrations at depth reaching non-negligible ^{10}Be concentrations at 2 m, attesting to an inheritance of $0.35\text{--}0.40 \text{ Mat g}^{-1}$. Following the process described by Braucher *et al.* (2003, 2009) and Siame *et al.* (2004), we modelled the ^{10}Be distribution taking into account the denudation and the inheritance. We use a chi-squared inversion (see Appendix A

for more details) to minimize the difference between observed and modelled ^{10}Be data, with the denudation and the inheritance free to vary. The best fit for S1 is for a denudation rate of $8.1 \pm 0.5 \text{ m Myr}^{-1}$, and an inheritance of 0.56 Mat g^{-1} (Fig. 6B) with a surface at steady state. Note that 0.56 Mat g^{-1} corresponds to the inheritance at the time of deposition before subsequent exposure and denudation. After exposure and denudation, the concentration at 2 m depth becomes lower than the initial inheritance. The best fit for S2 gives a denudation rate of $6.8 \pm 0.2 \text{ m Myr}^{-1}$, and an inheritance of 0.44 Mat g^{-1} with a surface at steady state (Fig. 6C). Note that the denudation rates estimated using the ^{10}Be concentrations from the profiles are consistent with denudation rates estimated from the surficial samples. Both surfaces being at steady state and assuming constant denudation rates, the best we can do is to determine the effective irradiation time (T_{eff}). The estimates of T_{eff} are 108 ka for S1 and 130 ka for S2.

Estimating slip rates. Both S1 and S2 surfaces are approaching steady state, therefore the apparent exposure age calculated for the youngest surface (S1) is closer to the true age. To determine the slip rate, we consider only the T_{eff} calculated for the S1 surface (108 ka). This yields a maximum slip rate of $1.3 \pm 0.1 \text{ mm yr}^{-1}$ using the S2/S1 riser offset ($136 \pm 14 \text{ m}$), and a minimum slip rate of $0.5 \pm 0.1 \text{ mm yr}^{-1}$ using the S1/S0 riser offset ($51 \pm 6 \text{ m}$). These results are consistent with slip rate estimates using the S1 apex offset ($46\text{--}139 \text{ m}$) that range between 0.4 mm yr^{-1} and 1.3 mm yr^{-1} .

2.2. North Ih Bogd segment

Along the North Ih Bogd segment, we estimated the horizontal slip rate of the Bogd Fault at Bitut, a site located 40 km east of the the Noyan Uul site (Fig. 1B). There, a system of alluvial terraces at the outlet of the Bitut valley is displaced by the N105°E striking Bogd Fault. The deformation is partitioned between a left-lateral strike-slip fault and a reverse fault located $\sim 1 \text{ km}$ north of the strike-slip fault (Fig. 7A). Vassallo *et al.* (2007b) estimated a $\sim 0.1 \text{ mm yr}^{-1}$ vertical slip rate along the reverse segment. The horizontal slip rate was not estimated.

Estimating offsets. As observed at the West Noyan Uul site, the number of terraces is differed on either side of the canyon (Fig. 7). On the west bank, we identified a series of three displaced terraces (T2, T1 and T0). T2 and T1 surfaces show a flat morphology with some boulders standing above the surface and still encased in it. The flatness of the two surfaces is due to erosion dominantly by wind deflation (Ritz *et al.* 2006; Vassallo *et al.* 2007b). T0 surface is very different showing a bar-and-swale morphology with well-preserved boulders bars, attesting to a much younger age.

Using our DEM combined with aerial photographs, we mapped the offset terrace risers and other morphotectonic features (Figs 7D and E). A tributary channel flows along the fault zone and has deposited two small alluvial cones on top of the T1 and T0 surfaces. The channel and the cones have obscured the risers in the area and, to estimate where the risers intersect the fault, we have to project these piercing lines from a distance of $\sim 20\text{--}30 \text{ m}$ to determine the offsets. We estimate cumulative displacements of $53.7 \pm 11.7 \text{ m}$ and $46.7 \pm 5.1 \text{ m}$ for T2/T1 riser and T1/T0 riser, respectively (Fig. 7E). On the right bank, only terraces T2 and T0 are preserved, while T1 has been removed by erosion. We estimate an offset of $9.1 \pm 0.5 \text{ m}$ for T2/T0 riser.

Dating markers. We used the ^{10}Be concentrations published in Vassallo *et al.* (2007b), which we recalculated using modern standards, to estimate the ages of the terraces. Additionally, we used

luminescence dating techniques (Lian & Roberts 2006; Wintle & Murray 2006) to have independent age control. In Fig. 8(A), we present ^{10}Be concentrations of surficial boulders. Although we did not survey the T3 (oldest) terrace (location in Fig. 7A) with kinematic GPS, we collected samples for ^{10}Be analysis on the top of the remnant surface. T3 terrace has ^{10}Be concentrations ranging between $1.35 \pm 0.10 \text{ Mat g}^{-1}$ and $1.69 \pm 0.13 \text{ Mat g}^{-1}$. T2 terrace, below T3, has concentrations ranging between $1.82 \pm 0.09 \text{ Mat g}^{-1}$ and $2.16 \pm 0.10 \text{ Mat g}^{-1}$ with one outlier (MO-10). In T1 terrace ^{10}Be concentrations range between $1.26 \pm 0.23 \text{ Mat g}^{-1}$ and $1.55 \pm 0.19 \text{ Mat g}^{-1}$. In T0 terrace, the ^{10}Be concentrations range from $0.11 \pm 0.06 \text{ Mat g}^{-1}$ to $0.42 \pm 0.05 \text{ Mat g}^{-1}$. There are no significant differences in ^{10}Be concentrations among T3, T2 and T1. T2 concentrations being slightly higher than T3 (which is an older surface) suggests that both surfaces are at steady state with a denudation rate slightly higher for T3 (e.g. Vassallo *et al.* 2007b). T1 concentrations are slightly lower than T2 concentrations suggesting that this surface is also approaching the steady state.

Figs 8(A) and (B) present the depth distribution of ^{10}Be concentrations within T1 and T2 terraces, respectively. At 250 cm depth, the ^{10}Be concentration is very small showing that pre-exposure of alluvial material before deposition is negligible (Vassallo *et al.* 2007b). Using the chi-squared inversion procedure described earlier, we model the data with denudation and inheritance free to vary. For terrace T2, the best fit is obtained with a steady-state denudation rate of 5.7 m Myr^{-1} and no inheritance, with a surface at steady state. Following Lal (1991), we calculate effective irradiation times T_{eff} of 135.2 ka for T2 terrace. For T1, the best fit is obtained with no denudation and no inheritance, yielding an age of $63.2 \pm 5 \text{ ka}$. However, as shown by the surface concentrations, the surface is approaching the steady state. Moreover, it is difficult to understand why erosion occurred on terrace T2 while no erosion would have affected terrace T1. As described earlier, both surfaces show evidence of denudation due to wind deflation. We therefore model the data using the same denudation rate determined for T2 (5.7 m Myr^{-1}) and allow the inheritance to vary. We obtain an age of $92.4 \pm 10 \text{ ka}$ for T1, which we consider as more realistic.

To estimate the age of T0 terrace from surface concentrations, we follow the same approach as developed in Vassallo *et al.* (2007b). We consider that no significant denudation has occurred on T0 surface because it presents a well-preserved morphology and because the samples were collected from the tops of boulders resting on well-preserved bars (see Ritz *et al.* 2006). We use the lowest concentration (M0-51) to minimize the pre-exposure, and assuming no denudation we calculate a minimum age of $5.1 \pm 0.9 \text{ ka}$.

For luminescence dating, we collected two samples. The first (M07-Bi-OSL-T2) was taken from a sand layer interstratified between T2 alluvium and the colluvium derived from the T3/T2 riser (Fig. 9) and the second was taken from beneath an alluvial surface 2 km to the west (Bi-OSL5, see location in Fig. 7A). At the second sample site, we measured a stream incision offset of $3.2 \pm 0.8 \text{ m}$, consistent with the offset of the riser between T0 surface and the riverbed (Fig. 10), and presumed to represent the 1957 displacement. M07-Bi-OSL-T2 was analysed by infrared stimulated luminescence (IRSL) using the Multiple Aliquot Additive Dose (MAAD) procedure (Aitken 1998) on the fine-grained (4–11 μm) polymineral sample (more details are provided in Appendix B). We used the IRSL dating technique because we anticipated older sample ages and quartz saturation: one sample had a very high dose rate, and also because the quartz within the area varies widely in its response to OSL dating due to possible albite contamination within the quartz grains (Hülle *et al.* 2010). Using the minimum equivalent

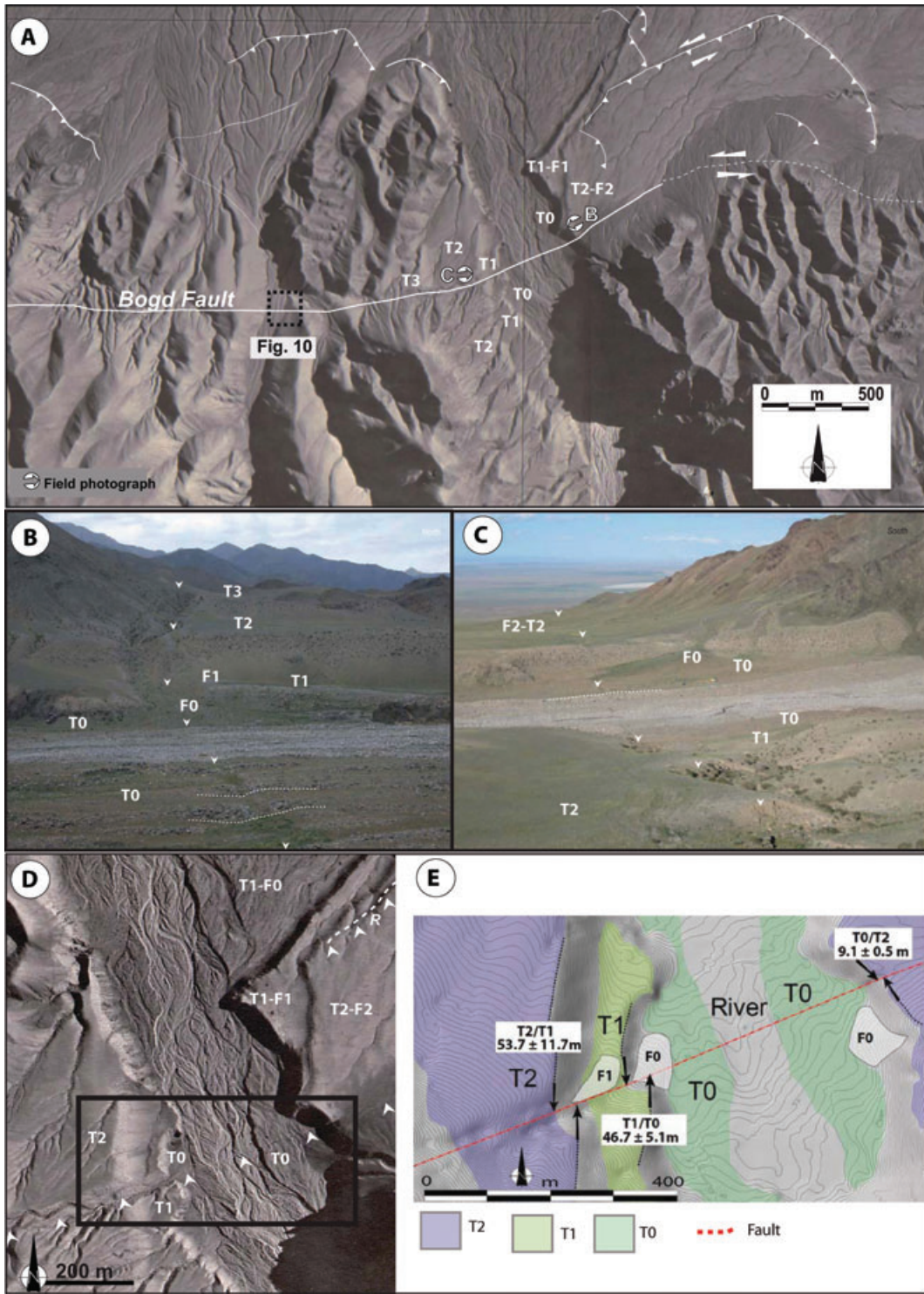


Figure 7. Morphological analyses at Bitut. (A) Aerial photo (1958) with a large view of the Bitut area. The surface ruptures associated with the 1957 event are mapped with white lines. (B) Aerial photography (1958) showing the terraces displaced along the Bogd Fault. The white triangles represent the 1957 surface rupture. (C) Digital elevation models and the morphotectonic interpretations with the piercing points used to measure the offset risers (black arrows). Note that T3 was not surveyed by the DEMs.

dose of 1100 ± 200 Grays (Gy), we calculate a minimum age of 210 ± 50 ka (Table 2). We consider this IRSL age estimate to be closer to the true age of T2 surface than the minimum age given from the steady state ^{10}Be concentration ($T_{\text{eff}} \sim 135$ ka). Bi-OSL5 was analysed using the blue light component of quartz and a Single Aliquot Regeneration protocol (Murray & Wintle 2000). The

central age model (CAM) gives an equivalent dose of 12.9 ± 1.6 Gy (details in Appendix B and in Table 2), which yields an age of 3.71 ± 0.46 ka, consistent with the ^{10}Be age of surface T0 within the Bitut valley.

Estimating slip rates. To estimate the minimum horizontal slip rates, we used the displacements of the offset risers divided by the

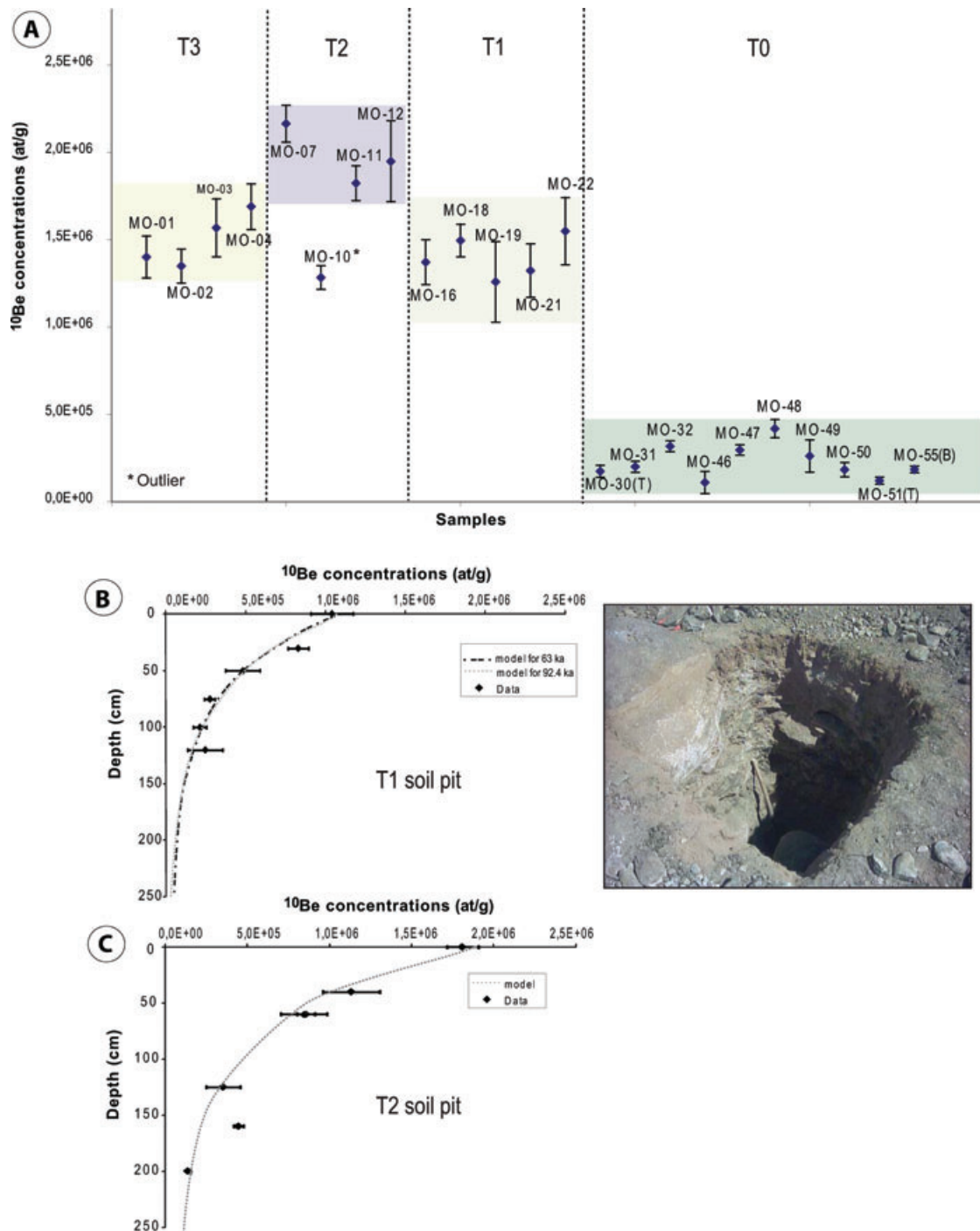


Figure 8. ^{10}Be analyses at Bitut. (A) ^{10}Be surface concentrations for the T3, T0, T2 and T1 terraces. (B) Depth distribution of ^{10}Be concentrations in T1 terrace. Black squares are for the data; the dashed lines represent the best fit using a chi-squared inversion. Field photography of the soil pit within the T1 terrace, (taken in 2005). (C) Depth distribution of ^{10}Be concentrations in T2 terrace. Black squares are for the data; the dashed lines represent the best fit using a chi-squared inversion.

ages of the upper surfaces. The offset of the T2/T1 riser (53.7 ± 11.7 m) divided by the IRSL age (210 ± 50 ka) of T2 yields a minimum slip rate of 0.28 ± 0.12 mm yr $^{-1}$. We can also obtain a maximum slip rate by dividing the offset of the riser by the age of the lower terrace. Dividing this same offset by the age (92.4 ± 10 ka) of T1 yields a maximum slip rate of 0.58 ± 0.21 mm yr $^{-1}$. These two estimates constrain the slip rate rather tightly. The offset of the T1/T0 riser (46.7 ± 5.1 m) divided by the ^{10}Be age of the upper

terrace T1, yields a minimum slip rate of 0.52 ± 0.12 mm yr $^{-1}$. In this case, unfortunately, the maximum slip rate obtained by dividing this offset by the age (5 ka) of the lower terrace T0 is ~ 9 mm yr $^{-1}$, which is valid as an upper bound but is not consistent with the maximum slip rate using T2/T1 riser offset or the maximum slip rate estimated at Noyan Uul (1.4 mm yr $^{-1}$). To summarize the results at Bitut, the average slip rate (over the past ~ 200 kyr) ranges from 0.16 to 0.79 mm yr $^{-1}$.

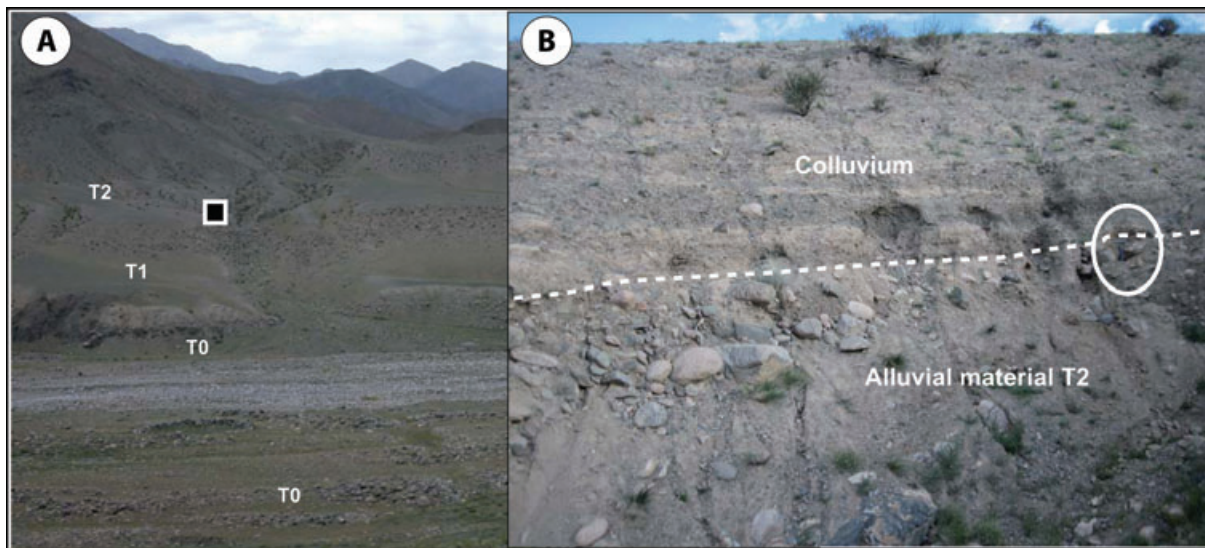


Figure 9. (A) Field photo showing the location of the IRSL sample M07-Bi-OSL-T2 in the Bitut terrace (Black square). (B) This sample is interstratified between the alluvial deposits of the T2 terrace and the colluviums from the T3 terrace. Sampling location is surrounded by the white circle.

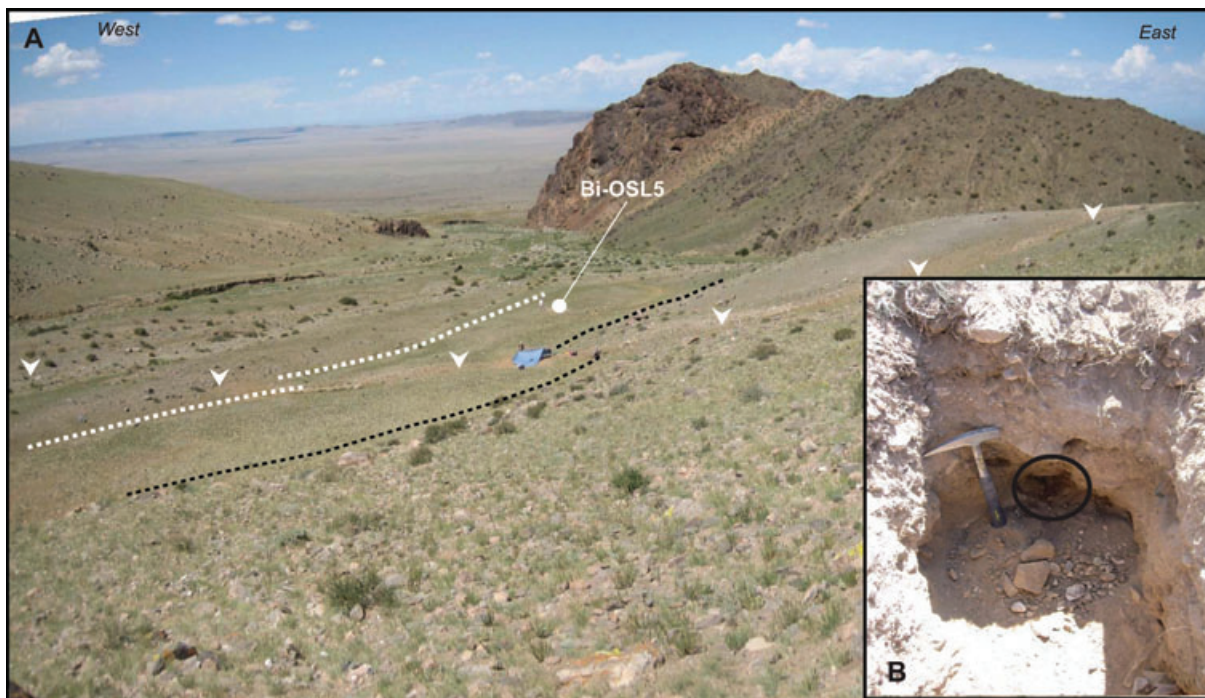


Figure 10. Field photo at the second site of Bitut area (looking north). (A) The white dashed lines represent the piercing lines used to define the offset (~ 3 m) within the T0 surface, corresponding to the 1957 coseismic offset. The black dashed lines are the piercing lines to measure the offset riser of S0/S1 surfaces. Location of the soil pit for Bi-OSL5 collecting is shown. (B) Field photo of the OSL collection pit with location of the sample (black open circle).

2.3 The East Ih Bogd segment

Along the East Ih Bogd segment, we estimated the horizontal slip rate of the Bogd Fault at Baast Uul, a site located 43 km east of the Bitut site (Fig. 1B). At this site, the Bogd Fault trends N095°E and separates the mountainous area to the south (the Dulaan Bogd Uul massif) from the alluvial plain to the north. In this area, the faulting pattern is complex showing evidence of partitioning between pure strike-slip faulting and reverse faulting (Fig. 11A). As has been described in several other places along the Bogd Fault system (i.e.

Bayasgalan *et al.* 1999b; Vassallo *et al.* 2007b), an old, inactive strike-slip fault is observed upslope within the Dulaan Bogd Uul massif.

Estimating offsets. The currently active trace of the Bogd Fault displaces two alluvial fan surfaces (S0 and S1) (Fig. 11C). The younger S0 surface shows bar-and-swale morphology with large boulders and shallow incision while the S1 is deeply incised with few remaining standing boulders. As described at Noyan Uul, the stream incisions are deeper and more broadly spaced on S1 than

Table 2. Elemental concentrations, dose rate data and ages for the Bogd samples.

Sample name	Per cent water content ^a	K (per cent) ^b	Th (ppm) ^b	U (ppm) ^b	Cosmic dose ^c additions (Gy/ka)	Total dose rate (Gy/ka)	Equivalent dose (Gy)	<i>n</i> ^d	Age (ka)
M07-Bi-OSL-T2	1 (24)	2.66 ± 0.09	11.09 ± 0.34	2.70 ± 0.34	0.31 ± 0.01	5.18 ± 0.13	>1100 ± 220	–	>210 ± 50 ^e
Bi-OSL5	1 (29)	2.30 ± 0.06	11.50 ± 0.35	2.79 ± 0.12	0.27 ± 0.01	3.47 ± 0.08	12.9 ± 1.60	17 (22)	3.71 ± 0.46 ^e
BG09–1	1 (17)	3.90 ± 0.50	21.64 ± 0.47	5.02 ± 0.20	0.28 ± 0.01	10.30 ± 0.16	311.3 ± 20.7	–	39.0 ± 2.66 ^{e,f}

Notes: ^aField moisture, with figures in parentheses indicating the full sample saturation per cent. Ages calculated using 30–35 per cent of saturation values (i.e. 10 per cent when full saturation is 34 per cent).

^bAnalyses obtained using laboratory Gamma Spectrometry (low resolution NaI detector).

^cCosmic doses and attenuation with depth were calculated using the methods of Prescott & Hutton (1994).

^dNumber of replicated equivalent dose (De) estimates used to calculate the mean equivalent dose. Figures in parentheses indicate total number of measurements made including failed runs with unusable data.

^eLinear + exponential fit used on equivalent dose data, errors to one sigma. Only Bi-OSL5 is a quartz OSL SAR age.

^fAnomalous Fading are reported on the calculated age at 5.5 per cent.

S0. The streams incised into S1 are displaced along the fault. S1 is bounded to the west by a deep stream incision showing a large cumulative left-lateral displacement. Along the eastern riser of the stream, we mapped two remnant bedrock hills on both sides of the Bogd Fault, against which the alluvium associated with the S1 surface is accumulated. Detailed field mapping of the surface, including a topographic survey, allows us to estimate an offset of 18.6 ± 1.9 m for the S1/S0 riser (Fig. 11D). The depositional contact between S1 surface and the bedrock hills, and the deeply incised stream channel are offset left-laterally by 116 ± 16 m (Fig. 11E).

Dating markers. We used luminescence dating to constrain the age of the S1 surface. We sampled the surface where its morphology is best preserved (see location, in Fig. 11C). The sampling protocol and preparation included manual extraction in a portable dark room (Rizza *et al.* 2011). We collected one sample (BG09–1) at 64 cm depth, within a stratified unit containing rounded gravels (<5 cm) in a sandy coarse matrix (Fig. 12). The sample was analysed using the fine-grained (4–11 μ m) polymineral sample. As we previously described, we anticipated that the quartz OSL would not provide a reliable equivalent dose analysis due to possible contamination in the quartz grains and also because the extremely high dose rate (see Table 2) would have saturated out the signal even if the sample had been deposited within the last 30–40 ka.

IRSL decay curves for the 4–11 μ m grain size show saturation, with a minimum equivalent dose of 311.3 ± 6.64 Gy (details in Appendix B). Moreover, we performed anomalous fading tests on the stability of the IRSL signal (see the appendix for more details) that indicate large signal instability, with recording fades of 5.5 ± 0.5 per cent (*g*-value of 4.16 per cent), which is in agreement with previously published fading tests in the Gobi-Altay area (i.e. Lehmkuhl & Lang 2001; Hülle *et al.* 2010). Using the minimum equivalent dose and the fading value, we estimated a corrected minimum age of 39.0 ± 2.66 ka for sample BG09–1.

Estimating slip rate. Dividing the S1 displacement (116 ± 16 m) by S1 minimum depositional age (39.0 ± 2.66 ka) yields a maximum horizontal slip rate of 3.01 ± 0.63 mm yr⁻¹. This corresponds to a maximum bound for the slip rate because the luminescence sample is saturated, and therefore represents a minimum age. However, similarities among fan surface chronologies between Baast Uul, Bitut and Noyan Uul sites, suggest that these surfaces were deposited contemporaneously. In Noyan Uul, S1 surface is 108 ka, while in Bitut, T1 surface is 92.4 ± 10 ka. These observations suggest that a major aggradational event occurred at ~ 100 ka within the Ih Bogd massif. Therefore, we suggest that the true age of surface S1 in Baast Uul is between 82.4 ka and 108 ka. If this assumption is correct, this yields a minimum horizontal slip rate of 1.26 ± 0.34 mm yr⁻¹.

2.4. Variations of the horizontal slip rate within the western Bogd Fault

Our study allows us to consider the evolution of the horizontal slip rate along the three western segments of the Bogd Fault. Along the WIB segment, at West Noyan Uul site, the horizontal slip rate increases from 0.57 ± 0.04 mm yr⁻¹ for the past 60 ka to 0.76 ± 0.1 mm yr⁻¹ for the past 15 ka, suggesting an increase of the horizontal slip rate through time. This result has to be considered in the local context of this site, where the deformation is distributed over two fault strands forming a releasing bend. Considered together with the fact that both fault strands are active (they both ruptured in 1957), our result suggests that the deformation is distributed over the two branches and is gradually migrating from the southern branch that forms the main topographic scarp to the younger northern branch cutting through the pull-apart. A similar feature was described along the Jid Fault in Mongolia Altay (i.e. Walker *et al.* 2006). At Noyan Uul site, the horizontal slip rate estimated over the past 110 ka ranges from 0.4 mm yr⁻¹ to 1.4 mm yr⁻¹, in agreement with previous estimates (Ritz *et al.* 1995; Ritz *et al.* 2006). Along the NIB segment, at Bitut site, the horizontal slip rate increases from 0.28 ± 0.12 mm yr⁻¹ for the past 210 ka to 0.52 ± 0.12 mm yr⁻¹ for the past 100 ka, also showing an increase of the horizontal slip rate through time, as observed at West Noyan Uul. This evolution is consistent with observations by Vassallo *et al.* (2007b) suggesting that the deformation was distributed over two branches during the past 200–100 ka and progressively migrated from the southern, older branch ~ 5 km within the Ih Bogd massif, to the northern younger branch that forms the present-day boundary of the massif. Along the EIB segment, at Baast Uul site, the horizontal slip rate is 1.26 ± 0.34 mm yr⁻¹ assuming that the 120 m offset S1 surface has the same age than the S1 surfaces in Noyan Uul and Bitut (82.4–108 ka).

Our results suggest that the horizontal slip rate decreases from the WIB segment to the NIB segment, then increases again along the EIB segment if we accept the slip rate determined at Baast Uul site based on an assumed age. The horizontal slip rate appears then to be constant at ~ 1 mm yr⁻¹ along the N095°E trending segments, while it decreases to ~ 0.5 mm yr⁻¹ along the N110°E trending segment. These differences in slip rate are consistent with the kinematic variations from purely left-lateral strike-slip faulting along the WIB and EIB segments to reverse-left-lateral faulting along the NIB segment, and consistent with the restraining bend geometry of the Bogd Fault. They also suggest that the total deformation is conserved along the western part of the Bogd Fault between Noyan Uul and Baast Uul.

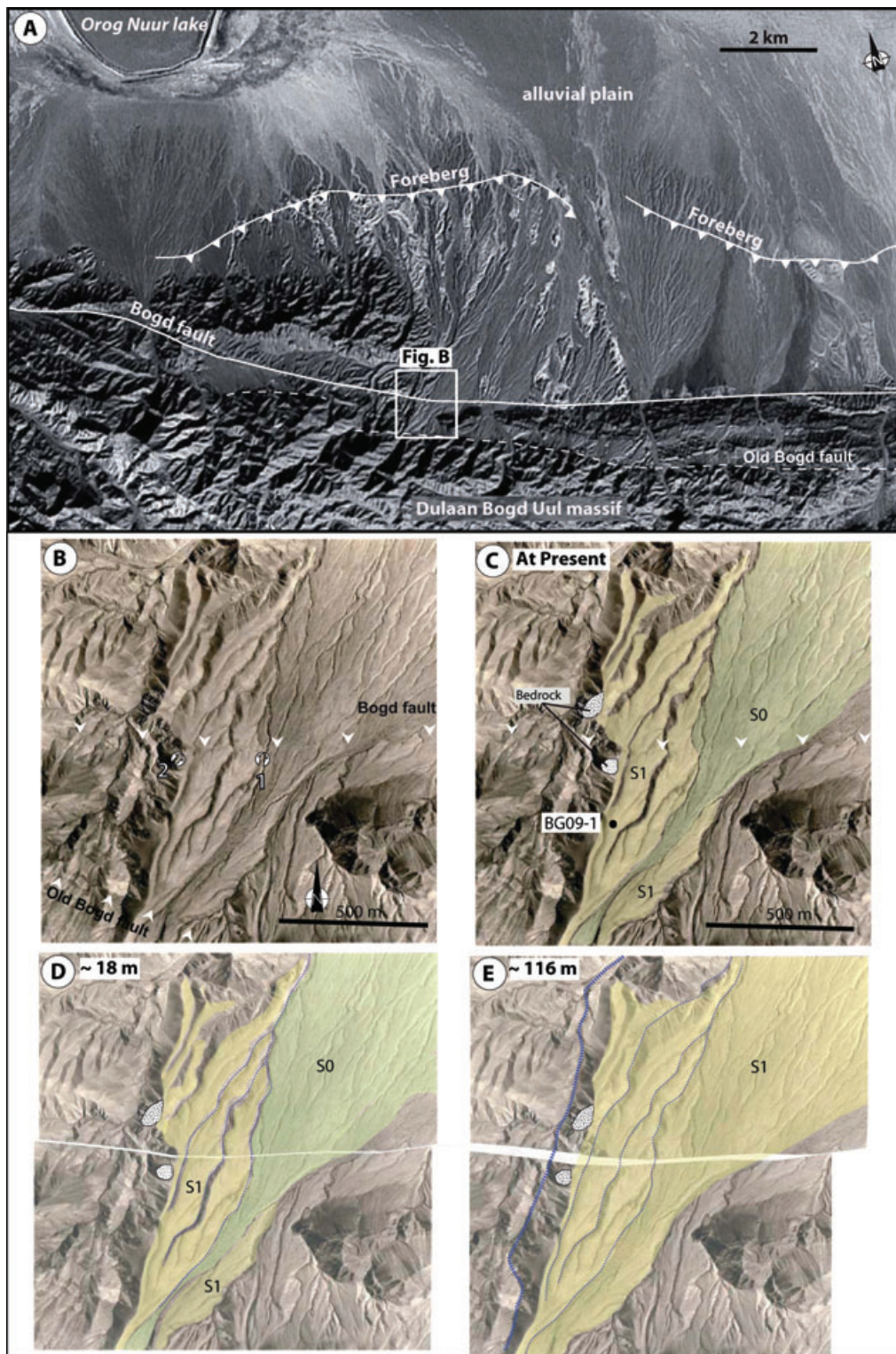


Figure 11. (A) Satellite image (source: <http://www.earth.google.com>) depicting the 1957 surface ruptures (white lines) in the Baast Uul area. The white dashed line represents the surface rupture of an old strand of the Bogd Fault, inactive during the 1957 earthquake. The white box is for the location of the study area. (B) Aerial photograph (1958) showing the S0 and S1 alluvial fans displaced along the Bogd Fault. The locations of field photos are shown: (1) Fig. 16(A) (2) Fig. 16(B). (C) Morphotectonic interpretations based on aerial photographs and reconnaissance field mapping. The black dot is for the location of BG09-1 luminescence sample. (D) Morphotectonic reconstruction for an offset of ~ 18 m using the riser S0/S1 of alluvial surfaces in the eastern part of the area. Note that thalwegs within S1 and S0 surfaces are connected across the fault. (E) Morphotectonic reconstruction for a compensation of ~ 116 m using the riser of S1 alluvial surface in the western part of the area. The uncertainty in these offset measurements come from the reconstruction of the piercing lines (thalwegs and risers) that range between a minimum (100 m) and maximum (132 m) values. In the western part of S1 riser, the large thalweg is well connected with the other side of the incision. Note also that thalwegs within the S1 surface are connected across the fault, and that the west edge of the fan, where it is deposited onto local bedrock highs is also aligned across the fault.

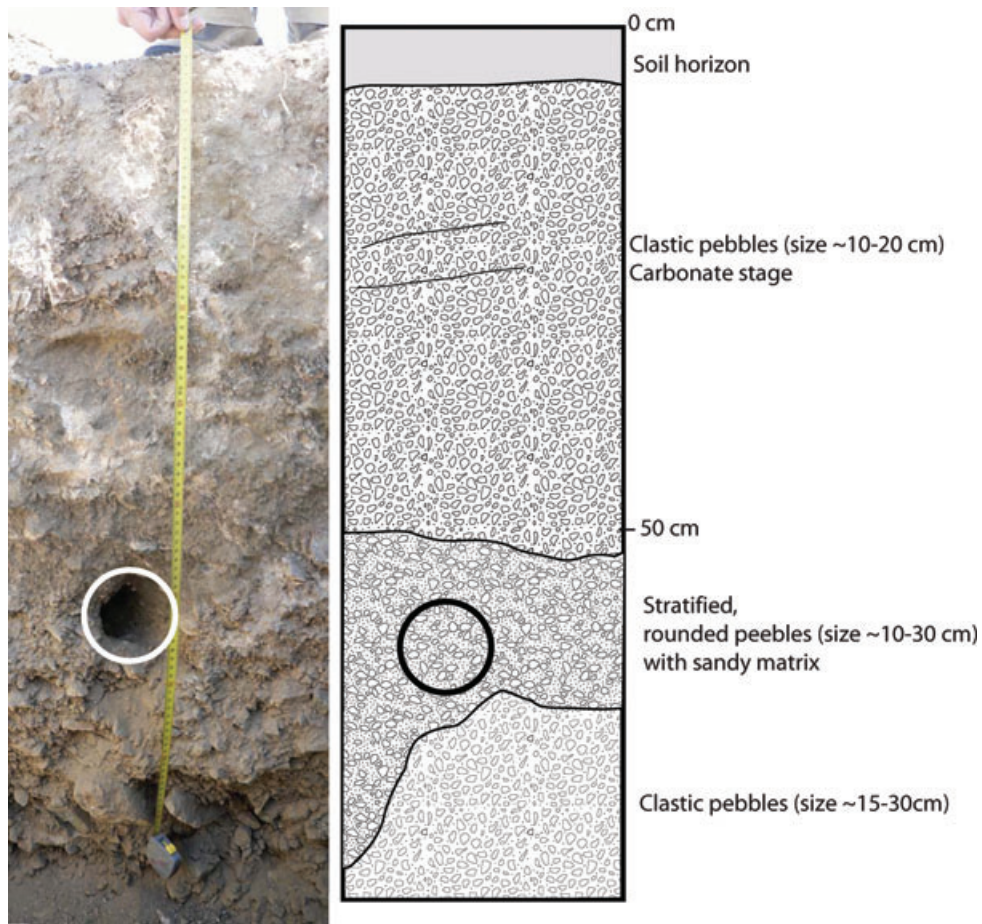


Figure 12. Field view showing the wall where BG09-1 was collected in a coarse matrix and the log interpretation of the profile. The OSL sample was collected within a palaeochannel incised into the fan deposits at 65 cm depth (black open dot is the location of the sampling).

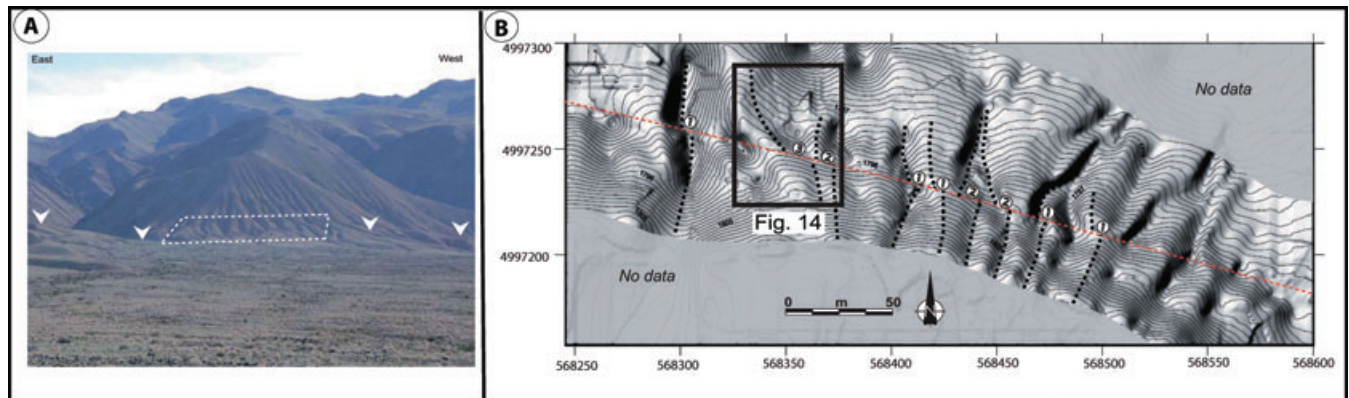


Figure 13. (A) Field view (looking south) showing the study area at Noyan Uul with the numerous thalwegs and ridgelines displaced along the Bogd Fault. The white dashed box represents the area surveyed with the kinematic GPS. (B) Digital elevation model showing the piercing lines (black dashed lines) used to calculate a mean horizontal offset along the Bogd Fault (red dashed line). (1) is for the 1957 offset, (2) and (3) are for larger offsets values. The coordinates are in Universal Transverse Mercator system, zone 48N. The contour interval is 1.0 m.

3 SLIP MAGNITUDES OF PAST EVENTS

In this section, we consider the slip magnitude of past events by analysing the distribution of displacements along the Bogd Fault using kinematic GPS surveys. Comparing 1957 offsets with cumulative displacements allows us to consider the distribution of slip associated with past earthquakes. We focus our analyses on three of the four sites previously described (Noyan Uul, Bitut and

Baast Uul) in addition to a new site (Zadgay Sayhr) located 11 km east of Baast Uul (Fig. 1B).

Noyan Uul. In Noyan Uul, we surveyed a 340-m-long section of the Bogd Fault, located 500 m west of the site described in Section 2.1.1, using RTKs GPS methods. At this location, a series of narrow thalwegs are offset left-laterally (Fig. 13). We use the axes of these thalwegs and the crest lines of ridges separating them (interfluves) to estimate the horizontal displacements associated with the 1957

Table 3. Details of the measured displacements along the Bogd Fault. Coordinates are in Universal Transverse, Mercator (UTM), zone 48. ($X1$; $Y1$) are the latitude and longitude coordinates for piercing point 1, and ($X2$; $Y2$) are the latitude and longitude coordinates for piercing point 2.

	Coordinate $X1$ (m)	Coordinate $Y1$ (m)	Coordinate $X2$ (m)	Coordinate $Y2$ (m)	Offset (m)
Noyan Uul					
Thalweg	568304,49	4997257,56	568299,82	4997257,83	4,68
Thalweg	568412,30	4997234,21	568407,35	4997236,13	5,30
Crest	568416,14	4997232,98	568420,26	4997231,60	4,34
Thalweg	568446,49	4997219,93	568442,10	4997220,89	4,50
Thalweg	568460,50	4997216,91	568456,52	4997218,97	4,48
Crest	568497,86	4997205,78	568493,33	4997205,92	4,53
				mean	4,6 ± 0,3
Crest	568331,13	4997251,65	568341,84	4997250,69	10,76
Crest	568365,60	4997244,92	568374,53	4997242,73	9,19
Thalweg	568435,23	4997225,70	568426,99	4997228,99	8,88
				mean	9,6 ± 1,0
thalweg	568342,27	4997250,16	568365,60	4997243,77	24,19
thalweg	568344,82	4997249,06	568366,63	4997241,87	22,97
thalweg	568347,75	4997248,28	568367,89	4997243,61	20,67
				mean	22,6 ± 1,8
Bitut					
Riser	607499,54	4988439,45	607496,23	4988438,00	3,61
Riser	607503,56	4988429,43	607505,96	4988430,43	2,60
Thalweg	607490,89	4988435,77	607494,83	4988437,42	4,27
Thalweg	607500,29	4988428,03	607502,65	4988429,02	2,56
				mean	3,2 ± 0,8
Riser S1/S0	607510,27	4988435,10	607516,89	4988437,67	7,10
Riser S1/S0	607507,94	4988439,67	607515,90	4988442,61	8,48
				mean	7,8 ± 0,9
Riverbed	608621,48	608624,11	4988727,20	4988728,07	2,77
Riser T0	608629,66	608632,87	4988730,56	4988731,43	3,33
Bar 1	608648,49	608651,27	4988737,42	4988737,86	2,81
Bar 2	608657,84	608660,91	4988740,92	4988741,80	3,19
Bar 3	608694,78	608698,00	4988755,23	4988756,55	3,47
Bar 4	608719,90	608723,26	4988766,48	4988767,65	3,56
				mean	3,2 ± 0,3
Baast Uul					
Top of riser S0 left bank	651557,51	4978725,38	651562,15	4978724,29	4,77
Top of riser S0 left bank	651558,20	4978726,49	651563,03	4978725,19	5,00
Base of riser S0 left bank	651562,81	4978725,01	651566,11	4978723,70	3,55
Thalweg – incision	651568,21	4978723,08	651565,43	4978723,69	2,85
Base of riser S0 right bank	651572,87	4978722,13	651577,43	4978721,07	4,68
Top of riser S0 right bank	651573,02	4978722,10	651577,72	4978721,06	4,81
				mean	4,2 ± 0,9
Riser S0/S1 – incision	651541,12	4978728,80	651560,76	4978724,87	20,04
Riser S0/S1 – incision	651544,69	4978728,20	651561,60	4978724,75	17,26
				mean	18,6 ± 1,9
Left riser	651300,62	4978786,25	651319,44	4978781,70	19,36
Incision thalweg	651307,42	4978784,00	651324,23	4978780,87	17,10
Incision thalweg	651308,00	4978784,28	651324,11	4978780,59	16,53
Incision thalweg maximum	651328,26	4978780,52	651307,95	4978784,00	20,60
				mean	18,4 ± 1,9
Zadgay Zayhr					
Thalweg	662517,71	4977946,67	662521,51	4977946,17	3,84
Thalweg	662518,09	4977949,74	662521,94	4977949,08	3,90
Thalweg (1)	662522,05	4977946,46	662518,31	4977947,49	3,88
Thalweg (1)	662522,32	4977947,52	662518,08	4977948,74	4,42
Thalweg	662549,01	4977944,34	662553,52	4977943,45	4,60
Thalweg	662634,97	4977929,02	662639,87	4977928,07	4,99
Thalweg	662639,34	4977927,13	662635,29	4977928,03	4,15
Thalweg	662656,64	4977925,12	662661,54	4977924,29	4,97
Thalweg	662672,02	4977922,39	662677,03	4977921,44	5,10
				mean	4,4 ± 0,5
Thalweg	662463,33	4977959,49	662471,41	4977959,49	8,08
Thalweg	662476,76	4977957,15	662485,34	4977955,59	8,72
Thalweg (2)	662544,31	4977944,11	662553,31	4977942,31	9,18
Thalweg (2)	662544,50	4977942,70	662553,27	4977940,88	8,95
Thalweg	662757,70	4977907,13	662766,78	4977905,51	9,22
				mean	8,8 ± 0,5
Thalweg (3)	662583,70	4977936,25	662596,10	4977934,76	12,49
Thalweg (3)	662589,32	4977934,72	662575,11	4977936,78	14,36
Thalweg (3)	662596,39	4977934,53	662583,60	4977936,07	12,89
				mean	13,2 ± 1,0

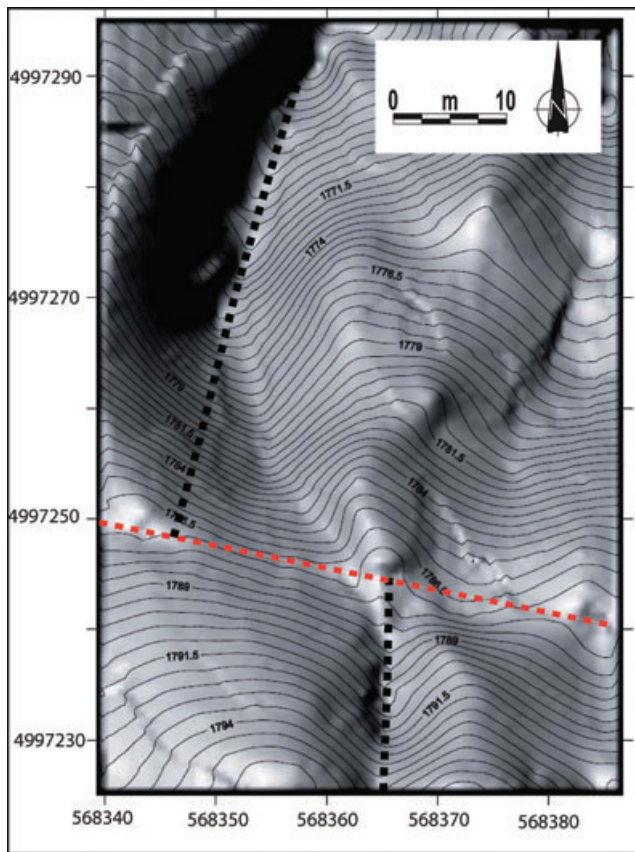


Figure 14. Digital elevation model showing the piercing lines (black dashed lines) used to calculate the mean cumulative horizontal offset (22.6 ± 1.8 m) along the Bogd Fault (red dashed line). The coordinates are in Universal Transverse Mercator system, zone 48. The contour interval is 50 cm.

earthquake and previous events. We calculate a mean value of 4.6 ± 0.3 m for the 1957 dislocation and a mean cumulative displacement of 9.6 ± 1.0 m (Table 3). Another topographic survey of an offset stream shows displacement of 22.6 ± 1.8 m (Figs 13 and 14), while two offsets of ~ 15 m are estimated farther west, using the aerial photography alone.

Bitut. In Bitut, we surveyed the offset bars associated with the lower terrace (T0) on the west bank of the Bitut River (Fig. 15). We measure a mean left-lateral horizontal offset of 3.2 ± 0.3 m, which is associated with the 1957 earthquake (Table 3). Two kilometres west of this site, we surveyed an offset stream (see Section 2.2) and an offset terrace with displacements of 3.2 ± 0.8 m and 7.8 ± 0.9 m, respectively (Fig. 10). Between this site and the Bitut valley, Vassallo *et al.* (2007b) reported offset values of 3.0 ± 0.5 m, 12 ± 0.5 m and 22.5 ± 2.5 m.

Baast Uul. At Baast Uul, our GPS survey shows a mean coseismic offset of 4.2 ± 0.9 m for the 1957 event using displaced streams incising the S0 surface, and cumulative offsets of 18.6 ± 1.9 m and 18.4 ± 1.9 m on the S0/S1 riser (see Section 2.3) and a stream incising S1 surface, respectively (Fig. 16).

Zadgay Sayhr. This site is located 11 km east of the Baast Uul site, just west of the intersection of the Toromhon thrust fault and the Bogd Fault. Numerous thalwegs are left-laterally offset with several ranges of magnitudes. We surveyed a 345-m-long section from which we measure a mean displacement of 4.4 ± 0.5 m, which we interpret to have resulted during the 1957 earthquake and larger offsets of 8.8 ± 0.5 m and 13.2 ± 1.0 m (Fig. 17 and Table 3).

Analysing the slip distribution. In Fig. 18, we present a compilation of our displacement data and published offset measurements (Kurushin *et al.* 1997; Vassallo *et al.* 2007b). This compilation shows the horizontal slip distribution associated with the 1957 earthquake along the Bogd Fault and allows us to compare the 1957 displacements with older ones (Fig. 18). Horizontal offsets associated with the 1957 earthquake range between 6.4 m and 0.7 m. The mean horizontal offset associated with the 1957 earthquake decreases from 5.2 m to 2.0 m from west to east. Along the WIB segment, the mean horizontal offset is 5.2 ± 0.9 m. Along the NIB segment, the mean 1957 horizontal offset is 3.1 ± 0.7 m, and varies from 3.9 m along the western part of this segment to 1.3 m along the eastern part. Along the EIB segment, the mean 1957 horizontal offset is 4.4 ± 1.3 m, increasing from 2.9 m on the western part of this segment to 4.6 m near the eastern end of the segment. Along the WBB segment, a single measurement of 3.0 ± 1.0 m represents the 1957 horizontal offset. Along the NBB segment the mean horizontal offset is 2.0 ± 0.1 m, ranging from 3.3 m along the western part of this segment to 0.7 m near the eastern end.

Fig. 18A shows the left-lateral offsets measured for the 1957 earthquake and the larger displacements for cumulative earthquakes. For instance, at Noyan Uul (WIB), Bitut (NIB) and Zadgay Sayhr (EIB), cumulative offsets appear to represent at least two and three times the 1957 horizontal dislocation. We measure until four and five times the 1957 horizontal dislocation at Noyan Uul and Bitut, respectively. The ratio between these measured cumulative offsets and the multiple values of the 1957 offset is comprised between 85 per cent and 100 per cent (between 1 and 2 sigma). This observation suggests that the cumulative displacements observed along the Bogd Fault are multiples of the 1957 offset, and is consistent with a characteristic slip distribution (e.g. Schwartz & Coppersmith 1984; Sieh 1996).

Fig. 18(B) shows that the distribution of the horizontal component is anticorrelated with the distribution of the vertical component. The horizontal component decreases along the NIB and NBB segments, while the vertical component increases. This evolution reflects the progressive change of kinematics from pure left-lateral strike-slip faulting to left-lateral-reverse faulting, consistent with the change of strike from the WIB segment to the NIB segment, and then from the EIB segment to the NBB segment.

4 RE-ESTIMATING THE MOMENT MAGNITUDE OF THE 1957 GOBI-ALTAY EARTHQUAKE

We use the coseismic offsets described in Section 3 to re-estimate the moment magnitude of the 1957 Gobi-Altay earthquake. We calculate the seismic moment using the relation $M_0 = \mu \cdot A \cdot SV$ with a shear modulus (μ) of 3.3×10^{10} N m⁻², and for each segment of the 1957 Bogd rupture, the mean slip vector (SV) with uncertainties, the surface rupture length and the fault dip (Table 4). We also assume that the rupture extended to a depth of 20 km as in Kurushin *et al.* (1997). We obtain a seismic moment M_0 constrained between 6.72×10^{20} N m and 1.21×10^{21} N m. Then using the relation $M_w = 2/3 \times \log(M_0) - 6.1$ (e.g. Hanks & Kanamori 1979), we estimate a moment magnitude M_w between 7.78 and 7.95.

5 RECURRENCE INTERVALS OF EARTHQUAKES

Combining slip rate estimates and the slip distribution per event allows us to estimate the mean recurrence intervals of past

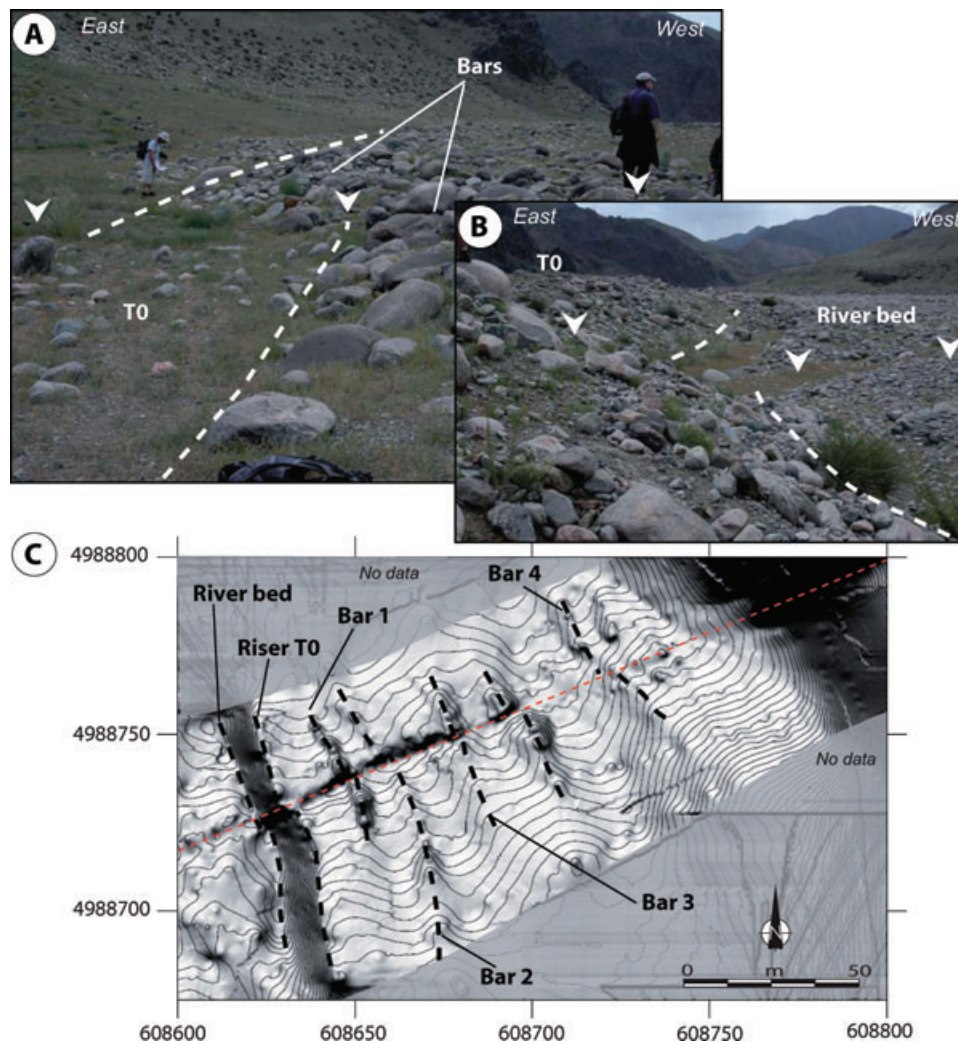


Figure 15. (A) Field view at Bitut (looking south) showing the bars and swale morphology in T0 terrace displaced along the Bogd Fault. We measured an offset of ~ 3 m using the piercing lines represented by the white dashed lines. (B) Field view (looking south) showing bottom of riser between T0 terrace and the riverbed (white dashed lines). (C) Digital elevation model showing the piercing lines (black dashed lines) used to calculate a mean horizontal offset along the Bogd Fault (red dashed line). The coordinates are in Universal Transverse Mercator system, zone 48N. The contour interval is 25 cm.

earthquakes along the different segments of the western Bogd Fault. Fig. 19 shows the mean recurrence interval estimates for large earthquakes along the WIB, NIB and EIB segments, determined by dividing the characteristic displacements observed at Noyan Uul, Bitut, Baast Uul and Gurvan Bulag site 2, by the long-term slip rate estimated at each of the three sites. These mean recurrence intervals are compared with the age of the penultimate event.

The mean recurrence intervals are between 3070 and 12 250 yr at Noyan Uul (WIB), between 3540 and 8750 yr at Bitut (NIB) and between 2060 and 5540 yr at Baast Uul (EIB), indicating that while the WIB and EIB segments have similar recurrence intervals, the NIB segment has slightly longer recurrence intervals. However, the mean interval at Bitut should be regarded as maximum bound because the slip rate there may be increasing through time (see Section 2.4). These intervals are similar to the age of the penultimate earthquake estimated from trenching investigations at Ulaan Bulag (Rizza 2010), a trench site located 33 km west of Noyan Uul, and from the ages of two surfaces (5.1 ± 0.9 ka ^{10}Be and 3.71 ± 0.46 ka $^{\text{OSL}}$) only displaced by the 1957 earthquake at Bitut. They are also similar to the age of the penultimate event (Gurvan Bulag site 1) along the Gurvan Bulag thrust fault that bounds the Ih Bogd

massif to the south (Prentice *et al.* 2002). The results are consistent with the idea that the segments of the western part of the Bogd Fault system (i.e. WIB, NIB, EIB segments + Gurvan Bulag thrust fault) have ruptured together in the past, and produces earthquakes with regular recurrence intervals. These results are in favour of a characteristic earthquake behaviour if the two eastern segments of the Bogd Fault system (WBB and NBB) show the same pattern.

6 CONCLUSION AND DISCUSSIONS

Our analysis of slip rate and slip magnitude of past events along the Bogd Fault leads us to draw a number of conclusions.

6.1 Measuring horizontal slip rate

Estimating horizontal slip rates from offset alluvial features can be difficult and interpretations of the age of offset features are often subject of debate. When offset markers are terrace risers, there is often ambiguity in whether the age of the offset corresponds to the age of the upper alluvial surface or the age of the lower surface or

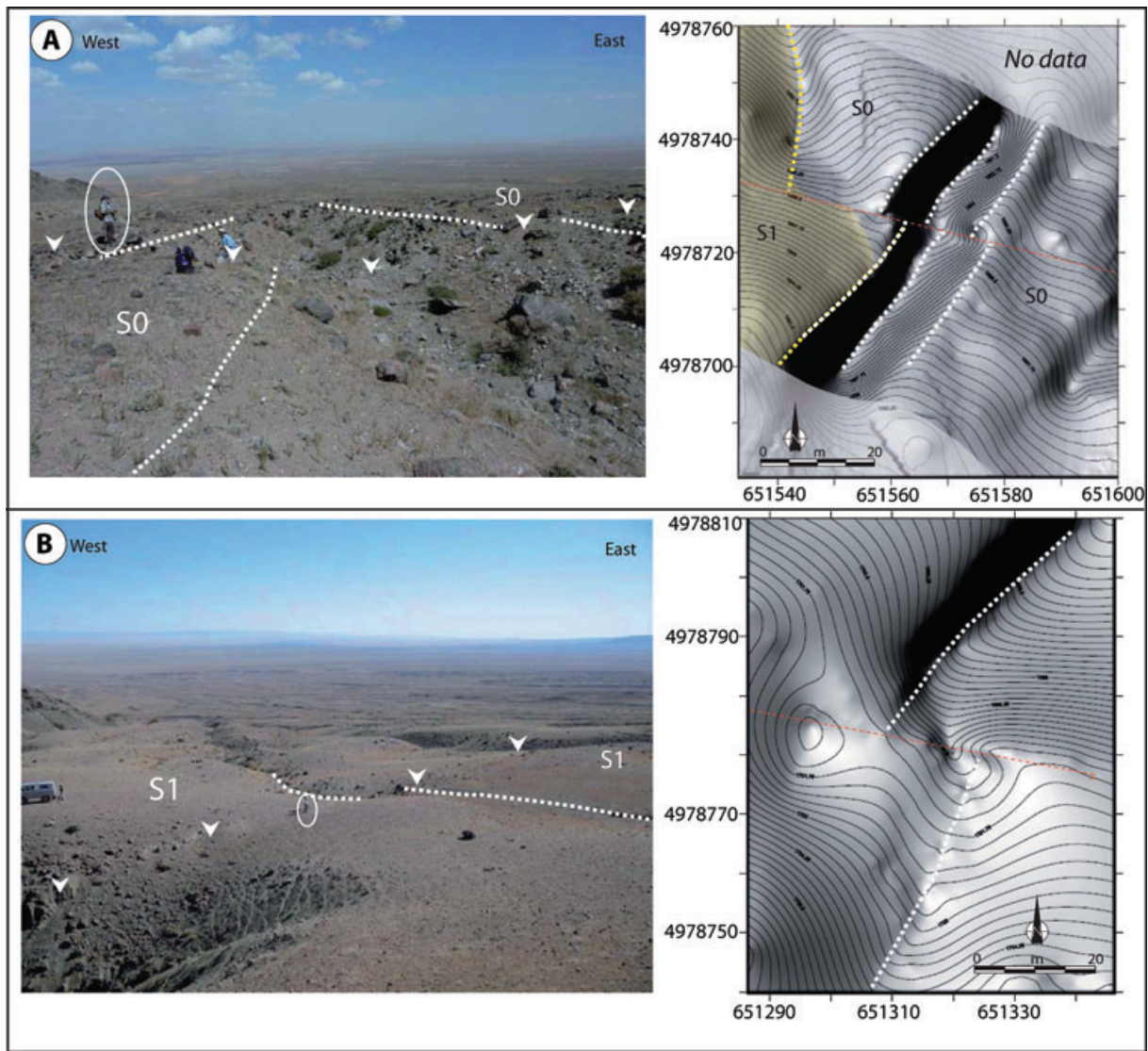


Figure 16. (A) Field view at Baast Uul (looking north) showing the piercing lines (white dashed lines) used to measure the thalweg within the S0 surface that has been displaced [state the amount of offset] along the Bogd Fault. Note that a person making scale is surrounded by the white circle. In the right part is the digital elevation model showing the piercing lines (white dashed lines) used to calculate the mean horizontal offset for the 1957 earthquake and one cumulative offset (yellow dashed lines, state the amount of offset) corresponding to the riser S0/S1. The coordinates are in Universal Transverse Mercator system, zone 48. The contour lines are spaced by 25 cm. (B) Field view at Baast Uul (looking north) showing the piercing lines (white dashed lines) used to measure the cumulative offset of a gully incised within the S1 surface. In the right part is the digital elevation model showing the piercing lines (white dashed lines) used to calculate the mean cumulative horizontal offset along the Bogd Fault (red dashed line). The coordinates are in Universal Transverse Mercator system, zone 48N. The contour interval is 25 cm.

something in between (e.g. Van der Woerd *et al.* 2000; Tapponnier *et al.* 2001; Meriaux *et al.* 2004; Hanks & Thatcher 2006; Cowgill *et al.* 2007; Gold *et al.* 2009).

In this paper, we estimated slip rates within two different faulted alluvial settings. Where the offset feature is a fan deposit that is largely preserved, we prefer using the age of abandonment of that deposit itself (e.g. for S1 at Noyon Uul and at Baast Uul). At Noyon Uul the fan apex is well preserved and the offset measurement comes from restoring that apex to its most likely position within the source drainage. At Baast Uul, we are restoring the western edge of the S1 fan where it is deposited against local bedrock highs. This feature also is a depositional feature that has not been subject to younger erosion that could have reduced the apparent offset. Therefore we are justified in using only the date of the offset deposit itself (S1)

and not the modern age of the alluvium within the deeply incised canyon to the west, which is also offset the same amount as the western edge of the S1 fan.

In a second setting, the fault is cutting through alluvial terraces within an alluvial piedmont. There the offset features are terrace risers, and we use the age of the upper terrace surface to obtain a minimum slip rate and the age of the lower terrace to obtain a maximum rate. In some cases these two methods provide relatively tight constraints on the slip rate (e.g. T2/T1 at Bitut, and T1/T0–2 at West Noyon Uul). In other cases, it is only possible to estimate either a minimum rate or a maximum rate (e.g. T3/T2 and T2/T1 at West Noyon Uul). In still other cases, the maximum and minimum rates are widely separated (e.g. T1/T0 at Bitut) and the geological context of the site as well as rates obtained from other sites along the fault

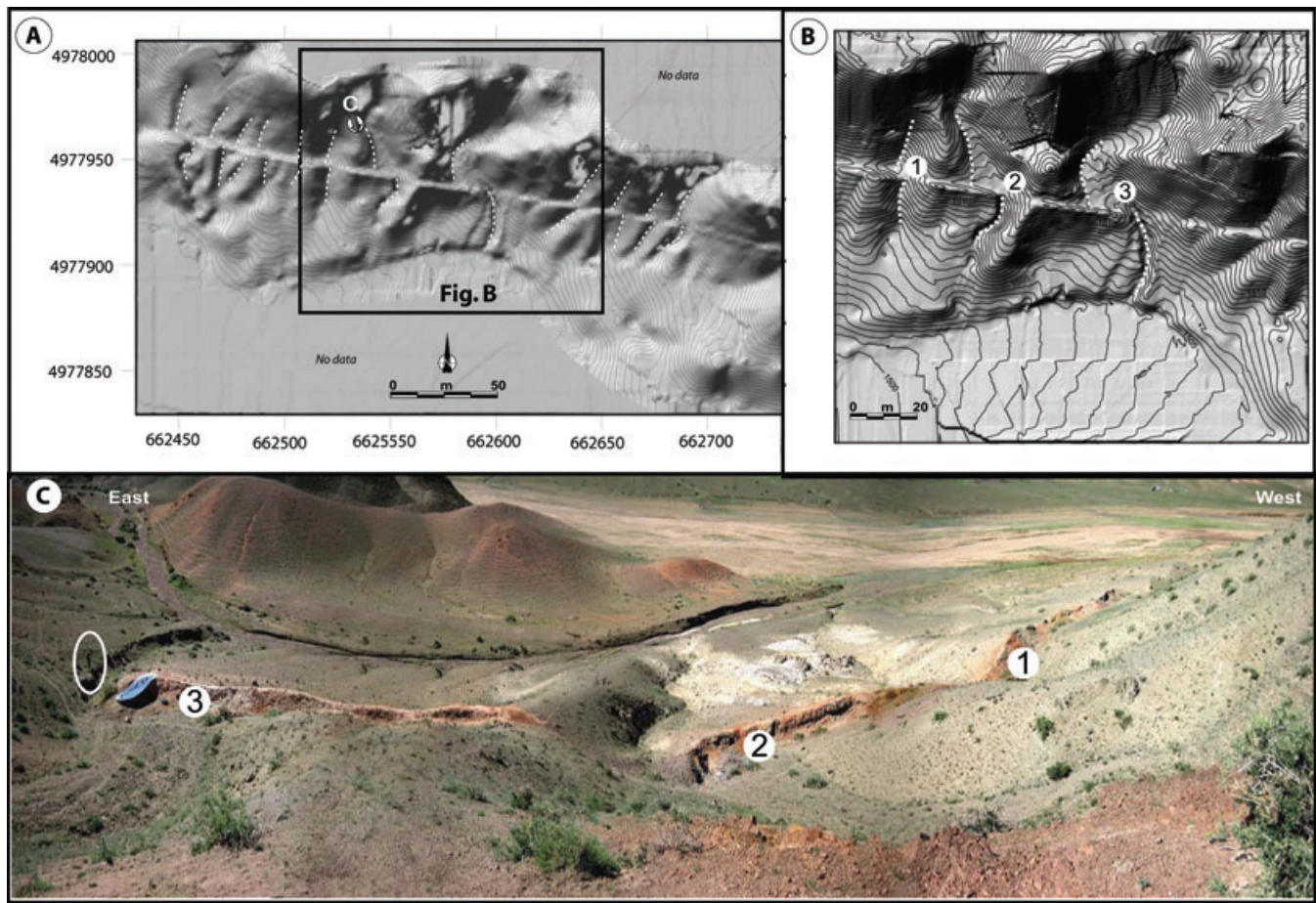


Figure 17. A) DEM, at Zadgay Zayhr site, showing the piercing lines (white dashed lines) used to calculate the mean horizontal offsets along the Bogd Fault. The coordinates are in Universal Transverse Mercator system, zone 48. The contour interval is 25 cm. The black box represents the detail mapping of the area presented in figure B. B) Digital elevation model showing the piercing lines (white dashed lines) used to calculate the horizontal offsets for the thalwegs labelled (1), (2) and (3) in Table 2. The coordinates are in Universal Transverse Mercator system, zone 48. The contour interval is 25 cm. C) Field panorama (looking south) showing the thalwegs labeled (1), (2) and (3), displaced along the Bogd Fault. Note that a person making scale is surrounded by the white circle.

must be used to assess whether the true slip rate lies closer to the minimum bound or to the maximum bound. As other studies have concluded, our study shows therefore that there is not a standard way for estimating slip rate along a strike-slip fault, and each site must be evaluated individually.

6.2 Slip rates

Our analysis of laterally faulted markers and ages of alluvial surfaces, yields a mean horizontal slip rate of $\sim 1 \text{ mm yr}^{-1}$ along the N095°E trending strike-slip WIB and EIB segments of the Bogd Fault. Between these two segments, along the N110°E trending NIB segment, the horizontal slip rate is $\sim 0.5 \text{ mm yr}^{-1}$. These variations are consistent with the geometry of the segments and their kinematic variation from mainly left-lateral strike-slip faulting along the WIB and EIB segments to reverse-left-lateral faulting along the NIB segment. They also suggest that the quantity of deformation is conserved along the three western segments of the Bogd Fault, between Noyan Uul and Baast Uul. The evolution of the horizontal slip rate through time at two sites (West Noyan Uul and Bitut) show also that the deformation can be distributed over several branches and gradually migrates from older branches to younger ones, as has been interpreted earlier (Bayasgalan *et al.* 1999b; Vassallo *et al.* 2007b).

6.3 Climate and erosion

The morphological analysis and the dating of alluvial surfaces along the western part of the Bogd Fault show that a major aggradational event occurred at $\sim 100 \text{ ka}$ at two different sites contemporaneously ($\sim 108 \text{ ka}$ at Noyan Uul and $\sim 92.4 \pm 10 \text{ ka}$ at Bitut). This confirms that a major climatic pulse occurred at the glacial-interglacial transition MIS 6/5e within the Gurvan Bogd region (Ritz *et al.* 1995; Hanks *et al.* 1997; Ritz *et al.* 2003; Vassallo *et al.* 2005; Ritz *et al.* 2006) and was probably contemporaneous with a major aggradational event in Altay as well (e.g. Nissen *et al.* 2009a, 2009b). The analysis of the distribution of ^{10}Be concentrations along depth profiles confirms that alluvial surfaces in this region undergo a denudation rate between 6 and 9 m Myr^{-1} as estimated elsewhere within the Gurvan Bulag mountain range (e.g. Vassallo *et al.* 2005; Vassallo 2006). This denudation rate is mainly due to wind deflation and is strong enough to limit the possibility of dating the alluvial surfaces in the Gobi-Altay area within the Upper Pleistocene period ($\sim 125 \text{ ka}$).

6.4 Slip distribution

Our study provides new estimates of the horizontal offset associated with the 1957 earthquake along the western part of the Bogd rupture, complementing Kurushin *et al.*'s (1997) study. Mean

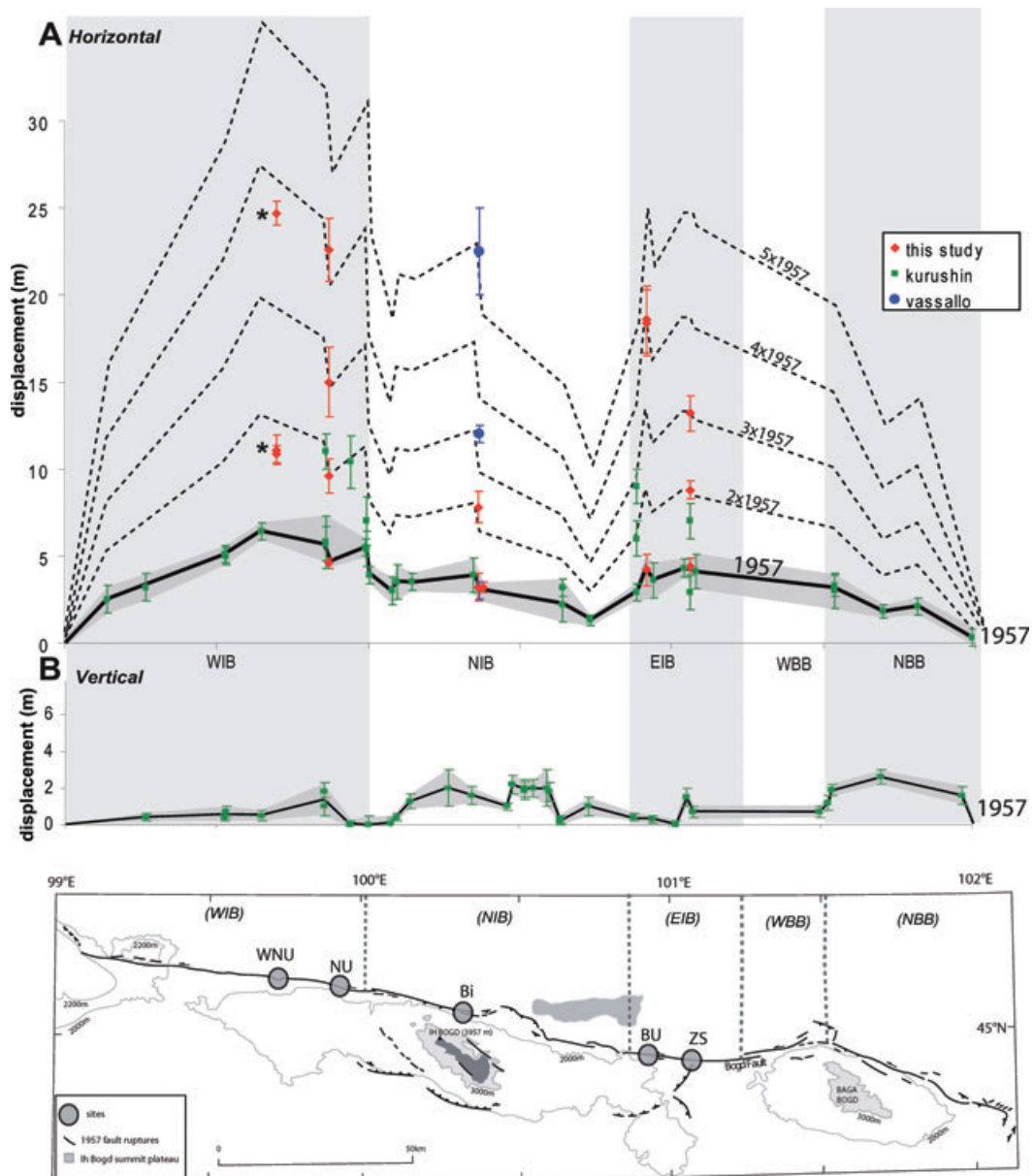


Figure 18. (A) Horizontal slip distribution along the WIB, NIB, EIB, WBB and NBB segments of the Bogd Fault. We combined our data, represented by the red diamonds (WNU : West Noyan Uul, NU : Noyan Uul, Bi: Bitut, BU: Baast uul, ZS: Zadgay Zayhr), together with Kurushin *et al.*'s (1997) data (green squares), and Vassallo *et al.*'s (2007b) data (blue dots). The black line represents the horizontal slip function for the 1957 earthquake along the Bogd Fault, and the grey area represents the uncertainty associated with each measure. (*) Note that at West Noyan Uul the cumulative offsets are in a pull apart basin, with distribution of the deformation between several branches. The black dashed lines are the multiples of the 1957 slip function. (B) Vertical slip distribution along the Bogd Fault using the vertical measures from Kurushin *et al.* (1997). The black line represents the vertical slip function for the 1957 earthquake along the Bogd Fault and the grey area, the uncertainty associated with each measure.

horizontal displacements at Noyan Uul (WIB), Bitut (NIB), Baast Uul (west EIB) and Zadgay Sayhr (east EIB) are 4.6 ± 0.3 m, 3.2 ± 0.4 m, 4.2 ± 0.9 m and 4.4 ± 0.5 m, respectively. Our data compiled together with published slip estimates (Kurushin *et al.* 1997; Vassallo *et al.* 2007b) provides a spectrum of the horizontal slip distribution along the Bogd Fault. The mean horizontal offset associated with the 1957 earthquake decreases progressively from 5.2 m to 2.0 m from west to east. The horizontal component is counter-correlated with the distribution of the vertical component reflecting the progressive change of kinematics from pure left-lateral strike-slip faulting to left-lateral-reverse faulting consistently with the geometry of the different Bogd Fault segments. Cumulative dis-

placements display multiples of the 1957 offset and are consistent with a characteristic slip distribution along the Bogd left-lateral strike-slip fault, as it has been observed along other large strike-slip faults (i.e. Schwartz *et al.* 1984; Sieh & Knopoff 1996; Rubin & Sieh 1997; Klinger *et al.* 2003; Haibing *et al.* 2005; Klinger *et al.* 2011).

6.5 Magnitude of the 1957 Gobi-Altay earthquake

Our data about 1957 dislocations combined with Kurushin *et al.*'s (1997) data allow refining the slip vector magnitude along the different segments of the Bogd Fault which evolves from 5.2 ± 1.0 m on the west part of the fault to 3.0 ± 1.1 m on the east part.

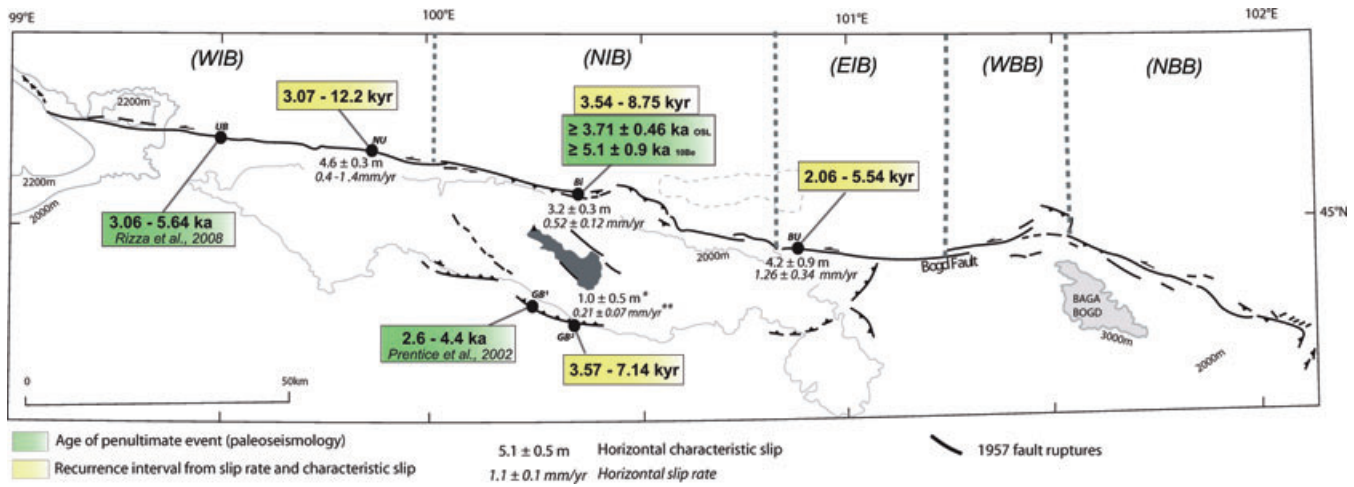


Figure 19. Recurrence intervals and ages of the penultimate earthquake along the three western segments of the Bogd Fault and the Gurvan Bulag thrust fault. UB, Ulaan Bulag; NU, Noyan Uul; Bi, Bitut; BU, Baast uul; GB¹, Gurvan Bulag site 1; GB², Gurvan Bulag site 2. Note that the slip rate at BU is estimated from an assumed age (see text for more details). The yellow boxes are the mean recurrence time calculated dividing the mean characteristic slip by the slip rate calculated at each site. Note that at GB² (Gurvan Bulag thrust fault), the mean recurrence time has been calculated dividing the characteristic slip of ~1 m (*Ritz *et al.* 2006) by the vertical slip rate of $0.21 \pm 0.07 \text{ mm yr}^{-1}$ (**Vassallo *et al.* 2005). The green boxes are for the ages of the penultimate earthquake along the Bogd Fault from palaeoseismic investigations. At Ulaan Bulag (UB), a trench opened in 2007 brackets the age of the penultimate earthquake using luminescence samples (Rizza *et al.* 2009). At Bitut (Bi), we used the age of the surfaces T0 only displaced by the 1957 earthquake (see Section 2.2). At Gurvan Bulag site 1 (GB¹), we used the mean age of the penultimate earthquake bracketed by luminescence and radiocarbon samples from palaeoseismic investigations (Prentice *et al.* 2002).

Table 4. Length of rupture, dip, mean horizontal offset (ΔH), mean vertical offset projected along the fault (ΔVf) and total coseismic displacement (SV) along the different segments of the Gobi-Altay earthquake. Note that ΔH is for the mean horizontal offset estimated from the horizontal slip distribution presented in Fig. 18(A); ΔVf is for the mean vertical offset estimated from the vertical slip distribution presented in Fig. 18(B), and projected along the fault. The slip vector (SV) is calculated using the relation $SV = \sqrt{(\Delta H^2 + \Delta Vf^2)}$.

Segment	Length (km)	Dip (°)	ΔH (m)	ΔVf (m)	SV (m)
WIB	74 ^a	70 ^b	5.2 ± 0.9^a	0.7 ± 0.6^a	5.2 ± 1.0^a
NIB	67 ^a	50 ^b	3.1 ± 0.7^a	1.6 ± 1.0^a	3.5 ± 1.0^a
EIB	23 ^a	85 ^b	4.4 ± 1.6^a	0.6 ± 0.6^a	4.7 ± 1.6^a
WBB	28 ^a	90 ^b	3.0 ± 1.0^a	0.9 ± 0.3^a	3.1 ± 1.0^a
NBB	50 ^a	60 ^b	2.0 ± 1.0^a	2.3 ± 0.6^a	3.0 ± 1.1^a
Toromhon north	12 ^b	45 ^b	–	–	5.0 ± 1.0^b
Toromhon central	4 ^b	60 ^b	–	–	5.0 ± 1.0^b
Toromhon south	5 ^b	45 ^b	–	–	0.3 ± 0.1^b
Tsagaan Uul east	7 ^b	45 ^b	–	–	1.5 ± 0.5^b
Tsagaan Uul central	9 ^b	45 ^b	–	–	3.0 ± 1.0^b
Tsagaan Uul west	10 ^b	45 ^b	–	–	1.0 ± 0.5^b
Gurvan Bulag	21 ^c	50 ^c	–	–	2.5 ± 0.5^c
Olziyt Uul	18 ^b	90 ^b	–	–	3.0 ± 1.0^b

Notes: ^aThis study.
^bKurushin *et al.* 1997.
^cPrentice *et al.* 2002.

Using these data, we estimated the moment magnitude of the Gobi-Altay earthquake at M_w 7.78–7.95. From seismic waves, the moment magnitude had been estimated between $M_w \sim 8.0$ and $M_w \sim 8.3$ (Okal 1976, Chen & Molnar 1977; Pacheco & Sykes 1992; Schlupp 1996). From surface ruptures, the moment magnitude had been estimated at $M_w \sim 8.1$ with a surface rupture of 270 km long, 20 km width and with a coseismic slip of 8 m (Molnar & Deng 1984). From their detailed rupture map, Kurushin *et al.* (1997) had re-estimated the moment magnitude at $M_w = 7.8-7.9$. Our result is consistent with this result and confirms that earlier M_w were probably over estimated, or that a greater amount of slip occurred at depth along some parts of the rupture than is observed at the surface, yielding higher moment estimates from seismologic methods than from geological methods.

6.6 Recurrence intervals

Combining our slip rate estimates and the slip distribution per event allowed us to determine mean recurrence intervals of ~2500–5200 yr for past earthquakes along the different segments of the western Bogd Fault. This is consistent with palaeoseismological data at Ulaan Bulag (west of Noyan Uul) where trenching investigation yielded a mean recurrence intervals of $4500 \pm 1400 \text{ yr}$ and a maximum average earthquake recurrence interval of 4200–5200 yr (Ritz *et al.* 2009; Rizza *et al.* 2009; Rizza 2010). These intervals are similar to the age of the penultimate event bracketed between 3060 and 5640 yr at Ulaan Bulag (Ritz *et al.* 2009; Rizza *et al.* 2009; Rizza 2010) and between 2600 and 4400 yr at Gurvan Bulag (GB¹), a thrust fault located on the southern flank of Ih Bogd massif (Prentice *et al.* 2002). These results are consistent with the idea that the three western segments and the Gurvan Bulag thrust fault ruptured contemporaneously during the penultimate earthquake and also during earlier events.

Concerning the penultimate event, note that at Zadgahy Sayhr (EIB segment), a trench investigation showed that an event occurred between 800 and 2350 yr cal BP (Bayasgalan 1999; Schwartz *et al.*

2007). This result raises the question of the magnitude of this event since a characteristic slip of 4.4 ± 0.5 m is observed along this fault segment. If this event produced a 4.4 m slip, as in 1957, the slip rate along this segment would be between two (i.e. 1.7 mm yr^{-1}) and five times (i.e. 6.2 mm yr^{-1}) larger than along the western part of the Bogd Fault. We believe therefore that this event was much smaller in magnitude than the 1957 event, and that the associated coseismic slip is included in the uncertainty (0.5 m) of the offset attributed to the 1957 earthquake (4.4 ± 0.5 m). This event may be associated with the seismic rupture along the 23-km-long EIB segment that would produce, if it breaks individually, an $M_w \sim 6.5$ earthquake with an average displacement of about 0.5 m following the statistical functions of Wells & Coppersmith (1994).

6.7 Characteristic earthquake

Our overall results suggest that similar horizontal dislocations and similar recurrence times for large earthquakes ($M_w \geq 7.5$) occurred along the three western segments of the Bogd Fault. These results combined with those along the Gurvan Bulag thrust fault suggest that the western Bogd Fault system, including the Gurvan Bulag thrust fault, is usually rupturing during a same seismic event. If the two eastern segments of the Bogd Fault (WBB and NBB) were showing the same pattern, then the Bogd Fault could be described as a system able of producing characteristic earthquakes.

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APPENDIX A: ^{10}Be DATING

A1. Sampling strategy

The concentration of *in situ* cosmogenic ^{10}Be is directly related to the time when alluviation ended, or when the subsequent incision led to the abandonment of the fan surface. The sampling strategy consisted in collecting pieces of large boulders (> 1 m), which allowed minimizing the effects of post-depositional surface deflation (Carretier 2000; Ritz *et al.* 2006; Vassallo *et al.* 2007b, Fig. A1).

A2. Experimental details

The samples from 1995 and 2003 fieldwork were analysed at Tandemron in Gif-sur-Yvette, France. Samples from 2005, 2007 and 2009 field collecting were prepared and analysed by AMS at the CEREGE, Aix-en-Provence, France (Table A1).

To estimate exposure ages from boulders as well as both denudation rates and exposure ages from depth profiles, the following eq. (A1) was used:

$$C(x, \varepsilon, t) = C_{\text{inh}} e^{-\lambda t} + \frac{P_0 P_n}{\frac{\varepsilon}{\Lambda_n} + \lambda} e^{-\frac{x}{\Lambda_n}} \left[1 - e^{-t(\frac{\varepsilon}{\Lambda_n} + \lambda)} \right] + \frac{P_0 P_{\mu s}}{\frac{\varepsilon}{\Lambda_n} + \lambda} e^{-\frac{x}{\Lambda_{\mu s}}} \left[1 - e^{-t(\frac{\varepsilon}{\Lambda_{\mu s}} + \lambda)} \right] + \frac{P_0 P_{\mu f}}{\frac{\varepsilon}{\Lambda_{\mu f}} + \lambda} e^{-\frac{x}{\Lambda_{\mu f}}} \left[1 - e^{-t(\frac{\varepsilon}{\Lambda_{\mu f}} + \lambda)} \right], \quad (\text{A1})$$

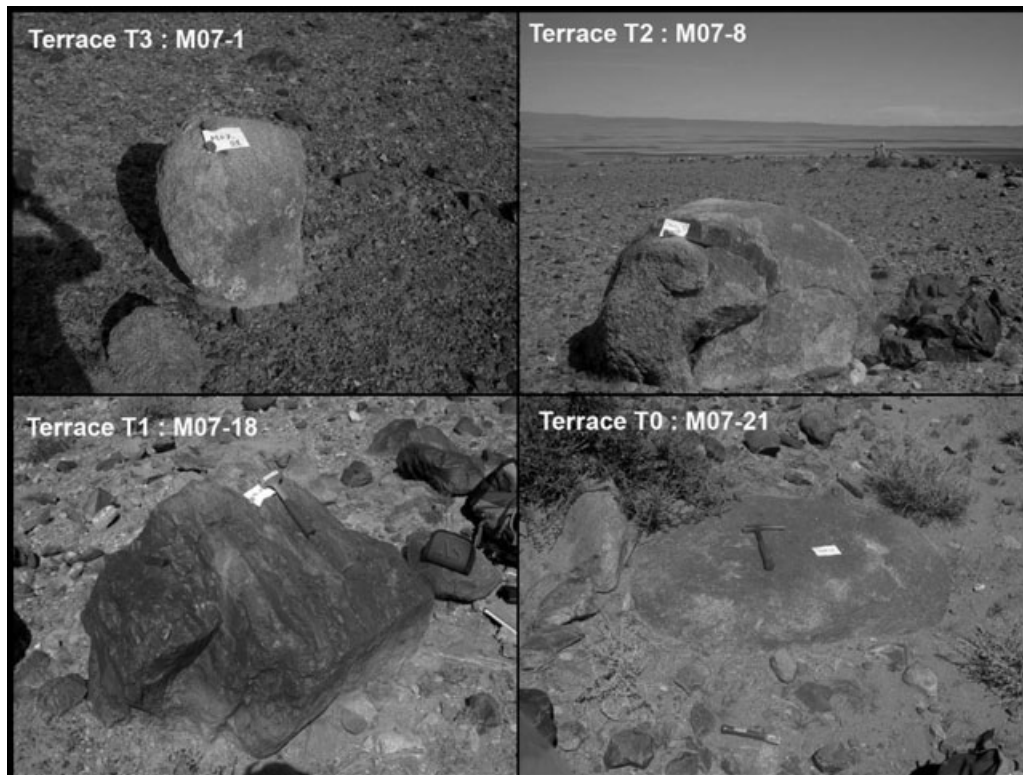


Figure A1. Boulders used for ^{10}Be dating during fieldwork in 2007 at West Noyan Uul.

Table A1. Location, elevation, pressure bar using CosmoCalc (Vermeesch 2007), density and thickness for samples used in this study. The samples shown in italics are data never published.

	Samples Names	Latitude WGS84 deg min	Longitude WGS84 deg min	Altitude m	Pressure mbar	Density g cm ⁻³	Thickness cm
NOYAN UUL site							
<i>S1 profile</i>	NU42	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU45	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU47	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU49 HF4	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU51	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU52A	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1 profile</i>	NU45HF4	45°8,24	99°53,68	1565	836	2.40	3.00
<i>S1</i>	DVI 1	45°7,48	99°53,14	1730	822	2.65	3.00
<i>S1</i>	DVI 3	45°7,48	99°53,14	1730	822	2.65	3.00
<i>S1</i>	DVI 4	45°7,48	99°53,14	1730	822	2.65	3.00
<i>S1</i>	DVI 5	45°7,48	99°53,14	1730	822	2.65	3.00
<i>S2</i>	DVII 1	45°7,48	99°53,14	1750	820	2.65	3.00
<i>S2</i>	DVII 2	45°7,48	99°53,14	1750	820	2.65	3.00
<i>S2</i>	DVII 3	45°7,48	99°53,14	1750	820	2.65	3.00
<i>S2</i>	DVII 4	45°7,48	99°53,14	1750	820	2.65	3.00
<i>S2</i>	DVII 5	45°7,48	99°53,14	1750	820	2.65	3.00
<i>Amalgam S1</i>	M07-33	45°7,91	99°53,28	1666	828	2.65	3.00
<i>Amalgam S2</i>	M07-34	45°7,51	99°53,28	1735	821	2.65	3.00
<i>S2 profile</i>	M07-35	45°7,51	99°53,13	1735	821	2.40	3.00
<i>S2 profile</i>	M07-37	45°7,52	99°53,13	1735	821	2.40	3.00
<i>S2 profile</i>	M07-39	45°7,53	99°53,13	1735	821	2.40	3.00
<i>S2 profile</i>	M07-40	45°7,54	99°53,13	1735	821	2.40	3.00
<i>S2 profile</i>	M07-41	45°7,55	99°53,13	1735	821	2.40	3.00
WEST NOYAN UUL site							
<i>T3</i>	M07-01	45°7,90	99°41,69	1618	833	2.60	3.00
<i>T3</i>	M07-02	45°7,96	99°41,75	1609	834	2.60	3.00
<i>T3</i>	M07-03	45°8,05	99°41,76	1606	835	2.60	3.00
<i>T3</i>	M07-04	45°8,09	99°41,79	1592	836	2.60	3.00
<i>T3</i>	M07-05	45°8,00	99°41,76	1608	834	2.60	3.00
<i>T3</i>	M07-06	45°8,01	99°41,71	1607	835	2.60	3.00
<i>T3 right bank</i>	M07-26	45°7,95	99°42,00	1607	835	2.60	3.00
<i>amalgam T3</i>	M07-07	45°8,00	99°41,82	1605	835	2.60	3.00
<i>T2</i>	M07-08	45°7,96	99°41,83	1598	835	2.60	3.00
<i>T2</i>	M07-09	45°7,96	99°41,83	1597	836	2.60	3.00
<i>T2</i>	M07-10	45°7,96	99°41,89	1597	836	2.60	3.00
<i>T2</i>	M07-11	45°8,01	99°41,81	1594	836	2.60	3.00
<i>T2</i>	M07-12	45°8,06	99°41,81	1587	837	2.60	3.00
<i>amalgam T2</i>	M07-13	45°7,60	99°41,84	1598	835	2.60	3.00
<i>T1</i>	M07-14	45°7,98	99°41,86	1592	836	2.60	3.00
<i>T1</i>	M07-15	45°7,99	99°41,86	1591	836	2.60	3.00
<i>T1</i>	M07-16	45°8,00	99°41,87	1593	836	2.60	3.00
<i>T1</i>	M07-17	45°8,00	99°41,87	1592	836	2.60	3.00
<i>T1</i>	M07-18	45°8,09	99°41,87	1587	837	2.60	3.00
<i>amalgam T1</i>	M07-19	45°7,80	99°41,87	1590	836	2.60	3.00
<i>amalgam T0</i>	M07-20	45°7,80	99°41,94	1583	837	2.60	3.00
<i>T0</i>	M07-21	45°7,98	99°41,89	1590	836	2.60	3.00
<i>T0</i>	M07-22	45°7,98	99°41,89	1593	836	2.60	3.00
<i>T0</i>	M07-23	45°8,05	99°41,91	1582	837	2.60	3.00
<i>T0</i>	M07-24	45°8,06	99°41,93	1583	837	2.60	3.00
<i>T0</i>	M07-25	45°8,06	99°41,83	1580	837	2.60	3.00
<i>T2 profile</i>	M07-27	45°8,00	99°41,83	1597	836	2.40	3.00
<i>T2 profile</i>	M07-28	45°8,00	99°41,84	1597	836	2.40	3.00
<i>T2 profile</i>	M07-29	45°8,00	99°41,84	1597	836	2.40	3.00
<i>T2 profile</i>	M07-30	45°8,00	99°41,84	1597	836	2.40	3.00
<i>T2 profile</i>	M07-31	45°8,00	99°41,84	1597	836	2.40	3.00

Table A1. (Continued.)

	Samples names	Latitude WGS84 deg min	Longitude WGS84 deg min	Altitude m	Pressure mbar	Density g cm ⁻³	Thickness cm	
BITUT site								
	T3 (boulder)	MO-03-01	45°2,78	100°22,28	1720	823	2.70	3.00
	T3 (boulder)	MO-03-02	45°2,79	100°22,28	1720	823	2.70	3.00
	T3 (boulder)	MO-03-03	45°2,79	100°22,28	1720	823	2.70	3.00
	T3 (boulder)	MO-03-04	45°2,80	100°22,28	1720	823	2.67	3.00
	T3 (boulder)	MO-03-05	45°2,80	100°22,28	1720	823	2.67	3.00
	T2 (boulder)	MO-03-07	45°2,70	100°22,44	1690	826	2.67	3.00
	T2 (boulder)	MO-03-10	45°2,60	100°22,59	1690	826	2.67	3.00
	T2 (boulder)	MO-03-11	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 (boulder)	MO-03-12	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 (boulder)	MO-03-12 0cm	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 (boulder)	MO-03-12 40cm	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 (boulder)	MO-03-12 60cm	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 profile	MO-03-T2-TOP	45°2,60	100°22,59	1690	826	2.62	3.00
	T2 profile	MO-03-T2-40	45°2,60	100°22,59	1690	826	2.60	3.00
	T2 profile	MO-03-T2-60	45°2,60	100°22,59	1690	826	2.60	3.00
	T2 profile	MO-03-T2-125	45°2,60	100°22,60	1690	826	2.60	3.00
	T2 profile	MO-03-T2-160	45°2,60	100°23,80	1690	826	2.60	3.00
	T2 profile	MO-03-T2-200	45°2,60	100°23,81	1690	826	2.60	3.00
	T2 (boulder)	MO-03-14	45°2,68	100°24,23	1713	824	2.60	3.00
	T1 (boulder)	MO-03-16	45°2,69	100°23,70	1650	830	2.60	3.00
	T1 (boulder)	MO-03-17	45°2,69	100°22,60	1650	830	2.60	3.00
	T1 (boulder)	MO-03-18	45°3,19	100°22,60	1650	830	2.60	3.00
	T1 (boulder)	MO-03-19	45°3,16	100°22,60	1650	830	2.60	3.00
	T1 (boulder)	MO-03-21	45°3,25	100°22,60	1650	830	2.60	3.00
	T1 (boulder)	MO-03-22	45°3,10	100°22,60	1650	830	2.60	3.00
	T1 profile	MO-03-T3-30	45°2,68	100°22,60	1650	830	2.60	3.00
	T1 profile	MO-03-T3-50	45°2,68	100°22,69	1650	830	2.60	3.00
	T1 profile	MO-03-T3-75	45°2,68	100°22,69	1650	830	2.60	3.00
	T1 profile	MO-03-T3-100	45°2,68	100°22,66	1650	830	2.60	3.00
	T1 profile	MO-03-T3-120	45°2,68	100°22,66	1650	830	2.60	3.00
	T0 (boulder)	MO-03-30(T)	45°2,54	100°22,56	1640	831	2.67	3.00
	T0 (boulder)	MO-03-31 (B)	45°2,54	100°22,56	1640	831	2.67	3.00
	T0 (boulder)	MO-03-32	45°2,51	100°22,56	1640	831	2.67	3.00
	T0 (valley)	MO-03-46	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-47	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-48	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-49	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-50	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-51(T)	44°59,64	100°22,56	2100	785	2.67	3.00
	T0 (valley)	MO-03-55(B)	44°59,64	100°22,56	2100	785	2.67	3.00

where:

(1) $C(x, \varepsilon, t)$ is the ¹⁰Be concentration function of depth x (g cm⁻²), denudation rate ε (g cm⁻² yr⁻¹) and exposure time t (yr).

(2) Λ_n (160), $\Lambda_{\mu s}$ (1500) and $\Lambda_{\mu f}$ (5300) are the effective apparent attenuation lengths (g cm⁻²), for neutrons, slow muons and fast muons, respectively.

(3) P_n , $P_{\mu s}$ and $P_{\mu f}$ are the relative contributions to the ¹⁰Be production rate of the three reactions ($P_n + P_{\mu s} + P_{\mu f} = 100$ per cent). The relative contributions to the total surface production rate are 97.85 per cent for the neutrons, 1.50 per cent for slow muons, and 0.65 per cent for the fast muons. These values are based on field-calibrated measurement by Braucher *et al.* (2003).

(4) P_0 is the production rate at the surface taken from Stone (2000).

(5) C_{inh} represents the ¹⁰Be concentration potentially acquired by the sample during exposure to cosmic rays prior to emplacement in their sampling position.

Because denudation processes are complex, in a first step, exposure ages were determined using negligible denudation. In that case, the exposure ages are minimum. However, when it was possible, exposure ages were better defined using depth profiles allowing the determination of denudation rate and inheritance using a chi-square inversion (Braucher *et al.* 2003).

$$\text{Chi - square} = \sum_{i=1}^n \left[\frac{C_i - C(x, \varepsilon, t)}{\sigma_i} \right]^2, \quad (\text{A2})$$

where

(1) C_i is the measured ¹⁰Be concentration at depth x_i .

(2) $C(x, \varepsilon, t)$ is the theoretical ¹⁰Be concentration determined using eq. (A1).

(3) σ_i is the analytical uncertainty at depth i , and n the total number of samples in the profile.

The ages obtained from surficial samples were used only if ages given by depth profiles were poorly constrained. When the surface

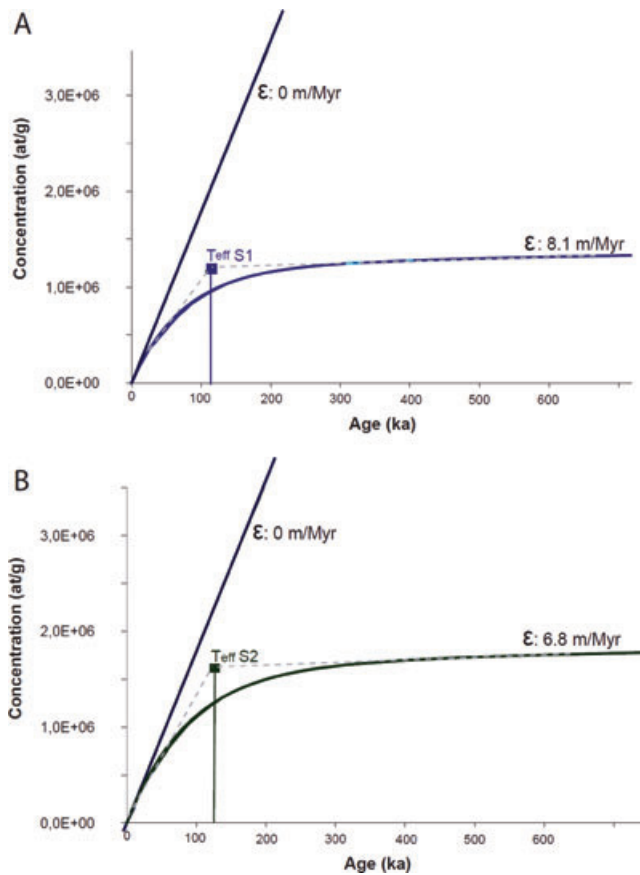


Figure A2. Effective irradiation times deduced from the ^{10}Be concentrations and the denudation rate calculated for S1 (A) and S2 (B) surfaces. The blue and green curves represent the modelled evolution of the ^{10}Be surface concentration versus the time at the measured denudation rates (ϵ) for each surface. The grey dashed lines define the linear increase in concentration and the linear evolution at steady state. The squares are the intersection between the straight lines and represent minimum time required to reach steady state (T_{eff}) for S1 and S2 surfaces. Since both of these surfaces are inferred to have reached steady state, these ages also represent minimum ages of abandonment.

concentration was at steady state, we determined a denudation rate and an effective irradiation time (Lal 1991). Fig. A2 shows the T_{eff} deduced from the variation of ^{10}Be concentrations for S1 and S2 versus time using the modelled denudation rate. Effective irradiation times T_{eff} are determined by intersecting straight lines defining the linear increase in concentration when production is larger than losses, and the asymptote at steady state.

APPENDIX B: LUMINESCENCE DATING

B1. OSL collecting protocol

B1.1. Bitut

At Bitut, the samples were carved into a pedestal fitting the shape of an aluminum container that was 6–7 cm in diameter. These containers were sealed and stored in dark plastic bags until transport to an OSL processing lab. We kept only the material in the centre of the block after removing the light-exposed edges. The block edges were painted in the lab, the paint was allowed to dry, and then the

edges were rigorously trimmed off to get to the centre. If any paint showed up in the sieves during the prep process, the sample fraction was thrown away and a new portion was used.

B1.2. Baast Uul

At Baast Uul, the main stratigraphic deposits in the alluvial surfaces are clastic limestone pebbles supported by sand matrix, stratified and poorly sorted. We decided that, to collect only the matrix, we would need to build ourselves a ‘portable’ dark room as the foundation area that would enable us to dig out suitable OSL samples. Our samples were collected by using the field protocol detailed in Rizza *et al.* (2011), covering the pits with several layers of black plastic and cover blankets, such that we could be confident that during the manual extraction and collection, there was not any light exposure at the time of sampling.

B2. Experimental details

B2.1. Dosimetry

Around 600 g of bulk sample were dried, sealed in a plastic planchets (Murray *et al.* 1987) and placed in a gamma-ray spectrometer to measure the concentration of radiogenic Thorium (Th), Potassium (K) and Uranium (U). Dose rates for each sample were calculated using radioisotope concentration, burial depth, elevation, geomagnetic latitude (Prescott & Hutton 1994) and present-day moisture. Alpha and beta attenuation contributions were corrected for grain-size attenuation (Aitken, 1985) if fine grains were used for dating (i.e. IRSL). All elemental concentrations and dose rates are shown in Table 1.

B2.2. Sample preparation protocol

In the OSL laboratory the middle portion of sediment within the aluminium cans were used for equivalent dose analyses (final mineral was either quartz sand or polymineral silt grains). After standard quartz separation, the samples were resieved to save the pure quartz grains and to discard the broken, and thus smaller, plagioclase grains that survived the hydrofluoric etch process. Within the samples, plagioclase was very plentiful and difficult to separate from the quartz. Only HF or intensive Frantz magnetic separation could achieve a reasonable grains separation. For this reason we believe there are probably inclusions of albite in some quartz grains. The largest sand size fraction was used for these samples because it was the dominant fraction (>50 per cent) within usable OSL grain sizes.

B2.2.1. Quartz dating The quartz grains from the Bitut sample (Bi-OSL5) were analysed with the Single Aliquot Regeneration (SAR) procedures (Murray & Wintle 2000) and run as aliquots with a size of 1–2 mm (Fig. B1). Extended descriptions of the machine and protocols used can be found in the Table B1.

B2.2.2. Feldspar dating We used IRSL, which is the emissions from the potassium-Feldspar silt fractions. The IRSL samples M07-Bi-OS1-T2 and BG09–1 were analysed using the MAAD procedure (Aitken 1998). The fine-grained (4–11 μm) polymineral sample splits for three samples were analysed using the total-bleach MAAD method for IRSL (Singhvi *et al.* 1982; Lang 1994; Richardson *et al.* 1997; Forman & Pierson 2002). Specific details for this

Table A2. Depth, elevation, ^{10}Be concentrations and their associated incertitude (two sigma error bars), production rate calculated using Stone (2000), and laboratory location for this study. The samples shown in italics are data never published.

	Samples	Depth (cm)	Altitude (m)	^{10}Be (at g^{-1})	^{10}Be error (at g^{-1})	Production (at $\text{g}^{-1} \text{yr}^{-1}$)	Measured at
<i>NOYAN UUL</i>							
S1 profile	NU42	205	1565	3,578E+05	5,600E+04	16,44	TANDETRON
S1 profile	NU45	160	1565	6,534E+05	1,308E+05	16,44	TANDETRON
S1 profile	NU47	110	1565	4,696E+05	5,935E+04	16,44	TANDETRON
S1 profile	NU49 HF4	80	1565	3,358E+05	2,772E+04	16,44	TANDETRON
S1 profile	NU51	20	1565	1,155E+06	1,184E+05	16,44	TANDETRON
S1 profile	NU52A	0	1565	1,917E+06	1,966E+05	16,44	TANDETRON
S1 profile	NU45HF4	160	1565	5,746E+05	1,818E+05	16,44	TANDETRON
S1	DVI 1	0	1730	1,420E+06	1,100E+05	18,22	TANDETRON
S1	DVI 3	0	1730	1,300E+06	1,700E+05	18,22	TANDETRON
S1	DVI 4	0	1730	1,040E+06	1,100E+05	18,22	TANDETRON
S1	DVI 5	0	1730	1,370E+06	1,400E+05	18,22	TANDETRON
S2	DVII 1	0	1750	1,100E+06	1,400E+05	18,49	TANDETRON
S2	DVII 2	0	1750	1,560E+06	1,700E+05	18,49	TANDETRON
S2	DVII 3	0	1750	1,830E+06	1,200E+05	18,49	TANDETRON
S2	DVII 4	0	1750	1,650E+06	1,400E+05	18,49	TANDETRON
S2	DVII 5	0	1750	1,550E+06	1,200E+05	18,49	TANDETRON
<i>Amalgam S1</i>	M07-33	0	1666	1,488E+06	4,277E+04	17,22	ASTER
<i>Amalgam S2</i>	M07-34	0	1735	2,078E+06	5,785E+04	18,33	ASTER
<i>S2 profile</i>	M07-35	200	1735	4,067E+05	3,640E+04	18,33	ASTER
<i>S2 profile</i>	M07-37	130	1735	5,108E+05	2,625E+04	18,33	ASTER
<i>S2 profile</i>	M07-39	75	1735	8,257E+05	1,533E+04	18,33	ASTER
<i>S2 profile</i>	M07-40	35	1735	1,556E+06	2,721E+04	18,33	ASTER
<i>S2 profile</i>	M07-41	10	1735	1,663E+06	1,896E+04	18,33	ASTER
<i>WEST NOYAN UUL</i>							
<i>T3</i>	M07-01	0	1618	1,143E+06	2,024E+04	17,24	ASTER
<i>T3</i>	M07-02	0	1609	1,587E+06	3,769E+04	17,13	ASTER
<i>T3</i>	M07-03	0	1606	1,611E+06	7,633E+04	17,09	ASTER
<i>T3</i>	M07-04	0	1592	1,742E+06	7,445E+04	16,92	ASTER
<i>T3</i>	M07-05	0	1608	1,666E+06	6,606E+04	17,12	ASTER
<i>T3</i>	M07-06	0	1607	1,749E+06	5,795E+04	17,10	ASTER
<i>T3 right bank</i>	M07-26	0	1607	1,571E+06	1,048E+05	17,10	ASTER
<i>amalgam T3</i>	M07-07	0	1605	1,644E+06	4,771E+04	17,08	ASTER
<i>T2</i>	M07-08	0	1598	1,022E+06	1,988E+04	16,99	ASTER
<i>T2</i>	M07-09	0	1597	1,516E+06	2,106E+04	16,98	ASTER
<i>T2</i>	M07-10	0	1597	1,414E+06	8,056E+04	16,98	ASTER
<i>T2</i>	M07-11	0	1594	2,269E+06	1,175E+05	16,94	ASTER
<i>T2</i>	M07-12	0	1587	1,153E+06	7,229E+04	16,85	ASTER
<i>amalgam T2</i>	M07-13	0	1598	1,614E+06	2,081E+04	16,99	ASTER
<i>T1</i>	M07-14	0	1592	9,429E+05	3,379E+04	16,91	ASTER
<i>T1</i>	M07-15	0	1591	9,603E+05	2,740E+04	16,90	ASTER
<i>T1</i>	M07-16	0	1593	6,882E+05	3,122E+04	16,94	ASTER
<i>T1</i>	M07-17	0	1592	5,082E+05	2,999E+04	16,91	ASTER
<i>T1</i>	M07-18	0	1587	4,910E+05	4,039E+04	16,85	ASTER
<i>amalgam T1</i>	M07-19	0	1590	8,163E+05	3,470E+04	16,89	ASTER
<i>amalgam T0</i>	M07-20	0	1583	6,578E+05	1,247E+04	16,80	ASTER
<i>T0</i>	M07-21	0	1590	2,487E+05	1,132E+04	16,89	ASTER
<i>T0</i>	M07-22	0	1593	4,485E+05	2,188E+04	16,93	ASTER
<i>T0</i>	M07-23	0	1582	2,998E+05	4,865E+04	16,79	ASTER
<i>T0</i>	M07-24	0	1583	4,861E+05	3,244E+04	16,80	ASTER
<i>T0</i>	M07-25	0	1580	5,295E+05	4,064E+04	16,80	ASTER
<i>T2 profile</i>	M07-27	0	1597	2,323E+05	1,930E+04	16,98	ASTER
<i>T2 profile</i>	M07-28	30	1597	6,363E+05	1,284E+04	16,98	ASTER
<i>T2 profile</i>	M07-29	50	1597	6,084E+05	1,059E+05	16,98	ASTER
<i>T2 profile</i>	M07-30	95	1597	6,502E+05	9,694E+04	16,98	ASTER
<i>T2 profile</i>	M07-31	130	1597	8,473E+05	4,123E+04	16,98	ASTER
<i>BITUT</i>							
T3 (boulder)	MO-03-01	0	1720	1,401E+06	1,205E+05	18,07	ASTER
T3 (boulder)	MO-03-02	0	1720	1,349E+06	9,779E+04	18,07	ASTER
T3 (boulder)	MO-03-03	0	1720	1,569E+06	1,661E+05	18,07	ASTER
T3 (boulder)	MO-03-04	0	1720	1,690E+06	1,311E+05	18,07	ASTER
T3 (boulder)	MO-03-05	0	1720	1,173E+06	8,101E+04	18,07	ASTER

Table A2. (Continued.)

	Samples	Depth (cm)	Altitude (m)	^{10}Be (at g^{-1})	^{10}Be error (at g^{-1})	Production (at $\text{g}^{-1} \text{yr}^{-1}$)	Measured at
T2 (boulder)	MO-03-07	0	1690	2,164E+06	1,053E+05	17,68	ASTER
T2 (boulder)	MO-03-10	0	1690	1,284E+06	6,773E+04	17,67	ASTER
T2 (boulder)	MO-03-11	0	1690	1,824E+06	9,974E+04	17,67	ASTER
T2 (boulder)	MO-03-12	0	1690	1,949E+06	2,322E+05	17,67	ASTER
T2 (boulder)	MO-03-12 0cm	0	1690	1,809E+06	9,576E+04	17,67	ASTER
T2 (boulder)	MO-03-12 40cm	40	1690	2,622E+06	8,295E+05	17,67	ASTER
T2 (boulder)	MO-03-12 60cm	60	1690	8,457E+05	6,867E+04	17,67	ASTER
T2 profile	MO-03-T2-TOP	0	1690	1,759E+06	1,452E+05	17,67	ASTER
T2 profile	MO-03-T2-40	40	1690	1,133E+06	1,410E+05	17,67	ASTER
T2 profile	MO-03-T2-60	60	1690	8,558E+05	1,714E+05	17,67	ASTER
T2 profile	MO-03-T2-125	125	1690	3,528E+05	5,331E+04	17,67	ASTER
T2 profile	MO-03-T2-160	160	1690	4,468E+05	1,026E+05	17,67	ASTER
T2 profile	MO-03-T2-200	200	1690	1,385E+05	3,181E+04	17,67	ASTER
T2 (boulder)	MO-03-14	0	1713	8,275E+05	8,005E+04	17,98	ASTER
T1 (boulder)	MO-03-16	0	1650	1,371E+06	1,281E+05	17,16	ASTER
T1 (boulder)	MO-03-17	0	1650	7,790E+05	2,755E+05	17,15	ASTER
T1 (boulder)	MO-03-18	0	1650	1,496E+06	9,360E+04	17,17	ASTER
T1 (boulder)	MO-03-19	0	1650	1,259E+06	2,302E+05	17,17	ASTER
T1 (boulder)	MO-03-21	0	1650	1,324E+06	1,534E+05	17,17	ASTER
T1 (boulder)	MO-03-22	0	1650	1,549E+06	1,928E+05	17,17	ASTER
T1 profile	MO-03-T3-30	30	1650	8,290E+05	6,339E+04	17,16	ASTER
T1 profile	MO-03-T3-50	50	1650	4,820E+05	1,079E+05	17,16	ASTER
T1 profile	MO-03-T3-75	75	1650	2,757E+05	3,330E+04	17,16	ASTER
T1 profile	MO-03-T3-100	100	1650	2,129E+05	4,263E+04	17,16	ASTER
T1 profile	MO-03-T3-120	120	1650	2,462E+05	1,102E+05	17,16	ASTER
T0 (boulder)	MO-03-30(T)	0	1640	1,745E+05	3,494E+04	17,04	ASTER
T0 (boulder)	MO-03-31 (B)	0	1640	2,004E+05	3,175E+04	17,04	ASTER
T0 (boulder)	MO-03-32	0	1640	3,179E+05	3,163E+04	17,04	ASTER
T0 (valley)	MO-03-46	0	2100	1,103E+05	6,368E+04	23,66	ASTER
T0 (valley)	MO-03-47	0	2100	2,963E+05	2,964E+04	23,66	ASTER
T0 (valley)	MO-03-48	0	2100	4,184E+05	5,247E+04	23,66	ASTER
T0 (valley)	MO-03-49	0	2100	2,622E+05	9,275E+04	23,66	ASTER
T0 (valley)	MO-03-50	0	2100	1,839E+05	4,117E+04	23,66	ASTER
T0 (valley)	MO-03-51(T)	0	2100	1,212E+05	2,053E+04	23,66	ASTER
T0 (valley)	MO-03-55(B)	0	2100	1,852E+05	1,870E+04	23,66	ASTER

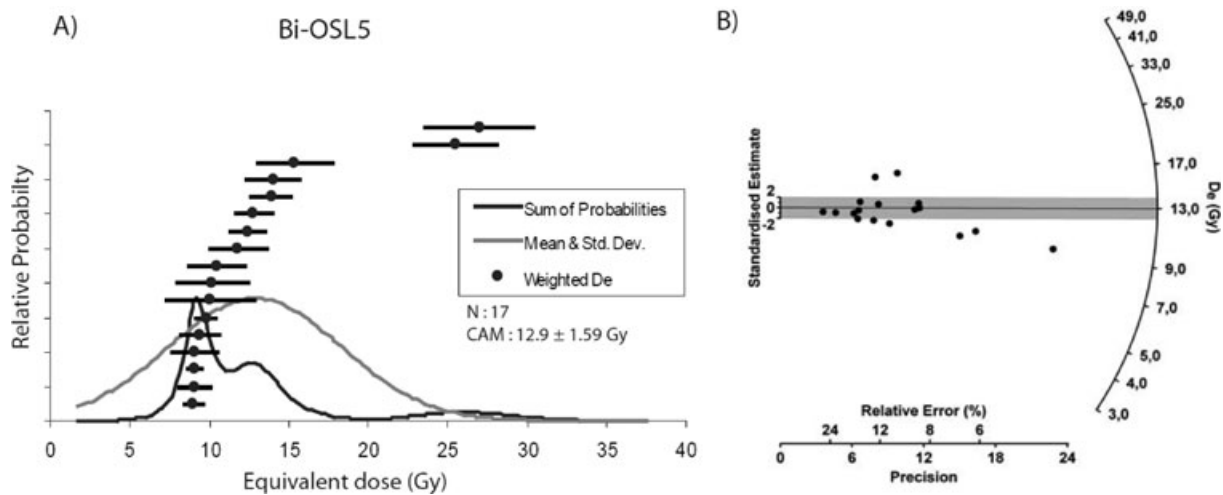


Figure B1. Results of OSL dating on the 250–180 μm quartz grains for Bi-OSL5 sample. (A) Relative probability of equivalent dose (D_e) distribution. The black filled dots are the individual D_e values for each aliquot, the grey line is the distribution of the Gaussian function, and the black line is the probability density function. N represents the total number of aliquots number. (B) Radial plots of the measured doses using Galbraith *et al.* (1999).

are addressed in this Table B2. Fading tests were carried out following the instructions of Auclair *et al.* (2003). The fading rate can be quantified by the g -value, which corresponds to the percentage fading loss per decade of time. These values are a ratio of lumi-

nescence emission after storage of 21 d divided by the immediate measurement (a ratio of 1.0 indicates stable luminescence), and we corrected to the “ g ” factor (i.e. Huntley & Lamothe 2001) using their formulae in a simple worksheet.

Table B1. Luminescence parameters for OSL used in preparation and analyses of samples

Measurement parameters:	
Mineral; grain size:	quartz: 250–180 μm
Stimulation source:	four clusters LED, emission centred at 470 nm
Power delivered to aliquot:	22 mW cm^{-2}
Duration of stimulation:	40 s
Photomultiplier:	TL/OSL DA 15A/B Reader, Minisys 14
Aliquot temperature:	125 °C
Detection filters:	2 Hoya U-340
Normalization:	natural (0.5 s)
Pre-heat:	260 °C for 10 s with same cut heat
Equivalent dose evaluation:	single aliquot regeneration (SAR)
Dose-rate evaluation:	lab and portable χ -spectrometer
Dose-rate range:	3.4 Gy ka^{-1}
Water content:	27 per cent
Cosmic-ray contribution:	3 per cent of total dose rate

Table B2. Luminescence parameters for IRSL used in preparation and analyses of samples.

Measurement parameters:	
Mineral; grain size:	polymineral: 4–11 μm
Stimulation source:	30 IR diodes, emission centred on 880 nm
Power delivered to aliquot:	17 m W cm^{-2}
Duration of stimulation:	100 s
Final signal level:	10 per cent of initial
Photomultiplier:	Thorn-EMI 9635Q
Aliquot temperature:	30 °C IRSL
Detection filters:	390–490 (blue)
Normalization:	natural (0.5 s)
Pre-heat:	124 °C for 64 hr
Delay before measurement:	24 hr or more
Equivalent dose evaluation:	additive method using integrated OSL/satisfactory plateau
Background evaluation:	after bleaching with natural sunlight and quartz window
Alpha effectiveness:	fine grains: $a = 0.07\text{--}0.09$
Dose-rate evaluation:	lab and portable χ -spectrometer
Dose-rate range:	5.18–10.30 Gy ka^{-1}
Water content:	17–29 per cent
Cosmic-ray contribution:	6–8 per cent of total dose rate

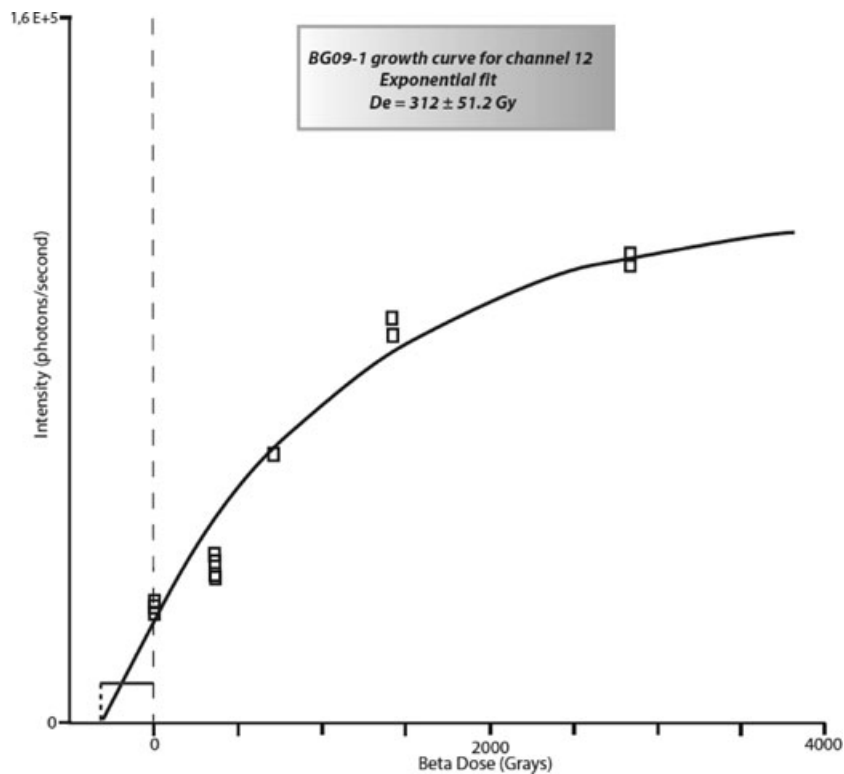


Figure B2. BG09–1 growth curve for IRSL, with the natural plotted on the intensity axis (photons/s). Multiple aliquot additive doses increase in response to increasing beta radiation and are fitted to an exponential. Dose is measured in Grays with 12 s of stimulation shown as 312 ± 51.2 Grays. To obtain the average doses all measurements from 1 to 30 s are used.