



Available online at www.sciencedirect.com

ScienceDirect

Geochimica et Cosmochimica Acta 333 (2022) 200-215

Geochimica et Cosmochimica Acta

www.elsevier.com/locate/gca

Effect and fate of Ni during aging and thermal-induced phyllomanganate-to-tectomanganate transformation

Hui Yin a,b,*, Bruno Lanson c, Shuang Zhang a,b, Lu Liu a,b, Caroline L. Peacock d, Jeffrey E. Post e, Mengqiang Zhu f, Wei Li g, Qian Wang h, Jing Zhang i, Yong Yang a,b, Guohong Qiu a,b, Xionghan Feng a,b, Wenfeng Tan a,b, Fan Liu a,b,*

^a Key Laboratory of Arable Land Conservation (Middle and Lower Reaches of Yangtse River), Ministry of Agriculture and Rural Affairs, College of Resources and Environment, Huazhong Agricultural University, Wuhan 430070, China

Received 17 October 2021; accepted in revised form 13 July 2022; Available online 19 July 2022

Abstract

Phyllomanganates are ubiquitous in a variety of environments and commonly enriched in transition metal elements, such as Ni. The effect of such foreign metal cations on phyllomanganate transformation is widely documented under aqueous conditions together with the induced modification of Ni geochemical behavior. A similar knowledge is lacking however on phyllomanganate transformation and on the induced fate of associated metal elements that may occur under dry conditions, that prevail in deserts and arid areas increasingly exposed to severe droughts or wildfires. The present study shows that crystallinity, morphology, Mn oxidation state, and Ni binding mechanisms are essentially unaffected when aging hexagonal birnessite (Mn oxidation state ~ 3.90 and Ni/Mn molar ratios of 0.00 and 0.13) in the dry state at room temperature for up to 8 years. In contrast, heating aged Ni-doped birnessite to 25–200 °C results in an increased proportion of edge-sharing Ni-Ni (Mn) pairs with increasing temperature induced by the migration of interlayer Ni to birnessite octahedral layers and/or by an increased sharing of coordination oxygens by interlayer Ni/Mn from adjacent layers. Further heating to 400 °C does not change this proportion, with birnessite layer structure being retained. Transformation of Ni-doped birnessite to cryptomelane is complete at 500 °C, while that of Ni-free birnessite is achieved at 400 °C, suggesting that Ni doping increases birnessite thermal stability. Birnessite-to-cryptomelane transformation comes with a strong increase of Mn oxidation state, whereas this parameter remains unchanged in heated birnessite samples. Ni incorporation in the cryptomelane framework, reduces its release during reductive acid dissolution by a factor of 396 ± 15 compared to initial birnessite. These results shed light on

E-mail addresses: yinhui666@mail.hzau.edu.cn (H. Yin), liufan@mail.hzau.edu.cn (F. Liu).

^b State Environmental Protection Key Laboratory of Soil Health and Green Remediation, Ministry of Ecology and Environment, Huazhong Agricultural University, Wuhan 430070, China

^c Univ. Grenoble Alpes, Univ. Savoie Mont Blanc, CNRS, IRD, Univ. Gustave Eiffel, ISTerre, F-38000 Grenoble, France
^d School of Earth and Environment, University of Leeds, Leeds LS2 9JT, UK

^e Department of Mineral Sciences, NHB 119, Smithsonian Institution, Washington, DC 20013-7012, USA

f Department of Ecosystem Science and Management, University of Wyoming, 1000 E. University Ave., Laramie, WY 82071, USA g Key Laboratory of Surficial Geochemistry, Ministry of Education, School of Earth Sciences and Engineering, Nanjing University,

Nanjing 210023, China

^h School of Earth and Atmospheric Sciences, Georgia Institute of Technology, 311 Ferst Dr., Atlanta, GA 30332-0340, USA

[&]quot;School of Earth and Atmospheric Sciences, Georgia Institute of Technology, 311 Ferst Dr., Atlanta, GA 30332-0340, USA Beijing Synchrotron Radiation Facility, Institute of High Energy Physics, Chinese Academy of Sciences, Beijing 100039, China

^{*} Corresponding authors.

mineral transformation affecting layered manganates under dry conditions and on the fate of associated transition metal elements.

© 2022 Elsevier Ltd. All rights reserved.

Keywords: Manganese oxide; Mineral transformation; Dry aging; Thermal treatment; Transition metals

1. INTRODUCTION

Manganese (oxyhydr)oxide minerals are widespread in geological settings such as soils, sediments, ocean nodules, and desert varnishes (Post, 1999; Lu et al., 2019). Minerals from the birnessite family are the most common naturally occurring Mn oxides and may form through both biogenic and abiotic processes (Villalobos et al., 2003; Morgan, 2005; Lanson et al., 2008). These phyllomanganates are composed of layers of edge-sharing [Mn(IV)O₆] octahedra with hydrated interlayers (Giovanoli, 1980). Existence of Mn vacancies and/or isomorphous substitutions of low valence cations such as Mn(III), Ni(II), Co(III), and Fe (III) for Mn(IV) in birnessite layers induce a net negative charge in the layers, which is compensated for by the presence of hydrated cations and/or protons in the interlayer region (Drits et al., 1997; Silvester et al., 1997). Owing to their unique structure (tiny particle sizes, vacancies, and mixed Mn valences), birnessites have strong adsorption and/or oxidation capacities for a variety of trace metals, and thus host a large number of metal cations (Peacock, 2009; Simanova and Pena, 2015). Natural birnessite minerals are generally enriched in metals such as Co, Ni, and Fe (Taylor et al., 1964; Manceau et al., 2014; Lee et al., 2019).

Birnessite-like minerals can be precursors of other naturally occurring Mn oxides such as tectomanganates. Natural cryptomelane, which has a 2×2 tunnel structure, is mainly found in supergene oxidation zones of Mn deposits, lateritic weathering profiles, and volcanic ashes, and may form through authigenic precipitation or transformation from birnessite minerals (Lu and Li, 2015). Birnessite transform to cryptomelane via dissolutionreprecipitation pathways (Jothiramalingam et al., 2006; Zhang et al., 2011, 2016; Sinha et al., 2013) or via topotactic transformation mechanisms (Tu et al., 1994; Ching et al., 2001; Grangeon et al., 2014, 2015). It was proposed that Mn(III) content in the octahedral layers plays a key role in the conversion from phyllomanganates to tectomanganates (Atkins et al., 2014; Grangeon et al., 2014, 2015; Yang et al., 2018). Most studies investigating the phyllomanganate-to-tectomanganate transformation were conducted in solution (Tu et al., 1994; Zhang et al., 2011; Atkins et al., 2014; Yang et al., 2018, 2019; Wu et al., 2019, 2021), consistent with the prevalence of wet conditions in soils and sediments. Dry conditions prevail however in desert soils, that cover more than 20 % of Earth's land area (Zhu et al., 2022), and in soils affected by seasonal droughts. As a result of the present climate change, the frequency of extreme weather patterns such

as severe droughts and extreme heat events has increased over the past two decades, a long-term tendency according to climate projections (Vicente-Serrano et al., 2020). As the phyllomanganate-to-tectomanganate transformation processes even in the dry state (Atkins et al., 2014; Grangeon et al., 2014, 2015; Yang et al., 2018), it is important to investigate such mineral transformation and the fate of associated metal cations in dry conditions.

Very few studies have reported dry-state transformation over long periods of time (Grangeon et al., 2014, 2015) and, to our knowledge, no study has investigated dry-state transformation at elevated temperatures linked for example to the increasing frequency of wildfires (Murphy and Timbal, 2008). During large wildfires, soil temperature may exceed 600 °C (Rein et al., 2008), possibly resulting in the transformation of soil minerals (Johnston et al., 2016; Li et al., 2019). In particular, laboratory experiments showed that birnessite can be thermally transformed to cryptomelane (Chen et al., 1986; Golden et al., 1986). Such transformation may significantly affect the fate of elements in the initial minerals. For example, heating As(V)-coprecipitated iron oxyhydroxysulfate mineral schwertmannite at temperatures higher than 400 °C results in the formation of hematite and causes the initially structurally incorporated As to be surface-complexed, thus enhancing As mobilization (Johnston et al., 2016). Similar processes have also been reported for transformation of naturally occurring goethite-rich materials upon heating (Perrier et al., 2006; Landers et al., 2009). However, no studies have yet investigated the fate of metal elements associated with Mn oxide minerals during their thermal-induced transformation.

The present study thus investigates the possible phase transformation affecting Ni-doped hexagonal birnessites during aging in the dry state at room temperature for 5 and 8 years and subsequent heating over a wide range of temperatures (50-500 °C) and the induced fate of associated Ni. Powder X-ray diffraction, electron microscopies, thermogravimetric analysis, and Ni and Mn K-edge X-ray absorption fine structure spectroscopy are combined with acid dissolution experiments for this purpose. Nickel was chosen as a model metal cation because of being 1) one important trace metal nutrient and contaminant, 2) a redox-insensitive element, and 3) because its binding mechanisms on birnessite-like minerals have been thoroughly studied in the literature. The present results provide insights into the mobility and bioavailability of metal nutrients and contaminants associated with the phase changes of host Mn oxide minerals, and may deepen our understanding of Mn-mediated metal geochemical behaviors in dry soils and during extreme heating events.

2. MATERIALS AND METHODS

2.1. Dry aging of Ni-doped birnessites at room temperature

Hexagonal birnessite, HB, and Ni-doped birnessite, Ni10, from our previous study (Yin et al., 2012) were used here. These samples were synthesized according to the protocol of McKenzie (1971). Briefly, this protocol involves the reduction of a boiling KMnO₄ solution (300 mL, 0.667 M) by an HCl solution (45 mL, 6 M) in the absence/presence of NiCl₂ (Ni/Mn molar ratio of 0.10). The pHs of the synthetic suspensions after mineral crystallization were ~ 7.2 and \sim 5.5, respectively. After washing thoroughly the obtained solids, they were dried at 40 °C for several days, ground carefully to particle sizes below 100 mesh, and reserved for further use. HB and Ni10 samples have specific surface areas (SSAs) of 19 and 71 m²·g⁻¹, respectively (Yin et al., 2012). The Ni content in Ni10 is 6.1 wt%, and the K contents in HB and Ni10 are 8.4 and 3.8 wt% respectively (Yin et al., 2012).

Aliquots of these birnessite powders were stored in sealed polystyrene tubes and kept in the dark at room temperature. After aging for 5 and 8 years, some of the samples were thoroughly characterized. As-obtained, aged samples were correspondingly named as HB_Ny and Ni10_Ny (N = 5 or 8; Table 1).

2.2. Thermal treatment of the 8-year aged samples

In addition, the 8-year aged samples (HB_8y and Ni10_8y) were heated to temperatures (50, 100, 150, 200, 300, 400 and 500 °C), spanning the common range of soil temperatures reported during wildfires (Rein et al., 2008). Typically, ~0.1 g of each sample was placed in a ceramic crucible and heated in a muffle furnace in air. After increasing the furnace temperature to the set value (rate of 10 °C·min⁻¹), the sample powder was kept at this temperature for 2 h, consistent with previous studies (Cai et al., 2001). The furnace door was then opened to let the powder cool down naturally to room temperature, and as-obtained solids were stored for further analysis. These samples were labelled HB_8y_T and Ni10_8y_T, where T stands for the temperature used in °C (Table 1).

2.3. Solid characterization

Powder X-ray diffraction (XRD) analysis was used to identify structural changes of Mn oxide samples. Data were collected on a Bruker D8 Advance diffractometer. As the XRD profile changes during drying aging at room temperature were expected to be minimal, the diffraction data of the 5- and 8-year aged samples were collected at a step size of 0.04 °20 Cu K α using a counting time of 10 or 40 sec per step. The XRD data of the thermal transformation products were collected using a scanning rate of 10° per min with a step size of 0.02 °20. Field emission scanning electron microscopy (FESEM) images of the samples were obtained using a Jeol JSM-6700F microscope after coating the minerals with a gold evaporated film for 30 s. High angle annular dark field (HAADF) scanning transmission electron

microscopy (STEM) images of Ni10_8y_500 were obtained by a JEOL JEM-ARM200F microscope equipped with a spherical aberration correction system. Thermogravimetric analysis (TGA) of the Ni-doped birnessites before and after 8-year aging was conducted from $\sim 35~^{\circ}\text{C}$ to 800 $^{\circ}\text{C}$ on a NETZSCH TG 209 thermal analyser in N2 or dry air at a flow rate of 20 mL·min $^{-1}$ with a heating rate of 10 $^{\circ}$ C·min $^{-1}$.

2.4. X-ray absorption fine structure (XAFS) spectra

The XAFS spectra were measured at room temperature on the 1W1B beamline at the Beijing Synchrotron Radiation Facility (Yin et al., 2012, 2015). Ni K-edge data were obtained in fluorescence mode over the 8157–9063 eV range. Monochromator energy was calibrated using a Ni metal foil ($E_0 = 8333$ eV). Mn K-edge data were collected in transmission mode over the 6363–7311 eV range for most samples (6340–7104 eV for 8-year aged HB and Ni10 samples heated to 50 and 150 °C). A Mn metal foil was used to calibrate the monochromator ($E_0 = 6539$ eV).

Data reduction and analysis were performed using Ifeffit/SIXPack (Ravel and Newville, 2005; Webb, 2005). Parameters used for background removal of Ni K-edge spectra were: $E_0 = 8343$ eV, Rbkg = 1.0 Å and kweight = 2. After background removal and normalization, spectra were converted into the k³-weighted functions and Fourier transforms (FTs) were calculated over 2.8-12.5 Å^{-1} (k range) and 1–6 Å (R range), respectively. Mn K-edge spectra were background-subtracted using the following parameters: $E_0 = 6553$ eV, Rbkg = 1.0 Å and kweight = 2. XAFS spectra were then converted into the k³-weighted functions, and FTs were calculated over 3.4- 12.1 Å^{-1} and 1-4 Å (k and R ranges, respectively). Structural parameters (distance, R; coordination number, CN; and Debye-Waller factor, σ^2) for Ni local environments were obtained by fitting the experimental k³-weighted EXAFS spectra to the standard EXAFS equation (Kelly et al., 2008) using several single-scattering paths, as described previously (Peña et al., 2010; Yin et al., 2012). Phase and amplitude functions for single-scattering paths were calculated using FEFF7 (Rehr et al., 1992). Clusters of atoms used for the analysis of Ni K-edge EXAFS spectra were based on chalcophanite crystal structure (Post and Appleman, 1988) and substituting Ni for interlayer Zn or layer Mn. An amplitude reduction factor (S_0^2) of 0.848 was adopted for Ni, consistent with our previous study (Yin et al., 2012).

To determine the relative proportions of heterovalent Mn cations (Mn⁴⁺, Mn³⁺ and Mn²⁺) in Mn oxides, Mn K-edge XANES spectra were analyzed using the Combo method (Manceau et al., 2012). In a typical procedure, the 17 standard spectra used to develop the method (Table S1) were first imported. 4–8 references with nonnegative fractions were then used for the linear combination fitting of normalized XANES spectra. Mn K-edge EXAFS data were fitted using birnessite (Grangeon et al., 2010) or cryptomelane (Yin et al., 2015) model. An amplitude reduction factor (S₀²) of 0.973 was adopted from our previous study (Yin et al., 2014).

Table 1

Nomenclature of hexagonal turbostratic birnessites and their derivatives

Sample	Dry aging (year	year)	Thermal treatm	atment of the 8 year aged sample at different temperatures (°C)	l sample at different te	emperatures (°C)			
	5	8	50	100	150	200	300	400	500
HB	HB_5y	HB_8y	HB_8y_50	HB_8y_100	HB_8y_150	HB_8y_200	HB_8y_300	HB_8y_400	HB_8y_500
Ni10	Ni10 5y	Ni10 8y	Ni10 8y 50	Ni10 8y 100	Ni10 8y 150	Ni10 8y 200	Ni10 8y 300	Ni10 8y 400	Ni10 8y 500

In all EXAFS fits, the number of independent variables used was smaller than the number of independent data points. During Fourier transform and EXAFS data fitting, a Hanning window was used.

2.5. Acid dissolution experiments

Consistent with previous studies (Beak et al., 2008; Oin et al., 2017) acid dissolution experiments were performed to investigate Ni crystal chemistry, stability, and bioaccessibility in these samples. Briefly, 0.0400 g of solid were dissolved at 37 °C in 250 mL of 2 M HCl solution while stirring at a constant rate. These conditions are similar to in vitro measurements of metal bioaccessibility in the gastric phase (Beak et al., 2008). Suspension aliquots (5 mL) were withdrawn at predetermined intervals and filtered immediately through 0.22 µm membrane filters. The concentrations of Mn²⁺ and Ni²⁺ in the filtrate were analyzed using atomic absorption spectrometry (AAS, Varian AAS 240FS). The metal (Me) dissolution ratio was calculated as $\chi_{Me} = Me$ (t)/Me(total), where Me(t) is the aqueous metal concentration at time t, and Me(total) is the final metal concentration after complete dissolution of the solid.

3. RESULTS

3.1. Changes in mineralogical phases

Powder XRD patterns of the Ni-doped birnessite samples before and after dry aging at RT for 5 and 8 years (Fig. 1a,b) are typical for turbostratic hexagonal birnessites, and can be indexed using a C-centered twodimensional unit cell (Drits et al., 1997, 2007). The unsplit (31, 02) bands and the ratios of (11, 20) to (31, 02) d values close to $\sqrt{3}$ both confirm the hexagonal layer symmetry (Drits et al., 2007). XRD profiles of aged samples show only minimal differences with those of fresh samples. For both HB and Ni10 samples, the position of the (31, 20) diffraction band shifts slightly towards lower angle for the 8-year aged sample compared to fresh and 5-year aged samples, indicating a slight increase of the layer lattice parameter b, without significant distortion of the hexagonal layer symmetry however (Grangeon et al., 2008). Consistent with the high values of Mn average oxidation state in the initial products (3.86 and 3.95 for HB and Ni10, respectively - Yin et al., 2012), similarity of the XRD profiles of fresh and aged samples supports the stability of birnessite layer and interlayer structures and the stability of these minerals in the dry state at room temperature.

Upon heating, powder XRD patterns of the 8-year aged samples undergo significant changes (Fig. 1c,d). When heated to temperatures ≤ 300 °C, all samples essentially retain their layer structure. After heating to 400 °C, HB_8y is fully transformed to cryptomelane (K(Mn₇⁴⁺Mn³⁺)O₁₆ – ICDD 74–1451), while Ni10_8y retains most of its layer structure, being only partly transformed to cryptomelane. After heating to 500 °C for 2 h, Ni10_8y is essentially transformed to cryptomelane, with a minor NiMn(IV)O₃ (ICDD 75–2089) impurity. Further

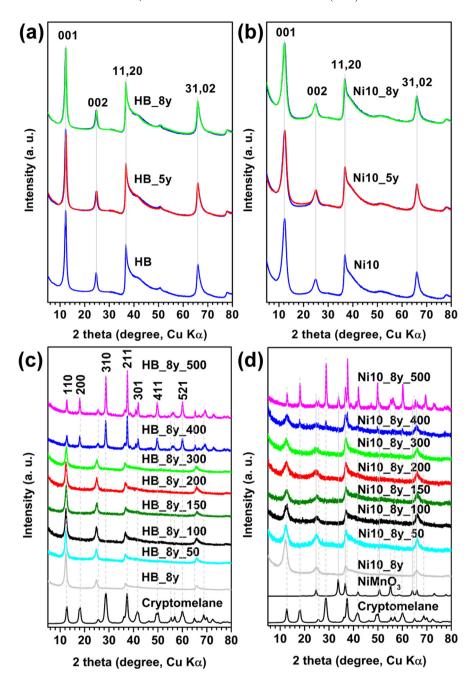


Fig. 1. Powder XRD patterns of pristine and aged turbostratic hexagonal birnessite samples HB (a) and Ni10 (b) for 5 and 8 years at 25 °C, and of thermal transformation products derived from 8-year aged samples HB_8y (c) and Ni10_8y (d). In panels (a)-(b), the high-angle region (30–80° 2θ) was scaled by a factor 4 (HB) or 2 (Ni10), and in each series the pristine sample was overlaid to emphasize the evolution of the XRD profiles upon aging. In panels (c)-(d), calculated XRD patterns for cryptomelane (ICDD 74–1451; dashed light gray lines) and NiMn (IV)O₃ (ICDD 75–2089; solid light gray lines) were overlaid to identify the transformation products at high temperatures.

linear combination fitting (LCF) analysis of Mn K-edge EXAFS spectra of thermal transformation products of HB_8y and Ni10_8y by using spectra of Ni10_8y_300 and HB_8y_500 as birnessite and cryptomelane standards, respectively, confirms the XRD results. For example, LCF analysis shows that Ni10_8y_400 is composed of 87 ± 1 % birnessite and 13 ± 1 % cryptomelane (Fig. S1 and Table S2).

3.2. Thermal stability of 8-year aged birnessites in N_2 or dry air

Eight-year aged birnessites are more stable than their fresh counterparts upon heating in N_2 atmosphere, although weight losses occur almost at the same temperatures (Fig. 2 and Table S3a). The amount of physisorbed water is identical in HB and HB_8y (35–150 °C – 8.8–

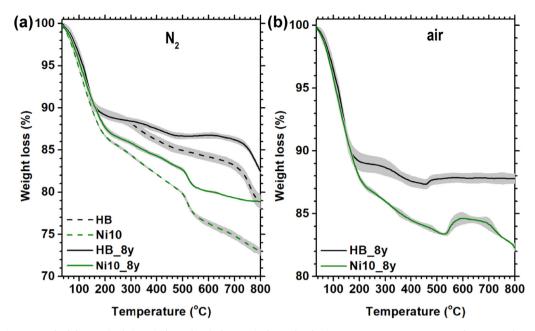


Fig. 2. TGA curves of Ni-free and Ni-doped birnessites before and after aging in dry state at room temperature for 8 years in N_2 atmosphere (a), and those of 8-year aged samples in dry air atmosphere (b). The data for the fresh Ni-doped samples (dashed lines) are from Yin et al. (2012). For all the samples, the TGA curves are averaged from 2 to 3 measurements, standard errors being shown as light gray shading.

9.3 %), while that in Ni10_8y (35–180 °C – 11.0 \pm 0.3 %) is slightly lower than that in Ni10 (12.0 \pm 0.1 %). The amount of structural water (150–700 °C), including hydroxyl groups and interlayer water molecules (Lee et al., 2007; Yin et al., 2012), is significantly lower in aged birnessites (4.8 \pm 0.5 % and 5.8 \pm 0.4 % in HB_8y and Ni10_8y, respectively) compared to corresponding fresh samples (7.5 \pm 0.9 % and 7. 2 \pm 0.2 %).The weight losses assigned to lattice oxygen (700–800 °C) are alike in HB and HB_8y (4.3 \pm 1.0 % and 3.8 \pm 0.5 %, respectively). Contrastingly, this weight loss is lower in Ni10_8y (4.1 \pm 0.2 %) compared to Ni10 (7.5 \pm 0.5 %). The lattice oxygen weight loss is related to Mn(IV) reduction and formation of phases such as Mn₂O₃ (Lee et al., 2007).

The profiles for the release of both physisorbed and structural water molecules from the 8-year aged samples upon heating in air are similar to those in N_2 (Fig. 2b and Table S3b). With increasing temperature, a weight increase is observed however for both HB_8y (460–800 °C – 0.4 ± 0.6 %) and Ni10_8y (520–660 °C – 1.0 ± 0.3 %). These weight gains are likely due to the oxidation of low valence $Mn^{2+/3+}$ by O_2 (Tang et al., 2014). Weight loss occurs again from 660 °C for Ni10_8y, owing to the Mn reduction and lattice O loss during thermal decomposition to phases such as Mn_2O_3 (Tang et al., 2014). These results are consistent with previous studies that indicate oxidation of $Mn^{2+/3+}$ at intermediate temperatures in air and reductive decomposition of MnO_2 at high temperatures both in air and N_2 (Lee et al., 2007; Tang et al., 2014).

3.3. Morphology changes

After aging for 8 years, birnessite morphology, 3D hierarchical microspheres composed of 2D disk-like plates, changes little, with the thickness of platy crystals remain-

ing also unchanged (Fig. 3). Heating HB 8y up to temperatures ≤ 300 °C slightly decreases microsphere size however, and even more so plate thickness (Fig. S2-4). HB 8y 400 particles appear as prisms, consistent with the transformation to cryptomelane. This morphology is different from commonly reported needles (Sinha et al., 2013; Grangeon et al., 2015; Zhang et al., 2016), but similar to that of cryptomelane obtained from thermal transformation of triclinic birnessite (Chen et al., 1986). After heating HB_8y further to 500 °C, cryptomelane prismatic crystals grow in all three dimensions compared to HB 8y 400. Similar to the corresponding fresh crystals, Ni10_8y crystals are smaller than HB_8y ones (Fig. 3, Fig. S2-4). Heating Ni10_8y to temperatures ≤ 400 °C, modifies crystal size and plate thickness only marginally. Although powder XRD analysis shows a minor amount of cryptomelane in Ni10 8y 400, no prismatic crystals were observed, whereas Ni10 8v 500 consists mainly of prismatic crystals (Fig. 3g). Both the HAADF-STEM image and the corresponding Fast Fourier transform (FFT) pattern obtained on a prismatic Ni10_8y_500 crystal confirm the nature of these crystals exhibiting cryptomelane [200] facets with a d spacing of 4.98 Å (Fig. 3h). Atomic image (Fig. 3i) shows the 2×2 tunnels of the cryptomelane framework. In this image, the uniform dark stripes represent the K atoms and the coordinated water molecules rather than Ni in the interlayer regions. This indicates that Ni is most likely structurally incorporated in the cryptomelane framework rather than adsorbed in tunnels and/or to crystal surface.

3.4. Changes in Mn Average Oxidation State (AOS)

Manganese K-edge XANES analysis using the Combo method (Manceau et al., 2012) shows that upon dry aging,

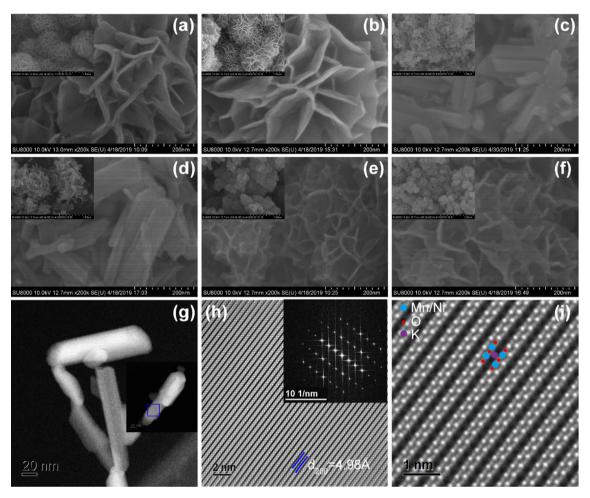


Fig. 3. High magnification FESEM images of HB_8y (a), HB_8y_300 (b), HB_8y_400 (c), HB_8y_500 (d), Ni10_8y (e), and Ni10_8y_400 (f) with the low magnification images in the inset of each panel. TEM image of Ni10_8y_500 (g) with an independent crystal in the inset with an area labeled using a blue bracket for the HAADF-STEM image with FFT in the inset (h) and the atomic resolution image (i). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

the proportion of low valence Mn (Mn $^{3+/2+}$) increases only slightly at the expense of Mn $^{4+}$ (Fig. 4 and Table S1). Mn AOS values calculated for HB_8y and Ni10_8y (3.81 and 3.90, respectively) are similar to those of HB and Ni10 (Yin et al., 2012), within the accuracy (± 0.04) of the method (Manceau et al., 2012). As birnessite Mn AOS is a key parameter for their reactivity towards various redox reactions (Wang et al., 2018), this similarity suggests that aged birnessites can effectively mediate the geochemical behavior of various redox-sensitive contaminants in the environment.

Heating of HB_8y and Ni10_8y modifies the relative proportions of various Mn species depending on temperature. Below 300 °C, the proportions of heterovalent Mn cations in HB_8y and Ni10_8y are essentially constant whereas increasing temperature to 400 °C for HB_8y and 500 °C for Ni10_8y induces the oxidation of low-valence Mn cations to Mn⁴⁺ (Fig. 4). Consistently, Mn AOS values calculated for HB_8y heated to 50–300 °C range 3.81–3.84, and increase to 3.91–3.93 after heating to temperatures \geq 400 °C. Similarly, Mn AOS values calculated for Ni10_8y heated to 50–150 °C remain con-

stant at 3.89-3.91. When heating Ni10_8y further to 200-400 °C, Mn AOS values decrease to 3.82-3.85, most likely owing to Mn⁴⁺ reduction. Over this temperature range, the weight loss due to hydroxyl groups probably results also in the loss of oxygen anions and thus in the reduction of high valence Mn cations. Contrasting evolutions of Mn AOS observed for HB 8y and Ni10 8y over this temperature range are probably ascribed to the larger loss of hydroxyl groups in Ni10_8y compared to HB_8y, possibly because of the lower crystallinity of the former (Fig. 1a,b and Fig. 2b). At 500 °C, the content of low valence Mn in Ni10 8y 500 is greatly decreased, and its Mn AOS is sharply increased to 3.98. The increase in Mn AOS at 400-500 °C is most likely due to the oxidation of low valence Mn by O₂ in the air, consistent with TGA results (Tang et al., 2014). In the latter case, weight gain related to absorption of O₂ by these samples occurs at 460-520 °C however, most likely owing to the steady temperature increase rate (Fig. 2b). Such an increase of Mn AOS after the birnessite-to-cryptomelane transformation is consistent with previous reports (Cai et al., 2001; Birkner and Navrotsky, 2017).

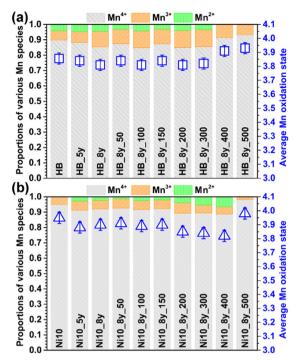


Fig. 4. Proportions of Mn species (Mn⁴⁺, Mn³⁺, and Mn²⁺) in Nifree HB (a) and Ni-doped Ni10 (b) birnessite samples after aging at room temperature for 5 and 8 years and in samples derived from heat treatment of the 8-year aged samples at different temperatures, and corresponding variations of the average Mn oxidation state. Data for fresh HB and Ni10 are from Yin et al. (2012).

3.5. Changes in local atomic environments of Mn

Consistent with the results of powder XRD analysis, dry aging at room temperature changes little the birnessite laver structure. Similarly, upon heating temperatures < 300 °C, the k^3 chi(k) functions and corresponding FTs (light gray bands in Fig. 5a,c) remain essentially unchanged, confirming the birnessite structure of heated HB_8y samples. After heating to 400 °C, the shape of the oscillation at 7.2–9.8 Å^{-1} (k³chi(k) function) changes to match that of cryptomelane (Yin et al., 2015) and the magnitude of the third peak (R + Δ R \sim 3.1 Å) in the corresponding FT is greatly increased, confirming the transformation of birnessite to cryptomelane. The oscillation profile and FT of Ni10_8y (light gray bands in Fig. 5b,d) indicate the same mineralogical evolution after heating to 500 °C.

Fittings of the Mn K-edge EXAFS spectra for heated birnessite samples based on a birnessite structure model, and on a cryptomelane structure model for HB_8y_400, HB_8y_500, and Ni10_8y_500, provide good agreements (Fig. 5 and Table S4). All Mn-O distances in the [MnO₆] octahedra of birnessite and cryptomelane are similar at 1. 904 ± 0.001 Å (n = 17) and 1.895 ± 0.001 Å (n = 3), respectively. Edge-sharing Mn-Mn distances in birnessite derived from HB (2.881 \pm 0.002 Å; n = 8) are slightly longer than those derived from Ni10 (2.876 \pm 0.003 Å; n = 9) at an α significance level of 0.05 (*t*-test). These longer distances are likely related to the higher content of layer Mn(III) in birnessite derived from HB compared to those derived from Ni10 (Fig. 4 and Table S1). Edge-sharing Mn-Mn distances

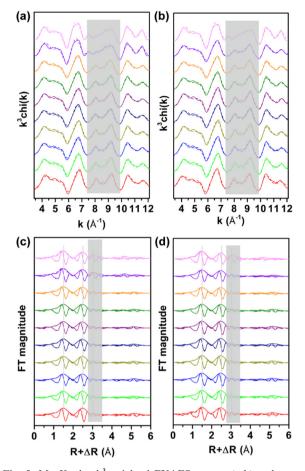


Fig. 5. Mn K-edge k³-weighted EXAFS spectra (a-b) and corresponding Fourier transforms (c-d) of aged and heat-treated HB (left) and Ni10 (right) birnessite samples. Samples are colored-coded: fresh samples (red), 5-year aged (green), 8-year aged (blue), HB_8y and Ni10_8y samples heated to 50 °C (dark yellow), 100 °C (navy), 150 °C (purple), 200 °C (olive), 300 °C (orange), 400 °C (violet), and 500 °C (light magenta). Solid and dashed lines are experimental data and their best fits, respectively. All spectra were fitted using a birnessite model derived from Grangeon et al. (2010), except for HB_8y_400, HB_8y_500, and Ni10_8y_500 spectra that were fitted using a cryptomelane model derived from Yin et al. (2015). Light grey bands indicate the main changes in these spectra. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

in cryptomelane (2.885 \pm 0.003 Å; n = 3) are statistically independent of the precursor and are slightly larger than those in birnessite layers (2.878 \pm 0.004 Å; n = 17).

Significant changes are observed for the distances between interlayer Mn(Ni) to the nearest layer Mn(Ni) (Me^E-Me^{TC}; Me = Mn, Ni) (Fig. 6a and Table S4). HB_8y, Ni10_8y and layered (birnessite) samples derived from them all exhibit similar Me^E-Me^{TC} distances (3.48–3.52 Å – Fig. 6b), consistent with values reported for hexagonal birnessites (Grangeon et al., 2010). This distance is significantly decreased to 3.44–3.45 Å for cryptomelane (Fig. 6c), consistent with previous studies (Lee et al., 2007; Yin et al., 2015). Shortening of Me^E-Me^{TC} distances in cryptomelane compared to phyllomanganates is consistent with the higher proportion of Mn⁴⁺ cations in the former.

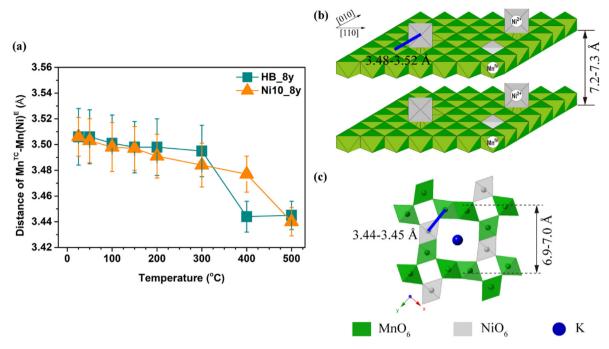


Fig. 6. Evolution as a function of temperature of the EXAFS-derived Mn-Mn(Ni) corner-sharing distances in HB and Ni10 (a), and illustrations of these distances (blue bold solid lines) in the birnessite (b) and cryptomelane (c) structures. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

3.6. Evolution of Ni binding mechanisms in the mineral

Nickel K-edge k^3 chi(k) oscillations and corresponding FTs of Ni10 and Ni10_8y are similar, indicating similar Ni local environments (Fig. 7). In the FTs, three main peaks located at $R + \Delta R \sim 1.6$ Å, ~ 2.4 Å and ~ 3.0 Å mainly correspond to the first Ni-O pair in [NiO₆] octahedron, edge-sharing Ni-Ni(Mn) pair (Ni^E) and triple corner-sharing pairs between interlayer Ni(Mn) and layer Ni(Mn) (Ni^{TC} – Peacock, 2009; Peña et al., 2010; Yin et al., 2012). Intensity ratios of these three peaks are constant indicative of similar proportions of Ni^E and Ni^{TC} (Fig. 7 and Table 2). In Ni10_8y, 74 \pm 4 % of the Ni content is located above/below vacant layer sites, which is statistically equivalent to Ni10 (Yin et al., 2012).

The k^3 -weighted Ni K-edge EXAFS functions of heated Ni10_8y samples exhibit significant changes (Fig. 8A). Upon heating, oscillations at $\sim 7.6~\text{Å}^{-1}$ and $\sim 8.2~\text{Å}^{-1}$ are gradually combined into one oscillation and a shoulder at $\sim 6.6~\text{Å}^{-1}$ increases in intensity. The corresponding FTs (Fig. 8B) indicate that these evolutions are related to the reversing of Ni^E and Ni^{TC} peak relative intensities. Quantitative fitting indicates Ni^E proportions of 27 \pm 4 %, 34 \pm 5 %, 37 \pm 5 %, 48 \pm 5 %, 48 \pm 7 %, and 51 \pm 10 % for Ni10_8y samples heated to 50, 100, 150, 200, 300, and 400 °C, respectively (Table 2).

3.7. Increased Ni stability during acid dissolution

Acid dissolution experiments were conducted to detect possible modifications of Ni availability in heat-treated Mn oxides (Fig. 9). For Ni10_8y and Ni10_8y_200, Ni²⁺ and Mn²⁺ release rates in solution are large over the first

few hours. With time, the metal release rates are gradually reduced till complete dissolution, because of the lower pool. In both samples, Ni²⁺ and Mn²⁺ releases are concurrent but Ni²⁺ is released slightly faster than Mn²⁺, consistent with the predominance of Ni^{TC}. Compared to Ni10_8y, the releases of Ni²⁺ and Mn²⁺ from Ni10_8y_200 are slower. $\sim 99 \%$ of both Ni²⁺ and Mn²⁺ are released after 5 and 7 h for Ni10_8y and Ni10_8y_200, respectively, possibly as a result of a minor crystal growth. Ni10 8y 500 dissolution is much slower as 50 days are needed to achieve complete dissolution. Transformation of the initial birnessite to cryptomelane is the most likely hypothesis for such a dramatic decrease of dissolution rate. In addition, Ni²⁺ and Mn²⁺ releases are not concurrent any more, Mn²⁺ being released to solution much faster than Ni^{2+} .

Nickel and Mn release kinetic curves can be fitted using the Kabai equation, $\chi_{Me}=1\text{-exp}(\text{-(kt)}^{\alpha})$, where k and α are coefficients related to the dissolution rate constant and the mineral structure, respectively (Fig. 9 – Kabai, 1973; Kaur et al., 2010). Parameters k for both Ni²⁺ and Mn²⁺ dissolution are steadily decreasing with increasing temperature. Ni and Mn release rates measured for Ni10_8y_200 are 67 \pm 4 % and 71 \pm 3 % those from Ni10_8y, respectively. Ni and Mn release rates measured forNi10_8y_500 are dramatically reduced to 0.25 \pm 0.01 % and 1.51 \pm 0.06 % of those from Ni10_8y, respectively.

The χ_{Ni} - χ_{Mn} curves obtained for Ni10_8y and Ni10_8y_200 follow the 1:1 line with a slight convexity however (Fig. 9d), and are essentially similar to that of the fresh Ni-doped birnessite (Qin et al., 2017). This suggests a uniform Ni distribution in these birnessite samples

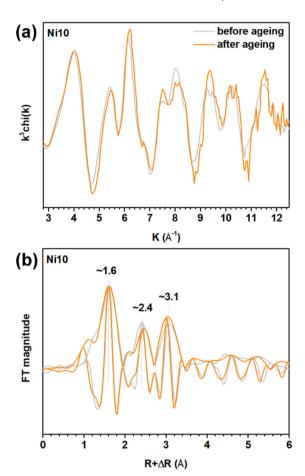


Fig. 7. Nickel K-edge k^3 -weighted functions (a) and corresponding Fourier transforms (b) of Ni-doped birnessite samples before (light gray lines) and after (orange lines) dry aging at room temperature for 8 years. Fourier transforms (b) were normalized to the first Ni-O shell (R $+\Delta R \sim 1.6~\text{Å})$. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

(Alvarez et al., 2006, 2007), consistent with elemental mapping (Fig. S5). Contrastingly, the χ_{Ni} - χ_{Mn} curve obtained for Ni10_8y_500 differs significantly from the previous two exhibiting a strong concavity, indicative of Ni location in the inner parts of cryptomelane framework (Alvarez et al., 2006, 2007).

4. DISCUSSION

4.1. Stability of (Ni-doped) birnessite during aging in dry state

When aged in the dry state at room temperature for up to 5 and 8 years, the structure and properties of investigated Ni-free and Ni-doped hexagonal birnessites are almost unchanged, preserving their hexagonal layer symmetry, morphology and Mn AOS. This suggests that layered Mn oxides such as birnessite may not as unstable as previously thought (Grangeon et al., 2014, 2015). Increased birnessite stability observed in the present study is most likely related

to their low content of Mn³⁺ as previously discussed for the conversion of δ-MnO₂, a poorly crystalline analogue of birnessite, to cryptomelane under similar experimental conditions (Grangeon et al., 2014, 2015), possibly reinforced by larger crystal sizes. Layer Mn(III) in phyllomanganates plays indeed a key role in the phyllomanganate-to-tecto manganate transformation (Cui et al., 2009; Atkins et al., 2014; Grangeon et al., 2014; Yang et al., 2018). Mn(III) contents of HB and Ni10 range 0.05-0.06 (Yin et al., 2012), much lower than those (0.30) of the reacting δ -MnO₂ of Grangeon et al. (2014). Increased crystal size of the present birnessites compared to δ-MnO₂ used by Grangeon et al. (2014) likely reinforces phyllomanganate stability. Crystals with small particle sizes generally have high total free energy indeed and are thus prone to phase transformation (Luo et al., 2008; Navrotsky et al., 2008). Theoretically, stability of hexagonal turbostratic birnessite should thus be more stable than δ-MnO₂, owing to its larger crystallite sizes. Powder XRD patterns of our Ni-doped birnessites have strong (00 l) reflections (Fig. 1a,b) that are absent in those of δ-MnO₂, suggesting a larger number of layers stacked coherently along the c*-axis. Further, the physical sizes of the present birnessite plates (diameters of 100-300 nm) are also much larger than those of the vernadite samples (mean diameter of ~ 10 nm – Yin et al., 2012; Grangeon et al., 2015).

Subject to a similar (micro-)structure, birnessite may thus be stable in arid environments such as upland soils and rock surfaces in deserts and represent a long-term sink for trace metal elements. Dark coatings commonly found on the upper surfaces of rocks in arid and semiarid regions, such as Gobi deserts in China and Mongolia and in karst areas, are identified to be mainly composed of birnessite (Lu et al., 2019), consistent with birnessite stability observed in our aging experiments. In the present study, birnessites were aged in the dark with limited O₂ in sealed tubes. On the other hand, rock surfaces exposed to both sun light and air, and birnessite from varnishes and coatings may undergo additional evolution keeping their layer structure anyhow. Field and laboratory studies have evidenced the reductive dissolution of birnesssite under photo illumination of aqueous suspensions however (Sunda and Huntsman, 1994; Kim et al., 2012; Liu et al., 2021), consistent with theoretical calculations (Sherman, 2005; Kwon et al., 2009). Further investigation is needed to assess birnessite stability of birnessites under dry light condition and underlying mechanisms at play.

4.2. Enhanced thermal stability of birnessite by Ni doping

The present results clearly suggest that Ni doping hampers the thermal-induced transformation of birnessite to cryptomelane, although thermodynamically favored. This mineral transformation may be affected by the presence of foreign cations into the birnessite structure however (Birkner and Navrotsky, 2017). It was for example reported that Cu²⁺ adsorption impedes Mn²⁺-induced transformation of vernadite to other Mn oxides at room temperature, ascribing the increased stability to the decrease of mineral

Table 2 Ni K-edge EXAFS structural parameters for thermal transformation products of the 8-year aged Ni10 sample.

Shells	Parameters	Ni10_8y	Ni10_8y_50	Ni10_8y_100	Ni10_8y_150	Ni10_8y_200	Ni10_8y_300	Ni10_8y_400
Ni-O ₁ ^a	R (Å)	2.040(0.006)	2.039(0.006)	2.038(0.007)	2.034(0.007)	2.035(0.007)	2.013(0.007)	2.015(0.009)
	$\sigma^2(\mathring{A})$	0.0051(0.0004)	0.0050(0.0004)	0.0051(0.0005)	0.0052(0.0005)	0.0050(0.0005)	0.0061(0.0005)	0.0065(0.0006)
$Ni-Mn_{1,E}$	CN	$6\times(1-f)^{b}$	$6\times(1-f)$	$6\times(1-f)$	$6\times(1-f)$	$6\times(1-f)$	$6\times(1-f)$	$6\times(1-f)$
	R (Å)	2.860(0.010)	2.846(0.010)	2.849(0.010)	2.845(0.009)	2.863(0.009)	2.849(0.009)	2.865(0.013)
	σ^2 (Å)	σ^2 (Ni-Mn _{1,TC})						
Ni-O ₂ ^a	R (Å)	3.340(0.065)	3.326(0.046)	3.317(0.039)	3.337(0.059)	3.357(0.103)	3.299(0.046)	3.293(0.052)
	σ^2 (Å)	0.0165(0.0086)	0.0103(0.0052)	0.0062(0.0035)	0.0114(0.0067)	0.0217(0.0143)	0.0105(0.0058)	0.0096(0.0062)
Ni-Mn _{1,TC}	CN	$6 \times f$						
	R (Å)	3.471(0.008)	3.484(0.009)	3.487(0.013)	3.481(0.011)	3.464(0.012)	3.457(0.016)	3.453(0.023)
	σ^2 (Å)	0.0057(0.0005)	0.0050(0.0006)	0.0046(0.0008)	0.0049(0.0007)	0.0063(0.0006)	0.0072(0.0009)	0.0084(0.0013)
Ni-O ₃	CN	$9 \times f + 12 \times (1-f)$						
	R (Å)	4.449(0.025)	4.457(0.025)	4.462(0.023)	4.459(0.026)	4.450(0.060)	4.395(0.030)	4.426(0.051)
	σ^2 (Å)	0.0075(0.0027)	0.0072(0.0026)	0.0054(0.0022)	0.0063(0.0026)	0.0146(0.0082)	0.0094(0.0035)	0.0132(0.0065)
Ni-Mn _{2,E}	CN	$6\times(1-f)$						
	R (Å)	4.976(0.039)	4.992(0.044)	4.988(0.046)	4.985(0.040)	5.001(0.028)	4.966(0.023)	4.962(0.042)
	σ^2 (Å)	σ^2 (Ni-Mn _{2,TC})						
Ni-Mn _{2,TC}	CN	$6 \times f$						
	R (Å)	5.472(0.020)	5.437(0.023)	5.448(0.032)	5.475(0.032)	5.545(0.022)	5.493(0.027)	5.457(0.057)
	σ^2 (Å)	0.0060(0.0018)	0.0065(0.0022)	0.0073(0.0030)	0.0068(0.0027)	0.0035(0.0011)	0.0058(0.0017)	0.0092(0.0039)
	f	0.735(0.036)	0.728(0.039)	0.656(0.050)	0.628(0.049)	0.519(0.045)	0.516(0.065)	0.489(0.104)
E0(eV)		-2.5(1.1)	-0.2(1.1)	-2.5(1.2)	0.1(1.2)	-1.8(1.2)	-3.5(1.1)	-3.3(1.5)
Chi Sq		183.54	289.48	256.49	447.79	360.36	249.14	276.41
R factor		0.0229	0.0215	0.0285	0.0291	0.0355	0.0304	0.0409

^a The coordination numbers of these Ni-O shells are 6. ^bf is the fraction of Ni adsorbed on vacancies as triple corner-sharing complexes (TC) of the total Ni in samples.

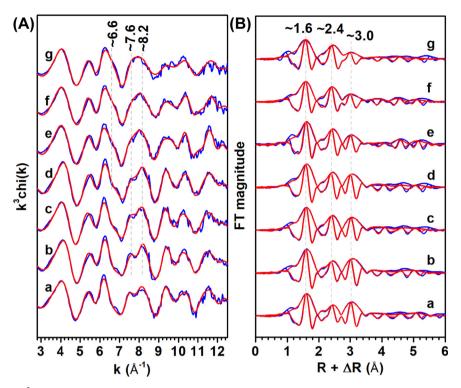


Fig. 8. Nickel K-edge k³-weighted EXAFS functions (A) and corresponding Fourier transforms (FTs, B) of heat-treated 8-year aged Nidoped birnessite (Ni10_8y). Experimental data and their best fits are shown as blue and red lines, respectively. Ni10_8y_01, Ni10_8y_50 (b), Ni10_8y_100 (c), Ni10_8y_150 (d), Ni10_8y_300 (f), and Ni10_8y_400 (g). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

surface energy induced by cation adsorption (Yang et al., 2019). Ni²⁺ adsorption on and affinity for birnessite surfaces are similar to those of Cu²⁺ (Kwon et al., 2013; Manceau and Steinmann, 2021) and decrease of Ni-doped birnessite surface energy likely contributes to its stability upon thermal treatment. Mn(III) content may also play an important role for birnessite thermal transformation as reported in aqueous (Cui et al., 2009; Atkins et al., 2014; Yang et al., 2018) and dry conditions (Grangeon et al., 2014, 2015). Consistently, Ni doping decreases birnessite Mn(III) content (0.06 and 0.10 in Ni10_8y and HB_8y, respectively - Table S1) and thus increases its thermal stability. Contrastingly, metal doping can significantly decrease birnessite particle size (Yin et al., 2012), thus increasing its total free energy, and easing its mineralogical transformation (Luo et al., 2008; Navrotsky et al., 2008). After Ni doping, birnessite crystallinity was greatly reduced. The temperature for the transformation of our Ni-doped birnessite to cryptomelane (500 °C) is lower than those reported for the transformation of well crystallized samples (Chen et al., 1986; Golden et al., 1986; Cai et al., 2001). In addition to its lower Mn(III) content, the increased stability of Ni10_8y compared to HB_8y may be related to the overall decrease in the mineral total free energy induced by Ni²⁺ adsorption on the vacancies and/ or edges. Foreign cation adsorption thus appears to increase birnessite stability and could be a key factor for phyllomanganate stability in natural environments (Yang et al., 2019).

4.3. Changes in Ni binding environments during thermal treatment

Nickel K-edge EXAFS fitting of Ni10 8y samples after heating to 25-400 °C evidences changes in Ni crystal chemistry. Increasing temperature from 25 °C to 200 °C, increases linearly the proportion of edge-sharing Ni-Ni (Mn) pairs ($R^2 = 0.9502$, n = 5, $\alpha = 0.05$; Fig. 10A), suggesting migration of interlayer Ni cations to the octahedral layer. Such rearrangement of cation configuration was previously observed during Ni²⁺ adsorption on hexagonal birnessites when increasing reaction pH (Peacock, 2009; Peña et al., 2010; Sorensen et al., 2020) and confirmed by density functional theory calculations (Peña et al., 2010; Kwon et al., 2013). In the present study, heating can provide the energy to overcome the energy barrier for this transformation, increasing temperature then favoring the transformation. After heating to 200 °C, the proportion of total Ni incorporated in birnessite octahedral layers, increases from 0.26 \pm 0.04 in Ni10 8y to 0.48 \pm 0.05 in Ni10 8y 200 (1-f value - Table 2). Increasing the temperature above 200 °C does not promote further Ni migration into birnessite octahedral layers however, possibly because remaining "interlayer" Ni cations sit on layer edge sites by forming double corner-sharing complexes, which possess a similar Ni-Mn distance of ~ 3.5 Å (Manceau et al., 2007).

Alternatively, the increased proportion of edge sharing Ni-Ni(Mn) pairs upon heating may result from the loss of

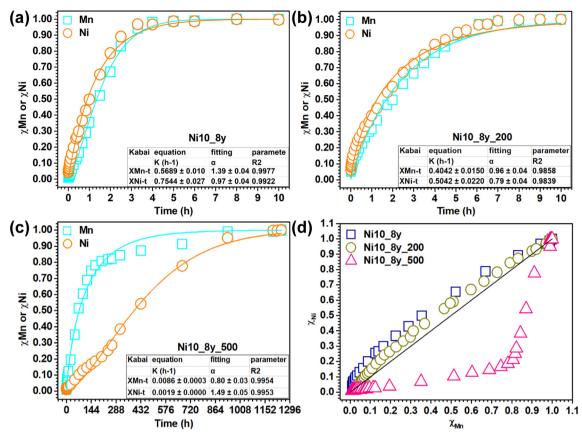


Fig. 9. Releases of Mn and Ni as a function of time along reductive acid dissolution (2 M HCl solution at 37 °C) of Ni10_8y (a), Ni10_8y_200 (b) and Ni10_8y_500 (c), and corresponding %Ni dissolved (χ_{Ni}) versus %Mn dissolved (χ_{Mn}) curves (d). χ_{Mn} and χ_{Ni} are normalized to the final concentration after complete dissolution. In panels a, b and c, the solid lines are the best fit to the dissolution curve using Kabai equation (Kabai, 1973). Fitting parameters are also listed in tables inserted in the panels. In panel d, the solid black line is the 1:1 line.

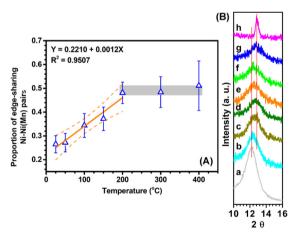


Fig. 10. Evolution of the proportion of edge-sharing Ni-Ni(Mn) pairs as a function of temperature for heat-treated Ni10_8y (A), and shift of the first peak position for heat-treated Ni10_8y (B): Ni10_8y (a), Ni10_8y_50 (b), Ni10_8y_100 (c), Ni10_8y_150 (d), Ni10_8y_200 (e), Ni10_8y_300 (f), Ni10_8y_400 (g), and Ni10_8y_500 (h).

the interlayer water molecules and subsequent rearrangement of adjacent layers. d-spacing of Ni10_8y first peak upon heating steadily decreases from birnessite (001) to cryptome-

lane (110) peak (Fig. 10B). Loss of interlayer water molecules decreases birnessite layer-to-layer distance (Fig. 6b), possibly allowing triple corner-sharing Ni and Mn atoms from adjacent layers to share their coordination O atoms owing to sample turbostratism (systematic random stacking of adjacent layers along the c*-axis) and creating local tunnel-like domains (Fig. 6c) (Grangeon et al., 2014). TGA data obtained for Ni10_8y in air consistently indicates a significant loss of water molecules over the 35-200 °C range (Fig. 2b), rearrangement of adjacent layers being also favored by increasing temperature. Both effects allow more coordination O atoms to be shared by interlayer Ni/Mn from adjacent layers. Departure of interlayer water molecules over the 35-200 °C range is consistent with previous reports indicating the departure of nonstructural water molecules (sorbed to mineral surfaces) at temperatures < 80–100 °C. and of interlayer water molecules from 80 to 100 °C to 205-250 °C (Villalobos et al., 2003; Zhu et al., 2012; Grangeon et al., 2014). Though TGA analysis indicates further loss of hydroxyl groups and interlayer water molecules during heating at 180-520 °C (Fig. 3b and Table S3b), the proportions of edge-sharing Ni-Ni(Mn) pairs are almost stable over this temperature range, the sample keeping its birnessite structure (Fig. 10A and Fig. 1d). Only when the heating temperature is increased to 500 °C, Ni10_8y is totally transformed to cryptomelane (Fig. 1d).

4.4. Mechanisms for enhanced Ni stability during acid dissolution

Along birnessite transformation to tectomanganates such as todorokite and nsutite incorporated foreign metal cations are released to solution to a large extent (Tu et al., 1994; Zhang et al., 2011; Atkins et al., 2014, 2016; Yang et al., 2018, 2019; Wu et al., 2019, 2021). These studies were conducted in aqueous solutions however, and no study is available on the fate of incorporated foreign metal cations during mineral transformation occurring in the dry state. The present study shows that long-term aging in dry state does not significantly modify Ni binding mechanisms to birnessite and thus its availability. On the contrary, thermal treatment of aged Nidoped birnessite significantly reduces Ni availability along reductive acid dissolution of the mineral matrix. Although retain their layer heated solids structure temperatures < 500 °C, Ni²⁺ release rate is gradually reduced with increasing temperature (Fig. 9). After heattreatment to temperatures \geq 500 °C, birnessite is transformed to cryptomelane and Ni²⁺ release rate is decreased by 99.75 \pm 0.01 % compared to the initial aged birnessite. Dissolution of Mn oxides in HCl solutions involves protonation, complexation, and reductive dissolution, with reductive dissolution by Cl being most efficient (Qin et al., 2017). Under similar conditions (H⁺ and Cl⁻ concentrations, temperature, stirring rate), mineral structure, microstructure, and crystal chemistry may further influence dissolution rate. Although poorly crystalline minerals usually dissolve more readily than their highly crystalline counterparts (Liu et al., 2019), Ni²⁺ release rates are reduced for Ni10 8y 200 compared to Ni10 8y despite powder XRD and SEM analysis being indicative of Ni10 8y higher crystallinity and larger particle size. Ni binding to the birnessite structure is thus likely responsible for the lower dissolution rates of Ni10_8y_200. At temperatures < 400 °C, increased proportion of edgesharing Ni-Ni(Mn) pairs (Fig. 10A), that may result both from Ni migration from interlayer to the layer and from interlayer Ni cations sharing their coordination O atoms with Ni(Mn) from adjacent layers, most likely accounts for this reduced availability. Increasing the temperature further to 500 °C induces the transformation of Ni10_8y into cryptomelane, that exhibits larger particle size and improved crystallinity compared to its Ni10 8y birnessite precursor (Figs. 1 and 3). Furthermore, Ni is thoroughly inserted in cryptomelane framework, as evidenced by HAADF-STEM analysis (Fig. 3i), thus providing a dramatically higher stability to this foreign cation. Though not directly relevant to acid leaching or weathering in natural environments, results from acid dissolution experiments might shed light on similar but much less harsh leaching processes in the environment.

5. CONCLUSIONS

Crystallinity, morphology, Mn AOS and Ni binding mechanisms are essentially unaffected when aging Ni-doped hexagonal birnessite in the dry state at room

temperature for up to 8 years. When heated, the proportion of edge-sharing Ni-Ni(Mn) pairs in aged birenssites increases with temperature increasing from 25 to 200 °C before reaching a steady state when increasing the temperature further from 200 to 400 °C. Over this whole temperature range, birnessite layer structure is retained and Mn AOS remains constant. At higher temperatures (400-500 °C), Ni-doped birnessites transform to cryptomelane, Ni doping increasing slightly birnessite stability and delaying the mineral transformation. At 500 °C, Ni-doped birnessite with an initial Ni/Mn molar ratio of 0.13 fully transforms to cryptomelane, with a minor NiMnO₃ impurity. Along this mineral transformation, Mn AOS is sharply increased, Mn cations being essentially Mn⁴⁺ in the cryptomelane framework, and corner-sharing Mn-Mn (Ni) bond length is decreased. The increasing proportion of edge-sharing Ni-Ni(Mn) pairs in heated birnessite decreases Ni release to solution during reductive acid dissolution, whereas incorporation of Ni in the cryptomelane framework essentially inhibits this release. These results provide additional insights in the stability of birnessitelike minerals under dry conditions that will most likely become more frequent owing to climate change, and in the geochemical behavior of associated trace metal elements.

DECLARATION OF COMPETING INTEREST

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

ACKNOWLEDGEMENTS

The authors gratefully thank the Natural Science Foundations of China (Grants 42077015 and 41771267), Key Science and Technology Projects of Inner Mongolia autonomous region (2019ZD001), and the Fundamental Research Funds for the Central Universities (Grant 103-510320036) for financial support.

APPENDIX A. SUPPLEMENTARY MATERIAL

Supplementary material to this article can be found online at https://doi.org/10.1016/j.gca.2022.07.014.

REFERENCES

Alvarez M., Rueda E. H. and Sileo E. E. (2006) Structural characterization and chemical reactivity of synthetic Mngoethites and hematites. *Chem. Geol.* 231, 288–299.

Alvarez M., Rueda E. H. and Sileo E. E. (2007) Simultaneous incorporation of Mn and Al in the goethite structure. *Geochim. Cosmochim. Acta* 71, 1009–1020.

Atkins A. L., Shaw S. and Peacock C. L. (2014) Nucleation and growth of todorokite from birnessite: Implications for tracemetal cycling in marine sediments. *Geochim. Cosmochim. Acta* 144, 109–125.

Atkins A. L., Shaw S. and Peacock C. L. (2016) Release of Ni from birnessite during transformation of birnessite to todorokite: Implications for Ni cycling in marine sediments. *Geochim. Cosmochim. Acta* 189, 158–183.

- Beak D. G., Basta N. T., Scheckel K. G. and Traina S. J. (2008) Linking solid phase speciation of Pb sequestered to birnessite to oral Pb bioaccessibility: implications for soil remediation. *Environ. Sci. Technol.* 42, 779–785.
- Birkner N. and Navrotsky A. (2017) Thermodynamics of manganese oxides: Sodium, potassium, and calcium birnessite and cryptomelane. *Proc. Natl. Acad. Sci. USA* 114, E1046–E1053.
- Cai J., Liu J., Willis W. S. and Suib S. L. (2001) Framework doping of iron in tunnel structure cryptomelane. *Chem. Mater.* 13, 2413–2422.
- Chen C. C., Golden D. C. and Dixon J. B. (1986) Transformation of synthetic birnessite to cryptomelane: An electron microscopic study. Clays Clay Miner. 34, 565–571.
- Ching S., Driscoll P. F., Kieltyka K. S., Marvel M. R. and Suib S. L. (2001) Synthesis of a new hollandite-type manganese oxide with framework and interstitial Cr(III). Chem. Commun., 2486–2487.
- Cui H., Qiu G., Feng X., Tan W. and Liu F. (2009) Birnessites with different average manganese oxidation states synthesized, characterized, and transformed to todorokite at atmospheric pressure. Clays Clay Miner. 57, 715–724.
- Drits V. A., Silvester E., Gorshkov A. I. and Manceau A. (1997) Structure of synthetic monoclinic Na-rich birnessite and hexagonal birnessite: I. Results from X-ray diffraction and selected-area electron diffraction. Am. Mineral. 82, 946–961.
- Drits V. A., Lanson B. and Gaillot A. C. (2007) Birnessite polytype systematics and identification by powder X-ray diffraction. *Am. Mineral.* 92, 771–788.
- Giovanoli R. (1980). Vernadite is random-stacked birnessite. Mineral. Deposita 15, 251–253.
- Golden D., Dixon J. and Chen C. (1986) Ion exchange, thermal transformations, and oxidizing properties of birnessite. *Clays Clay Miner.* **34**, 511–520.
- Grangeon S., Lanson B., Lanson M. and Manceau A. (2008) Crystal structure of Ni-sorbed synthetic vernadite: a powder X-ray diffraction study. *Mineral. Mag.* 72, 1279–1291.
- Grangeon S., Lanson B., Miyata N., Tani Y. and Manceau A. (2010) Structure of nanocrystalline phyllomanganates produced by freshwater fungi. *Am. Mineral.* **95**, 1608–1616.
- Grangeon S., Lanson B. and Lanson M. (2014) Solid-state transformation of nanocrystalline phyllomanganate into tectomanganate: influence of initial layer and interlayer structure. *Acta. Crystallogr. B* 70, 828–838.
- Grangeon S., Fernandez-Martinez A., Warmont F., Gloter A., Marty N., Poulain A. and Lanson B. (2015) Cryptomelane formation from nanocrystalline vernadite precursor: a high energy X-ray scattering and transmission electron microscopy perspective on reaction mechanisms. *Geochem. Trans.* 16, 12.
- Johnston S. G., Burton E. D. and Moon E. M. (2016) Arsenic Mobilization Is Enhanced by Thermal Transformation of Schwertmannite. *Environ. Sci. Technol.* 50, 8010–8019.
- Jothiramalingam R., Viswanathan B. and Varadarajan T. K. (2006) Synthesis and structural characterization of copper incorporated manganese oxide OMS-2 materials synthesized via potassium birnessite. *Mater. Chem. Phys.* 100, 257–261.
- Kabai J. (1973) Determination of specific activation energies of metal oxides and metal oxide hydrates by measurement of the rate of dissolution. Acta Chim. Acad. Sci. Hung 78, 57–73.
- Kaur N., Singh B. and Kennedy B. J. (2010) Dissolution of Cr, Zn, Cd, and Pb single- and multi-metal-substituted goethite: relationship to structural, morphological, and dehydroxylation properties. Clays Clay Miner. 58, 415–430.
- Kelly S. D., Hesterberg D. and Ravel B. (2008) Analysis of soils and minerals using X-ray absorption spectroscopy. In *Methods* of Soil Analysis, Part 5-Mineralogical Methods (eds. A. L. Ulrey and R. L. Drees). Soil Science Society of America, Wisconsin, pp. 387–463.

- Kim K., Yoon H. I. and Choi W. (2012) Enhanced dissolution of manganese oxide in ice compared to aqueous phase under illuminated and dark conditions. *Environ. Sci. Technol.* 46, 13160–13166.
- Kwon K. D., Refson K. and Sposito G. (2009) On the role of Mn (IV) vacancies in the photoreductive dissolution of hexagonal birnessite. *Geochim. Cosmochim. Acta* **73**, 4142–4150.
- Kwon K. D., Refson K. and Sposito G. (2013) Understanding the trends in transition metal sorption by vacancy sites in birnessite. *Geochim. Cosmochim. Acta* **101**, 222–232.
- Landers M., Gilkes R. J. and Wells M. A. (2009) Rapid dehydroxylation of nickeliferous goethite in lateritic nickel ore: X-ray diffraction and TEM investigation. *Clays Clay Miner.* 57, 751–770.
- Lanson B., Marcus M. A., Fakra S., Panfili F., Geoffroy N. and Manceau A. (2008) Formation of Zn-Ca phyllomanganate nanoparticles in grass roots. *Geochim. Cosmochim. Acta* 72, 2478–2490.
- Lee S. H., Kim T. W., Park D. H., Choy J.-H., Hwang S.-J., Jiang N., Park S.-E. and Lee Y.-H. (2007) Single-step synthesis, characterization, and application of nanostructured K_x $Mn_{1-y}Co_yO_{2-\delta}$ with controllable chemical compositions and crystal structures. *Chem. Mater.* **19**, 5010–5017.
- Lee S., Xu H., Xu W. and Sun X. (2019) The structure and crystal chemistry of vernadite in ferromanganese crusts. *Acta. Crystallogr. B* **75**, 591–598.
- Li M., Liu H., Chen T., Wei L., Wang C., Hu W. and Wang H. (2019) The transformation of α-(Al, Fe)OOH in natural fire: Effect of Al substitution amount on fixation of phosphate. *Chem. Geol.* 524, 368–382.
- Liu Y., Li Y., Li Y., Chen N., Ding H., Jin S., Wang C. and Lu A. (2021) Photo-stimulated anoxic reduction of birnessite (δ-MnO₂) by citrate and its fine structural responses: Insights on a proton-promoted photoelectron transfer process. *Chem. Geol.* **561** 120029.
- Liu L., Wang X., Zhu M., Ma J., Zhang J., Tan W., Feng X., Yin H. and Liu F. (2019) The speciation of Cd in Cd-Fe coprecipitates: Does Cd substitute for Fe in goethite structure? ACS Earth Space Chem. 3, 2225–2236.
- Lu A., Li Y., Ding H., Xu X., Li Y., Ren G., Liang J., Liu Y., Hong H., Chen N., Chu S., Liu F., Li Y., Wang H., Ding C., Wang C., Lai Y., Liu J., Dick J., Liu K. and Hochella, Jr., M. F. (2019) Photoelectric conversion on Earth's surface via widespread Fe- and Mn-mineral coatings. *Proc. Natl. Acad.* Sci. USA 116, 9741–9746.
- Lu A. and Li Y. (2015) Reactivity of natural Mn oxide cryptomelane. In *Advances in the Environmental Biogeochemistry of Manganese Oxides* (eds. X. Feng, W. Li, M. Zhu and D. L. Sparks). American Chemical Society, Washington, pp. 89–106.
- Luo W., Hu W. and Xiao S. (2008) Size effect on the thermodynamic properties of silver nanoparticles. J. Phys. Chem. C 112, 2359–2369.
- Manceau A., Marcus M. A. and Grangeon S. (2012) Determination of Mn valence states in mixed-valent manganates by XANES spectroscopy. *Am. Mineral.* **97**, 816–827.
- Manceau A. and Steinmann S. N. (2021) Nature of high- and low-affinity metal surface sites on birnessite nanosheets. *ACS Earth Space Chem.* **5**, 66–76.
- Manceau A., Lanson M. and Geoffroy N. (2007) Natural speciation of Ni, Zn, Ba, and As in ferromanganese coatings on quartz using X-ray fluorescence, absorption, and diffraction. *Geochim. Cosmochim. Acta* 71, 95–128.
- Manceau A., Lanson M. and Takahashi Y. (2014) Mineralogy and crystal chemistry of Mn, Fe Co, Ni, and Cu in a deepsea Pacific polymetallic nodule. *Am. Mineral.* **99**, 2068–2083.
- McKenzie R. M. (1971) The synthesis of birnessite, cryptomelane, and some other oxides and hydroxides of manganese. *Mineral. Mag.* **38**, 493–503.

- Morgan J. J. (2005) Kinetics of reaction between O2 and Mn(II) species in aqueous solutions. *Geochim. Cosmochim. Acta* **69**, 35–48.
- Murphy B. F. and Timbal B. (2008) A review of recent climate variability and climate change in southeastern Australia. *Int. J. Climatol.* **28**, 859–879.
- Navrotsky A., Mazeina L. and Majzlan J. (2008) Size-Driven Structural and Thermodynamic Complexity in Iron Oxides. *Science* **319**, 1635–1638.
- Peacock C. L. (2009) Physiochemical controls on the crystalchemistry of Ni in birnessite: genetic implications for ferromanganese precipitates. *Geochim. Cosmochim. Acta* 73, 3568–3578.
- Peña J., Kwon K. D., Refson K., Bargar J. R. and Sposito G. (2010) Mechanisms of nickel sorption by a bacteriogenic birnessite. *Geochim. Cosmochim. Acta* **74**, 3076–3089.
- Perrier N., Gilkes R. J. and Colin F. (2006) Heating Fe oxide-rich soils increases the dissolution rate of metals. *Clays Clay Miner.* **54**, 165–175.
- Post J. E. (1999) Manganese oxide minerals: Crystal structures and economic and environmental significance. *Proc. Natl. Acad. Sci.* USA 96, 3447–3454.
- Post J. E. and Appleman D. E. (1988) Chalcophanite, ZnMn3O7.3-H2O: New crystal-structure determinations. *Am. Mineral.* 73, 1401–1404.
- Qin Z., Xiang Q., Liu F., Xiong J., Koopal L. K., Zheng L., Ginder-Vogel M., Wang M., Feng X., Tan W. and Yin H. (2017) Local structure of Cu²⁺ in Cu-doped hexagonal turbostratic birnessite and Cu²⁺ stability under acid treatment. *Chem. Geol.* 466, 512–523.
- Ravel B. and Newville M. (2005) ATHENA, ARTEMIS, HEPHAESTUS: data analysis for X-ray absorption spectroscopy using IFEFFIT. *J. Synchrotron Radiat.* 12, 537–541.
- Rehr J. J., Zabinsky S. I. and Albers R. C. (1992) High-order multiple scattering calculations of X-ray-absorption fine structure. *Phys. Rev. Lett.* 69, 3397–3400.
- Rein G., Cleaver N., Ashton C., Pironi P. and Torero J. L. (2008) The severity of smouldering peat fires and damage to the forest soil. *Catena* **74**, 304–309.
- Sherman D. M. (2005) Electronic structures of iron(III) and manganese(IV) (hydr)oxide minerals: Thermodynamics of photochemical reductive dissolution in aquatic environments. *Geochim. Cosmochim. Acta* **69**, 3249–3255.
- Silvester E. J., Manceau A. and Drits V. A. (1997) The structure of monoclinic Na-rich birnessite and hexagonal birnessite. Part 2. Results from chemical studies and EXAFS spectroscopy. *Am. Mineral.* 82, 962–978.
- Simanova A. A. and Pena J. (2015) Time-resolved investigation of cobalt oxidation by Mn(III)-rich delta-MnO₂ using quick X-ray absorption spectroscopy. *Environ. Sci. Technol.* **49**, 10867–10876.
- Sinha A. K., Pradhan M. and Pal T. (2013) Morphological evolution of two-dimensional MnO₂ nanosheets and their shape transformation to one-dimensional ultralong MnO₂ nanowires for robust catalytic activity. *J. Phys. Chem. C* 117, 23976– 23986.
- Sorensen J. V., Gueguen B., Stewart B. D., Peña J., Rouxel O. and Toner B. M. (2020) Large nickel isotope fractionation caused by surface complexation reactions with hexagonal birnessite. *Chem. Geol.* 537 119481.
- Sunda W. G. and Huntsman S. A. (1994) Photoreduction of manganese oxides in seawater. *Mar. Chem.* **46**, 133–152.
- Tang Q., Jiang L., Liu J., Wang S. and Sun G. (2014) Effect of surface manganese valence of manganese oxides on the activity of the oxygen reduction reaction in alkaline media. ACS Catal. 4, 457–463.
- Taylor R., McKenzie R. and Norrish K. (1964) The mineralogy and chemistry of manganese in some Australian soils. *Soil Res.* 2, 235–248.

- Tu S., Racz G. J. and Goh T. B. (1994) Transformations of Synthetic Birnessite as Affected by pH and Manganese Concentration. Clays Clay Miner. 42, 321–330.
- Vicente-Serrano S. M., Quiring S. M., Peña-Gallardo M., Yuan S. and Domínguez-Castro F. (2020) A review of environmental droughts: Increased risk under global warming? *Earth-Sci. Rev.* 201 102953.
- Villalobos M., Toner B., Bargar J. and Sposito G. (2003) Characterization of the manganese oxide produced by pseudomonas putida strain MnB1. Geochim. Cosmochim. Acta 67, 2649–2662.
- Wang Y., Benkaddour S., Marafatto F. F. and Pena J. (2018) Diffusion- and pH-dependent reactivity of layer-type MnO₂: Reactions at particle edges versus vacancy sites. *Environ. Sci. Technol.* 52, 3476–3485.
- Webb S. M. (2005) SIXPACK: a graphical user interface for XAS analysis using IFEFFIT. *Physica Scripta* T115, 1011–1014.
- Wu Z., Peacock C. L., Lanson B., Yin H., Zheng L., Chen Z., Tan W., Qiu G., Liu F. and Feng X. (2019) Transformation of Cocontaining birnessite to todorokite: Effect of Co on the transformation and implications for Co mobility. *Geochim. Cosmochim. Acta* 246, 21–40.
- Wu Z., Lanson B., Feng X., Yin H., Tan W., He F. and Liu F. (2021) Transformation of the phyllomanganate vernadite to tectomanganates with small tunnel sizes: Favorable geochemical conditions and fate of associated Co. *Geochim. Cosmochim.* Acta 295, 224–236.
- Yang P., Lee S., Post J. E., Xu H., Wang Q., Xu W. and Zhu M. (2018) Trivalent manganese on vacancies triggers rapid transformation of layered to tunneled manganese oxides (TMOs): Implications for occurrence of TMOs in low-temperature environment. *Geochim. Cosmochim. Acta* 240, 173–190.
- Yang P., Post J. E., Wang Q., Xu W., Geiss R., McCurdy P. R. and Zhu M. (2019) Metal adsorption controls stability of layered manganese oxides. *Environ. Sci. Technol.* 53, 7453–7462.
- Yin H., Tan W. F., Zheng L. R., Cui H. J., Qiu G. H., Liu F. and Feng X. H. (2012) Characterization of Ni-rich hexagonal birnessite and its geochemical effects on aqueous Pb²⁺/Zn²⁺ and As(III). *Geochim. Cosmochim. Acta* **93**, 47–62.
- Yin H., Li H., Wang Y., Ginder-Vogel M., Qiu G. H., Feng X. H., Zheng L. R. and Liu F. (2014) Effects of Co and Ni co-doping on the structure and reactivity of hexagonal birnessite. *Chem. Geol.* 381, 10–20.
- Yin H., Dai X. X., Zhu M. Q., Li F. H., Feng X. H. and Liu F. (2015) Fe-doped cryptomelane synthesized by refluxing at atmosphere: Structure, properties and photocatalytic degradation of phenol. J. Hazard. Mater. 296, 221–229.
- Zhang Q., Xiao Z., Feng X., Tan W., Qiu G. and Liu F. (2011) α-MnO₂ nanowires transformed from precursor δ-MnO₂ by refluxing under ambient pressure: The key role of pH and growth mechanism. *Mater. Chem. Phys.* **125**, 678–685.
- Zhang Q., Cheng X. D., Qiu G. H., Liu F. and Feng X. H. (2016) Size-controlled synthesis and formation mechanism of manganese oxide OMS-2 nanowires under reflux conditions with KMnO₄ and inorganic acids. *Solid State Sci.* **55**, 152–158.
- Zhu M., Farrow C. L., Post J. E., Livi K. J. T., Billinge S. J. L., Ginder-Vogel M. and Sparks D. L. (2012) Structural study of biotic and abiotic poorly-crystalline manganese oxides using atomic pair distribution function analysis. *Geochim. Cos-mochim. Acta* 81, 39–55.
- Zhu Y., Luo Q. and Liu Y. (2022) Deserts and Desertification. IntechOpen, London.

Associate editor: Colleen M Hansel